(1981) Evidence for chamber-dependent radical sources: impact on kinetic computer models for air pollution. Int. J. Chem. Kinet. 13, 735.

- Carter W. P. L., Atkinson R., Winer A. M. and Pitts J. N., Jr. (1982) Reply to "Comments on a smog chamber and modeling study of the gas phase NO_x-air photooxidation of toluene and the cresols". Int. J. Chem. Kinet. 14, 813.
- Joshi S. B., Dodge M. C. and Bulfalini J. J. (1982) Reactivities of selected organic compounds and contamination effects. *Atmospheric Environment* 16, 1301.
- Killus J. P. and Whitten G. Z. (1981) Comments on "a smog chamber and modeling study of the gas phase NO_x -air photooxidation of toluene and the cresols". Int. J. Chem. Kinet. 13, 1101.
- Kyle E. and Orchard S. W. (1977) The photolysis of methylglyoxal vapour at 436 nm. J. Photochemistry 7, 305.
- Nieboer H. and Duyzer J. H. (1978) Experimental and mathematical simulation of photochemical air pollution. In *Photochemical Smog Formation in the Netherlands* (Edited by R. Guicherit), p. 89, TNO-'s-Gravenhage.
- Parmenter C. S. (1964) Primary photochemical processes in glyoxal at 4358 A. J. Chem. Phys. 411, 658.

AN EDDY-CORRELATION MEASUREMENT OF NO₂ FLUX TO VEGETATION AND COMPARISON TO O₃ FLUX*

This paper is potentially very misleading, because the title suggests that they specifically measured the flux of NO_2 . Indeed, their hot molybdenum converter also has an efficiency of unity for HNO₃ (D. H. Stedman, private communication), so the measured flux was at *least* ($NO_2 + HNO_3$), and may have included other compounds as well.

In Champaign, Illinois, over grass, I measured HNO₃ fluxes by a profile method, and found them to be of the same magnitude as the NO₂ fluxes measured by Wesely *et al.* HNO₃ concentrations were frequently around 2 ppbv in the daytime. Therefore, it seems likely to me that a substantial fraction of the flux measured by Wesely *et al.*, was really HNO₃, and that their published V_d for NO₂ (0.6 cm s⁻¹) is significantly *larger* than the actual V_d for NO₂.

Thank you for bringing this to the attention of potential users of this data. My own measurements will be submitted for publication in the very near future.

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AUTHORS' REPLY

Huebert correctly identifies one of the flaws concerning interpretation by Wesely *et al.* (1982) of their data. As stated by Huebert, nitric acid vapor HNO₃ was indeed measured with an efficiency of unity and the resulting possible contribution to NO_x flux was not noted in the paper. The possibility that concentrations of HNO₃ can be rather high where the experiment was performed was brought to our attention only very recently. All values of NO₂ flux, deposition velocity and surface resistance estimated by Wesely *et al.* should be considered as those for NO_x rather than NO₂. For future eddy-correlation experiments, fast-response NO_x measuring system could be easily modified to allow determination of the

* Wesely M. L., Eastman J. A., Stedman D. H. and Yalvac E. D. (1982) Atmospheric Environment 16, 815–820.

flux both with and without HNO_3 , perhaps on alternate half hours, so that the effects of HNO_3 could be determined.

Another difficulty is the statement by Wesely *et al.* that NO_2 is removed efficiently by water due to rapid reactions that form nitrous and nitric acid. These reactions are actually very slow for pure water and the typically small values of $[NO_2]$ found in ambient atmospheric conditions (Lee and Schwartz, 1981). Thus for NO_2 to be taken up quickly by vegetation, some rapid reactions with plant materials are needed. Atmospheric ozone is subject to similar constraints but still has a large deposition velocity over lush vegetation during the daytime, apparently because of access via leaf stomatal opening to reactive substances in solution inside plant leaves. It is possible that NO_2 is also taken up rather rapidly when stomatal apertures are open but the most likely chemical reactions are not known with certainty.

Finally, it has been pointed out (O. T. Denmead, private communication) that the daytime fluxes and deposition velocities for NO_x and O₃ are slightly overestimated by Wesely *et al.* because of the neglect of small corrections to the eddy-correlation estimates needed for density effects due to heat and water vapor transfer (Webb *et al.*, 1980). Such corrections would increase the surface resistance by about 10% during the daytime but leave the main conclusion of Wesely *et al.* unaffected.

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REFERENCES

- Lee Y.-N. and Schwartz S. E. (1981) Evaluation of the rate of uptake of nitrogen dioxide by atmospheric and surface liquid water. J. geophys. Res. 86, 11, 971-11, 983.
- Webb E. K., Pearman G. I. and Luenig R. (1980) Correction of flux measurements for density effects due to heat and water vapor. Q. Jl R. met. Soc. 106, 85-100.
- Wesely M. L., Eastman J. A., Stedman D. H. and Yalvac E. D. (1982) An eddy-correlation measurement of NO_2 flux to vegetation and comparison to O_3 flux. Atmospheric Environment 16, 815–820.

BACKGROUND POLLUTANT MEASUREMENTS IN AIR MASSES AFFECTING THE EASTERN HALF OF THE UNITED STATES—I. AIR MASSES ARRIVING FROM THE NORTHWEST

Measurements of background concentrations of particulate and gaseous species in the atmosphere, such as those recently reported by Kelly *et al.* (1982), provide an important reference point for understanding anthropogenic pollutants. For this reason, several points in this paper, with respect to particulate organic carbon, require comment.

Kelly et al. (1982) compared organic carbon data for a U.S. continental rural site in South Dakota with extractable particulate organic matter measured at a remote site in Barrow, Alaska, in March 1979 (Daisey et al., 1981). In comparing the South Dakota measurements to the Barrow

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