

1 Stochastic surrogate model for meteotsunami early warning system in the 2 eastern Adriatic Sea

Cléa Denamiel^{1*}, Jadranka Šepić¹, Xun Huan², Célia Bolzer^{3,4}, Ivica Vilibić¹

¹ Institute of Oceanography and Fisheries, Šetalište I. Meštrovića 63, 21000 Split, Croatia

² University of Michigan, 2033 AL, 1231 Beal Ave, Ann Arbor, MI 48109-2133, U.S.A.

³ SeaTech, Ecole d'ingénieurs - Université de Toulon, CS 60584, 83041 Toulon Cedex 9, France

⁴ EniProgetti Eni House, Basing View, Basingstoke, Hampshire, RG21 4YY, U.K.

*Corresponding author: Cléa Denamiel (cdenamie@izor.hr)

5 Key Points

- 6 • Design and evaluation of an innovative meteotsunami early warning system prototype
7 using stochastic surrogate approach
- 8 • Forecast of the atmospheric internal gravity waves driving meteotsunami events with
9 deterministic state-of-the-art models
- 10 • Stochastic surrogate model based on generalized polynomial chaos expansion methods
11 and running at nearly no computational cost

12 Abstract

13 The meteotsunami early warning system prototype using stochastic surrogate approach and
14 running operationally in the eastern Adriatic Sea is presented. First, the atmospheric Internal
15 Gravity Waves (IGWs) driving the meteotsunamis are either forecasted with state-of-the-art
16 deterministic models at least a day in advance or detected through measurements at least 2-h
17 before the meteotsunami reaches sensitive locations. The extreme sea-level hazard forecast at
18 endangered locations is then derived with an innovative stochastic surrogate model –
19 implemented with generalized Polynomial Chaos Expansion (gPCE) method and synthetic IGWs

20 forcing a barotropic ocean model – used with the input parameters extracted from deterministic
21 model results and/or measurements. The evaluation of the system, both against five historical
22 events and for all the detected potential meteotsunamis since late 2018 when the early warning
23 system prototype became operational, reveals that the meteotsunami hazard is conservatively
24 assessed but often overestimated at some locations. Despite some needed improvements and
25 developments, this study demonstrates that gPCE-based methods can be used for
26 atmospherically-driven extreme sea-level hazard assessment, and in geosciences in wide.

27 **Plain Language Summary**

28 Atmospherically-driven extreme sea-level events are one of the major threats to people
29 and assets in the coastal regions. Assessing the hazard associated with such events together with
30 uncertainty quantification in a precise and timely manner is thus of primary importance in
31 modern societies. In this study, an early warning system for the eastern Adriatic meteotsunamis –
32 destructive long waves with periods from few minutes up to an hour generated by traveling
33 atmospheric disturbances, is presented and evaluated. The system is based on state-of-the-art
34 deterministic atmospheric and ocean models as well as an innovative statistical model developed
35 to forecast the meteotsunami hazard. The evaluation reveals that the meteotsunami hazard is
36 conservatively assessed but often overestimated. This study demonstrates that the presented
37 methodology can be used for extreme sea-level hazard assessment and in general for hazard
38 studies in geosciences.

39 **Key Words**

40 Meteotsunami early warning system, extreme sea-level hazard assessment, eastern Adriatic

41 **1 Introduction**

42 During the past decade, meteorological tsunamis or meteotsunamis – destructive long
43 waves in the tsunami frequency band generated by traveling atmospheric disturbances
44 (Monserrat et al., 2006), have become the object of an increasing number of studies all over the
45 globe (Tanaka, 2010; Šepić et al., 2012; Cho et al., 2013; Okal et al., 2014; Pattiaratchi
46 & Wijeratne, 2014; Pellikka et al., 2014; Whitmore & White, 2014; Olabarrieta et al., 2017,
47 Masina et al., 2017; Dusek et al., 2019). These extreme events have the potential to produce
48 substantial damages to houses, goods and infrastructures (Hibiya & Kajiura, 1982; Salaree et al.,
49 2018; Linares et al., 2019) – e.g. more than seven million US dollar losses in Vela Luka harbor,
50 Croatia during the 21st of June 1978 meteotsunami (Vučetić et al., 2009; Orlić et al., 2010), but
51 also to claim human lives – e.g. seven people killed during a sunny day in 1954 (Ewing et al.,
52 1954) in the Great Lakes near Chicago, USA. Rather than addressing a particular catastrophic
53 event, this work focuses on the design and evaluation of an innovative meteotsunami early
54 warning system tested in operational mode, since late 2018, in the eastern Adriatic. As fully
55 preventing meteotsunami impact is, for now, close to impossible (Vilibić et al., 2016), the
56 principal goal of such a system is to allow the local communities to better prepare for these
57 destructive events (e.g. set temporary protection against flooding and waves, avoid swimming,
58 etc.) in order to minimize the losses. However, deterministically forecasting the atmospheric
59 disturbances responsible for meteotsunamis is challenging (Renault et al., 2011; Denamiel et al.,
60 2019) and the uncertainties in anticipating their location and intensity as well as their relationship
61 to flood in sensitive harbor locations must be taken into account. In addition, as meteotsunamis
62 are rare events which require specific model setup – e.g. for the ocean, a 1-min atmospheric
63 forcing and a resolution below 50m in the harbors where resonance occurs, the available forecast

64 results are generally not designed to capture them (Denamiel et al., 2019). For the Adriatic Sea, a
65 specific numerical suite was thus implemented to deterministically forecast the atmospheric
66 disturbances – e.g. the Internal Gravity Waves (IGWs; Vilibić & Šepić, 2009; Denamiel et al.,
67 2019), driving the meteotsunamis along the Croatian coastline.

68 In order to quantify the uncertainties linked to the meteotsunami extreme sea-levels, the
69 origin, propagation and sources of uncertainty of the complex ocean-atmosphere system must be
70 described (Arnst & Ponthot, 2014; Ghanem et al., 2017; Bulthuis et al., 2019). In the Adriatic
71 Sea, the location, speed, period, amplitude and direction of the forecasted atmospheric
72 disturbances are the primary sources of uncertainties linked to the meteotsunami events and can
73 thus be seen as random variables characterized by their prior distributions. In the field of
74 uncertainty quantification (Le Maître & Knio, 2010; Ghanem et al., 2017), generalized
75 Polynomial Chaos Expansion (gPCE) methods (Xiu & Karniadakis, 2002; Soize & Ghanem,
76 2004) have been widely used to build surrogate models that propagate, at nearly no
77 computational cost, the uncertainties of a given stochastic forcing to the results of a deterministic
78 model. Furthermore, in the past decade, gPCE methods have been applied with success in
79 geosciences: Formaggia et al. (2013) built a surrogate model of basin-scale geochemical
80 compaction, Wang et al. (2016) studied the acoustic uncertainty predictions, Sraj et al. (2014)
81 estimated the wind drag parameter forcing an ocean model, Giraldi et al. (2017) documented the
82 propagation of earthquake ocean floor displacement uncertainty to the tsunami wave parameters
83 and Bulthuis et al. (2019) used a surrogate model to quantify the uncertainty of the multi-
84 centennial response of the Antarctic ice sheet to climate change. Following the footsteps of these
85 recent studies, the newly developed meteotsunami surrogate model was thus designed to
86 propagate the known uncertainties of the atmospheric disturbances to the forecast of extreme

87 sea-levels at five sensitive locations along the Croatian coastline: Vela Luka, Vrboska, Stari
88 Grad, Rijeka dubrovačka and Ston (Fig. 1).

89 In this paper, the setup of the Croatian early warning system prototype, which provides
90 meteotsunami hazard assessments depending on the deterministically forecasted and measured
91 atmospheric pressure waves and the stochastically deduced maximum elevation distributions
92 derived with the surrogate model, is first described in details in Section 2. In section 3, its
93 evaluation for five different locations along the Croatian coastline is performed first, against five
94 different historical events, and then for automatically detected events since the system became
95 operational in late 2018. Finally, the methodological choices made to design this first
96 meteotsunami early warning system as well as its performance and the improvements needed to
97 increase its reliability are discussed in Section 4.

98 **2 Design of the meteotsunami early warning system**

99 **2.1 Data and models**

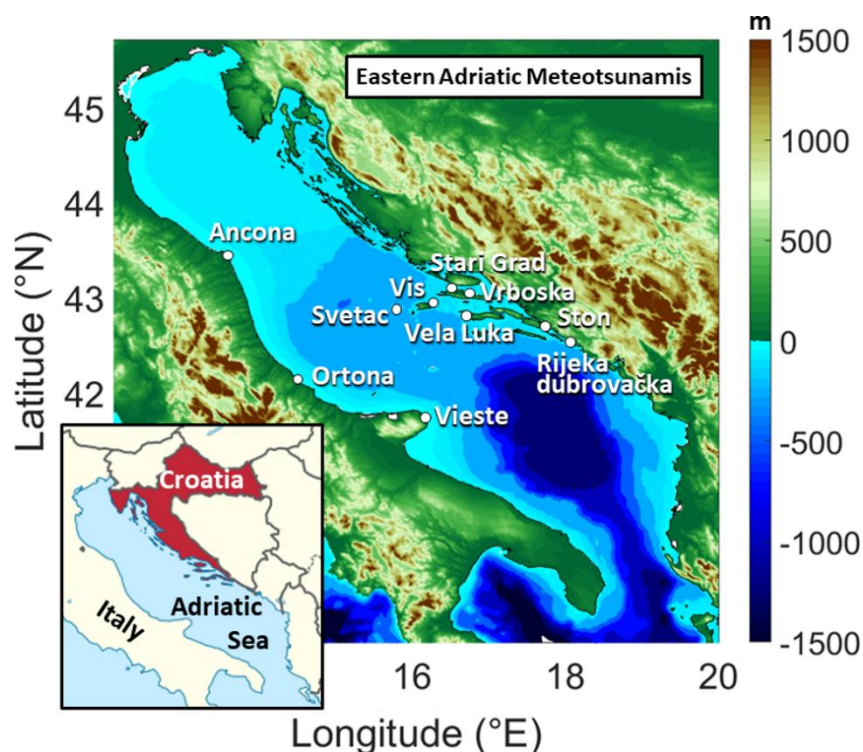
100 The Croatian Meteotsunami Early Warning System (CMeEWS, Šepić et al., 2017) –
101 developed within the framework of the project MESSI (“Meteotsunamis, destructive long ocean
102 waves in the tsunami frequency band: from observations and simulations towards a warning
103 system”; <http://www.izor.hr/messi>), receives three different kind of data: (1) synoptic conditions
104 from the Croatian Meteorological and Hydrological Service (DHMZ) operational atmospheric
105 products, (2) high-resolution atmospheric and ocean model results provided by the Adriatic Sea
106 and Coast (AdriSC) modelling suite (Denamiel et al., 2019), and (3) measurements from the
107 MESSI observational network along the Adriatic coast. The synoptic data are used for a long-
108 term qualitative forecast (at least a week) of meteotsunamigenic conditions through assessment
109 of the synoptic meteotsunami index (Šepić et al., 2016). However, such an approach cannot be

110 used in quantitative meteotsunami hazard assessment and forecast, and is not further discussed in
111 this paper.

112 The AdriSC modelling suite is composed of a basic module providing high-resolution
113 regional atmospheric and ocean results for the entire Adriatic Sea and a dedicated meteotsunami
114 module. The basic module uses a modified version of the Coupled Ocean-Atmosphere-Wave-
115 Sediment-Transport (COAWST) modelling system developed by Warner et al. (2010), which
116 couples (online) (1) the Regional Ocean Modeling System (ROMS) (Shchepetkin &
117 McWilliams, 2005, 2009), with nested grids of 3-km (covering the entire Adriatic and Ionian
118 Seas) and 1-km (covering the Adriatic Sea only), and (2) the Weather Research and Forecasting
119 (WRF) model (Skamarock et al., 2005), with nested grids of 15-km (covering the central
120 Mediterranean basin) and 3-km (identical to the 3-km ROMS grid). The dedicated meteotsunami
121 module couples (offline) the WRF model – which downscales the hourly 3-km WRF results of
122 the basic module to a 1.5-km resolution for a grid covering the entire Adriatic Sea, with the
123 2DDI ADvanced CIRCulation (ADCIRC) model (Luettich et al., 1991) using a mesh of up to 10
124 m resolution in the areas sensitive to meteotsunami hazard. In this deterministic configuration,
125 the ADCIRC model is forced (1) every minute by the WRF 1.5-km wind and pressure fields, and
126 (2) every hour by the ROMS 1-km sea-level fields (including tides). Every day at midnight, the
127 next 48h hourly-forecast results from the COAWST run, as well as the 15min-forecast results
128 from WRF 1.5-km and ADCIRC simulations for the next day, are published at
129 <http://www.izor.hr/adriSC>.

130 The MESSI observational system currently encompasses a network of sensors set-up with
131 a 1-min sampling rate and installed in areas where either the generation or the amplification of
132 meteotsunamis are known to occur: eight air pressure sensors located in (1) Ancona, Ortona and

133 Vieste on the Italian coast, up to 200 km from any endangered location along the Croatian
 134 coastline, (2) Vis and Svetac in the middle of the Adriatic Sea and (3) Vela Luka, Stari Grad and
 135 Vrboska which are known sensitive harbors (Fig. 1) as well as two tide gauges located in the
 136 harbors of Vela Luka and Stari Grad (Fig. 1).



137
 138 Figure 1. Locations of interest including measurement network along the Italian coast and in the
 139 middle of the Adriatic Sea (Ancona, Ortona, Vieste, Svetac and Vis) and sensitive harbor
 140 locations along the Croatian coast (Vela Luka, Stari Grad, Vrboska, Ston and Rijeka
 141 dubrovačka).

142 Within the CMeEWS, the extreme sea-level hazard assessment relies on the newly
 143 developed meteotsunami stochastic surrogate model. This model is based on generalized
 144 Polynomial Chaos Expansion (gPCE) methods (Xiu & Karniadakis, 2002; Soize & Ghanem,
 145 2004) which, compared to sampling approaches (e.g. Monte Carlo simulations), are highly
 146 efficient for propagating the uncertainties of model inputs to outputs (e.g., Knio & Le Maître,

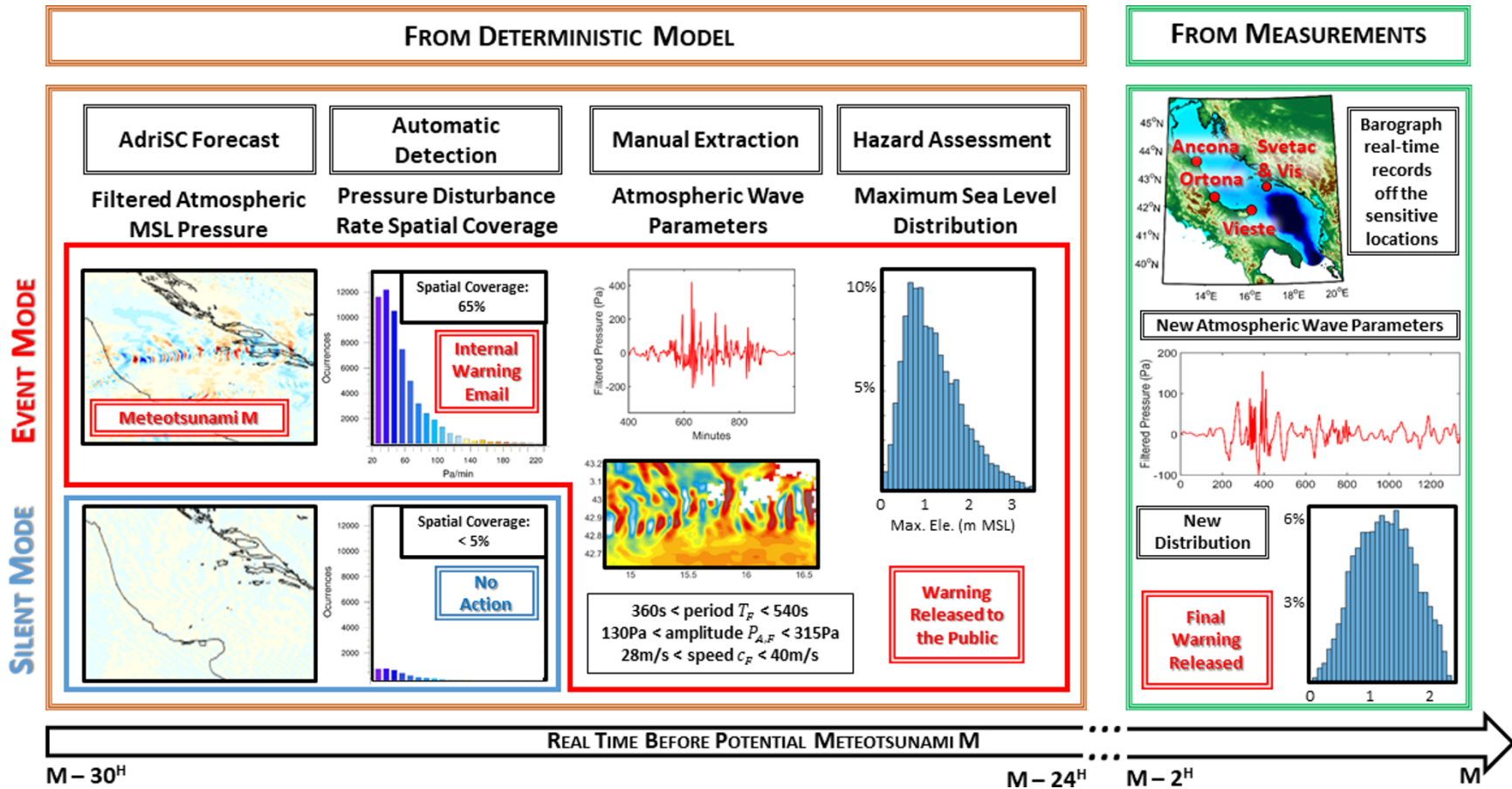
147 2006 and Najm et al., 2009 provide detailed discussions in the context of computational fluids
 148 applications). In this study, the stochastic surrogate model propagates the uncertainties from the
 149 meteorological input (i.e. the IGWs responsible for the meteotsunami generation) to the
 150 maximum sea-levels at different locations along the Croatian coastline. The surrogate model is
 151 based on polynomials expansions that decompose into deterministic coefficients and random
 152 orthogonal bases. The coefficients – which are the projection of the maximum meteotsunami
 153 elevation distribution onto each polynomial basis, are derived from a quadrature based
 154 approximation using numerical simulations undertaken with the ADCIRC model (identical to the
 155 one used in the AdriSC modelling suite) forced only by synthetic pressure disturbances (no wind,
 156 no tide). As described in Denamiel et al. (2018), the synthetic atmospheric pressure forcing is
 157 split into (1) a mean atmospheric pressure component (P_0) assumed constant over the entire
 158 Adriatic Sea and (2) a stochastic gravity wave component (P_{GW}) depending on 6 stochastic
 159 parameters – start location (y_0), direction (θ), speed (c), period (T), amplitude (P_A) and
 160 width (d) of the disturbance. These 6 parameters are assumed to have uniform distributions and
 161 are defined on the following intervals: $y_0 \in [41.25^\circ, 43.65^\circ]$, $\theta \in \left[-\frac{\pi}{3}, \frac{\pi}{2}\right]$,
 162 $c \in [15\text{m s}^{-1}, 40\text{m s}^{-1}]$, $T \in [300\text{s}, 1800\text{s}]$, $P_A \in [50\text{Pa}, 400\text{Pa}]$ and $d \in [30\text{km}, 150\text{km}]$. Examples
 163 of synthetic gravity wave spatial and temporal properties can be visualized as supporting
 164 information (Fig. S1). Practically, as the input parameters are assumed to be uniformly
 165 distributed, (1) the delayed Gauss-Patterson sparse grid method (Smolyak, 1963; Novak et al.,
 166 1999; Burkardt, 2014) is applied to automatically select all the combined values of the 6
 167 stochastic parameters of the synthetic pressure forcing and thus to define the number of
 168 simulations (in this study 4161 as the gPCE is defined for polynomial degrees up to 6) used to

169 derive the polynomial coefficients, while (2) the random orthogonal bases are built with
 170 Legendre polynomials. The meteotsunami hazard forecast is illustrated in Figure 2 and is based
 171 on the meteotsunami stochastic surrogate model receiving atmospheric pressure field input from
 172 both (1) the WRF 1.5-km next day forecast results (brown box, Fig. 2) and (2) the real-time
 173 transmitted observations from Ancona, Ortona, Vieste, Svetac and Vis stations (green box, Fig.
 174 2).

175 2.2 Operational mode

176 Every day, as soon as the WRF 1.5-km 1-min forecast results are available – which is at
 177 least 30h before any potential meteotsunami event (M) can occur, the high-pass filtered (with a
 178 2h cutoff period) mean sea-level pressure (i.e. P_{GW} for meteotsunami events) is automatically
 179 extracted (AdriSC Forecast step, Fig. 2). Then the maximum temporal rate of change (over a 4-
 180 min interval T_4) of this filtered pressure – i.e. $R_M = \max_{T_4} \frac{\partial P_{GW}}{\partial T_4}$, is derived at each WRF 1.5-km
 181 grid sea point. Such a condition has been proven to be efficient for the detection of
 182 meteotsunamigenic disturbances (Vilibić et al., 2016). No later than 28h before any
 183 meteotsunami event, the spatial coverage (in percentage) of the WRF 1.5-km grid sea points with
 184 a maximum temporal rate above 20Pa per 4-min interval ($R_M \geq 20$) is calculated (Automatic
 185 Detection step, Fig. 2). If this coverage exceeds 5%, a potential meteotsunami has been detected
 186 (event mode of the warning) and an automatic email – including a figure of the distribution of
 187 $R_M \geq 20$, is sent to the AdriSC team (red box, Fig. 2). The threshold of 5% is prescribed, being
 188 based on the analysis of recent meteotsunami events in which reproduction by the AdriSC
 189 modelling suite has been included (Denamiel et al., 2019). Otherwise (silent mode of the

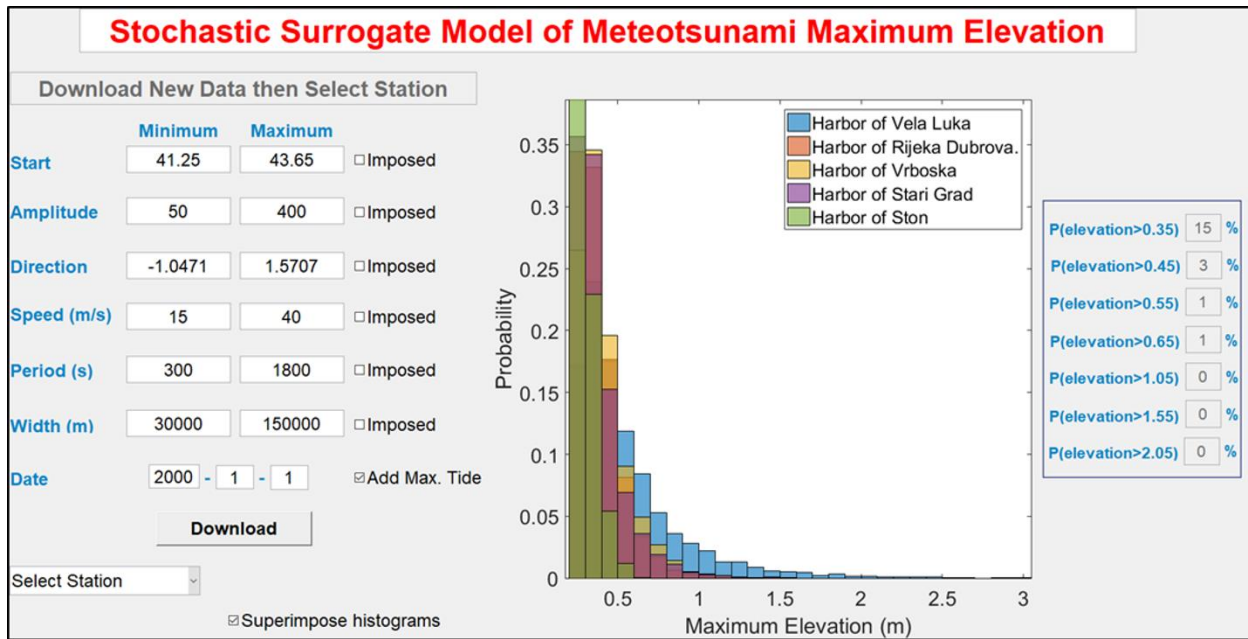
190 warning), at this stage, it is assumed that no meteotsunami will occur in the next forecasted day
191 (blue box, Fig. 2).



192

193 Figure 2. Operational meteotsunami hazard forecast within the CM_eEWS, based on atmospheric pressure field input from both (1) the
 194 deterministic model results (brown box) and (2) the measurements (green box). Every day, at least 30h before any meteotsunami event, the high-
 195 pass filtered pressure is extracted from the AdriSC forecast and used to automatically detect meteotsunamis by checking the spatial coverage of the
 196 values above 20Pa per 4-min interval of the maximal pressure temporal rate. If this coverage is below 5% then no meteotsunami is forecasted
 197 (blue box) – “silent” warning mode, otherwise a potential meteotsunami M is foreseen to occur (red box) – “event” warning mode, and an email is
 198 sent to the AdriSC team. At least 24h before the potential meteotsunami M occurs, the first forecast of hazard assessment is derived from the
 199 stochastic surrogate model used with ranges of pressure wave parameters manually extracted from the modelled filtered pressure. Finally, when
 200 the real-time observations become available, the hazard assessment is updated with new parameters extracted from the measurements.

201 In case of automatic meteotsunami detection, no later than **27h before any**
202 **meteotsunami event**, the filtered pressure field is visualized and analyzed by the AdriSC team.
203 If the detected pressure disturbance is recognized as an atmospheric pressure gravity wave, the
204 ranges of variation of the forecasted wave parameters including a $\pm 10\%$ of the parameter interval
205 of definition – latitude of origin ($y_{0,F} \pm 0.24^\circ N$), direction of propagation ($\theta_F \pm 0.26 rad$),
206 amplitude ($P_{A,F} \pm 35 Pa$), period ($T_F \pm 150 s$) and width ($d_F \pm 12000 m$), are manually estimated
207 from the WRF 1.5-km 1-min filtered air pressure results (Manual Extraction step, Fig. 1). To the
208 best of the author knowledge, the technology to automatically detect and extract the parameters
209 of the atmospheric disturbances driving the Adriatic Sea meteotsunamis is yet to be developed
210 and thus, for the moment, human intervention is unfortunately required in the early warning
211 system. As the errors associated with manually deriving the speed of the gravity waves (c_F) are
212 quite large, this parameter is always taken on its full range of definition [$15 m s^{-1}, 40 m s^{-1}$]. At
213 least 24h before the forecasted meteotsunami event, the meteotsunami stochastic surrogate model
214 – based on generalized Polynomial Chaos Expansion (gPCE), is used to deduce the
215 meteotsunami maximum elevation distributions at different locations of interest (Hazard Forecast
216 step, Fig. 2) via the user friendly interface developed in Matlab (Fig. 3). These distributions are
217 derived from 20000 random combinations of the six uniformly distributed input variables
218 selected in the range of the extracted parameters. In order to produce a conservative estimate of
219 the final maximum elevation expected at the locations of interest, (1) the surrogate model results
220 below 0.1m are ignored as irrelevant for meteotsunami hazard, and (2) the maximum tidal
221 elevation of the forecasted 24h period is added to the results of the stochastic surrogate model.



222

223 Figure 3. User friendly interface of the stochastic surrogate model of meteotsunami maximum

224 elevation developed in Matlab.

225 The maximum elevation distribution depending on the interval of definition of the atmospheric

226 wave parameters is generated and a first warning provides, at each location of interest, (1) the

227 probability of the expected maximum elevation derived from the surrogate model, and (2) the

228 deterministic maximum elevation from the ADCIRC model run, which is also taken into account

229 during the decision process for timely managing the hazard. Finally, it is planned that once the

230 early warning system will be fully operational (24/7 watch or fully automated procedure), in the

231 2h period before the forecasted meteotsunami event (i.e. the estimated time for the atmospheric

232 disturbances to cross the Adriatic Sea from the Italian coast), the 1-min air pressure measurements

233 from Ancona, Ortona, Vieste, Svetac and Vis will be analyzed by the AdriSC team and, if any

234 pressure gravity wave is detected, amplitude and period will be extracted from the observations

235 $(P_{A,M}, T_M)$. These parameters will then be used as constant values in the stochastic surrogate

236 model and new maximum elevation distributions will be produced with 20000 random

237 combinations of the three remaining uniformly distributed input variables selected in the range of
238 the parameters extracted from the model results $(y_{0,F}, \theta_F, d_F)$. The final meteotsunami warning
239 using the updated distribution of the maximum elevation (including maximum tidal elevation)
240 will then be ready to be published and accessible to users.

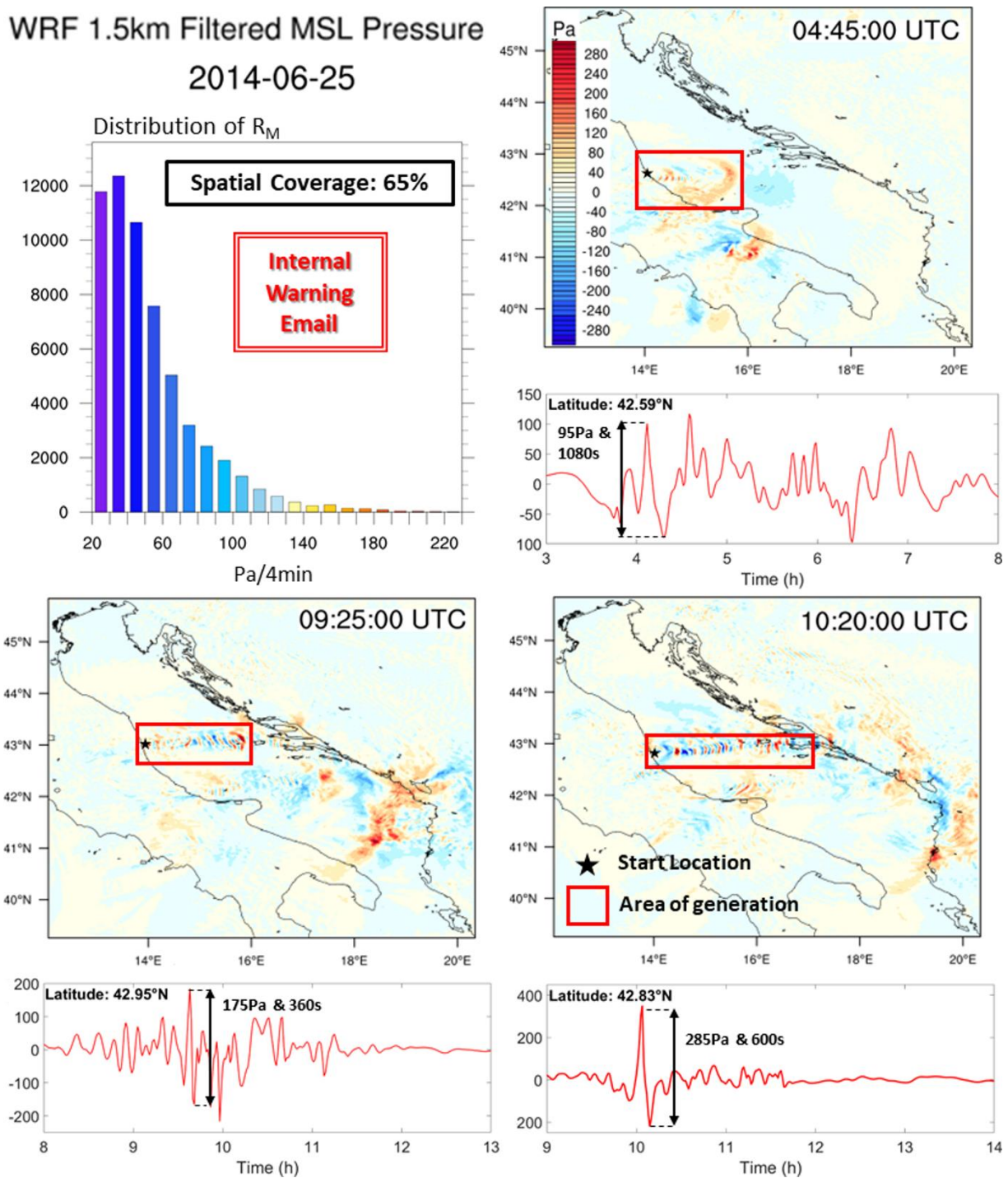
241 **3 Evaluation of the meteotsunami early warning system**

242 3.1 Evaluation against historical events

243 The first evaluation of the CMeEWS is performed against well-recorded events that took
244 place, before the early warning system became operational, at five locations of interest: Vela
245 Luka, Rijeka dubrovačka, Stari Grad, Vrboska and Ston (Fig. 1). In 2014, two strong events
246 happened at the end of June (Šepić et al., 2016), with reported maximum elevations of 1.5m in
247 Vela Luka, 0.5m in Stari Grad, 0.75m in Vrboska and 1.75m in Rijeka dubrovačka on the 25th of
248 June, and of 0.5m in Ston on the 26th of June. In summer 2017, tsunami-like waves were also
249 generated and observed in Stari Grad on the 28th of June (maximum elevation of 0.75m;
250 Denamiel et al., 2019) as well as the 30th of June during the night
251 (<http://www.izor.hr/meteotsunami>; maximum elevation of 0.32m measured at 18:30 UTC) and in
252 Vrboska on the 1st of July (maximum elevation of about 0.75m). Finally, on the 31st of March
253 2018, a meteotsunami wave with maximum reported sea elevation of 0.5m flooded Stari Grad
254 (Denamiel et al., 2019). For five of these events, the deterministic results of the AdriSC
255 Meteotsunami Forecast component have already been evaluated against a set of 48 air pressure
256 sensors and 19 tide gauges (Denamiel et al., 2019). This evaluation highlighted that the WRF
257 1.5-km model used in the AdriSC modelling suite presents some skills in forecasting the internal
258 gravity waves (IGWs) responsible for the observed meteotsunamis (i.e. the IGWs were always
259 forecasted by the model but their intensity or direction of propagation may not have been

260 reproduced perfectly). However, it also revealed that the slightest shift in location of the
261 modelled atmospheric disturbances resulted in the incapability of the ADCIRC model to
262 reproduce the observed meteotsunamis in the deterministic mode of the forecast. The stochastic
263 approach was thus developed to counter these shortcomings.

264 In this study, the stochastic surrogate model of the CMeEWS is also tested against these
265 five events in order to assess its capability to provide relevant warning to the public. In addition,
266 as the pressure sensors only became operational at the end of 2017, the atmospheric wave
267 parameters used in the stochastic surrogate model are only extracted from the WRF 1.5-km 1-
268 min high-pass filtered atmospheric pressure results. Finally, the meteotsunami impact highly
269 depends on the location of interest because (1) observations have shown that extreme
270 meteotsunami elevations present significant spatial variations in the eastern Adriatic Sea (Šepić
271 et al, 2016), and (2) flooding – the main hazard caused by meteotsunamis, depends on the
272 geomorphology/harbor design (Denamiel et al., 2018). In addition, due to the design of the
273 surrogate model (i.e. uniform prior distribution of the parameters), a majority of the stochastic
274 combinations lead to small oscillations (maximum elevations below 0.2m as seen in fig. 3) while
275 only about 10% lead to meteotsunami conditions. In this study it is thus assumed that flooding
276 occurs when at least 10% of the stochastic surrogate model maximum elevations reach more than
277 1.05m in Vela Luka, 0.65m in Rijeka dubrovačka, 0.55m in Vrboska, 0.45m in Stari Grad and
278 0.35m in Ston. These threshold values are prescribed considering the resilience of the coastline
279 in these locations (e.g. the salt plant located in Ston is the least resilient to strong sea-level
280 changes and meteotsunami waves), which in turn is largely defined by the real meteotsunami
281 hazard (e.g. the community of Vela Luka is the most resilient to meteotsunami hazard, as they
282 were hit by the strongest meteotsunami events along the Croatian coastline, Orlić, 2015).

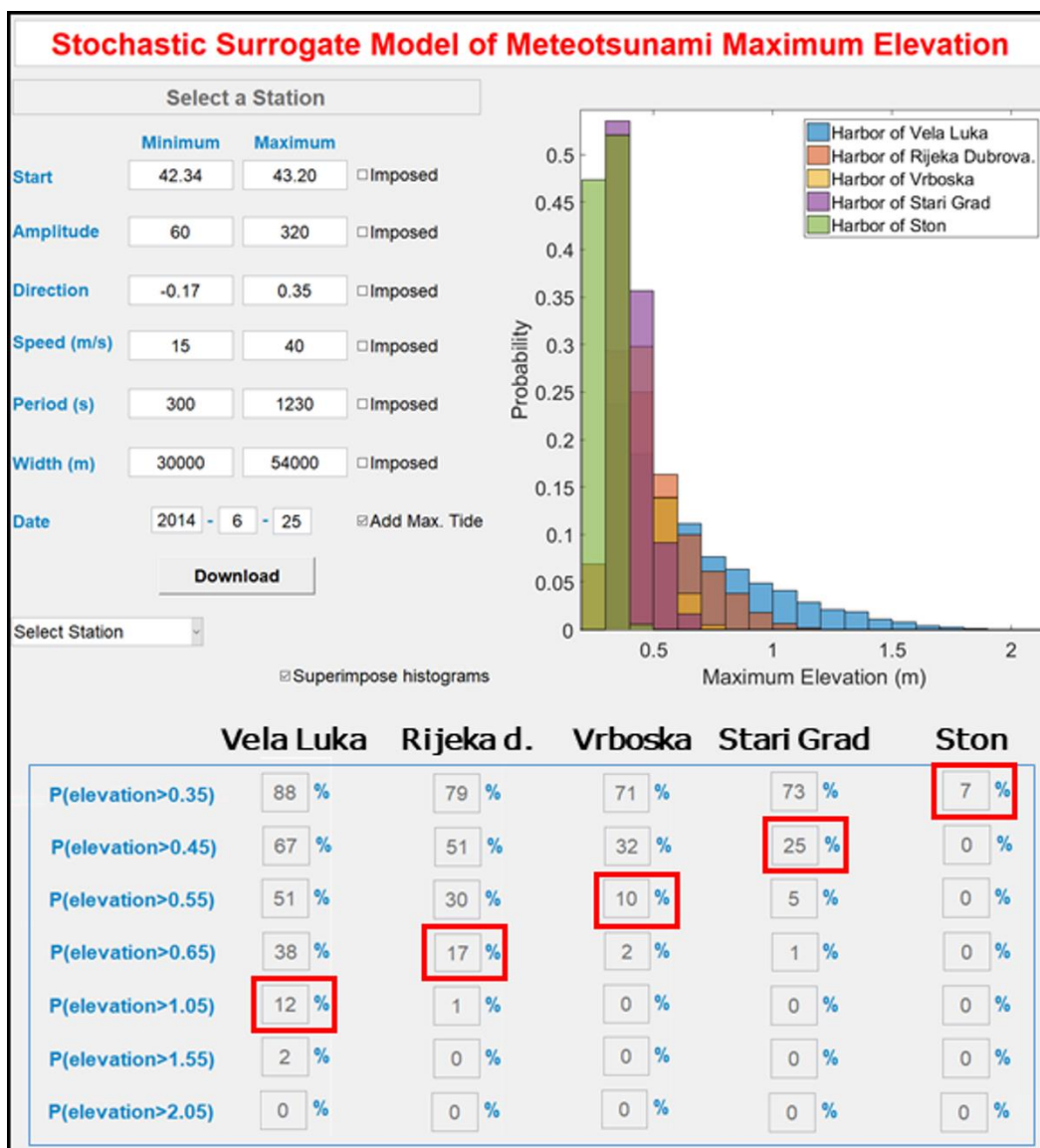


284 Figure 4. Meteotsunami event of the 25th of June 2014: distribution and spatial coverage of the
 285 maximum temporal rate of change (R_M) and associated spatial and temporal variations of the
 286 three atmospheric gravity waves extracted from the WRF 1.5-km forecast model. Time series of
 287 filtered MSL pressure are extracted at the start location of the three different disturbances (black
 288 stars) and direction of propagation is given by the orientation of the red boxes representing the
 289 area of generation of the meteotsunami waves.

290 Table 1. Input and output of the surrogate model during the five events used in the evaluation
 291 against historical events: (1) range of atmospheric gravity wave parameters (start location,
 292 amplitude, direction, period and width) extracted from the WRF 1.5-km forecast model results
 293 and (2) probability (in percent) of the maximum meteotsunami elevation surpassing the flooding
 294 threshold defined at five different locations (Vela Luka, Rijeka dubrovačka – R. dubro., Stari
 295 Grad, Vrboska and Ston). When the probabilities are above or equal to 10% (highlighted in
 296 bold), the meteotsunami warning is triggered. In addition, probabilities at locations at which
 297 flooding has been reported by eye-witnesses during the events are highlighted in red.

			25/06/14	26/06/14	28/06/17	01/07/17	31/03/18
Range of the input parameters	Latitude (°N)	Minimum	42.34	41.25	42.24	41.25	42.03
		Maximum	43.20	41.70	43.13	42.81	42.79
	Amplitude (Pa)	Minimum	60	255	85	100	85
		Maximum	320	340	185	275	215
	Direction (rad)	Minimum	-0.17	0.08	-0.17	0.35	0.26
		Maximum	0.35	0.60	0.70	1.04	0.78
	Period (s)	Minimum	300	330	1290	300	330
		Maximum	1230	630	1800	1410	1350
Width (km)	Minimum	30	30	88	30	30	
	Maximum	54	54	112	92	54	
Probability (%)	Vela Luka	$P(\xi_{\max} \geq 1.05\text{m})$	12	10	20	7	19
	R. dubro.	$P(\xi_{\max} \geq 0.65\text{m})$	17	1	5	3	12
	Stari Grad	$P(\xi_{\max} \geq 0.45\text{m})$	25	0	15	2	25
	Vrboska	$P(\xi_{\max} \geq 0.55\text{m})$	10	16	50	10	23
	Ston	$P(\xi_{\max} \geq 0.35\text{m})$	7	27	7	2	11

298



299

300 Figure 5. Maximum elevation distribution derived with the meteotsunami surrogate model at the
 301 five locations of interest (Vela Luka, Rijeka dubrovačka, Stari Grad, Vrboska and Ston) for the
 302 25th of June 2014 event.

303 In addition, as the thresholds dependent on the meteotsunami impact at the five studied locations,
 304 their values will most probably be re-evaluated in the near-future when more well-documented
 305 meteotsunami events in the eastern Adriatic will become available.

306 For each of the five meteotsunami events used in the evaluation, the distribution and spatial
307 coverage of the maximum temporal rate of change above 20Pa per 4 minutes (R_M) as well as
308 the different IGWs generated by the WRF 1.5-km model are analyzed. An example of this data is
309 presented in Figure 4 for the 25th of June 2014 event (figures for other events are given as
310 supporting information S2 to S5). As the spatial coverage of $R_M \geq 20$ is above 5% for all the
311 events, the switch of the warning system to the event mode would have been triggered in
312 operational conditions. The intervals of the atmospheric disturbance parameters
313 ($y_{0,F}, \theta_F, P_{A,F}, T_F, d_F$) defined with a $\pm 10\%$ margin to cover all possible IGW conditions
314 forecasted during the 24-h period of the event are thus presented in Table 1. The probabilities of
315 the maximum meteotsunami elevation (ξ_{\max}) surpassing the flooding threshold defined at the
316 five locations of interest are extracted from the surrogate model results and also presented in
317 Table 1. In addition, an example of the surrogate model results is presented Figure 5 for the 25th
318 of June 2014 event (figures for other events are given as supporting information S6 to S9). Given
319 the flooding criteria chosen in this study, in operational mode, the meteotsunami warnings would
320 have been triggered as follow:

- 321 • the 25th of June 2014: for Vela Luka, Rijeka dubrovačka, Vrboska and Stari Grad, which
322 all have been reported to be flooded (Šepić et al., 2016); this is in accordance with the
323 forecasted deterministic ADCIRC maximum elevation results (1.45m in Vela Luka,
324 0.80m in Rijeka dubrovačka, 0.65m in Stari Grad and 0.55m in Vrboska),
- 325 • the 26th of June 2014: for Vela Luka, Vrboska and Ston but, following eyewitness
326 reports, only Ston experienced flooding which was accurately forecasted with the
327 deterministic ADCIRC maximum elevation of 0.55m,

- 328 • the 28th of June 2017: for Vela Luka, Stari Grad and Vrboska but, following eyewitness
329 reports, only Stari Grad experienced flooding; the deterministic results obtained with the
330 ADCIRC model forecasted an elevation of only 0.35m in Stari Grad which would not
331 have been enough to cause flooding,
- 332 • the 1st of July 2017: for Vrboska, which was the only place flooded during this event; the
333 deterministic ADCIRC model forecasted 1m maximum elevation in Vela Luka but did
334 not captured proper meteotsunami amplification in Vrboska,
- 335 • the 31st of March 2018: for all the five locations but, following eyewitness reports, only
336 Stari Grad experienced flooding; the deterministic ADCIRC model did not reproduce at
337 all this event (only 0.25m forecasted in Stari Grad).

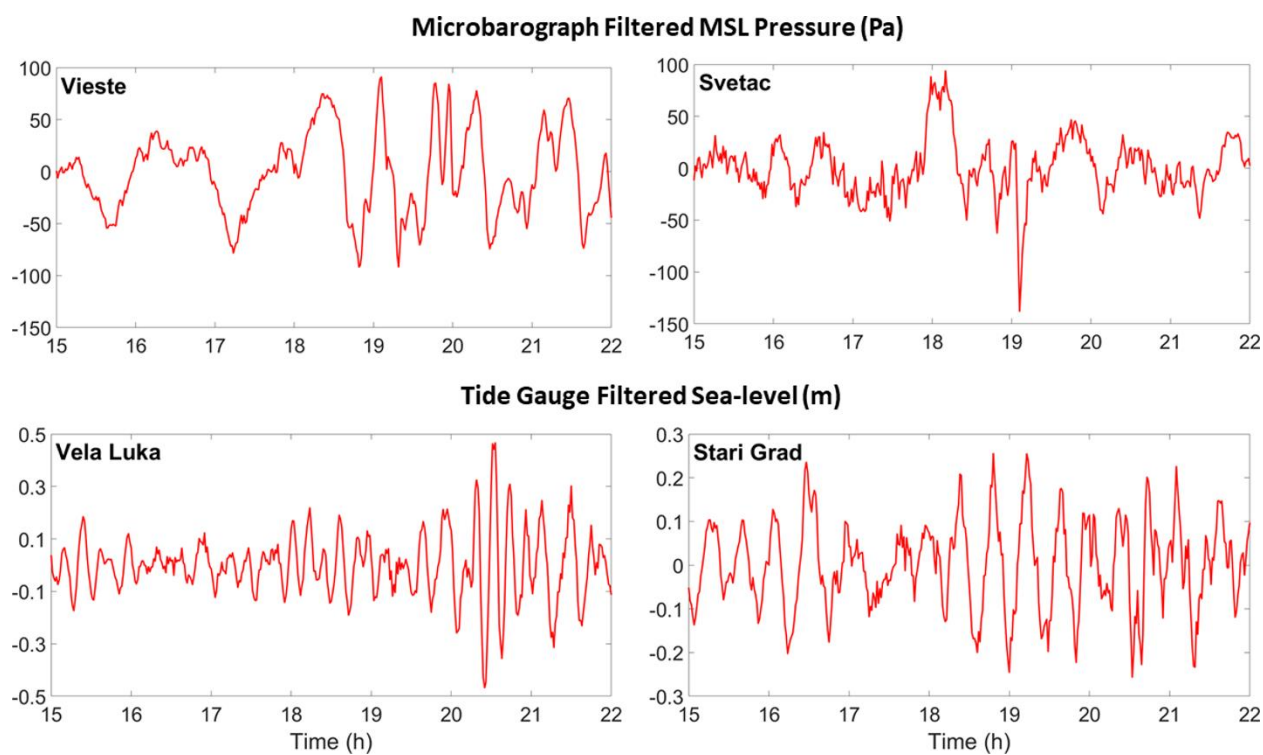
338 In summary, for the five studied historical events, the surrogate model of meteotsunami
339 maximum elevation is capable of forecasting the meteotsunami hazard in the areas that were
340 flooded, which was not always the case of the deterministic ADCIRC model (Denamiel et al.,
341 2019). Unfortunately, for many events, it also predicts flooding in areas where no meteotsunami
342 impact was reported.

343 3.2 Evaluation in operational mode

344 Since September 2018, the CMeEWS is tested in operational mode but meteotsunami
345 warnings are not yet released to the public. After nearly a year of run, meteotsunami hazard
346 forecasts were performed with the surrogate model forced by both deterministic model results
347 and measurements, for several events presenting the required meteotsunamigenic conditions
348 (Table 2).

349 The first event occurred on the 29th of October 2018 in the evening during the Vaia
350 storm, but was not publicly reported as a meteotsunami. The switch of the warning system to

351 event mode was triggered by (1) a 32% spatial coverage for $R_M \geq 20$ and (2) the analysis of the
 352 WRF 1.5-km filtered MSL pressure which revealed the presence of several high-frequency
 353 atmospheric disturbances travelling northwards from Vieste to the Croatian coastline (as can be
 354 seen in Figure S10 of the supporting information). However, only relatively small sea-level
 355 oscillations were deterministically forecasted with the ADCIRC model in the studied harbors
 356 along the track of the pressure disturbance (Vela Luka, Stari Grad and Vrboska). The first hazard
 357 forecast, based on the numerical results (Fig. S9), triggered the meteotsunami warning for all the
 358 locations except Rijeka dubrovačka (R1, Table 2 and Fig. S11 of the supporting information).
 359 The analysis of the filtered pressure measured at Vieste and Svetac (Fig. 6) – which were the
 360 stations the closest to the forecasted track of the pressure disturbances, showed that several
 361 IGWs of about 80Pa of amplitude and 10min of period were recorded between 18:00 and 22:00
 362 UTC.



363
 364 Figure 6. Available 1-min measurements (high-pass filtered with a 2h cutoff period) along the

365 forecasted track of the atmospheric disturbances during the 29th of October 2018: mean sea-level
366 pressure at Vieste and Svetac and sea-level at Vela Luka and Stari Grad.

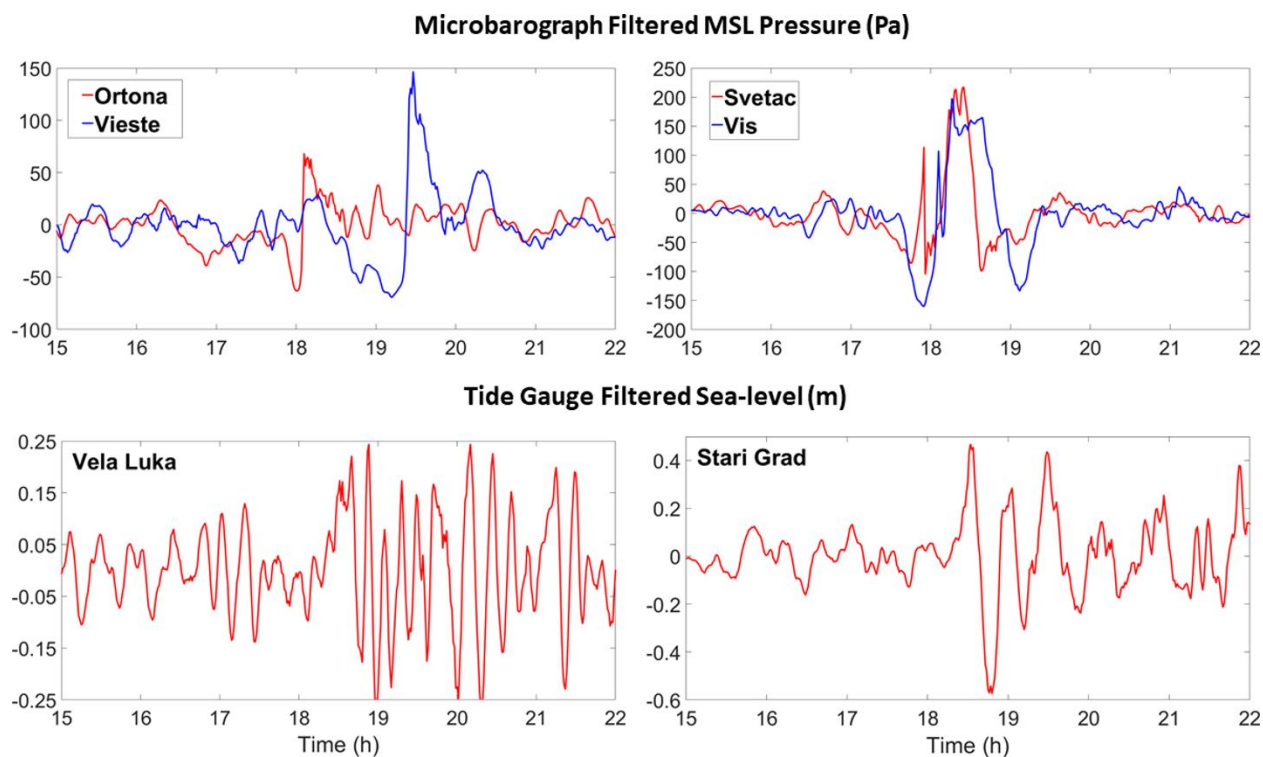
367 Table 2. As Table 1 but for the events that were detected since the warning system became
 368 operational in late 2018. R1 stands for a meteotsunami hazard forecast forced with input
 369 parameters extracted from the WRF-1.5km numerical model, while R2 hazard forecast uses air
 370 pressure amplitude and period extracted from the measurements and imposed as constant values,
 371 if a pressure disturbance is captured by the microbarographs. providing the final meteotsunami
 372 hazard.

			29/10/18		09/07/19		10/07/19	02/08/19
			R1	R2	R1	R2	R1	R1
Range of the input parameters	Latitude (°N)	Minimum	41.25		43.40		43.17	42.54
		Maximum	41.49		43.65		43.65	43.02
	Amplitude (Pa)	Minimum	86	80	175	135	172	53
		Maximum	173		245		400	123
	Direction (rad)	Minimum	1.31		-0.26		-0.26	-0.26
		Maximum	1.57		0.26		0.26	0.26
	Period (s)	Minimum	390	600	1530	1800	750	450
		Maximum	870		1800		1230	750
	Width (km)	Minimum	30		38		48	38
		Maximum	54		62		72	62
Probability (%)	Vela Luka	$P(\xi_{\max} \geq 1.05\text{m})$	10	6	0	0	0	1
	R. dubro.	$P(\xi_{\max} \geq 0.65\text{m})$	7	1	1	1	1	26
	Stari Grad	$P(\xi_{\max} \geq 0.45\text{m})$	29	14	19	29	2	21
	Vrboska	$P(\xi_{\max} \geq 0.55\text{m})$	29	25	8	20	1	0
	Ston	$P(\xi_{\max} \geq 0.35\text{m})$	11	5	0	0	0	10

373

374 After the hazard forecast was updated based on these measured values, the warning only
375 remained for Stari Grad and Vrboska (R2, Table 2 and Fig. S12 of the supporting information).
376 After the event, filtered sea-level measured at Vela Luka and Stari Grad (Fig. 6) revealed that
377 high-frequency oscillations with the respective periods of about 12min and 25min occurred at
378 both locations and generated the respective maximum elevations of 0.48m at 20:30 UTC and
379 0.26m at approximately 18:45 UTC. If the maximum tidal elevation (about 0.16m for both
380 locations during this event) is added, then the total elevation reached 0.64m in Vela Luka, which
381 is not enough to generate flooding, and 0.42m in Stari Grad which is slightly below the 0.45m
382 threshold that is used for the meteotsunami hazard warning. Unfortunately, no sea-level
383 measurements were available in Vrboska and, similarly to Stari Grad, even if a small
384 meteotsunami had occurred, it is unlikely that its effect could be visually distinguished from the
385 impact of the Vaia storm.

386 The next events all took place during summer storms in July and August 2019, when
387 unfortunately, the Ancona microbarograph stopped transmitting data. Between the 9th and the
388 10th of July 2019, the Adriatic region experienced severe storms which brought heavy rains,
389 hurricane force downbursts, tornadoes and the largest hailstorm ever recorded to date along the
390 Italian coast. For both days the event mode of the early warning system was triggered as (1) the
391 spatial coverage for $R_M \geq 20$ reached 22% and 44%, mostly due to the passage of the storm, and
392 (2) the analysis of the WRF 1.5-km filtered MSL pressure showed the presence of high-
393 frequency atmospheric disturbances with amplitudes greater than 150Pa travelling eastwards
394 from Ancona to the Croatian coastline (as can be seen in Figures S13 and S14 of the supporting
395 information).



396

397 Figure 7. Available 1-min measurements (high-pass filtered with a 2h cutoff period) along the
 398 forecasted track of the atmospheric disturbances during the 9th of July 2019: mean sea-level
 399 pressure at Ortona, Vieste, Svetac and Vis and sea-level at Vela Luka and Stari Grad.

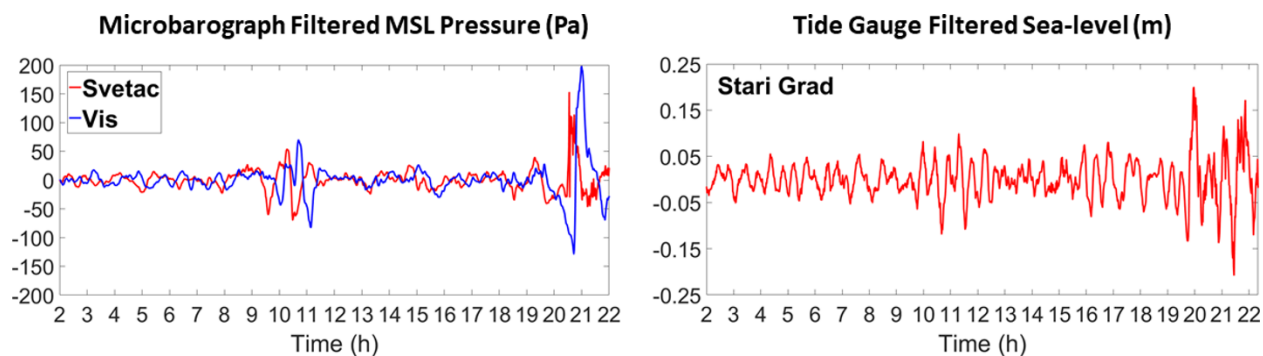
400 However, for these two days, similarly to the Vaia storm, the deterministic ADCIRC model only
 401 forecasted relatively small oscillations in the harbors of Vela Luka, Stari Grad and Vrboska
 402 located along the track of the pressure disturbances. For the 9th of July 2019, the first hazard
 403 forecast, based on numerical model results, triggered the meteotsunami warning in Stari Grad
 404 (R1, Table 2 and Fig. S15 of the supporting information). In addition, the analysis of the filtered
 405 pressure measured at Ortona, Vieste, Svetac and Vis stations (Fig. 7) clearly showed an
 406 atmospheric disturbance of about 135Pa and 30min period travelling eastward from Svetac to
 407 Vis between 17:30 and 18:30 UTC. As both Ortona and Vieste are located south from the
 408 forecasted track of the pressure disturbances, the pressure waves recorded at these stations were

409 assumed to be incapable to affect Stari Grad harbor where the warning was issued. Based on the
410 final hazard assessment (R2, Table 2 and Fig. S16 of the supporting information) updated with
411 the values extracted from the Svetac and Vis stations, the Stari Grad warning was confirmed and
412 an additional warning was triggered for Vrboska. During the evening of the 9th of July 2019, a
413 meteotsunami occurred in the harbor of Stari Grad, where the promenade was flooded
414 ([https://www.dalmacijadanas.hr/meteoroloski-tsunami-na-hvaru-more-se-povuklo-za-vise-od-](https://www.dalmacijadanas.hr/meteoroloski-tsunami-na-hvaru-more-se-povuklo-za-vise-od-metra)
415 [metra](https://www.dalmacijadanas.hr/meteoroloski-tsunami-na-hvaru-more-se-povuklo-za-vise-od-metra)). The analysis of the filtered sea-levels in Stari Grad (Fig. 7) confirmed the presence of a
416 1.05m height and 25min period meteotsunami wave just before 19:00 UTC. During the event,
417 the measured maximum elevation reached 0.47m which is, even without adding the maximum
418 tidal elevation, beyond the threshold value of 0.45m defined for meteotsunami warning. Sea-
419 level oscillations were also recorded in Vela Luka (Fig. 7), but the maximum elevation never
420 surpassed 0.25m. Finally, no meteotsunami was reported in Vrboska and thus the warning was
421 most probably too conservative for this location. For the 10th of July 2019, the forecasted
422 meteotsunami conditions were similar to the ones obtained from the previous day, except
423 concerning the periods of the disturbances which were all below 18min instead of the measured
424 30min. As meteotsunami are extremely sensitive to the period of the atmospheric disturbances,
425 no warning was triggered by the hazard forecast based on these numerical results (R2, Table 2
426 and Fig. S17 of the supporting information). In addition, the monitoring of the air pressure
427 measurements did not show any disturbance with period greater than 18min and no
428 meteotsunami was reported in the studied locations.

429 Two more storms took place in the Adriatic Sea during the 13th and the 28th of July 2019
430 (not presented in this study) and both triggered the event mode of the warning system, but
431 conditions for these storms were extremely similar to the 10th of July 2019 event and the hazard

432 forecast based on both numerical results and measurements did not trigger any meteotsunami
433 warning.

434 Finally, the last event occurred the 2nd of August 2019 just before a storm that swept the
435 eastern Adriatic coast, where falling trees blocking roads, damaged power distribution lines and
436 flooding were reported in the media. The event mode was triggered by (1) a 19% spatial
437 coverage for $R_M \geq 20$ and (2) the analysis of the WRF 1.5-km filtered MSL pressure which
438 revealed that a high-frequency atmospheric disturbance was travelling eastwards around 10:00
439 UTC in the middle of the Adriatic (about 42.77°N of latitude), from the Italian to the Croatian
440 coasts (as can be seen in Figure S18 of the supporting information). The forecasted
441 meteotsunami hazard based on these numerical results was quite high and warnings were
442 triggered for Rijeka dubrovačka, Stari Grad and Ston (R1, Table 2 and Fig. S19 of the supporting
443 information). Similarly to the other events, the deterministic results of the ADCIRC model only
444 forecasted some oscillations of small amplitude in the harbors of interest. Due to technical
445 problems the Ortona and Vela Luka stations were not transmitting data during this event, thus the
446 analysis of the filtered pressure was based on measurements at Svetac and Vis (Fig. 8).
447 Interestingly, some disturbances were indeed travelling eastwards during the 2nd of August
448 between 10:00 and 12:00 UTC. However, their amplitude was below 50Pa and they were not
449 capable of generating strong oscillations and/or flooding along the Croatian coast. The warnings
450 were thus canceled and in fact no meteotsunami was reported for this event. Finally, the biggest
451 atmospheric disturbance – which generated some moderate oscillations (about 0.15m of
452 amplitude) in the harbor of Stari Grad, as can be seen in the filtered sea-level data (Fig. 8) – was
453 recorded between 20:00 and 22:00 UTC during the peak of the storm.



454
455 Figure 8. Available 1-min measurements (high-pass filtered with a 2h cutoff period) along the
456 forecasted track of the atmospheric disturbances during the 2nd of August 2019: mean sea-level
457 pressure at Svetac and Vis and sea-level at Stari Grad.

458 For this event, the assessment of the meteotsunami hazard was first largely overestimated due to
459 the deterministic forecast of pressure disturbances capable of generating strong sea-level
460 oscillations in the eastern Adriatic but, as the measured pressure disturbances were far smaller
461 than expected, no meteotsunami occurred.

462 The evaluation of the CMeEWS in operational mode highlights that the microbarograph
463 network plays a crucial role in terms of delivering the final warnings and confirms that the
464 surrogate model forecasts the meteotsunami hazard in a conservative way even during storms
465 events which, in the eastern Adriatic, are not the classical generation mechanism of the
466 meteotsunamigenic pressure disturbances.

467 **4 Discussion and conclusions**

468 Notwithstanding major research efforts, the scarcity of the measurements and the
469 reliability of the numerical models in meteotsunami studies are still major restrictions for hazard
470 assessment and forecast, and even more for risk management (e.g. for the determination of a
471 100-year meteotsunami event). Based on lessons from river flooding hazard warning systems

472 designed and evaluated in hydrological studies (e.g. Beven, 2006; Sivakumar, 2008), two major
473 conclusions can be drawn: (1) the promotion of uncertainty analysis of measurements and
474 modelled results is of crucial importance for hazard assessment and forecast, and (2) the
475 effectiveness of the warning systems is not determined only by the predictive accuracy of the
476 models, but also by the lead time and the available social response set.

477 The presented prototype of meteotsunami early warning system combining deterministic
478 and stochastic hazard assessment was designed to address such concerns. In particular, the very
479 first use of a gPCE-based surrogate model to derive atmospherically-driven extreme sea-level
480 hazard was motivated by the successful application of such methods for uncertainty
481 quantification in a wide range of areas including mechanics, engineering, water resources and
482 geosciences (e.g. Foo et al., 2007; Rupert & Miller, 2007; Giraldi et al., 2017). The main
483 advantages of this kind of approach are (1) the propagation of the uncertainties associated with
484 the atmospheric disturbances (e.g. location, direction, speed) to the maximum elevation results,
485 (2) the potentiality of using both deterministic forecast results and measurements to provide the
486 surrogate model input parameters, and (3) the few minutes of computation needed to assess, with
487 a large number of samples and no additional deterministic simulation, the hazard of any studied
488 event (e.g. meteotsunami). However, the main disadvantages are that the surrogate model (1)
489 only relies on ocean numerical results forced by synthetic atmospheric disturbances (e.g.
490 idealized pressure waves), and (2) requires a large number of synthetic simulations to be built
491 with good enough accuracy (e.g. in this study, 4161 simulations were used to build the model
492 with approximately 80% accuracy). Additionally, in operational mode, the early warning system
493 currently presents three major weaknesses. First, due to the high-resolution of the deterministic
494 models and thus the relative slowness of the system, the early forecast of the meteotsunami

495 hazard (at least 24-h prior to any event) is only derived once from numerical results obtained 2
496 days in advance. This means that the first warnings are always based on conditions forecasted
497 from a 72-h old assimilation cycle which can lack of accuracy, particularly during extreme
498 events. Second, human intervention is still required in the present set-up of the early warning
499 system in order to extract the IGW parameters from the deterministic forecast. And third, to be
500 able to provide the final meteotsunami warnings derived from hazards forecasted with input
501 parameters extracted from the measured mean sea-level air pressures along the Italian coast and
502 the middle Adriatic, the microbarograph data should be analyzed in a timely manner with
503 efficient operational tools which, in the CMeEWS are still under development.

504 On one hand, the evaluation of the early warning system with five well-recorded events,
505 demonstrates that (1) the IGWs driving the eastern Adriatic meteotsunamis are always forecasted
506 and well detected, and (2) the meteotsunami hazard derived only from the deterministic model
507 results is conservative but tends to be largely overestimated in certain locations such as Vela
508 Luka or Vrboska. On the other hand, the evaluation in operational mode highlights the
509 importance of (1) taking into account the uncertainties associated with the forecasted
510 meteotsunamigenic atmospheric disturbances particularly during storm events when the
511 deterministic model lacks of accuracy, (2) updating the final warnings using meteotsunami
512 hazards based on input parameters extracted from the measured pressure disturbances, and (3)
513 extending and maintaining the measurement network (microbarographs and tide gauges) along
514 the Italian and Croatian coastlines in order to produce more accurate hazard assessments and to
515 better understand how and where the system failed. Following these conclusions, to improve the
516 accuracy of the warnings, for all potential future events, (1) the system should be thoroughly re-
517 evaluated, (2) the measurements recorded by the microbarographs should be used in a timely

518 manner to derive the final hazard assessment, (3) the flooding criteria and the input parameter
519 ranges of the surrogate model should be finely tuned as more data will become available, and (4)
520 ultimately, once the prototype will be fully tested, the meteotsunami warnings will not only be
521 triggered when more than 10% of the maximum elevations surpass the thresholds defined at the
522 sensitive locations, but their strength (yellow, orange, red) will also be defined depending on the
523 detailed statistical information (maximum, 75th-percentile, mean, median, etc.) extracted from
524 the extreme sea-level distributions.

525 Finally, the CMeEWS combining 1-min air pressure measurements – accurate but
526 scarcely spread along the Italian coast and the middle Adriatic Sea, state-of-the-art deterministic
527 models – dedicated to meteotsunami forecast but computationally costly and slow, and a newly
528 developed stochastic surrogate model – running at nearly zero cost but yet to be fully tested,
529 highlights the need to use real time high-temporal resolution observational networks for regional
530 early warning systems in the Mediterranean and presents an alternative way to deal with
531 atmospherically-driven extreme sea-level hazard assessment.

532 Acknowledgments and data

533 Acknowledgement is made for the support of the ECMWF staff, in particular Xavier Abellan, as
534 well as for ECMWF's computing and archive facilities used in this research, which has been
535 supported by projects MESSI (UKF Grant 25/15), ADIOS (Croatian Science Foundation Grant
536 IP-2016-06-1955) and ECMWF Special Project (The Adriatic decadal and inter-annual
537 oscillations: modelling component). The authors would also like to thank the two anonymous
538 reviewers for their valuable contributions. The Matlab interface developed for the stochastic
539 surrogate model – including the gPCE coefficients needed to create the meteotsunami maximum
540 elevation distributions at the five studied locations, as well as the WRF 1.5-km filtered results of
541 sea surface pressure can be obtained under the Open Science Framework (OSF) FAIR data
542 repository at <https://osf.io/jysqu/> (doi:10.17605/OSF.IO/JYSQU).

543

544 **References**

- 545 Arnst, M., & Ponthot, J.-P. (2014). An overview of nonintrusive characterization, propagation,
546 and sensitivity analysis of uncertainties in computational mechanics. *Int. J. Uncertain. Quan.*, 4,
547 387–421. <https://doi.org/10.1615/int.j.uncertaintyquantification.2014006990>
- 548 Beven, K. J. (2006). On undermininEg the science? *Hydrological Processes*, 20(3), 141–146.
- 549 Bulthuis, K., Arnst, M., Sun, S., & Pattyn, F. (2019). Uncertainty quantification of the multi-
550 centennial response of the Antarctic ice sheet to climate change. *The Cryosphere*, 13, 1349-1380.
551 <https://doi.org/10.5194/tc-13-1349-2019>
- 552 Burkardt J. (2014). Slow Exponential Growth for Gauss Patterson Sparse Grids,
553 http://people.sc.fsu.edu/~jburkardt/presentations/sgmga_gps.pdf.
- 554 Cho, K.-H., Choi, J.-Y., Park K.-S., Hyun S.-K., & Park J.-Y. (2013). A synoptic study on
555 tsunami-like sea level oscillations along the west coast of Korea using an unstructured-grid
556 ocean model. In Conley, D. C., Masselink, G., Russell, P. E., & O'Hare, T. J. (Eds.), *Proc. 12th*
557 *Int. Coastal Symposium* (Vol. 1, pp. 678–683). Plymouth, MN: Coastal Processes Research
558 Group, School of Marine Science and Engineering, Plymouth University.
- 559 Denamiel, C., Šepić, J., & Vilibić, I. (2018). Impact of Geomorphological Changes to Harbor
560 Resonance During Meteotsunamis: The Vela Luka Bay Test Case. *Pure Appl. Geophys.*, 175,
561 3839. <https://doi.org/10.1007/s00024-018-1862-5>
- 562 Denamiel, C., Šepić, J., Ivanković, D., & Vilibić, I. (2019). The Adriatic Sea and Coast
563 modelling suite: Evaluation of the meteotsunami forecast component. *Ocean Modell.*, 135, 71–
564 93. doi:10.1016/j.ocemod.2019.02.003

- 565 Dusek, G., DiVeglio, C., Licate, L., Heilman, L., Kirk, K., Paternostro, C., & Miller,
566 A. (2019). A meteotsunami climatology along the U.S. East Coast. *Bull. Amer. Meteor. Soc.*, 0.
567 <https://doi.org/10.1175/BAMS-D-18-0206.1>
- 568 Ewing, M., Press, F., & Donn, W. L. (1954). An explanation of the Lake Michigan wave of 26
569 June 1954. *Science*, 120, 684–686. doi:10.1126/science.120.3122.684
- 570 Foo, J., Yosibash, Z., & Karniadakis, G. E. (2007). Stochastic simulation of riser-sections with
571 uncertain measured pressure loads and/or uncertain material properties. *Comput. Method. Appl.
572 Mech. Eng.*, 196, 4250–4271.
- 573 Formaggia, L., Guadagnini, A., Imperiali, I., Lever, V., Porta, G., Riva, M., Scotti, A. &
574 Tamellini, L. (2013). Global sensitivity analysis through polynomial chaos expansion of a basin-
575 scale geochemical compaction model. *Computational Geosciences*, 17(1), 25–42.
576 doi:10.1007/s10596-012-9311-5
- 577 Giraldi, L., Le Maître, O. P., Mandli, K. T., Dawson, C. N., Hoteit, I., & Knio, O. M. (2017).
578 Bayesian inference of earthquake parameters from buoy data using a polynomial chaos-based
579 surrogate. *Comput. Geosci.*, 21(4), 683–699. <https://doi.org/10.1007/s10596-017-9646-z>
- 580 Ghanem, R., Higdon, R., & Owhadi, H. (2017). Handbook of Uncertainty Quantification,
581 Springer. <https://doi.org/10.1007/978-3-319-12385-1>
- 582 Hibiya, T., & Kajiura, K. (1982). Origin of the *Abiki* phenomenon (a kind of seiche) in Nagasaki
583 Bay. *J. Oceanogr. Soc. Jpn.*, 38, 172–182. doi:10.1007/BF02110288

- 584 Knio, O. M., & Le Maître, O. P. (2006). Uncertainty propagation in CFD using polynomial
585 chaos decomposition. *Fluid Dynamics Research*, 38(9),616–640.
586 doi:10.1016/j.fluidyn.2005.12.003
- 587 Le Maître, O. P. & Knio, O. M. (2010). Spectral Methods for Uncertainty Quantification: With
588 Applications to Computational Fluid Dynamics, *Springer Science & Business Media*.
589 <https://doi.org/10.1007/978-90-481-3520-2>
- 590 Linares, A., Wu, C. H., Bechle, A. J., Anderson, E. J., & Kristovich, D. A. R. (2019). *Sci. Rep.*,
591 9, 2105. doi: 10.1038/s41598-019-38716-2
- 592 Luettich, R. A., Birkhahn, R. H., & Westerink, J. J. (1991). Application of ADCIRC-2DDI to
593 Masonboro Inlet, North Carolina: A brief numerical modeling study. Contractors Report to the
594 US Army Engineer Waterways Experiment Station, August, 1991.
- 595 Masina, M., Archetti, R., Besio, G. & Lamberti, A. (2017). Tsunami taxonomy and detection
596 from recent Mediterranean tide gauge data. *Coastal Engineering*, 127, 145–169.
- 597 Monserrat, S., Vilibić, I., & Rabinovich, A. B. (2006). Meteotsunamis: atmospherically induced
598 destructive ocean waves in the tsunami frequency band. *Nat. Hazards Earth Syst. Sci.*, 6, 1035-
599 1051.
- 600 Najm, H. N., Debusschere, B. J., Marzouk, Y. M., Widmer, S. & Le Maître, O. P. (2009).
601 Uncertainty quantification in chemical systems. *Int. J. Numer. Meth. Engng.*, 80, 789–814.
602 doi:10.1002/nme.2551
- 603 Novak, E., Ritter, K., Schmitt, R., Steinbauer, A. (1999). Simple cubature formulas with high
604 polynomial exactness. *Constructive Approximation*, 15(4), 499–522.

- 605 Okal, E. A., Visser, J. N. J., & de Beer, C. H. (2014). The Dwarskersbos, South Africa local
606 tsunami of August 27, 1969: field survey and simulation as a meteorological event. *Nat.*
607 *Hazards*, *74*, 251–268. doi:10.1007/s11069-014-1205-5
- 608 Olabarrieta, M., Valle-Levinson, A., Martinez, C. J., Pattiaratchi, C., & Shi, L. (2017).
609 Meteotsunamis in the northeastern Gulf of Mexico and their possible link to El Niño Southern
610 Oscillations *Nat. Hazards*, *88*(3), 1325-1346. <https://doi.org/10.1007/s11069-017-2922-3>
- 611 Orlić, M., Belušić, D., Janeković, I., & Pasarić, M. (2010). Fresh evidence relating the great
612 Adriatic surge of 21 June 1978 to mesoscale atmospheric forcing. *J. Geophys. Res.*, *115*,
613 C06011. doi:10.1029/2009JC005777
- 614 Orlić, M. (2015). The first attempt at cataloguing tsunami-like waves of meteorological origin in
615 Croatian coastal waters. *Acta Adriat.*, *56*, 83–95.
- 616 Pattiaratchi, C., & Wijeratne, E. M. S. (2014). Observations of meteorological tsunamis along
617 the south-west Australia. *Nat. Hazards*, *74*, 281–303. doi:10.1007/s11069-014-1263-8
- 618 Pellikka, H., Rauhala, J., Kahma, K. K., Tapani, S., Boman, H., & Kangas, A. (2014). Recent
619 observations of meteotsunamis on the Finnish coast. *Nat. Hazards*, *74*, 197–215.
620 doi:10.1007/s11069-014-1150-3
- 621 Renault, L., Vizoso, G., Jansà, A., Wilkin, J., & Tintoré, J. (2011). Toward the predictability of
622 meteotsunamis in the Balearic Sea using regional nested atmosphere and ocean models.
623 *Geophys. Res. Lett.*, *38*, L10601. doi:10.1029/2011gl047361
- 624 Rupert, C., & Miller, C. (2007). An analysis of polynomial chaos approximations for modeling
625 single-fluid-phase flow in porous medium systems. *J. Comput. Phys.*, *226*, 2175–2205.

- 626 Salaree, A., Mansouri, R., & Okal, E. A. (2018). The intriguing tsunami of 19 March 2017 at
627 Bandar Dayyer, Iran: field survey and simulations. *Nat. Hazards*, *90*(3), 1277–1307. doi:
628 10.1007/s11069-017-3119-5
- 629 Šepić, J., Vilibić, I., & Strelec Mahović, N. (2012). Northern Adriatic meteorological tsunamis:
630 observations, link to the atmosphere, and predictability. *J. Geophys. Res.*, *117*, C02002.
631 doi:10.1029/2011JC007608
- 632 Šepić, J., Međugorac, I., Janeković, I., Dunić, N. & Vilibić, I. (2016). Multi-meteotsunami event
633 in the Adriatic Sea generated by atmospheric disturbances of 25–26 June 2014. *Pure Appl.*
634 *Geophys.*, *173*, 4117-4138. <https://doi.org/10.1007/s00024-016-1249-4>
- 635 Šepić, J., Vilibić, I., Beg Paklar, G., Dadić, V., Denamiel, C., Dunić, N., et al. (2017). Towards
636 understanding and operational early warning of the Adriatic meteotsunamis: Project MESSI.
637 *Revue Paralia*, *10*, 177–182. doi: 10.5150/cmcm.2017.033
- 638 Shchepetkin, A. F., & McWilliams, J. C. (2005). The regional oceanic modeling system: A split-
639 explicit, free-surface, topography-following-coordinate ocean model. *Ocean Modell.*, *9*, 347–
640 404.
- 641 Shchepetkin, A. F., & McWilliams, J. C. (2009). Correction and commentary for “Ocean
642 forecasting in terrain-following coordinates: Formulation and skill assessment of the regional
643 ocean modeling system” by Haidvogel et al., *J. Comput. Phys.*, *227*, pp. 3595–3624. *J. Comput.*
644 *Phys.*, *228*, 8985–9000. doi:10.1016/j.jcp.2009.09.002
- 645 Sivakumar, B. (2008). Undermining the science or undermining Nature? *Hydrological*
646 *Processes*, *22*, 893–897.

- 647 Skamarock, W. C., Klemp, J. B., Dudhia, J., Gill, D. O., Barker, D. M., Wang, W., & Powers, J.
648 G. (2005). A Description of the Advanced Research WRF Version 2. NCAR Technical Note
649 NCAR/TN-468+STR. doi:10.5065/D6DZ069T
- 650 Smolyak, S. A. (1963). Quadrature and interpolation formulas for tensor products of certain
651 classes of functions. *Dokl. Akad. Nauk SSSR*, 148(5), 1042–1045.
- 652 Soize, C., & Ghanem, R. G. (2004). Physical Systems with Random Uncertainties: Chaos
653 Representations with Arbitrary Probability Measure. *SIAM Journal on Scientific*
654 *Computing*, 26(2), 395–410. <https://doi.org/10.1137/S1064827503424505>
- 655 Sraj, I., Mandli, K., Knio, O., Dawson, C. N., & Hoteit, I. (2014). Uncertainty quantification and
656 inference of Manning's friction coefficients using DART buoy during the Tohoku tsunami.
657 *Ocean Model.*, 83, 82–97. doi:10.1016/j.ocemod.2014.09.001
- 658 Tanaka, K. (2010). Atmospheric pressure-wave bands around a cold front resulted in a meteo-
659 tsunami in the East China Sea in February 2009. *Nat. Hazards Earth Syst. Sci.*, 10, 2599–2610.
- 660 Vilibić, I., & Šepić, J. (2009). Destructive meteotsunamis along the eastern Adriatic coast:
661 Overview. *Phys. Chem. Earth*, 34, 904–917.
- 662 Vilibić, I., Šepić, J., Rabinovich, A.B. & Monserrat, S. (2016). Modern approaches in
663 meteotsunami research and early warning. *Front. Mar. Sci.*, 3(57). doi:
664 10.3389/fmars.2016.00057
- 665 Vučetić, T., Vilibić, I., Tinti, S., & Maramai, A. (2009). The Great Adriatic flood of 21 June
666 1978 revisited: An overview of the reports. *Phys. Chem. Earth*, 34, 894–903.

- 667 Wang, M., Wan, Z., & Huang, Q. (2016). A New Uncertain Analysis Method for the Prediction
668 of Acoustic Field with Random and Interval Parameters. *Shock and Vibration*, 16
669 pp. <https://doi.org/10.1155/2016/3693262>
- 670 Warner, J. C., Armstrong, B., He, R., & Zambon, J. B. (2010). Development of a Coupled
671 Ocean-Atmosphere-Wave-Sediment Transport (COAWST) modeling system. *Ocean Modell.*,
672 35, 230-244.
- 673 Whitmore, P., & White, B. (2014). Meteotsunami forecasting: sensitivities demonstrated by the
674 2008 Boothbay, Maine, event. *Nat. Hazards*, 74, 11–24. doi:10.1007/s11069-014-1056-0
- 675 Xiu, D., & Karniadakis, G. E. (2002). The Wiener--Askey Polynomial Chaos for Stochastic
676 Differential Equations. *SIAM Journal on Scientific Computing*, 24(2), 619–644.
677 <https://doi.org/10.1137/S1064827501387826>

Figure 1.

Author Manuscript

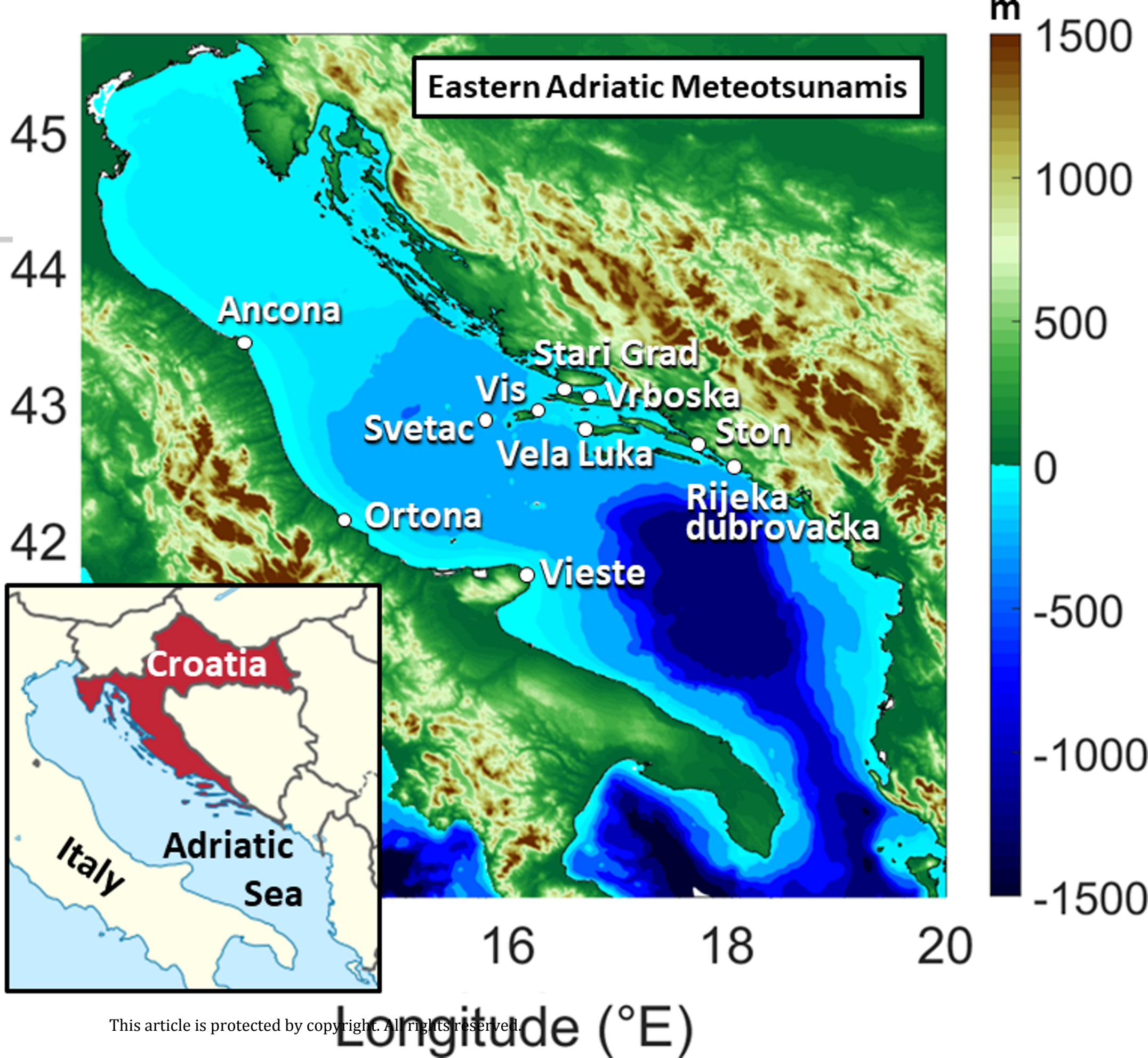
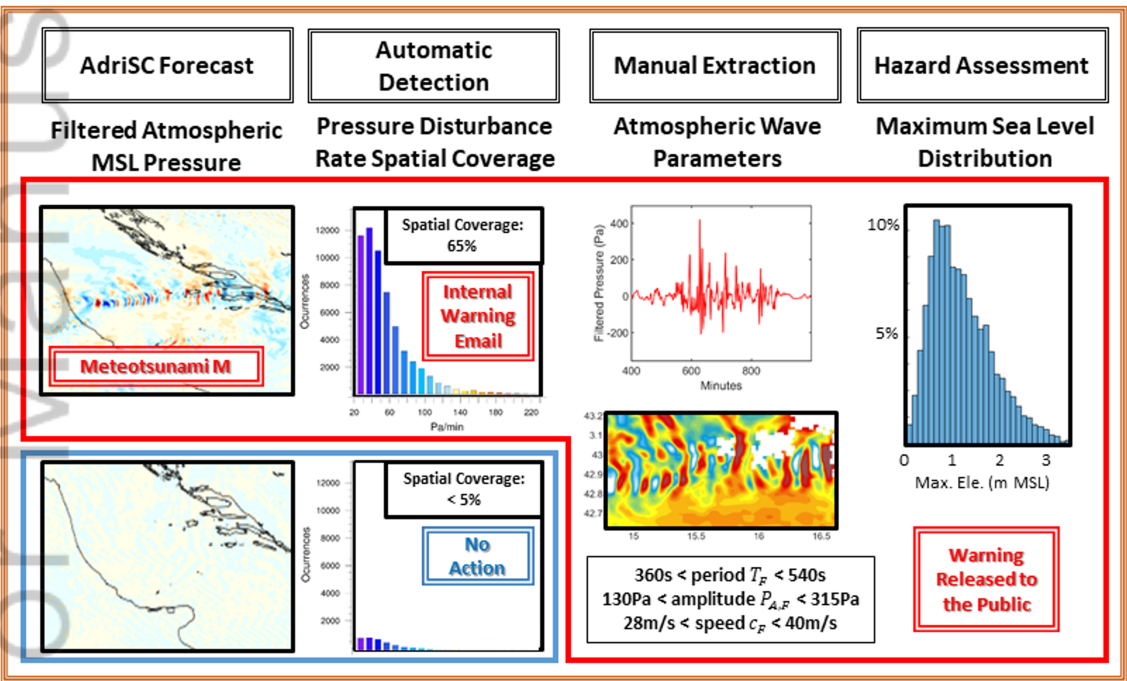


Figure 2.

Author Manuscript

FROM DETERMINISTIC MODEL



FROM MEASUREMENTS

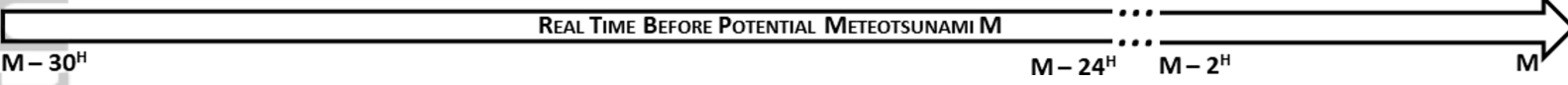
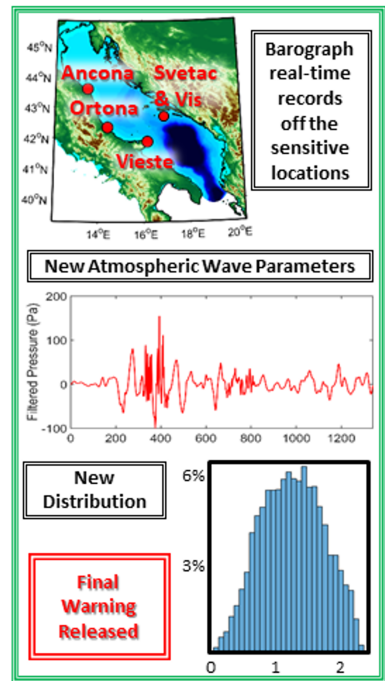


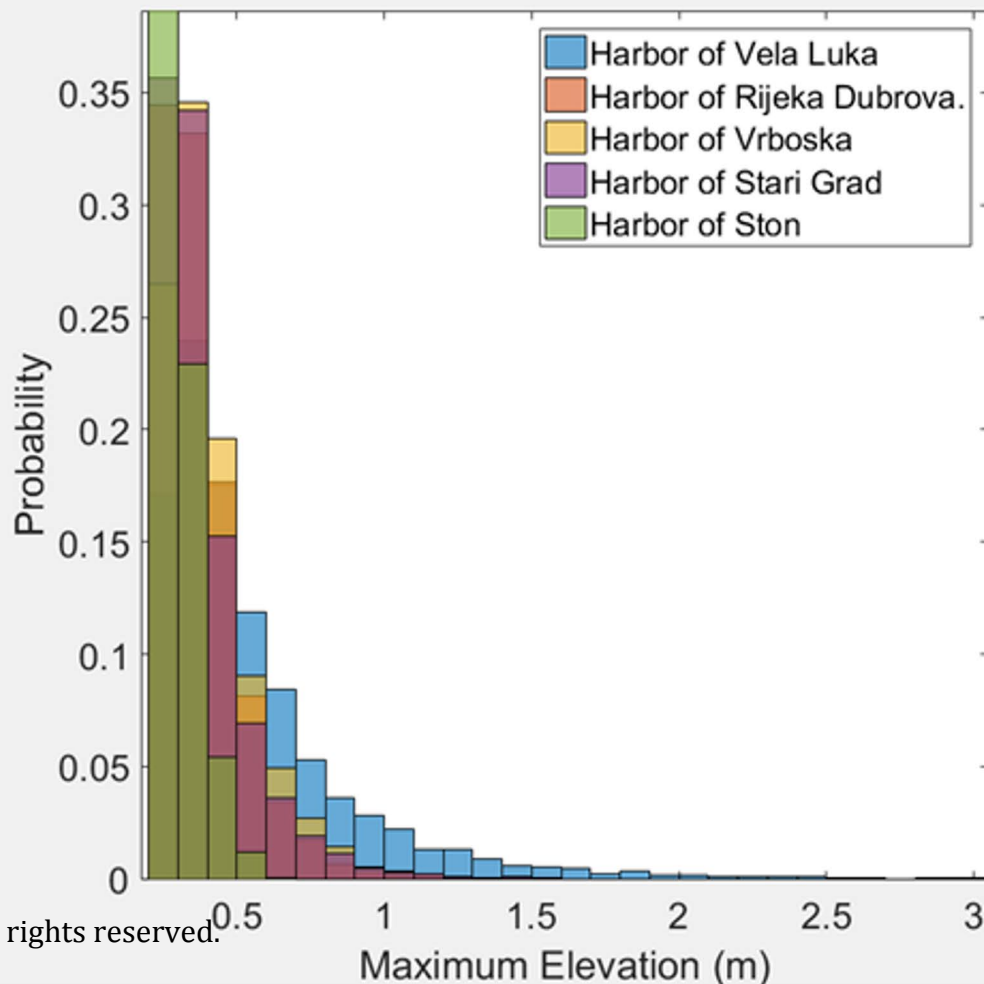
Figure 3.

Author Manuscript

Stochastic Surrogate Model of Meteotsunami Maximum Elevation

Download New Data then Select Station

	Minimum	Maximum	
Start	<input type="text" value="41.25"/>	<input type="text" value="43.65"/>	<input type="checkbox"/> Imposed
Amplitude	<input type="text" value="50"/>	<input type="text" value="400"/>	<input type="checkbox"/> Imposed
Direction	<input type="text" value="-1.0471"/>	<input type="text" value="1.5707"/>	<input type="checkbox"/> Imposed
Speed (m/s)	<input type="text" value="15"/>	<input type="text" value="40"/>	<input type="checkbox"/> Imposed
Period (s)	<input type="text" value="300"/>	<input type="text" value="1800"/>	<input type="checkbox"/> Imposed
Width (m)	<input type="text" value="30000"/>	<input type="text" value="150000"/>	<input type="checkbox"/> Imposed
Date	<input type="text" value="2000"/> - <input type="text" value="1"/> - <input type="text" value="1"/>		<input checked="" type="checkbox"/> Add Max. Tide
<input type="button" value="Download"/>			
Select Station	<input type="text"/>		



P(elevation>0.35)	<input type="text" value="15"/>	%
P(elevation>0.45)	<input type="text" value="3"/>	%
P(elevation>0.55)	<input type="text" value="1"/>	%
P(elevation>0.65)	<input type="text" value="1"/>	%
P(elevation>1.05)	<input type="text" value="0"/>	%
P(elevation>1.55)	<input type="text" value="0"/>	%
P(elevation>2.05)	<input type="text" value="0"/>	%

This article is protected by copyright. All rights reserved.

Superimpose histograms

Figure 4.

Author Manuscript

WRF 1.5km Filtered MSL Pressure

2014-06-25

Distribution of R_M

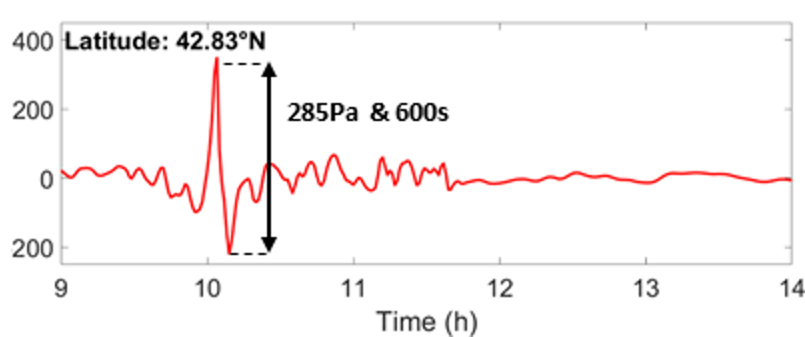
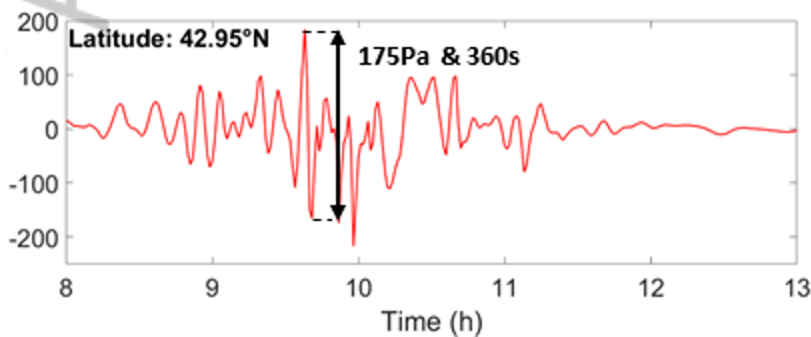
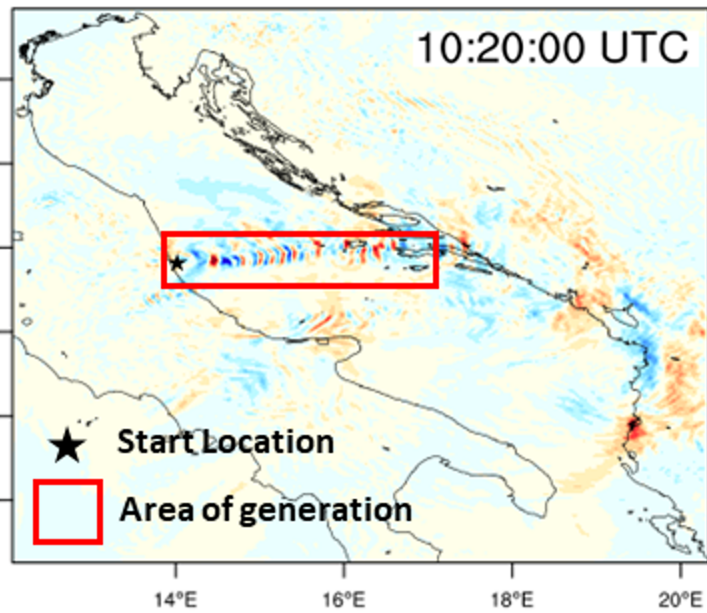
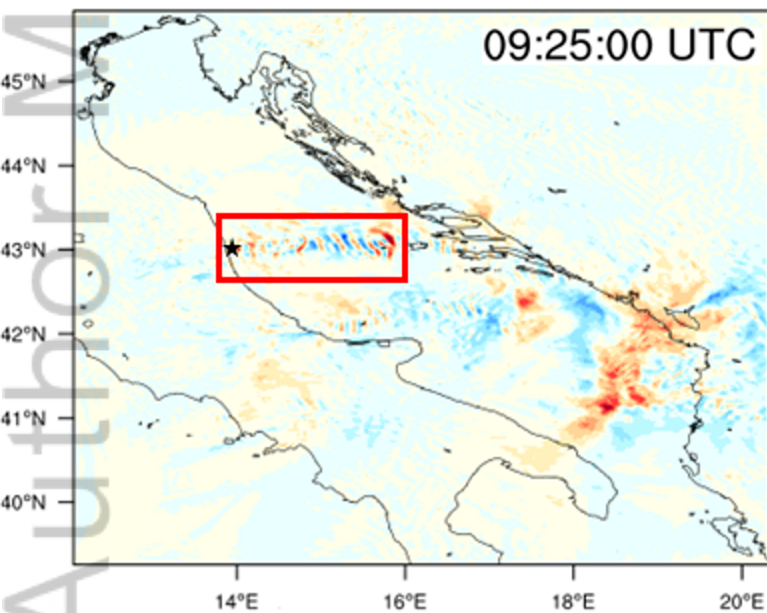
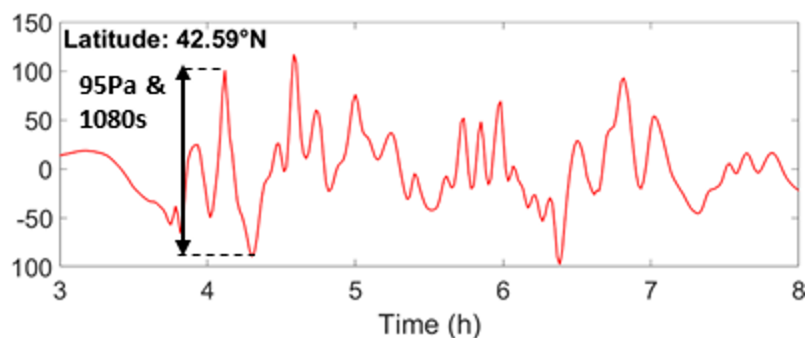
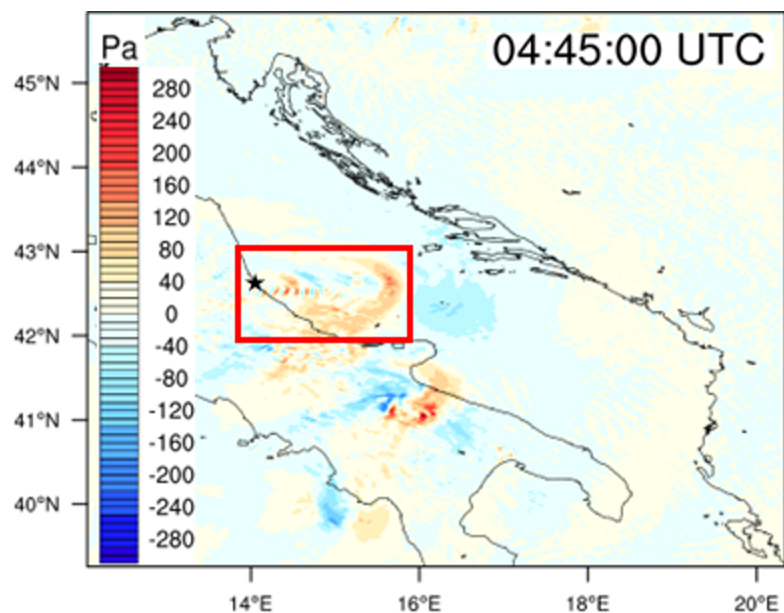
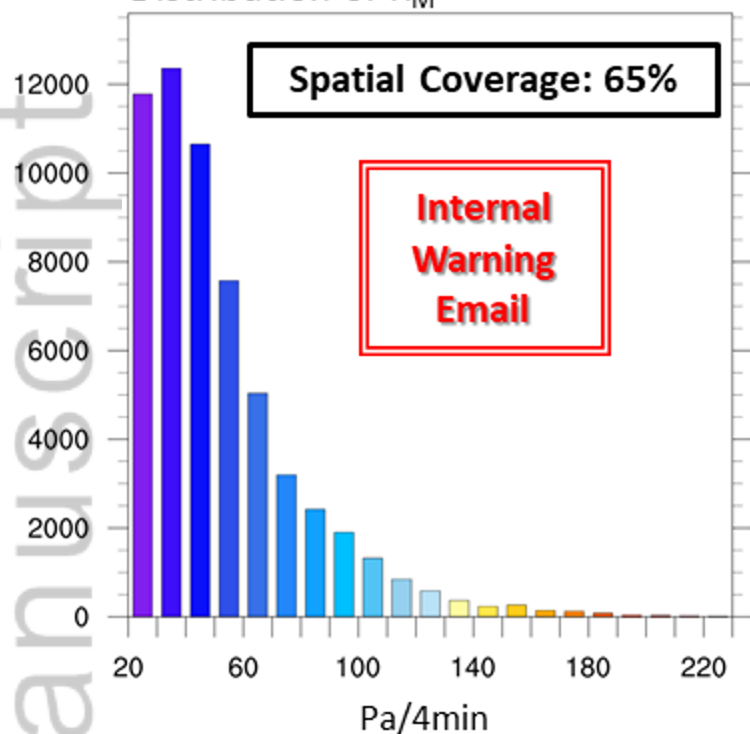


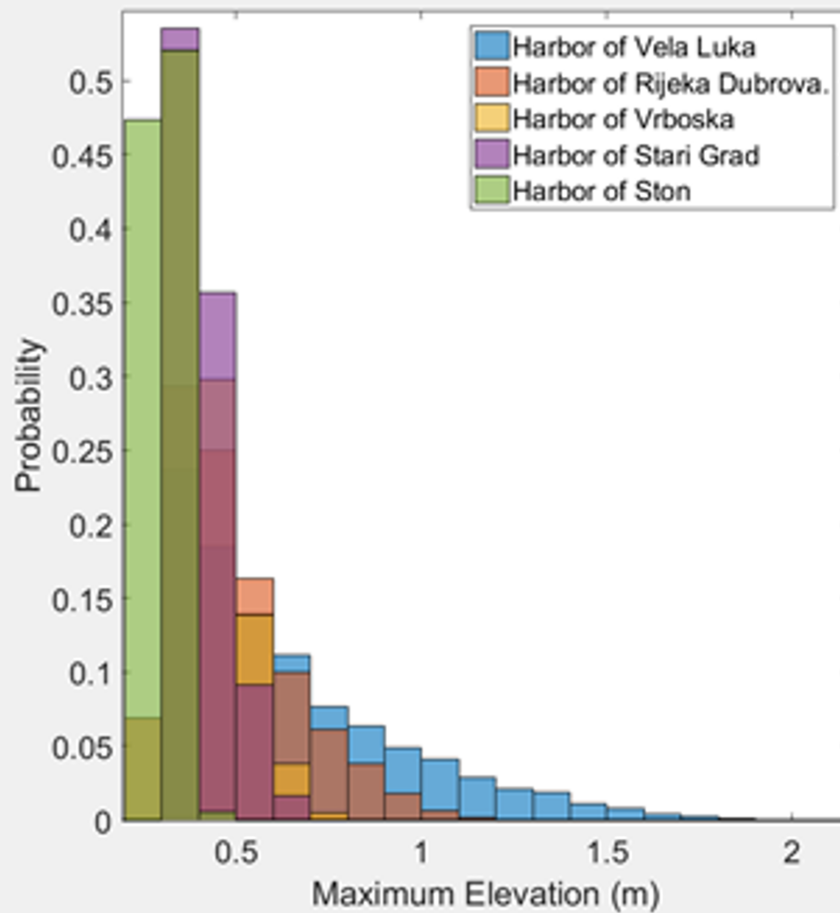
Figure 5.

Author Manuscript

Stochastic Surrogate Model of Meteotsunami Maximum Elevation

Select a Station

	Minimum	Maximum	
Start	42.34	43.20	<input type="checkbox"/> Imposed
Amplitude	60	320	<input type="checkbox"/> Imposed
Direction	-0.17	0.35	<input type="checkbox"/> Imposed
Speed (m/s)	15	40	<input type="checkbox"/> Imposed
Period (s)	300	1230	<input type="checkbox"/> Imposed
Width (m)	30000	54000	<input type="checkbox"/> Imposed
Date	2014 - 6 - 25		<input checked="" type="checkbox"/> Add Max. Tide
<input type="button" value="Download"/>			



Superimpose histograms

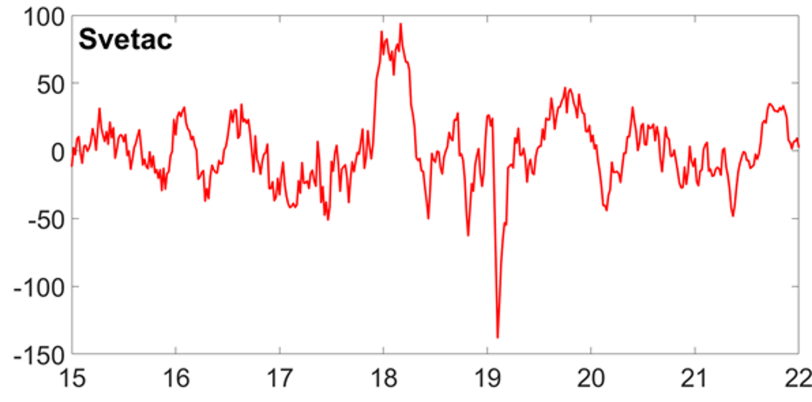
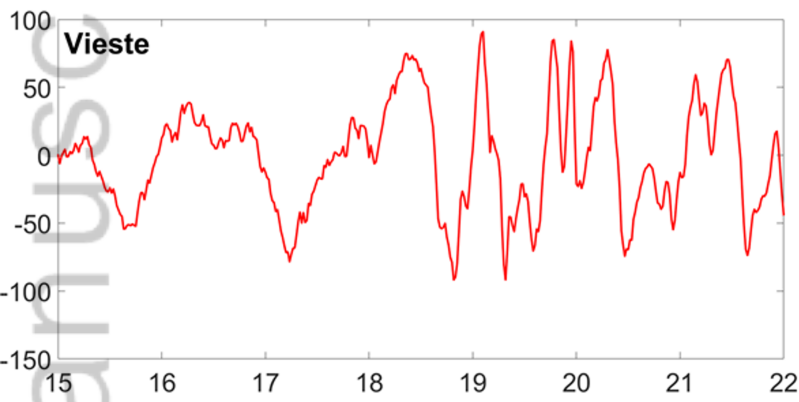
Vela Luka Rijeka d. Vrboska Stari Grad Ston

P(elevation>0.35)	88 %	79 %	71 %	73 %	7 %
P(elevation>0.45)	67 %	51 %	32 %	25 %	0 %
P(elevation>0.55)	51 %	30 %	10 %	5 %	0 %
P(elevation>0.65)	38 %	17 %	2 %	1 %	0 %
P(elevation>1.05)	12 %	1 %	0 %	0 %	0 %
P(elevation>1.55)	2 %	0 %	0 %	0 %	0 %
P(elevation>2.05)	0 %	0 %	0 %	0 %	0 %

Figure 6.

Author Manuscript

Microbarograph Filtered MSL Pressure (Pa)



Tide Gauge Filtered Sea-level (m)

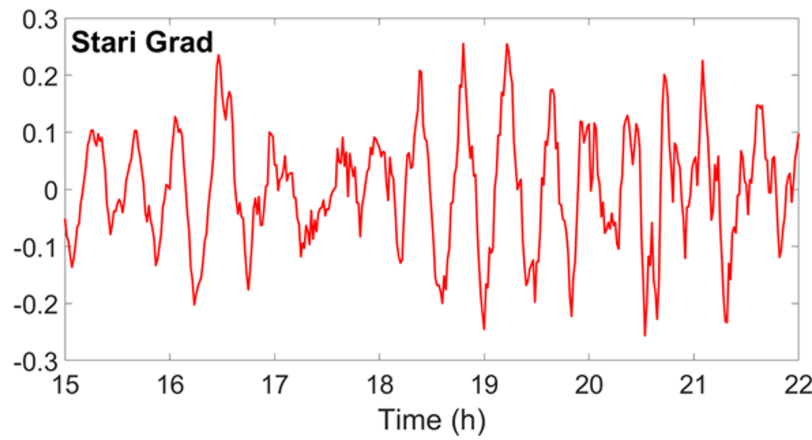
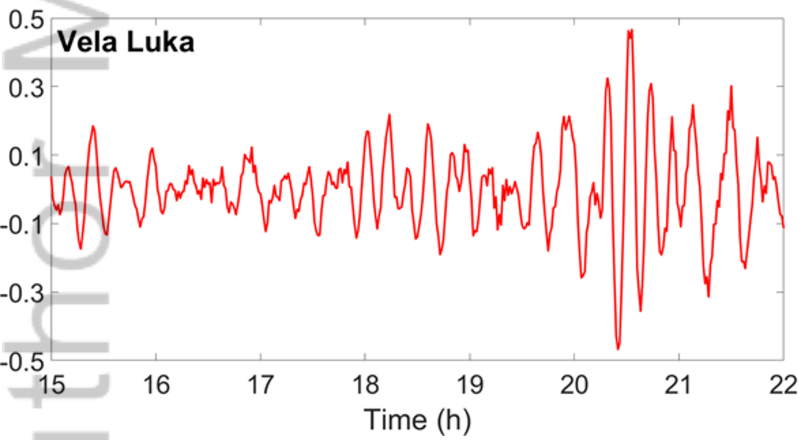
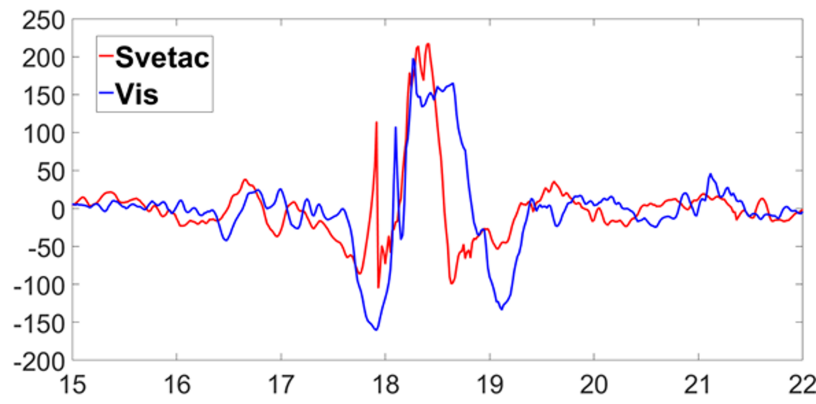
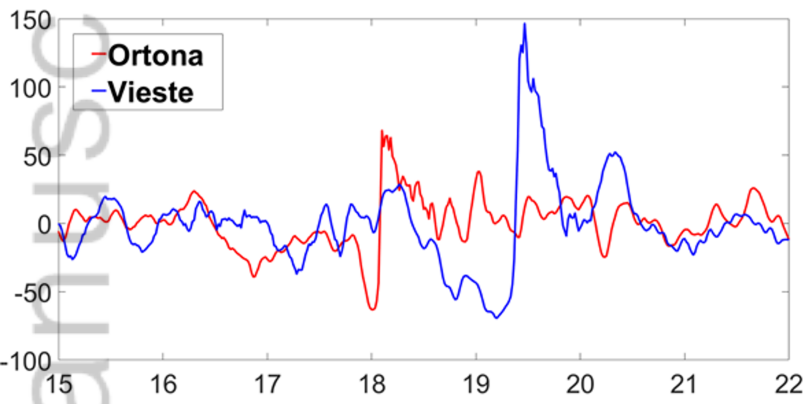


Figure 7.

Author Manuscript

Microbarograph Filtered MSL Pressure (Pa)



Tide Gauge Filtered Sea-level (m)

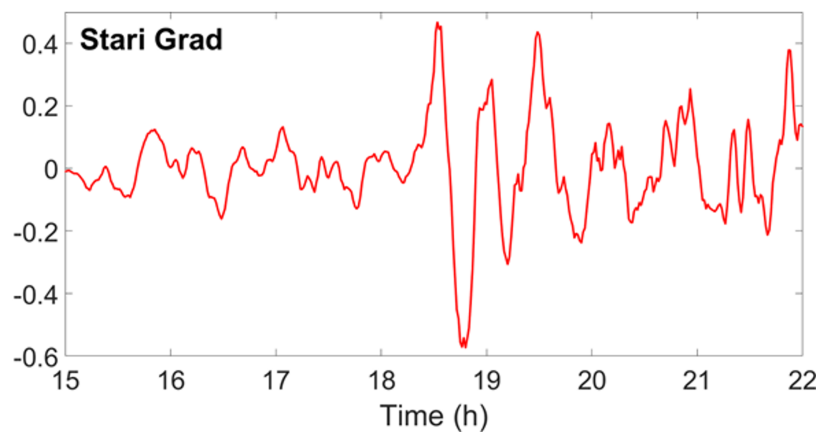
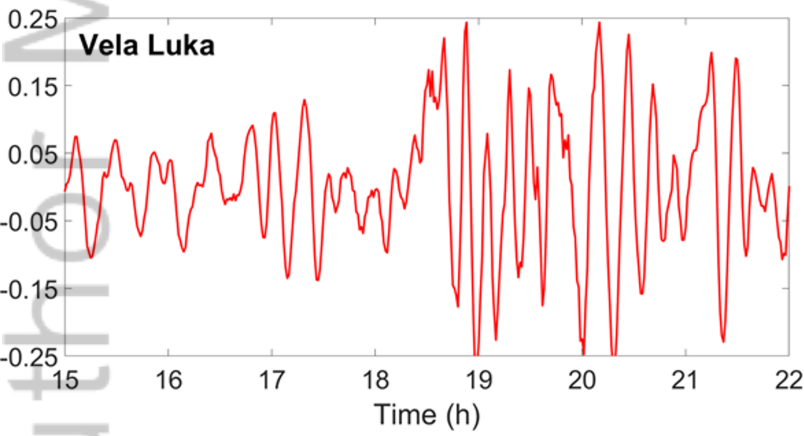
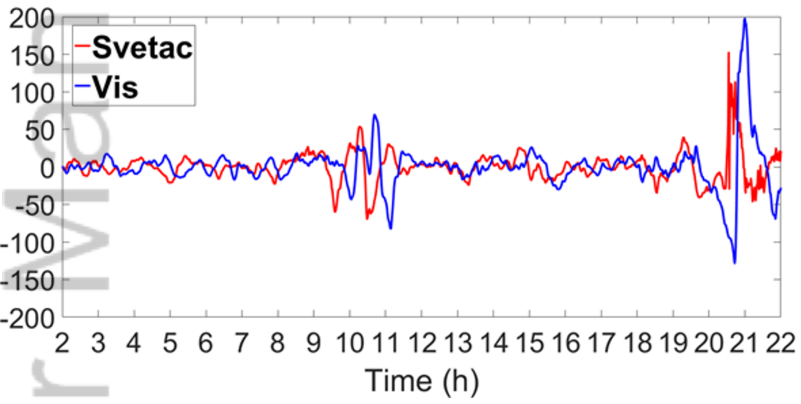


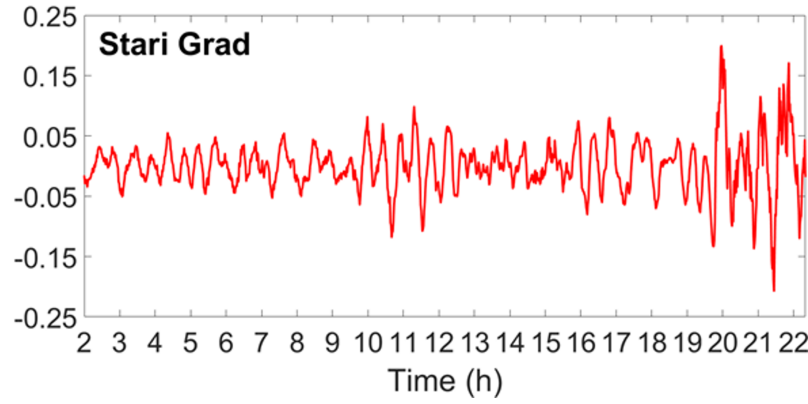
Figure 8.

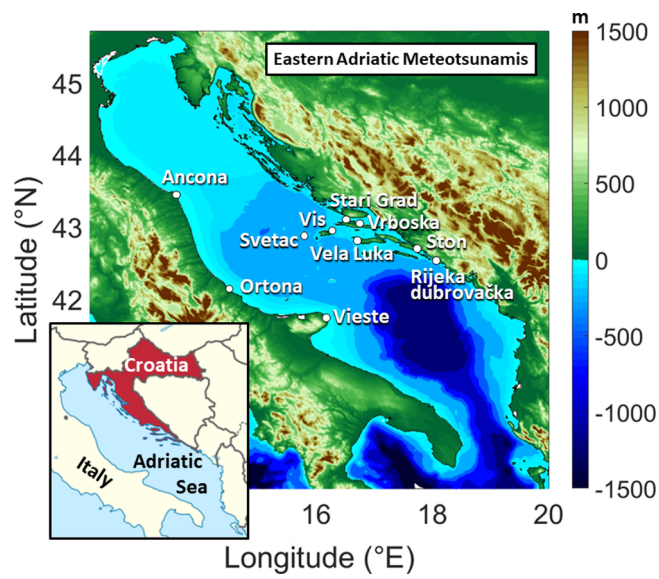
Author Manuscript

Microbarograph Filtered MSL Pressure (Pa)

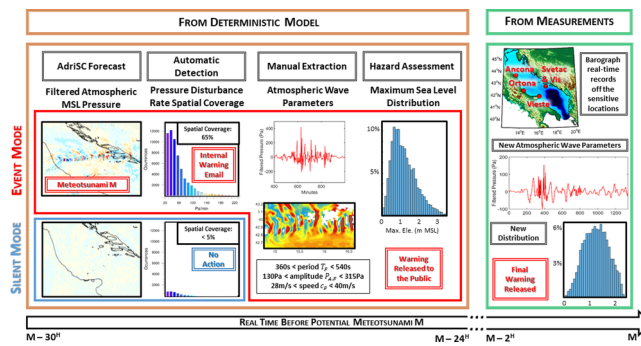


Tide Gauge Filtered Sea-level (m)

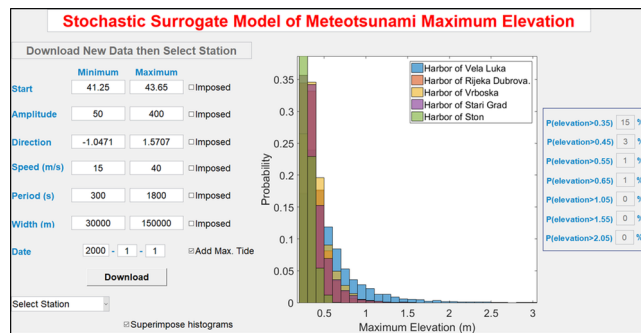




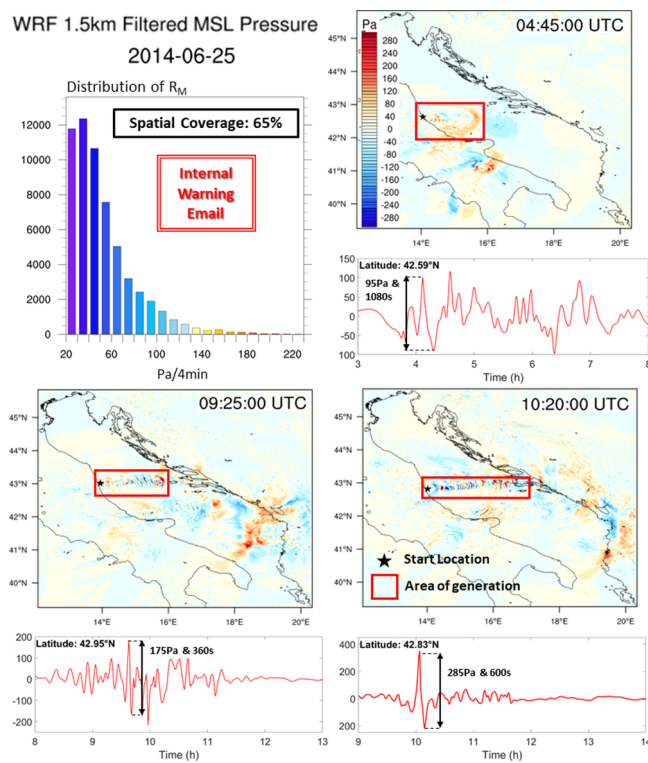
2019JC015574-f01-z-.tif



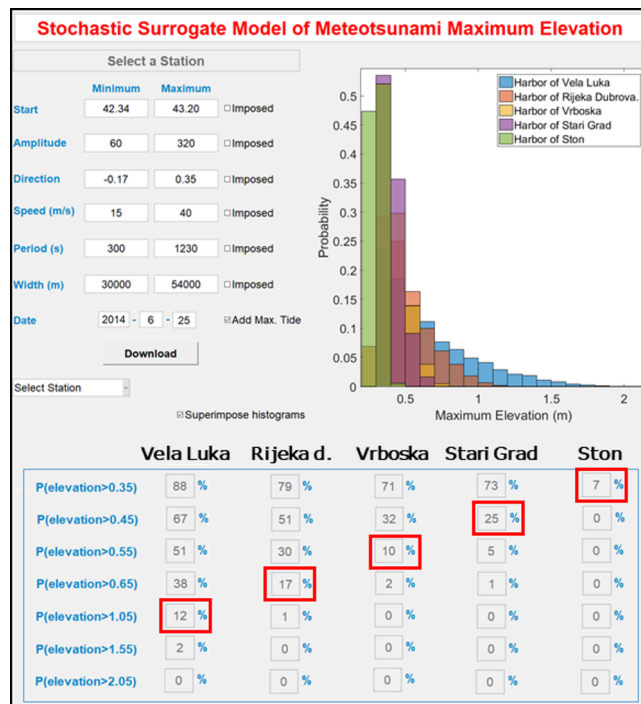
2019JC015574-f02-z.tif



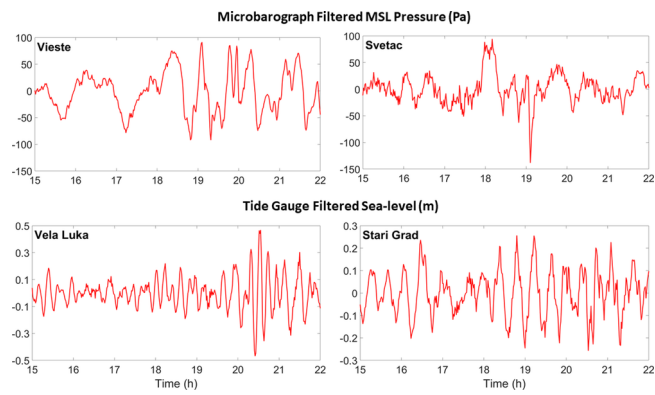
2019JC015574-f03-z-.tif



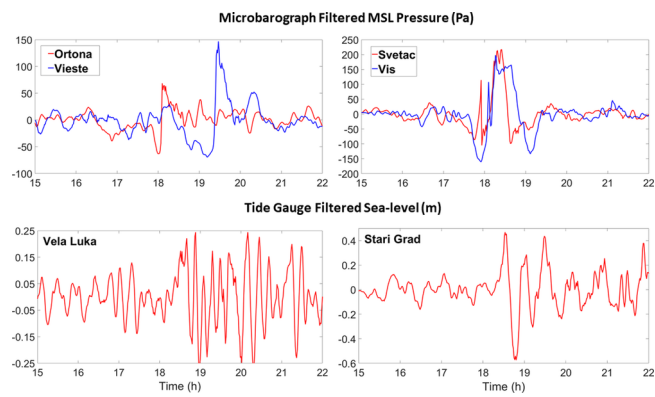
2019JC015574-f04-z.tif



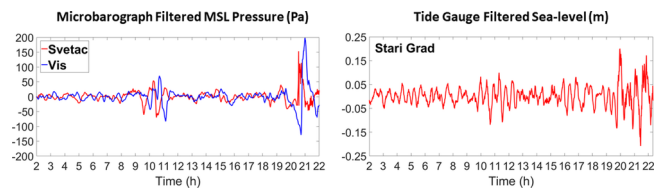
2019JC015574-f05-z-.tif



2019JC015574-f06-z-.tif



2019JC015574-f07-z-.tif



2019JC015574-f08-z-.tif