

COMMENT ON THE PAPER "ON THE INFLUX OF SMALL COMETS INTO THE EARTH'S UPPER ATMOSPHERE II. INTERPRETATION" BY L. A. FRANK, J. B. SIGWARTH AND J. D. CRAVEN

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Abstract. The suggestion that the Earth is being bombarded by cometoids that subject it to an influx of 3×10^{11} water molecules/cm²-sec encounters severe aeronomical problems. There is a serious disposal problem for water molecules deposited near 100 km. The implications for Venus and Mars also present serious problems. It is argued that the proposable is untenable.

In an effort to explain transitory decreases in ultraviolet dayglow emission rates that they observed from the DE-1 spacecraft (Frank, et al., 1986a) L. A. Frank, J. B. Sigwarth and J. D. Craven (1986b) have made a startling suggestion. They propose that water is injected into the thermosphere by a stream of comets and absorbs the radiation. They require a flux of water of 3×10^{11} molecules/cm²-sec averaged over the entire earth. The authors argue that the water and its dissociation products will penetrate to the lower thermosphere before they come to rest. There are a large number of problems with the proposed mechanism.

The authors recognize that the flux of water — 1.34×10^{15} gm/yr — is so large that it would supply all of the water in the oceans, cryosphere, and crust — 2×10^{24} gm — in 1.5 billion years if the inflow rate is steady. Since disposing of the oxygen associated with two extra oceans would require that the interior of the earth become oxidized to a depth of several hundred kilometers to a much greater degree than it is (0.5% Fe₂O₃ to 10% FeO), a steady influx of comets over geological time is ruled out. Consequently, Frank, et al. (1986b) propose that the comet showers are episodic. They suggest that a periodic perturbation of the Oort cloud once every 250 million years sends clouds of small comets into the inner solar system. Thus the Earth's ocean would have been acquired in about 15 steps over 4.5 billion years.

Although current wisdom favors outgassing of most of the ocean soon after the Earth was accreted, almost all of the supporting evidence for this view bears on outgassing of volatiles in general. It is presumed that water was one of the volatiles (in fact the quintessential volatile) that was originally accreted and then outgassed. Data concerning ocean volume and sea level are too imprecise and limited in time to rule out accession of the ocean in 6.6% increments once every 250 million years. Although there is strong evidence that enough liquid water to create sedimentary rocks was present on the Earth's surface 3.8 billion years ago (Holland, 1984), it is difficult to say how

much liquid water was needed to produce them. That the material from which the Earth accreted was rich in water is attested to by the great abundance of hydrated minerals in the interior. The notion that enough water to provide the planet with an ocean was not released during accretion is a strange one, but possibly not one that can be rigorously excluded. It should be noted, for example, that it may be difficult to explain why the D/H ratios in sea water is elevated so much above the general solar system value if the comets in question were accreted beyond the orbit of Saturn. But until the isotope ratio is measured for comets this argument, while strong, is not decisive.

It is difficult to see how Venus and Mars could avoid this cometary bombardment and thus be accreting at about the same rate per unit area as Earth. The escape flux of hydrogen from the three planets is about 10^7 /cm²-sec from Venus (Kumar, et al., 1983), 3×10^8 /cm²-sec from Earth (Donahue, 1977), 1.8×10^8 /cm²-sec from Mars (McElroy and Donahue, 1972; Anderson and Hord, 1973; Liu and Donahue, 1976). In the case of Venus there are numerous measurements of the mole fraction of water vapor in the lower atmosphere (vonZahn, et al., 1983). Taking 200 ppm by volume to be a generously large value for this quantity would put the water content of the atmosphere at about 5×10^{22} molecules/cm². The presumed cometary supply would double this quantity in 5500 years. So there is a severe disposal problem on Venus. If the surface of Venus is assumed to be somehow a sink for 3×10^{11} water molecules/cm²-sec the oxidation rate would have to be about 0.1 mm/year over the entire planet if the FeO concentration in material freshly exposed on the crust of Venus is the same as the terrestrial value. More troublesome is the consequence that the surface would be highly oxidized. This condition appears to be incompatible with the presence of sulfur gases in the atmosphere (Lewis and Prinn, 1984).

If a significant variation with heliocentric distance of cometary lifetime in the inner solar system is invoked to reduce the influx at Venus so that it becomes compatible with the low hydrogen escape rate, the problem with Mars would become even more serious. The question is how to reconcile an influx of water greater (probably much greater) than 3×10^{11} molecules per cm²-sec at Mars, an escape flux of only 1.8×10^8 per cm²-sec with the absence of a reservoir of water on Mars that contains more than the Martian equivalent of a terrestrial ocean. The proposed cometary source would have supplied a large fraction of an ocean during the past few hundred million years. If that water is sequestered, where is it? If not, how can an escape scenario be developed in which the present escape flux is three or four orders of magnitude below the value

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required to get rid of the hydrogen associated with the cometary supply of water. Obviously, the problem of the oxygen sink also needs to be addressed.

There are even deeper problems associated with producing an escape flow of hydrogen at Mars that is of the order of 5×10^{11} atoms/cm²-sec (or much larger if there is a sizeable heliocentric variation). One of these is a fundamental energy limit on the rate at which hydrogen powered by extreme ultraviolet heating can escape from a planet (Watson, et al., 1981). This limit is 9×10^{11} atoms/cm²-sec for Mars and applies to a "blowoff" form of escape in which the atmosphere is dominated by water vapor and its dissociation products. The limit is set by the rate at which heat can be supplied to offset the adiabatic cooling of the rapidly expanding gas. In practice the actual flux would fall below this limit for a number of reasons. Hence, if the rate of water influx at Earth averaged over the lifetime of the planet is 10^{11} molecules/cm² sec, the influx at Mars could not exceed five times that at Earth at the extreme, and in practice probably could not even exceed the terrestrial influx rate itself without causing this fundamental constraint to imply the presence of something like the equivalent of a terrestrial ocean of water somewhere on Mars.

One reason that the escape flux of hydrogen would not attain the theoretical energy limit is that there are not enough solar photons available at Mars to dissociate the incoming water. The diurnally averaged solar flux of photons whose energy is above the dissociation limit of water is only 1.5×10^{11} photons/cm² sec. Hence no more than about 3×10^{11} hydrogen atoms/cm² sec can be generated and escape, and the water must accumulate if it arrives at a rate greater than 1.5×10^{11} molecules/cm²-sec.

Frank et al. (1986b) also recognize that there is a difficulty disposing of the water molecules and their dissociation products when they arrive in the lower thermosphere of the Earth. If they are to be transported away by vertical molecular and eddy diffusion the flux required (Hunten and Donahue, 1976; Donahue, 1977) is

$$\phi_i = \frac{b_i f_i}{H} - (K + D_i) n \frac{df_i}{dz} \quad (1)$$

where D_i is the molecular diffusion coefficient, n is the total atmospheric number density, b_i is the molecular diffusion parameter $D_i n$, f_i the mixing ratio for total hydrogen, H the atmospheric scale height, and K the eddy diffusion coefficient. If it is assumed that most of the hydrogen is transported downward and very little upward, and K is taken to be 10^6 cm²/sec at 100 km, a flux of 6×10^{11} /cm²-sec requires a mixing ratio gradient of about 2000 ppm/km at 100 km, decreasing to 150 ppm/km at 50 km. The mixing ratio for hydrogen in all forms near 100 km would need to be of the order of 5%. The assumption of diffusive equilibrium at higher altitudes then calls for atomic hydrogen densities between 10^{11} and 10^{12} /cm³ at the exobase, more than a factor of a million higher than those that are observed. This would result in an escape flux due to Jeans escape alone of about 10^{14} /cm²-sec, which violates the basic

assumption. Therefore, much of the flux of hydrogen would have to go upward. Limiting flux ($b_i f_i / H$) then requires that f_i be 3×10^{-2} . If the downward flux is to be vanishingly small this mixing ratio must be maintained throughout the mesosphere and stratosphere. Such large amounts of hydrogen compounds would remove all of the ozone from the middle atmosphere. The frost point would be exceeded globally near the mesopause and a global cloud layer blanket the Earth. Furthermore, the observed mixing ratios for "total hydrogen" are close to 15 ppm, a factor of 5×10^{-4} smaller than those that would be called for. At higher altitudes hydrogen densities would be needed that are larger than those observed by a factor of 2000. Finally the upward flux of hydrogen has been inferred by Brieg, et al., (1976) from H^+ , O^+ and O densities. The value obtained was about 3×10^8 /cm²-sec, a factor of 2000 smaller than the flux required by the cometary source. These difficulties would be overcome if K were several orders of magnitude larger. Not only is this difficult to justify physically, it would shift the homopause upward by about 50 km and so is precluded.

To solve these problems Frank, et al. (1986b) suggest that advection in the middle atmosphere with a vertical velocity of about 2 m/sec at 90 km may transport the hydrogen compounds downward. It is not at all clear what kind of global circulation pattern is being proposed. In any event the vertical velocity needed is more like 5 m/sec than 2 m/sec because the total hydrogen mixing ratio is about 15 ppm rather than 30 ppm. These speeds are extraordinarily large. They would lead to transport lifetimes $\sim z/w$ where z is a vertical scale, for which 5 km is an appropriate value, of 1000 sec at 90 km, 4500 sec at 80 km, rising to 2.3×10^5 sec at 50 km. Such short transport times would change the entire photochemical structure of the region. In particular they would move the atomic oxygen maximum down toward 85 km. The associated heating would also change the thermal structure of the mesosphere drastically. But the most serious problem with the proposal is that measured vertical winds at these altitudes measure at most 25 cm/sec and usually are far smaller (Balsley and Riddle, 1984). These measurements have been made over a period of several years at Poker Flats, Alaska. Other, less sensitive, techniques employed at lower latitudes, nevertheless set upper limits to the vertical mesospheric wind well below the values needed here.

I conclude that water influxes of the magnitude invoked by Frank, et al. (1986b) are aeronomically unacceptable if the water is deposited, as they assert, in the lower thermosphere. There is no way to dispose of the water without seriously violating observational constraints. Transport by eddy and molecular diffusion would require mixing ratios for water and its products orders of magnitude too high. Transport by vertical advection, keeping the mixing ratio acceptably low, calls for vertical winds two orders of magnitude too large. If the water vapor were to somehow travel coherently downward as far as 55 km the mixing ratio of

water produced by the dispersion of the jets would be some tens of ppm, still too high but perhaps not outrageously so. Bevilacqua, et al. (1983) and Olivero et al (1986) measure 6 ppm by volume at 55 km. (It should be kept in mind that the source of water in the stratosphere is mainly oxidation of CH₄ and not surface water.) It seems possible that these impulses occurring at any one place with a frequency of about once per week should produce observable signatures for Doppler radars such as the Mesosphere, Stratosphere, Troposphere (MST) radar at Poker Flats.

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References

- Anderson, D. E., Jr. and C. W. Hord, Mariner 6 and 7 Ultraviolet Spectrometer Experiment: Analysis of Hydrogen Lyman Alpha Data, *J. Geophys. Res.*, **28**, 6666-6673, 1971.
- Balsley, B. B. and A. C. Riddle, Monthly Mean Values of the Mesospheric Wind Field over Poker Flat, Alaska, *J. Atmos. Sci.*, **41**, 2368-2375, 1984.
- Bevilacqua, R. M., J. J. Olivero, P. R. Schwartz, C. J. Gibbins, J. M. Bologna and D. J. Thacker, An Observational Study of Water Vapor in the Mid-latitude Mesosphere using Ground-based Microwave Technique, *J. Geophys. Res.*, **88**, 8523, 1983.
- Breig, E. L., W. B. Hanson, J. H. Hoffman, and D. C. Kayser, *In situ* Measurements of Hydrogen Concentration and Flux Between 160 and 300 km in the Thermosphere, *J. Geophys. Res.*, **81**, 2677-2685, 1976.
- Donahue, T. M., Hydrogen. In, The Upper Atmosphere and Magnetosphere, Studies in Geophysics, National Academy of Sciences, 72-84, 1977.
- Frank, L. A., J. B. Sigwarth, and J. D. Craven, On the Influx of Small Comets into the Earth's Upper Atmosphere I. Observations, *Geophys. Res. Lett.*, **13**, 303-306, 1986a.
- Frank L. A., J. B. Sigwarth, and J. D. Craven, On the Influx of Small Comets into the Earth's Upper Atmosphere II. Interpretation, *Geophys. Res. Lett.*, **13**, 307-310, 1986b.
- Holland, H. D. The Chemical Evolution of the Atmospheres and Oceans, Princeton University Press, 1984.
- Hunten, D. M. and T. M. Donahue, Hydrogen Loss from the Terrestrial Planets, *Ann. Rev. Earth Planet. Sci.*, **4**, 265, 1976.
- Kumar, S., D. M. Hunten, and J. B. Pollack, Nonthermal Escape of Hydrogen and Deuterium from Venus, *Icarus*, **52**, 1982.
- Lewis, J. S. and R. G. Prinn, Planets and their Atmospheres, Academic Press, Inc., 1984.
- Liu, S. C. and T. M. Donahue, The Regulation of Hydrogen and Oxygen Escape from Mars, *Icarus*, **28**, 231-246, 1976.
- McElroy, M. B. and T. M. Donahue, Stability of the Martian Atmosphere, *Science*, **177**, 986-988, 1972.
- Olivero J. J., J. J. Tsou, C. L. Hale, and R. G. Joiner, Solar Absorption Microwave Measurement of the Upper Atmospheric Water Vapor, *Geophys. Res. Lett.* **13**, 197-200, 1986.
- Watson, A. J., T. M. Donahue and J. C. G. Walker, The Dynamics of a Rapidly Escaping Atmosphere: Applications to the Evolution of Earth and Venus, *Icarus*, **48**, 150-165, 1981.

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