

1 Mars Nightside Electrons Over Strong Crustal Fields

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D R A F T

2 Abstract.

3 We investigated seven years worth of data from the electron reflectome-
4 ter and magnetometer aboard Mars Global Surveyor to quantify the depo-
5 sition of photoelectron and solar wind electron populations on the nightside
6 of Mars, over the strong crustal field region located in the southern hemi-
7 sphere. Just under 600,000 observations, each including energy and pitch an-
8 gle distributions, were examined. For solar zenith angles (SZA) less than 110° ,
9 photoelectrons have the highest occurrence rate; beyond that, plasma voids
10 occur most often. In addition, for $SZA > 110^\circ$, energy deposition of electrons
11 mainly occurs on vertical field lines with median pitch angle averaged en-
12 ergy flux values on the order of 10^7 - 10^8 eV cm⁻² s⁻¹. The fraction of down-
13 ward flux that is deposited at a given location was typically low (16% or smaller),
14 implying the majority of precipitated electrons are magnetically reflected or
15 scattered back out. The average energy of the deposited electrons is found
16 to be 20-30 eV, comparable to typical energies of photoelectrons and unac-
17 celerated solar wind electrons. Median electron flux values, from near ver-
18 tical magnetic field lines past solar zenith angle of 110° , calculated in this
19 study produced a total electron content of 4.2×10^{14} m⁻² and a correspond-
20 ing peak density of 4.2×10^3 cm⁻³.

1. Introduction

21 Unlike the Earth, which has a global magnetic field, Mars has localized crustal magnetic
22 fields [Acuña *et al.*, 1998, 2001]. These crustal fields complicate the interaction with the
23 interplanetary magnetic field (IMF), resulting in a sophisticated magnetic topology [e.g.,
24 Mitchell *et al.*, 2001; Nagy *et al.*, 2004; Brain *et al.*, 2007; Liemohn *et al.*, 2007]. The
25 strongest crustal magnetic fields are located in the southern hemisphere [e.g., Connerney
26 *et al.*, 2001]. Brain *et al.* [2007] used electron pitch angle distributions (PAD) from Mars
27 Global Surveyor (MGS) to infer the magnetic field topology of Mars. On the nightside,
28 Brain *et al.* [2007] classified two-sided loss cones (trapped populations) and plasma voids
29 (locations where the observations are at or near the instrument background level), as
30 indicators of closed magnetic field lines corresponding to the Martian crustal fields. In
31 contrast, nightside one-sided loss cones often are related to open/draped field lines, sug-
32 gesting connection to the IMF, allowing planetward streaming electrons on one end and
33 atmospheric absorption on the other. In particular, radial crustal fields form magnetic
34 cusps generally located between magnetic loop structures. These cusps are ideal locations
35 for solar wind/magnetosheath electrons to precipitate [Mitchell *et al.*, 2001; Liemohn *et*
36 *al.*, 2003; Sfaeini *et al.*, 2007; Xu *et al.*, 2014b]. Němec *et al.* [2010] found that the
37 occurrence rate of nightside ionosphere patches observed by Mars Express is four times
38 larger over cusp regions rather than where the magnetic field is horizontal.

39 Day-to-night plasma transport and electron precipitation are both important mecha-
40 nisms for the creation of the nightside ionosphere of Mars [e.g., Fox *et al.*, 1993]. Verigin
41 *et al.* [1991] used an analysis of HARP measurements from PHOPOS 2 to propose that

42 a characteristic omnidirectional electron flux of $\sim 10^8 \text{ cm}^{-2} \text{ s}^{-1}$ is sufficient to create the
43 nightside ionosphere. *Haider et al.* [1992] confirmed this to be true given that the elec-
44 trons are precipitating. *Liemohn et al.* [2007] showed that some closed field lines straddle
45 the terminator, allowing photoelectrons to precipitate into the nightside ionosphere. The
46 nightside precipitation variability due to solar wind was investigated by *Lillis et al.* [2013]
47 and found that plasma voids vary significantly with solar wind pressure.

48 *Haider et al.* [2007] calculated that solar wind electron precipitation creates a peak ion
49 layer at $\sim 130 \text{ km}$. *Němec et al.* [2011] observed enhanced ionization over magnetic cusp
50 regions. This localized ionization, especially when enhanced, can cause density gradients
51 up to $600 \text{ cm}^{-3} \text{ km}^{-1}$ [*Fillingim et al.*, 2010]. They found that density gradients can lead
52 to plasma transport resulting in currents and Joule heating. *Němec et al.* [2010] concluded
53 that most of the nightside ionosphere has a peak electron density lower than $5 \times 10^3 \text{ cm}^{-3}$
54 and that over strong crustal field regions, the peak density does not vary with solar zenith
55 angle (SZA), implying electron precipitation is the main mechanism for the formation of
56 nightside ionosphere over such regions.

57 Excitation, the cause for aurora on the nightside, is another consequence of precipitat-
58 ing electrons. Here on Earth there are primarily two types of aurora, diffuse and discrete,
59 a review of terrestrial aurora can be viewed in *Swift* [1981]. The diffuse aurora is caused
60 by scattering of electrons in the plasma sheet into the loss cone that precipitate into our
61 atmosphere [e.g., *Ashour-Abdalla and Kennel*, 1978]. Field aligned acceleration mecha-
62 nisms are the primary cause for the discrete aurora. Characteristic energy of precipitating
63 electrons on Earth are $\sim 2\text{-}3 \text{ keV}$, but auroral electrons can range from $0.5\text{-}40 \text{ keV}$. Au-
64 rora on Venus are thought to be produced by electrons with energies less than 300 eV

65 [*Fox and Stewart*, 1991]. The first reports of aurora emission on Mars came from the
66 ultraviolet spectrometer (SPICAM) onboard Mars Express (MEX) [*Bertaux et al.*, 2005].
67 They reported aurora different from any other seen in our solar system. The Martian
68 aurora reported was controlled by the crustal fields and was localized and highly con-
69 centrated. *Leblanc et al.* [2008] found a strong connection between auroral events and
70 magnetic cusps. Several others have investigated auroral electron spectra from MGS [e.g.
71 *Brain et al.*, 2006] and Mars Express [e.g., *Lundin et al.*, 2006a, b]. *Leblanc et al.* [2006]
72 and *Dubinin et al.* [2008a] proposed that the auroral electron energy distribution seen
73 by Mars Express peaked at a few tens of eV. The accelerated electrons in these papers
74 have enough energy to produce auroral emissions. Recently, *Soret et al.* [2015] used a
75 Monte-Carlo model to find that the height of observed auroral emission from SPICAM
76 could be recreated with electrons between 50-1000 eV. *Gérard et al.* [2015] did not find a
77 correlation between the observed UV aurora and downward electron flux measurements
78 at the Mars Express altitude. This could be due to the field lines where the aurora is
79 occurring are not vertical but tilted so the spacecraft is not measuring them. Another
80 possibility is an acceleration process occurring below the spacecraft.

81 *Brain et al.* [2005] investigated magnetosheath plasma intrusions below 400 km on the
82 dayside and analyzed its dependence on IMF orientation and on seasons. They found that
83 crustal magnetic fields raise the magnetic pileup boundary and that cusps allow sheath
84 electrons to enter the atmosphere. *Brain et al.* [2007] also examined the dayside magnetic
85 field structure using PAD. They recorded isotropic distributions near strong crustal field
86 regions. A new method of separating photoelectrons from solar wind electrons (classifying
87 by energy spectra rather than PAD) was used by *Xu et al.* [2014b] in their statistical

88 study of dayside solar wind precipitation on the magnetic cusps. They investigated the
89 occurrence rates of both populations and dependence on solar zenith angle, magnetic
90 elevation angle, and seasonal variation. They compared the solar wind energy deposition
91 to solar EUV flux input and found it was 0.1%-2% of the solar EUV flux.

92 This study is concerned with the nightside ($\text{SZA} > 90^\circ$) of Mars and will use the same
93 population classification method used by *Xu et al.* [2014b]. The net energy deposition
94 of electrons over the strong crustal field region will be investigated, something which has
95 not been calculated yet on the nightside, as a function of solar zenith angle and magnetic
96 elevation angle. The fractional deposition rate and the average energy of deposited parti-
97 cles will be analyzed. Finally, the consequences of electron deposition, i.e excitation and
98 ionization, will be examined.

2. Methodology

99 The electron reflectometer (ER) onboard MGS recorded superthermal electron angular
100 distributions ranging from energies of 10 eV to 20 keV. The field of view spanned 360°
101 $\times 14^\circ$ and was divided into sixteen 22.5° sectors. Measurements from sectors 8, 10,
102 and 11 were discarded due to high fluxes being recorded frequently, regardless of field line
103 orientations [*Xu et al.*, 2014a] With the magnetic field information from the magnetometer
104 (MAG), these angular distributions can be converted into PAD's [*Mitchell et al.*, 2001].
105 Important to note is the uneven sampling by the ER due to the orientation of the magnetic
106 field with respect to the ER. If the magnetic field was perpendicular to the plane of the
107 field of view, field aligned pitch angles were not sampled [see Fig. 9 of *Liemohn et al.*,
108 2006]. If the magnetic field was parallel to the plane than the entire PAD was sampled.
109 "Modified pitch angles" were used similar to *Xu et al.* [2014b] to identify if the electrons

110 were moving toward or away from the planet. If the magnetic field is positive (directed
111 away from the planet) then the pitch angles were flipped, for example, pitch angles of 10°
112 becomes pitch angles of 170° . In the end, pitch angles are organized into 10 degree bins
113 with 0° - 90° directed towards the planet (i.e. downward) and 90° - 180° directed away from
114 the planet (upward). From this point forward in the text, all pitch angles are modified.

115 MGS orbit was locked to 0200/1400 local time at an altitude of 405 ± 36 km. All
116 measurements with a solar zenith angle of less than 90° were excluded to limit this study
117 to the night side. *Xu et al.* [2014a, b] conducted the analogous studies for the dayside
118 strong crustal field regions. Another stipulation added was to focus this study on the
119 strong crustal fields located in the southern hemisphere, as the magnetic cusps in this
120 region are more likely to allow electron precipitation to enter the thermosphere below
121 200 km and cause enhanced ionization and excitation. The data was filtered to only
122 include data within a box spanning from 160° to 200° east longitude and 30° to 70° south
123 latitude. To make sure the crustal fields and not piled up IMF were being measured, *Xu*
124 *et al.* [2014b] used a minimum magnetic field strength of 35 nT in their dayside study. On
125 the night side, the IMF does not build up, therefore the minimum magnetic field strength
126 used is 5 nT (removal of 74 observations), to ensure pitch angle accuracy. It is still
127 possible, however, that non crustal field lines are included in the remaining observations.
128 Overall, this yielded $\sim 600,000$ observations over seven years of collected data, each with
129 pitch angle and energy distributions.

130 Many energy spectra had values under the background flux levels or resembled the
131 background curve. The night side does not have a constant plasma source term, therefore
132 plasma voids are common. Plasma voids occur on closed magnetic field lines that have

133 lost their photoelectron population and are isolated from solar wind plasma [Mitchell
134 *et al.*, 2001]. Two conditions were used on each energy spectrum to filter out voids.
135 The 190, 115, 79, 61, 47, and 36 eV energy channels were chosen and if more than
136 one of these flux values fell below the background level then it was classified as a void
137 (Figure 1a). If the flux is close to the background, the signal-to-noise of the measurement
138 is low, making the observation unreliable. Even if they are distinguishable from the
139 background, the fluxes will be too low to cause significant impact to the Martian nightside
140 ionosphere/thermosphere. However, many spectra existed that have the same shape as
141 the background curve similar to Figure 1b. The “cliff” in the background curve is from
142 a change in the instrument geometry factor at higher energies. This type of spectra does
143 not meet our above criteria but is also suggestive of a void. To filter such observations
144 out, the same six energy channels were chosen and the measured flux was divided by
145 the background flux and the standard deviation was computed. These same six flux
146 channels envelop the background cliff. If the ratio of background flux to measured flux
147 is similar in all six channels, this means the measured flux as a curve is close to the
148 background. Therefore, if the standard deviation of these six ratios is small, we know
149 that all the channels have similar ratios. If the standard deviation was below 2, indicating
150 a rather “flat” distribution mimicking the background values, then the spectrum was also
151 considered a void.

152 An important value is the solar zenith angle which divides sunlit areas from darkness or
153 the terminator. Photoelectrons measured by MGS transport from the main source regions
154 at 100-200 km to 400 km along closed magnetic field lines. While the production peak
155 of photoelectrons is roughly located around 130 km, these electrons mostly lose energy

156 locally due to collisions with the neutral atmosphere. Only above a certain altitude, or
157 supposed “superthermal electron exobase”, can photoelectrons transport to high altitudes
158 [*Nagy and Banks, 1970*]. This superthermal electron exobase varies from an altitude of
159 140 km to 180 km, as reported by several studies [e.g., *Mantas and Hanson, 1979; Lillis*
160 *et al., 2008; Steckiewicz, 2015; Xu et al., 2016a, manuscript submitted*]. To ensure no
161 source, we used an altitude of 200 km to calculate this photoelectron source terminator.
162 A base altitude of 90 km was chosen instead of the surface of Mars as we are assuming
163 the light will be attenuated by the atmosphere below this altitude. The terminator was
164 computed to be $SZA = \sim 104^\circ$ at an altitude of 200 km. Beyond this solar zenith angle,
165 it is considered that there is no source term for photoelectrons.

166 The modified pitch angle approach allowed us to take the absolute value of the magnetic
167 elevation angle data. Doing so shrinks the range from the usual -90° to $+90^\circ$ to a reduced
168 range of just $0^\circ - 90^\circ$. Elevation angles with values of 0° are tangential with respect to
169 the planet and angles with values of 90° are radial with respect to the planet.

3. Results

3.1. Determination of Populations

170 *Xu et al. [2014b]* used the electron flux ratio of multiple energy pairs (26eV/115eV,
171 36eV/115eV, 47eV/115eV) to identify the populations of photoelectrons and solar wind
172 electrons. Photoelectrons have a characteristic “knee” in their spectra near 60-70 eV (due
173 to a drop of solar photons below 15 nm), while solar wind (magnetosheath) electrons do
174 not (Figure 1c). Therefore, taking the flux of an energy channel above and below the
175 knee will give a larger ratio for photoelectrons than solar wind electrons. For the flux
176 ratio of 47eV/115eV, the population of solar wind electrons had an upper bound of 14

177 and the photoelectrons had a lower bound of 19 in *Xu et al.* [2014b]. In this study, a
178 flux ratio of 16 will be used as a hard cutoff in which samples with a ratio above 16 are
179 considered photoelectrons and below as solar wind electrons. Doing so allows us to classify
180 all observations as either photoelectron, solar wind electron, or void.

181 Histograms of the flux ratio as a function of solar zenith angle are shown in Figure
182 2. All histograms are at pitch angles of 90° - 100° , because this pitch angle range has the
183 most points due to the instrument field of view limitation. These histograms show either
184 one or two population distributions. As shown in Figure 2a, the histogram has a one-
185 population distribution with a peak ratio ~ 28 suggesting photoelectrons dominating for
186 $SZA = 90^\circ$ - 105° . It is expected since the atmosphere is still sunlit, even though $SZA >$
187 90° . As the solar zenith angle increases, the distribution becomes bimodal with another
188 peak at ~ 7 (Figure 2b), suggesting access of both photoelectrons and solar wind electrons
189 to this location. Abruptly at the 110° - 115° solar zenith angle bin, the histograms revert
190 back into a single-population distribution, as the photoelectron peak decays, with peak
191 ratio ~ 5 , indicative of a solar wind population (Figure 2c).

192 To analyze the dependence on magnetic elevation angle, we divided the dataset into
193 three solar zenith angle ranges, 90° - 105° , 105° - 130° , and 130° - 155° . Figure 3 displays
194 histograms of these three ranges split into multiple magnetic elevation angle bins. Again,
195 all histograms are at pitch angles of 90° - 100° . Note that the y-axis is not constant across
196 the plots. Figures 3a, 3b, and 3c show that the relative size of the solar wind electron
197 population to the photoelectron population increases with increasing magnetic elevation
198 angle. Solar wind electrons enter through crustal field lines connected to the IMF and,
199 at 400 km altitude over the strong crustal field regions, these field lines are more likely to

200 be vertical. Figures 3d, 3e, and 3f show the relative size of the photoelectron population
201 to the solar wind electron population decreases with elevation angle. Photoelectrons at
202 these solar zenith angles are more likely to be trapped on closed field lines. Figures 3g, 3h
203 and 3i reiterate that a photoelectron population does not exist at high solar zenith angles.
204 The sample number of each population, however, cannot be directly compared, because
205 MGS measured each elevation angle bin unevenly for a given solar zenith angle due to
206 the seasonal effect. Thus we determine the occurrence rate of each electron population
207 by normalizing the sample number by the total observations in each bin.

3.2. Occurrence Rate of Electron Populations

208 Every measurement can be labeled as either void, photoelectron, or solar wind electron.
209 Figure 4 shows the occurrence rates of the three classifications against pitch angle and
210 solar zenith angle. The occurrence rate is the fraction of data points in that bin with
211 that particular classification. Each bin has at least 1000 data points with the average
212 being around 32,000 data points. Voids are most prevalent once the spacecraft is past
213 the photoelectron source terminator and for field aligned pitch angles (in the loss cone).
214 Diagrams of this loss cone can be seen in Figure 3 of *Liemohn et al.* [1997]. Electrons
215 with downward field aligned pitch angles, i.e. near 0° , are more likely to make it further
216 into the ionosphere and be lost and since this study is solely on the nightside, there are
217 not many electrons escaping with upward directed pitch angles, i.e. near 180° . Electrons
218 with perpendicular pitch angles will mirror at higher altitudes, being less exposed to the
219 denser ionosphere at lower altitudes. Photoelectrons populate sunlit areas but once past
220 the terminator the population decays due to lack of a source term. The photoelectron's
221 energy degrades and the electron becomes part of the thermal population, which has low

222 enough energy to recombine. Solar wind electrons have an increase in their occurrence
223 rate past the terminator and remain relatively constant throughout the night side.

224 More insight can be gained by plotting the occurrence rate as a function of pitch angle
225 and elevation angle for a range of solar zenith angles. In Figure 5, the occurrence rate
226 is plotted for both photoelectrons and solar wind electrons to see their behavior with
227 increasing solar zenith angle. The average sample size is 3000 data points but never drops
228 below 300. Photoelectrons are completely dominant for solar zenith angles of 96° - 99°
229 (Figure 5a), which is still magnetically connected to a sunlit source region. As solar zenith
230 angle increases, the occurrence rate of photoelectrons decreases from Figure 5a to 5b, as
231 the source weakens while crossing the photoelectron source terminator. In contrast, the
232 photoelectrons at perpendicular pitch angles are trapped populations, bouncing at high
233 altitudes where collision frequency is much lower. The occurrence rate becomes lower at
234 small elevation angles, i.e. more horizontal magnetic fields, from Figure 5b to 5c and
235 on to 5d. At the altitude of MGS, shorter loop structures tend to be horizontal, thus
236 easier for photoelectron energies to degrade due to more frequent collisions with neutral
237 particles, compared to more vertical/taller magnetic field structures. The solar wind
238 population occupies vertical field lines that are more easily connected to the IMF. The
239 solar wind occurrence rate eventually becomes constant once the photoelectron population
240 has sufficiently degraded below the instrument detection threshold as shown in Figure
241 5i and 5j. At $SZA = 117^{\circ}$ - 120° (Figure 5e and 5j), the photoelectron population is a
242 “shadow” of the solar wind population, same shape but small fraction of the values.
243 Figure 3h is indicative of a solar wind population yet the measurements in the tail of the

244 distribution are being classified as photoelectrons. This “shadow” is an artifact of our
245 classifying method and the photoelectrons have degraded by this point.

3.3. Energy Deposition

246 The energy deposition due to superthermal electrons precipitating into the Martian
247 ionosphere is important for ionization, local heating, and excitation (aurora on the night-
248 side, dayglow on the dayside). Figure 6 displays the calculated median pitch angle aver-
249 aged upward, downward, and net deposited (downward-upward) energy fluxes by elevation
250 angle and solar zenith angle. The deposited flux is not a subtraction of the two medians
251 but first the deposited flux was calculated for each observation and then the median found.
252 This calculation was done with voids and without voids, the first and second columns re-
253 spectively. Furthermore, the energy fluxes are integrals across the entire ER energy range
254 from ~ 10 eV to ~ 20 keV, not just the six energy channels used to classify voids in the
255 earlier section. Sample sizes for Figures 6a, 6b, and 6c have an average of 6000 while the
256 average sample size for 6d, 6e, and 6f is 4000 and neither drop below 100. For Figure 6d,
257 only measurements with voids across all upward pitch angle bins were discarded, and the
258 same for Figure 6e with downward bins. Figure 6f had measurements discarded only if
259 all 18 pitch angle bins were classified as voids. Gray boxes have negative values on the
260 order of 10^6 - 10^7 eV/cm²/s while a few get as low 10^4 or as high as 10^8 , indicating a net
261 upward flow. The black boxes in Figure 6a, 6b, and 6c have median values of 0 eV/cm²/s,
262 i.e. the median is within the void population of spectra classifications. For comparison,
263 6.2×10^{11} eV/cm²/s is equivalent to 1.0 mW/m².

264 From these plots we can identify two regions, one where the amount of voids is significant
265 enough to obscure calculations and another where they are not. The void populated region

266 in Figure 6c, i.e. where the deposited flux is zero, now contains net upward fluxes when
267 the voids are removed (Figure 6f). Histograms of the net energy flux in this region show
268 a bimodal distribution with two peaks on either side of zero for Figure 6f. This region in
269 both Figures 6d and 6e are so similar that the subtraction between the two appears to be
270 zero. Also, the pitch angle averaged energy spectra for both the upward and downward
271 direction in this region are very close to each other, probably within measurement error.
272 The median is probably not a good value to represent this type of distribution (bimodal)
273 found in the gray bins and the upward net flux is likely to be noise. The inclusion of voids
274 gives us an idea of what we might observe with any random measurement constrained to
275 the orbit of MGS. What we might observe when electrons are precipitating is given by
276 the right column of Figure 6, i.e. where we have excluded voids.

277 Energy deposition past $SZA = 110^\circ$ is due to solar wind electrons. Figures 2 and 4b
278 show the disappearance of the photoelectron population at at this solar zenith angle.
279 Typical post-photoelectron source terminator pitch angle averaged deposited flux values
280 occur primarily on near vertical field lines with an average value of 2.0×10^8 eV/cm²/s.

281 Not all downward flux will be deposited into the ionosphere to cause ionization, heating,
282 and excitation. The fractional deposition rate is calculated by dividing the net deposited
283 flux by the downward flux. It is unitless and is the percentage of downward flux that
284 is deposited. Voids have been excluded in this calculation. Figure 7 shows the median
285 value for each solar zenith angle, elevation angle bin. Gray boxes have negative values
286 with magnitude ~ 0.001 - 0.05 , suggesting that these negative values are indeed noise. The
287 fractional deposition rate provides insight as to where there is a higher rate of magnetic
288 reflection and/or back scattering. The rate is low for the first solar zenith angle bin,

289 90°-95°, where the flux tube is fully sunlit and the upward and downward fluxes are very
290 similar. As we move to higher solar zenith angles, the locally generated (i.e. upward)
291 fluxes are reduced and therefore there is a greater net downward flux. Once past $SZA =$
292 110° , where the photoelectrons have vanished, the deposition rate decreases due to the
293 lack of a magnetic footprint in the sunlit ionosphere, restricting electron transport into the
294 nightside. The rate is fairly constant on the nightside, due to only solar wind electrons
295 precipitating. The decrease in neutral density from day to night (about a factor of 3
296 [Keating *et al.*, 2007]) could also be a reason for the decrease in deposition rate with solar
297 zenith angle. A decrease in neutral density will lower the frequency of collisions, allowing
298 more electrons to be reflected/back scattered before depositing their energy. Vertical field
299 lines correspond to longer field structures, therefore the electron travels a longer path and
300 has more opportunities to deposit its energy. The highest median fractional deposition
301 rate is 0.16 meaning 84% of the downward flux is reflected/scattered out. Each bin had
302 at least one instance where the deposition rate was at least 0.85. There are times where
303 the majority of electrons are deposited, across all solar zenith angles and elevation angles,
304 but it does not happen often.

3.4. Average Energy

305 The energy of an average electron is an important calculation due to implications in-
306 volving the ionization of neutrals and the depth it will cause ionization at [Banks *et al.*,
307 1974]. The average energy of an electron was calculated and the median plotted against
308 solar zenith angle and elevation angle in Figure 8. The average energy was calculated by
309 dividing the energy flux by the number flux, both integrated over energy and pitch angle.
310 The rows of Figure 8 are for upward directed electrons, downward directed electrons, and

311 for deposited electrons, from top to bottom, respectively. In order to calculate the upward
312 average energy, only the upward pitch angle bins for the number and energy fluxes are
313 integrated over and likewise for the downward direction. The average deposited energy
314 is a subtraction of the two opposite-directed fluxes and then the median found from the
315 resulting values as before. Again, voids have been excluded from the calculation as the
316 inclusion of them would skew the results and have non-physical values.

317 Where photoelectrons are the primary species ($90^\circ < \text{SZA} < 110^\circ$), the average energies
318 of electrons is roughly 15-25 eV, characteristic energies of photoelectrons. Past $\text{SZA} =$
319 110° , at higher elevation angles ($B_{elev} > 50^\circ$), the average upward energy is higher than
320 the average downward energy, implying that low energy electrons are being deposited and
321 higher energy electrons are being magnetically reflected and/or scattered back out. The
322 average energy for the deposited electrons is 20-30 eV, lower than the downward electrons
323 in this region, affirming this conclusion. The average energy for all electrons is higher for
324 lower elevation angles ($B_{elev} < 50^\circ$) than for greater elevation angles. High energy trapped
325 electrons are more likely to survive over lower energy trapped electrons.

326 The energy of the deposited electrons past the photoelectron source terminator is rather
327 low, 20-30 eV. The depth of ionization due to these particles will occur around the pho-
328 toelectron exobase, 140-180 km [e.g., *Nagy and Banks*, 1970; *Lillis et al.*, 2008; *Xu et*
329 *al.*, manuscript submitted]. They will not have the energy to penetrate deeper into the
330 ionosphere, where *Haider et al.* [2007] found the peak ion layer to be ~ 130 km. The
331 nightside ionosphere may have an ion peak shifted higher in altitude over areas where
332 electrons are precipitating.

3.5. Consequences of Energy Deposition

3.5.1. Excitation (Aurora)

Brain et al. [2006] investigated the possibility of aurora on Mars. They looked at peaked electron energy spectra and found peaks of 100 eV-2.5 keV. They noted that many of the examined spectra, including the MEX event in *Bertaux et al.* [2005], occurred during solar energetic particle (SEP) events. *Brain et al.* [2006] used a typical auroral-like energy spectra for analysis. The fluxes recorded were observed by MGS on April 21, 2001 over the strong crustal field region at SZA $\sim 125^\circ$. The downward flux at this time was 8.1×10^9 eV/cm²/s (1.3×10^{-2} mW/m²) and this flux was used as input to a model to estimate the amount of emission produced from the deposition of electrons. They note that half of this flux was deposited producing ~ 4 R of emission from the CO₂⁺ (289.7 nm) line, which is ~ 17 times weaker than the MEX event.

Here we examined how many deposited energy flux occurrences in our analysis were greater than this deposited energy flux value of the *Brain et al.* [2006] study, 4×10^9 eV/cm²/s (6.4×10^{-3} mW/m²). The fraction of values that exceed this number is plotted in Figure 9a. We then looked specifically at the SZA/elevation angle bin that had the highest median net energy flux past 110° . This occurred at SZA = 110° - 115° and at $B_{elev} = 80^\circ$ - 90° . Figures 9b and 9c show the histogram of net energy flux values in this bin and the cumulative distribution function (CDF) respectively. A red dashed line denotes the *Brain et al.* [2006] deposited energy flux value. The blue dashed line in Figure 9c is the median flux for this bin. The deposited energy flux exceeded the *Brain et al.* [2006] value only six percent of the time for this bin and throughout the nightside along near-vertical field lines, it varies from one to seven percent. There are values that reach up to 20%

355 but these are before the photoelectron production terminator where the ionosphere is still
 356 sunlit. The energy flux needed to cause substantial emission does not occur all that often
 357 on the nightside.

358 3.5.2. Ionization

359 The creation of the nightside ionosphere of Mars is caused by day-to-night transport
 360 of electrons, precipitating electrons, or a combination thereof. Electron density profiles
 361 have been analyzed using radio occultations and predicted using models [e.g., Table 1
 362 in *Fillingim et al., 2007; Zhang et al., 2015*] and estimates of the peak density are on
 363 the order of 10^3 - 10^4 cm^{-3} . The average deposited flux in Figure 6f above SZA of 110°
 364 along near-vertical field lines is 2.0×10^8 $\text{eV}/\text{cm}^2/\text{s}$. We can divide this energy flux
 365 by the average energy to ionize a particle at Earth, 35 eV, and divide by the depth of
 366 the ionosphere, ~ 200 km [*Schunk and Nagy, 2000*]. This calculation produces a volume
 367 production rate of electrons of ~ 0.3 $\text{cm}^{-3} \text{s}^{-1}$. The collisional ionization cross section for
 368 CO_2 peaks at roughly 100 eV [*Strickland and Green, 1969*], however, so the ionization
 369 may be inefficient, due to the average energy of deposited electrons being 20-30 eV, and
 370 the presumed average ionization energy, 35 eV, may not be high enough. Therefore, the
 371 volume production rate is an upper bound for the flux value used.

With this production rate, the average density in the ionosphere can be estimated by

$$0.3 \text{ cm}^{-3} \text{ s}^{-1} = (n_{\text{O}_2^+})(n_{e^-})k \quad (1)$$

372 where $k = 6.4 \times 10^{-8} \text{ cm}^3 \text{ s}^{-1}$ [*Peverall et al., 2001*], which is the dissociative recombination
 373 rate for O_2^+ , the most common ion in the ionosphere [e.g., *Hanson et al., 1977*], when
 374 $T_e \sim 2000$ K. If we assume that the densities of O_2^+ and e^- are equal, i.e photochemical
 375 equilibrium, then the average density produced by a net energy flux of 2.0×10^8 $\text{eV}/\text{cm}^2/\text{s}$

376 is $2.1 \times 10^3 \text{ cm}^{-3}$. Note that this calculation only accounts for near-vertical field lines
377 over the strong crustal field region in the southern hemisphere.

378 However, we assumed a uniform electron distribution in a 200 km thick ionosphere, and
379 our density value is an average throughout. A calculation of the total electron content
380 (TEC) would be a better way to define this ionosphere. TEC values have been estimated
381 at Mars [e.g., *Safaenili et al.*, 2007; *Lillis et al.*, 2009] and nightside values are on the
382 order of 10^{14} m^{-2} [Figure 1a of *Lillis et al.*, 2010]. An average density of $2.1 \times 10^3 \text{ cm}^{-3}$
383 over 200 km corresponds to a TEC of $4.2 \times 10^{14} \text{ m}^{-2}$, which is agreeable with previous
384 estimates.

385 Using a simple triangle distribution, a peak density can be calculated using $\text{TEC} = 4.2$
386 $\times 10^{14} \text{ m}^{-2}$. A peak layer of 160 km is used [e.g. *Fillingim et al.*, 2007], and this produces
387 a peak density of $4.2 \times 10^3 \text{ cm}^{-3}$. *Fillingim et al.* [2007] used a model to do this same
388 calculation and our TEC calculations and maximum electron density calculations are on
389 the same order of magnitude. In this study, pitch angle averaged net energy flux was used,
390 while *Fillingim et al.* [2007] used differential downward energy flux. Our assumption of
391 the average energy to ionize a particle is probably low, but these are quick calculations
392 to understand the ionosphere our deposited flux values may cause. These values may be
393 enough to support the nightside ionosphere in some areas but other mechanisms such as
394 day-to-night transport may be needed in other regions to sustain the ionosphere.

4. Conclusion

395 The same method utilized by *Xu et al.* [2014b] was used to classify the populations
396 of photoelectrons and solar wind electrons and investigate superthermal electron energy
397 deposition as measured by MGS on the nightside of Mars over the strong crustal field

398 region. The ratio of two energy flux values at 47 eV and 115 eV will be greater for
399 photoelectrons because of the “knee” in the energy spectra around 60-70 eV. Histograms
400 as a function of solar zenith angle show a single population distribution at low solar zenith
401 angles ($SZA \leq 100^\circ$) centered around a flux ratio of 30. The distribution becomes bimodal
402 with another peak around a flux ratio of 10 from $SZA = 100^\circ$ - 110° , and at $SZA = 110^\circ$, the
403 histogram changes back into a single population distribution losing the higher flux ratio
404 peak. This demonstrates the photoelectron population being the dominant population
405 while still in sunlit areas and eventually degrading in energy below instrument detection
406 with increasing solar zenith angle leaving only solar wind electrons.

407 The occurrence rate was calculated as functions of solar zenith angle, elevation angle,
408 and pitch angle. Voids are prominent at $SZA > 110^\circ$, past the photoelectron source termi-
409 nator and away from magnetic loops with a sunlit footprint. This also is the solar zenith
410 angle where the photoelectron population has sufficiently degraded. In the sunlit sectors,
411 photoelectrons are the main population over all pitch angles and elevation angles. Once
412 past the terminator, the field aligned pitch angles are the first to lose their photoelectron
413 population. The perpendicular pitch angles are trapped thus having a longer lifespan.
414 The occurrence rate then drops for horizontal elevation angles. These field structures
415 tend to be shorter and photoelectrons degrade in energy more quickly. The solar wind
416 occurrence rate is highest at near vertical field lines, more easily connected to the IMF.

417 The energy deposition on the nightside occurs primarily on vertical field lines and before
418 the terminator on horizontal field lines. Typical nightside values for pitch angle averaged
419 deposited flux is $\sim 2.0 \times 10^8$ eV/cm²/s. Past $SZA = 110^\circ$, it is safe to assume that all
420 energy deposition is due to solar wind electrons since histograms and occurrence rates

421 show photoelectrons to be depleted by this point, especially from a statistical point of
422 view. Note that this study used only MGS mapping phase data, at 2 AM local time,
423 and so this cutoff is specific to that orbital constraint. The region on the nightside just
424 past the terminator also has high flux values most likely due to a magnetic loop with
425 one footprint on the dayside and the other on the nightside. More research into these
426 regions could help answer questions about electron transport to the nightside and are
427 probable regions for aurora. The maximum median fractional deposition rate found was
428 0.16. Most of the precipitating electrons are magnetically reflected or scattered back out.
429 The average energy of deposited electrons was found to be 20-30 eV, perhaps creating
430 an ion peak shifted upward in altitude in areas where electrons are precipitating on the
431 nightside.

432 One consequence of electron deposition is excitation and emission. *Brain et al.* [2006]
433 used a model and found that a deposited energy flux of 4×10^9 eV/cm²/s (6.4×10^{-3}
434 mW/m²) will create ~ 4 R of emission from CO₂⁺ (289.7 nm). We found that deposited
435 electron energy flux values that exceed the *Brain et al.* [2006] value occur one to seven
436 percent of the time along near-vertical field lines.

437 Another consequence of electron deposition is ionization. Estimates of the TEC and
438 peak density were calculated using an average deposited electron energy flux found in this
439 study, 2.0×10^8 eV/cm²/s. The TEC was found to be 4.2×10^{14} m⁻² with a corresponding
440 peak density of 4.2×10^3 cm⁻³. We note that this is limited to near vertical field lines over
441 the strong crustal field region past SZA of 110°.

442 It should be noted that the probabilities calculated in this study are based on a lower
443 limit of the counts that is linked to the ER instrumental background threshold. There

444 certainly could be photoelectron or solar wind electron fluxes below this limit that are
445 neglected in the statistics presented above. Therefore, all of the probabilities for these
446 two populations are lower limits, and the "void" probabilities are upper limits. That said,
447 these neglected components of the electron populations are, by definition, at very low
448 fluxes, and are therefore not likely to cause an appreciable level of ionization or excitation
449 in the thermosphere.

450 There is still further research to do from this study. A calculation of the decay rate
451 of photoelectrons as a function of time in darkness and not solar zenith angle has not
452 been done. This could be done through data analysis and compared to model results. A
453 superthermal electron transport (STET) model has been developed for Mars [*Liemohn*
454 *et al.*, 2003, 2006; *Xu and Liemohn*, 2015] and could be employed to ascertain the decay
455 rate for photoelectrons.

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459 tary Data System (PDS) and the full MGS MAG/ER dataset used in the paper is available
460 upon request to Dr. David Mitchell (mitchell@ssl.berkeley.edu).

References

461 Acuña, M. H., et al. (1998), Magnetic field and plasma observations at Mars: Ini-
462 tial results of the Mars Global Surveyor mission, *Science*, *279*(5357), 1676-1680,
463 doi:10.2307/2894341.

- 464 Acuña, M. H., et al. (2001), Magnetic field of Mars: Summary of results from
465 the aerobraking and mapping orbits, *J. Geophys. Res.*, *106*(E10), 23403-23417,
466 doi:10.1029/2000JE001404.
- 467 Ashour-Abdalla, M., and C. F. Kennel (1978), Diffuse auroral precipitation, *J. Geomagn.*
468 *Geoelectr.*, *30*(3), 239-255, doi:10.5636/jgg.30.239.
- 469 Banks, P. M., C. R. Chappell, and A. F. Nagy (1974), A new model for
470 the interaction of auroral electrons with the atmosphere: Spectral degradation,
471 backscatter, optical emission, and ionization, *J. Geophys. Res.*, *79*(10), 1459-1470,
472 doi:10.1029/JA079i010p01459.
- 473 Bertaux, J.-L., F. Leblanc, O. Witasse, E. Quemerais, J. Lilensten, S. A. Stern, B. Sandel,
474 and O. Korabely (2005), Discovery of an aurora on Mars, *Nature*, *435*(7043), 790-794,
475 doi:10.1038/nature03603.
- 476 Brain, D., J. Halekas, R. Lillis, D. Mitchell, R. Lin, and D. Crider (2005), Vari-
477 ability of the altitude of the Martian sheath, *Geophys. Res. Lett.*, *32*(18), L18203,
478 doi:10.1029/2005GL023126.
- 479 Brain, D. A., J. S. Halekas, L. M. Peticolas, R. P. Lin, J. G. Luhmann, D. L. Mitchell, G.
480 T. Delory, S. W. Bougher, M. H. Acuña, and H. Rème (2006), On the origin of aurorae
481 on Mars, *Geophys. Res. Lett.*, *33*(1), L01201, doi:10.1029/2005GL024782.
- 482 Brain, D. A., R. J. Lillis, D. L. Mitchell, J. S. Halekas, and R. P. Lin (2007), Electron
483 pitch angle distributions as indicators of magnetic field topology near Mars, *J. Geophys.*
484 *Res.*, *112*(A9), A09201, doi:10.1029/2007JA012435.
- 485 Connerney, J. E. P., M. H. Acuña, P. J. Wasilewski, G. Kleteschka, N. F. Ness, H. Reme, R.
486 P. Lin, and D. L. Mitchell (2001), The global magnetic field of Mars and implications for

- 487 crustal evolution, *Geophys. Res. Lett.*, *28*(21), 4015-4018, doi:10.1029/2001GL013619.
- 488 Dubinin, E., M. Fraenz, J. Woch, J. Winnigham, R. Frahm, R. Lundin, and S. Barabash
489 (2008a), Suprathermal electron fluxes on the nightside of Mars: Aspera-3 observations,
490 *Planetary and Space Science*, *56*(6), 846-851, doi:10.1016/j.pss.2007.12.010.
- 491 Fillingim, M. O., L. M. Peticolas, R. J. Lillis, D. A. Brain, J. S. Halekas, D. L. Mitchell, R.
492 P. Lin, D. Lummerzheim, S. W. Bougher, and D. L. Kirchner (2007), Model calculations
493 of electron precipitation induced ionization patches on the nightside of Mars, *Geophys.*
494 *Res. Lett.*, *34*(12), L12101, doi:10.1029/2007GL029986.
- 495 Fillingim, M. O., L. M. Peticolas, R. J. Lillis, D. A. Brain, J. S. Halekas, D. Lum-
496 merzheim, and S. W. Bougher (2010), Localized ionization patches in the nighttime
497 ionosphere of Mars and their electrodynamic consequences, *Icarus*, *206*(1), 112-119.
498 doi:10.1016/j.icarus.2009.03.005.
- 499 Fox, J. L., and A. I. F. Stewart (1991), The Venus ultraviolet aurora: A soft electron
500 source, *J. Geophys. Res. Space Physics*, *96*(A6), 9821-9828, doi:10.1029/91JA00252.
- 501 Fox, J. L., J. F. Brannon, and H. S. Porter (1993), Upper limits to the nightside ionosphere
502 of Mars, *Geophys. Res. Lett.*, *20*(13), 1339-1342, doi:10.1029/93GL01349.
- 503 Gérard, J.-C., L. Soret, L. Libert, R. Lundin, A. Stiepen, A. Radioti, and J.-
504 L. Bertaux (2015), Concurrent observations of ultraviolet aurora and energetic
505 electron precipitation with Mars Express, *J. Geophys. Res. Space Physics*, *120*,
506 doi:10.1002/2015JA021150.
- 507 Haider, S. A., J. Kim, A. F. Nagy, C. N. Keller, M. I. Verigin, K. I. Gringauz, N. M.
508 Shutte, K. Szego, and P. Kiraly (1992), Calculated ionization rates, ion densities, and
509 airglow emission rates due to precipitating electrons in the nightside ionosphere of Mars,

- 510 *J. Geophys. Res. Space Physics*, 97(A7), 10,637-10,641, doi:10.1029/92JA00317.
- 511 Haider, S. A., V. Singh, V. R. Choksi, W. C. Maguire, and M. I. Verigin (2007), Calculated
512 densities of $\text{H}_3\text{O}^+(\text{H}_2\text{O})_n$, $\text{NO}_2^-(\text{H}_2\text{O})_n$, $\text{CO}_3^-(\text{H}_2\text{O})_n$ and electron in the nighttime
513 ionosphere of Mars: Impact of solar wind electron and galactic cosmic rays, *J. Geophys.*
514 *Res. Space Physics*, 112(A12), A12309, doi:10.1029/2007JA012530.
- 515 Hanson, W. B., S. Sanatani, and D. R. Zuccaro (1977), The Martian ionosphere as ob-
516 served by the Viking retarding potential analyzers, *J. Geophys. Res.*, 82(28), 4351-4363,
517 doi:10.1029/JS082i028p04351.
- 518 Keating, G. M., S. W. Bougher, M. E. Theriot, R. H. Tolson, R. W. Zurek, R. C. Blan-
519 chard, J. R. Murphy, and J.-L. Bertaux, Mars Neutral Upper Atmosphere Temporal
520 and Spatial Variations Discovered from the Accelerometer Science Experiment Aboard
521 Mars Reconnaissance Orbiter, in Lunar and Planetary Institute Conference Abstracts,
522 vol. 38 of Lunar and Planetary Institute Conference Abstracts, pp. 2074-+, 2007.
- 523 Krymskii, A., T. K. Breus, N. F. Ness, M. H. Acuña, J. E. P. Connerney, D. H. Crider,
524 D. L. Mitchell, and S. J. Bauer (2002), Structure of the magnetic field fluxes connected
525 with crustal magnetization and topside ionosphere at Mars, *J. Geophys. Res. Space*
526 *Physics*, 107(A9), 1245, doi:10.1029/2001JA000239.
- 527 Leblanc, F., J. Y. Chaufray, J. Lilensten, O. Witasse, and J.-L. Bertaux (2006), Martian
528 dayglow as seen by the SPICAM UV spectrograph on Mars Express, *J. Geophys. Res.*
529 *Planets*, 111(E9), E09511, doi:10.1029/2005JE002664.
- 530 Leblanc, F., et al. (2008), Observations of aurorae by SPICAM ultraviolet spectrograph
531 on board Mars Express: Simultaneous ASPERA-3 and MARSIS measurements, *J. Geo-*
532 *phys. Res. Space Physics*, 113(A8), A08311, doi:10.1029/2008JA013033.

- 533 Liemohn, M. W., G. V. Khazanov, T. E. Moore, and S. M. Guiter (1997), Self-consistent
534 superthermal electron effects on plasmaspheric refilling, *J. Geophys. Res. Space Physics*,
535 *102*(A4), 7523-7536, doi:10.1029/96JA03962.
- 536 Liemohn, M. W., D. L. Mitchell, A. F. Nagy, J. L. Fox, T. W. Reimer, and Y. Ma (2003),
537 Comparisons of electron fluxes measured in the crustal fields at Mars by the MGS
538 Magnetometer/Electron Reflectometer instrument with a B-field-dependent transport
539 code, *J. Geophys. Res. Planets*, *108*(E12), 5134, doi:10.1029/2003JE002158.
- 540 Liemohn, M. W., et al. (2006), Numerical interpretation of high-altitude photoelectron
541 observations, *Icarus*, *182*(2), 383-395, doi: 10.1016/j.icarus.2005.10.036.
- 542 Liemohn, M. W., Y. Ma, A. F. Nagy, J. U. Kozyra, J. D. Winningham, R. A. Frahm, J.
543 S. Sharber, S. Barabash, and R. Lundin (2007), Numerical modeling of the magnetic
544 topology near Mars auroral observations, *Geophys. Res. Lett.*, *34*(24), L24202, doi:
545 10.1029/2007GL031806.
- 546 Lillis, R. J., D. L. Mitchell, R. P. Lin, and M. H. Acuña (2008), Electron reflectometry in
547 the Martian atmosphere, *Icarus*, *194*, 544-561, doi:10.1016/j.icarus.2007.09.030
- 548 Lillis, R. J., M. O. Fillingim, L. M. Peticolas, D. A. Brain, R. P. Lin, and S. W. Bougher
549 (2009), Nightside ionosphere of Mars: Modeling the effects of crustal magnetic fields
550 and electron pitch angle distributions on electron impact ionization, *J. Geophys. Res.*,
551 *114*(E11), E11009, doi:10.1029/2009JE003379.
- 552 Lillis, R. J., D. A. Brain, S. L. England, P. Withers, M. O. Fillingim, and A. Safaeinili
553 (2010), Total electron content in the Mars ionosphere: Temporal studies and dependence
554 on solar EUV flux, *J. Geophys. Res.*, *115*(A11), A11314, doi:10.1029/2010JA015698.

- 555 Lillis, R. J., and D. A. Brain (2013), Nightside electron precipitation at Mars: Geographic
556 variability and dependence on solar wind conditions, *J. Geophys. Res. Space Physics*,
557 *118*(6), 3546-3556, doi:10.1002/jgra.50171.
- 558 Lundin, R., et al. (2006a), Ionospheric plasma acceleration at Mars: ASPERA-3 results,
559 *Icarus*, *182*(2), 308-319, doi:10.1016/j.icarus.2005.10.035.
- 560 Lundin, R., et al. (2006b), Plasma acceleration above Martian magnetic anomalies, *Sci-*
561 *ence*, *311*(5763), 980-983, doi:10.2307/3843558.
- 562 Mantas, G. P., and W. B. Hanson (1979), Photoelectron fluxes in the Martian iono-
563 sphere, *J. Geophys. Res.*, *84*(A2), 369-385, doi:10.1029/JA084iA02p00369.
- 564 Mitchell, D. L., R. P. Lin, C. Mazelle, H. Rme, P. A. Cloutier, J. E. P. Connerney, M. H.
565 Acuña, and N. F. Ness (2001), Probing Mars' crustal magnetic field and ionosphere with
566 the MGS Electron Reflectometer, *J. Geophys. Res. Planets*, *106*(E10), 23419-23427,
567 doi:10.1029/2000JE001435.
- 568 Nagy, A. F., and P. M. Banks (1970), Photoelectron fluxes in the ionosphere, *J. Geophys.*
569 *Res.*, *75*(31), 6260-6270, doi:10.1029/JA075i031p06260.
- 570 Nagy, A., et al. (2004), The plasma environment of Mars, in *Mars Magnetism and Its*
571 *Interaction with the Solar Wind*, pp. 33-114, Springer, Netherlands.
- 572 Němec, F., D. D. Morgan, D. A. Gurnett, and F. Duru (2010), Nightside ionosphere of
573 Mars: Radar soundings by the Mars Express spacecraft, *J. Geophys. Res.*, *115*(E12),
574 E12009, doi:10.1029/2010JE003663.
- 575 Němec, F., D. D. Morgan, D. A. Gurnett, and D. A. Brain (2011), Areas of enhanced
576 ionization in the deep nightside ionosphere of Mars, *J. Geophys. Res.*, *116*(E6), E06006,
577 doi:10.1029/2011JE003804.

- 578 Peverall, R., S. Rosén, J. R. Peterson, M. Larsson, A. Al-Khalili, L. Viktor, J. Semaniak,
579 R. Bobbenkamp, A. Le Padellec, A. N. Maurellis, and W. J. van der Zande(2001), Dis-
580 sociative recombination and excitation of O_2^+ : Cross sections, product yields and impli-
581 cations for studies of ionospheric airglows, *The Journal of Chemical Physics*, *114*(15),
582 6679-6689, doi:10.1063/1.1349079.
- 583 Safaeinili, A., W. Kofman, J. Mouginot, Y. Gim, A. Herique, A. B. Ivanov, J. J.
584 Plaut, and G. Picardi (2007), Estimation of the total electron content of the Martian
585 ionosphere using radar sounder surface echoes, *Geophys. Res. Lett.*, *34*(23), L23204,
586 doi:10.1029/2007GL032154.
- 587 Schunk, R., and A. Nagy (2000), *Ionospheres: Physics, Plasma Physics, and Chemistry*,
588 *Cambridge Atmospheric and Space Science Series*, *59*, Cambridge Univ. Press, New
589 York.
- 590 Soret, L., J.-C. G'erald, L. Libert, V. I. Shematovich, D. V. Bisikalo, A. Stiepen, J.-L.
591 Bertaux, SPICAM observations and modeling of Mars aurorae, *Icarus*, *264*, 398-406.
592 doi:10.1016/j.icarus.2015.09.023.
- 593 Steckiewicz, M., et al. (2015), Altitude dependence of nightside Martian suprathermal
594 electron depletions as revealed by MAVEN observations, *Geophys. Res. Lett.*, *42*, 8877-
595 8884, doi:10.1002/2015GL065257.
- 596 Strickland, D. J., and A. E. S. Green (1969), Electron impact cross sections for CO_2 , *J.*
597 *Geophys. Res.*, *74*(26), 6415-6424, doi:10.1029/JA074i026p06415.
- 598 Swift, D. W. (1981), Mechanisms for auroral precipitation: A review, *Rev. Geophys.*,
599 *19*(1), 185-211, doi:10.1029/RG019i001p00185.

- 600 Verigin, M. I., K. I. Gringauz, N. M. Shutte, S. A. Haider, K. Szego, P. Kiraly, A.
601 F. Nagy, and T. I. Gombosi (1991), On the possible source of the ionization in the
602 nighttime Martian ionosphere: 1. Phobos 2 HARP electron spectrometer measurements,
603 *J. Geophys. Res. Space Physics*, *96*(A11), 19,307-19,313, doi:10.1029/91JA00924.
- 604 Xu, S., M. W. Liemohn, D. L. Mitchell, and M. D. Smith (2014a), Mars photoelectron en-
605 ergy and pitch angle dependence on intense lower atmospheric dust storms, *J. Geophys.*
606 *Res. Planets*, *119*, 1689-1706, doi:10.1002/2013JE004594.
- 607 Xu, S., M. W. Liemohn, and D. L. Mitchell (2014b), Solar wind electron precipitation
608 into the dayside Martian upper atmosphere through the cusps of strong crustal fields,
609 *J. Geophys. Res. Space Physics*, *119*(12), 10,100-10,115, doi:10.1002/2014JA020363.
- 610 Xu, S., and M. W. Liemohn (2015), Superthermal electron transport model for Mars,
611 *Earth and Space Science*, *2*(3), 47-64, doi:10.1002/2014EA000043.
- 612 Xu, S., M. W. Liemohn, S. Bougher, D. L. Mitchell (2016a), Enhanced carbon dioxide
613 causing the dust storm-related increase in high-altitude photoelectron fluxes at Mars,
614 *Geophys. Res. Lett.*, Article in Press, do: 10.1002/2015GL066043.
- 615 Xu, S., M. W. Liemohn, S. Bougher, and D. L. Mitchell (2016b), Effects of Solar Zenith
616 Angle on Martian High-Altitude Photoelectrons, *J. Geophys. Res. Space Physics*, sub-
617 mitted, ms# 2015JA022149.
- 618 Zhang, M. H. G., J. G. Luhmann, and A. J. Kliore (1990), An observational study of the
619 nightside ionosphere of Mars and Venus with radio occultation methods, *J. Geophys.*
620 *Res. Space Physics*, *95*(A10), 17,095-17,102, do:10.1029/JA095iA10p17095.
- 621 Zhang, S., J. Cui, P. Guo, J. L. Li, J. S. Ping, N. C. Jian, K. F. Zhang
622 (2015), Martian electron density profiles retrieved from Mars Express dual-frequency

Figure 1. These are three different plots of energy spectra at different times, pitch angle, solar zenith angle, and elevation angle. Pitch angles are modified. (a). An example matching our first criteria for a void-like spectrum. (b). An example matching our second criteria for a void-like spectrum. (c). The solid line is at $SZA = 90^\circ$ and $B_{elev} = 24^\circ$ with a flux ratio of 43, denoting photoelectrons. The dotted line is at $SZA = 145^\circ$ and $B_{elev} = 85^\circ$ with a flux ratio of 8, denoting solar wind electrons.

Figure 2. Histograms of the flux ratio (47eV/115eV) for modified pitch angles of 90° - 100° and (a) $SZA = 90^\circ$ - 95° , (b) $SZA = 105^\circ$ - 110° , (c) $SZA = 110^\circ$ - 115° . A red dashed line marks the flux ratio, 16, in which above, samples are considered photoelectrons and below, as solar wind electrons.

Figure 3. The columns are for different solar zenith angle ranges, from left to right: 90° - 105° , 105° - 130° , and 130° - 155° . The rows are for different magnetic elevation angle ranges of 0° - 10° , 40° - 50° , and 80° - 90° , from top to bottom. All histograms are sampled from modified pitch angles 90° - 100° . A red dashed line marks the flux ratio, 16, in which above, samples are considered photoelectrons and below, as solar wind electrons.

Figure 4. Occurrence rates for (a) voids, (b) photoelectrons, and (c) solar wind electrons as functions of solar zenith angle and modified pitch angle.

Figure 5. Occurrence rates for photoelectrons and solar wind electrons, the left and right columns respectively. The rows are for solar zenith angles 96° - 99° , 102° - 105° , 108° - 111° , 111° - 114° , and 117° - 120° (top to bottom).

623 radio occultation measurements, *Advances in Space Research*, 55(9), 2177-2189,

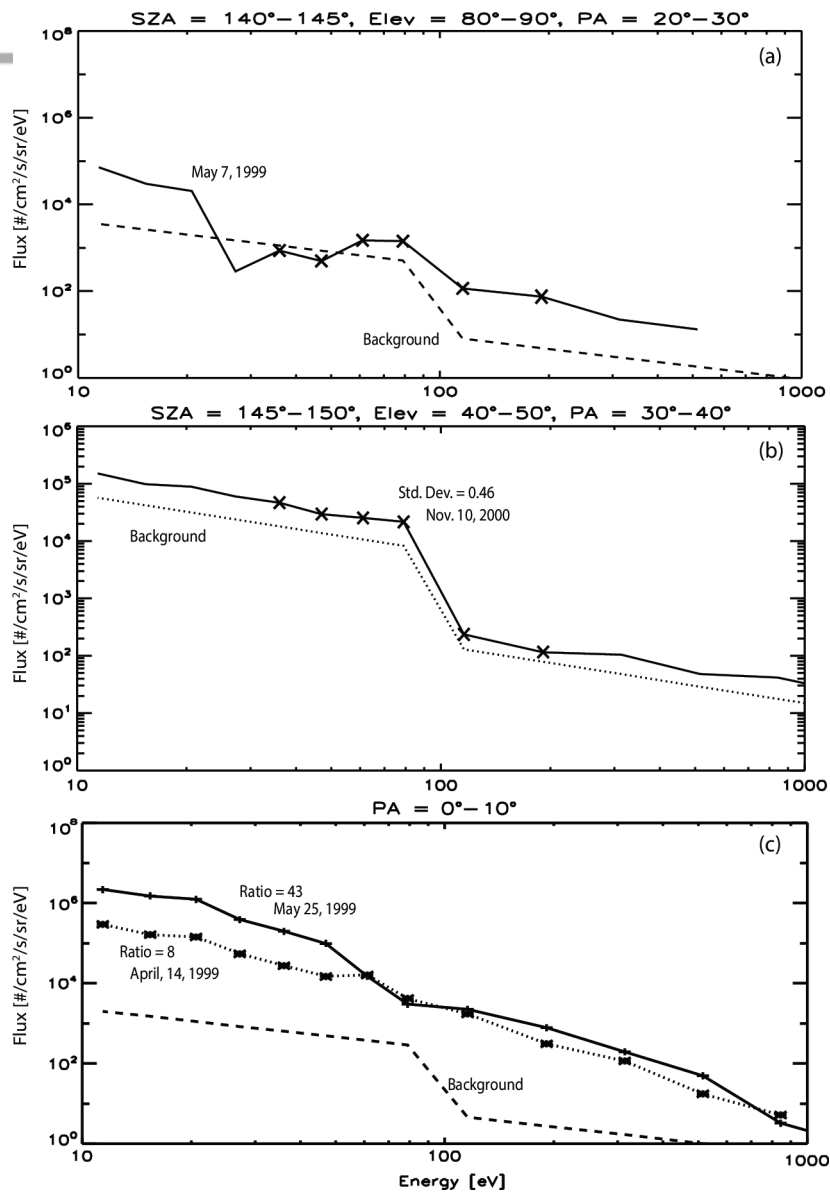
624 doi:10.1016/j.asr.2015.01.030.

Figure 6. Median pitch angle averaged energy fluxes for (a,d) upward directed electrons, (b,e) downward directed electrons, and the (c,f) net deposited (downward-upward) energy flux. The first column includes voids in the calculations and the second column does not. Gray boxes have negative values on the order of 10^6 - 10^7 eV/cm²/s while a few get as low 10^4 and as high as 10^8 .

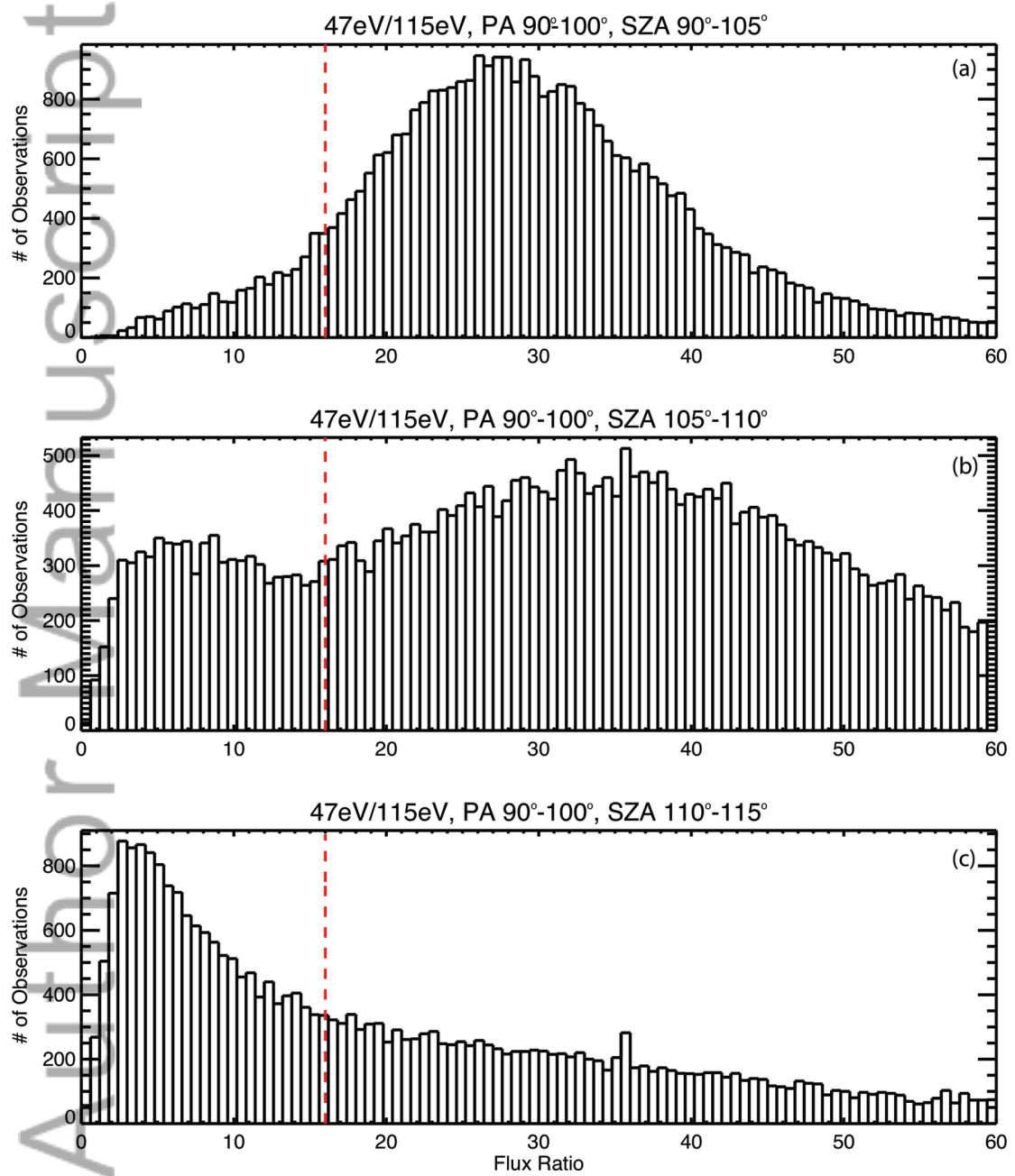
Figure 7. The median of the fractional deposition rate (net flux divided by downward flux), which is the fraction of downward flux that is deposited. Voids are excluded in this calculation. Gray boxes have negative values with magnitudes around ~ 0.001 - 0.05 .

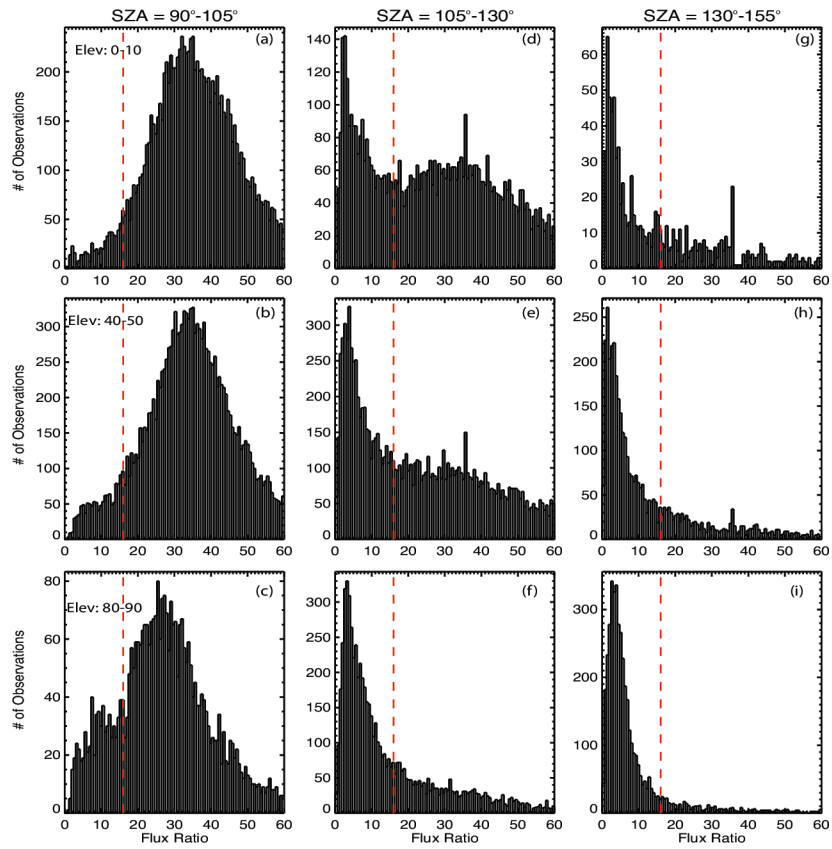
Figure 8. The median average energy for upward directed electrons (a), downward directed electrons (b), and the deposited electrons (c). Voids are excluded in these calculations.

Figure 9. (a). The fraction of net energy flux values that exceed 4×10^9 eV/cm²/s. (b). Histogram of the net energy flux values for $SZA = 110^\circ$ - 115° and $B_{elev} = 80^\circ$ - 90° . (c). CDF of the distribution in 9b. The red dashed line denotes the *Brain et al.* [2006] value and the blue dashed line is the median net energy flux.

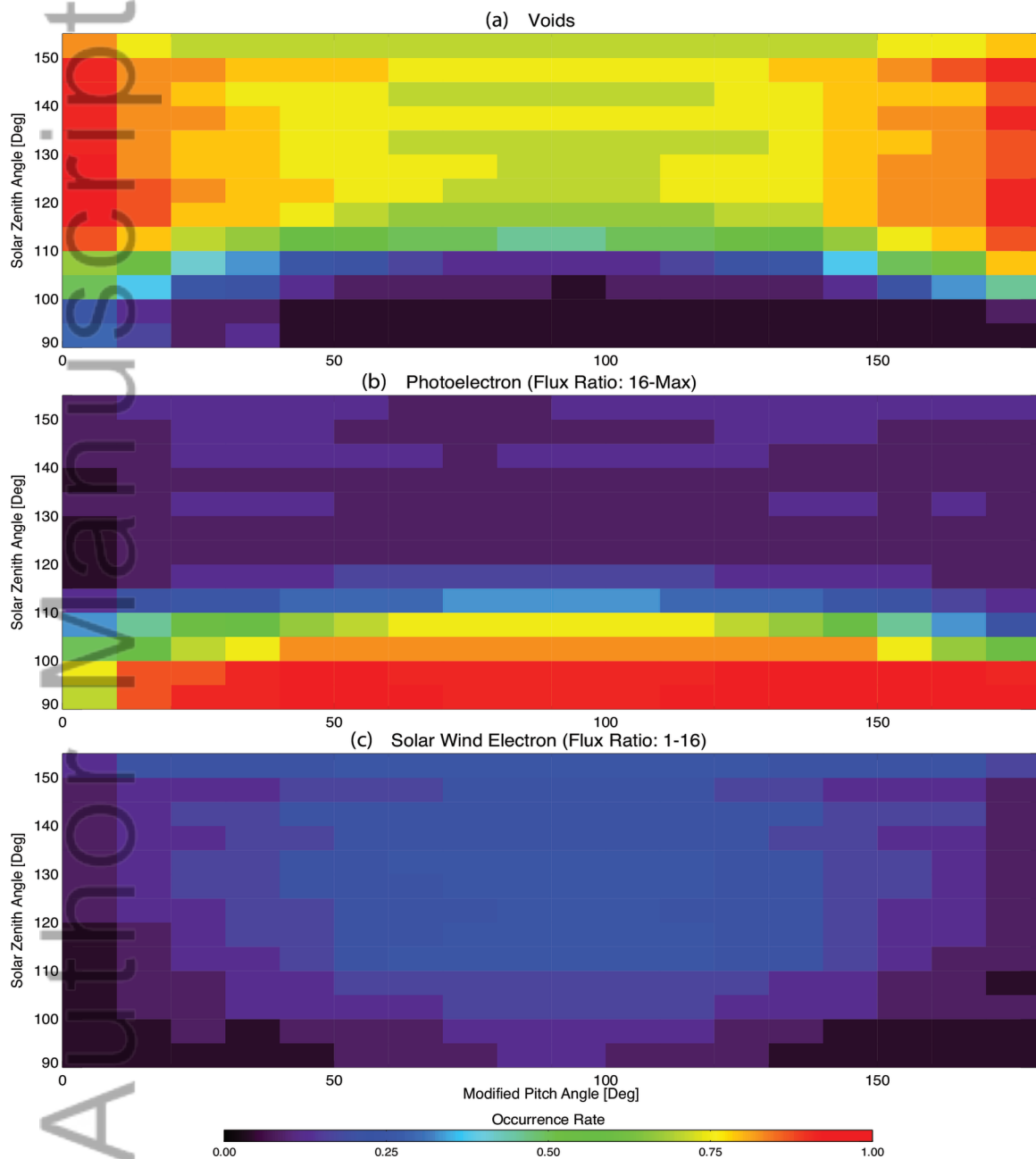


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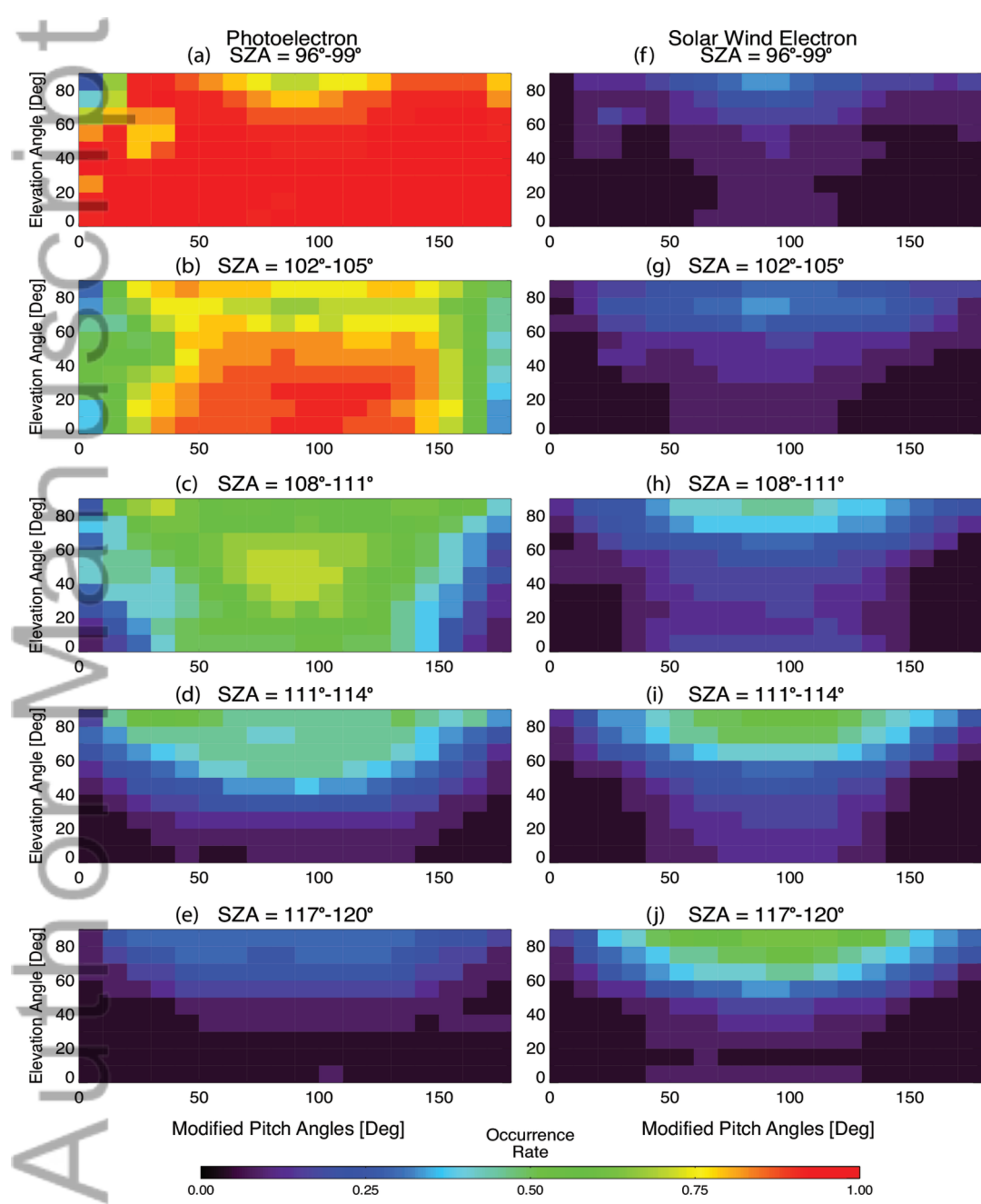




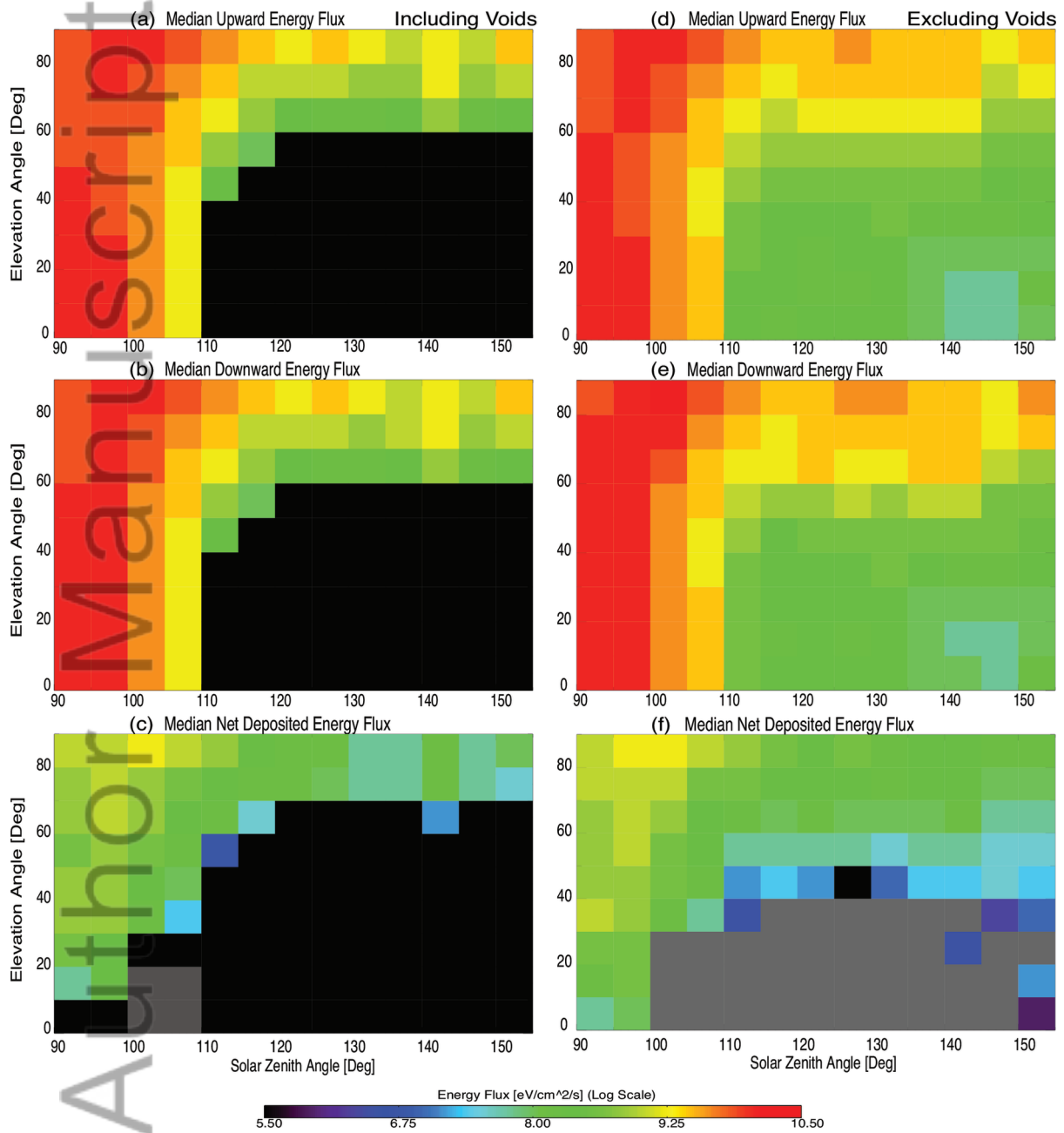
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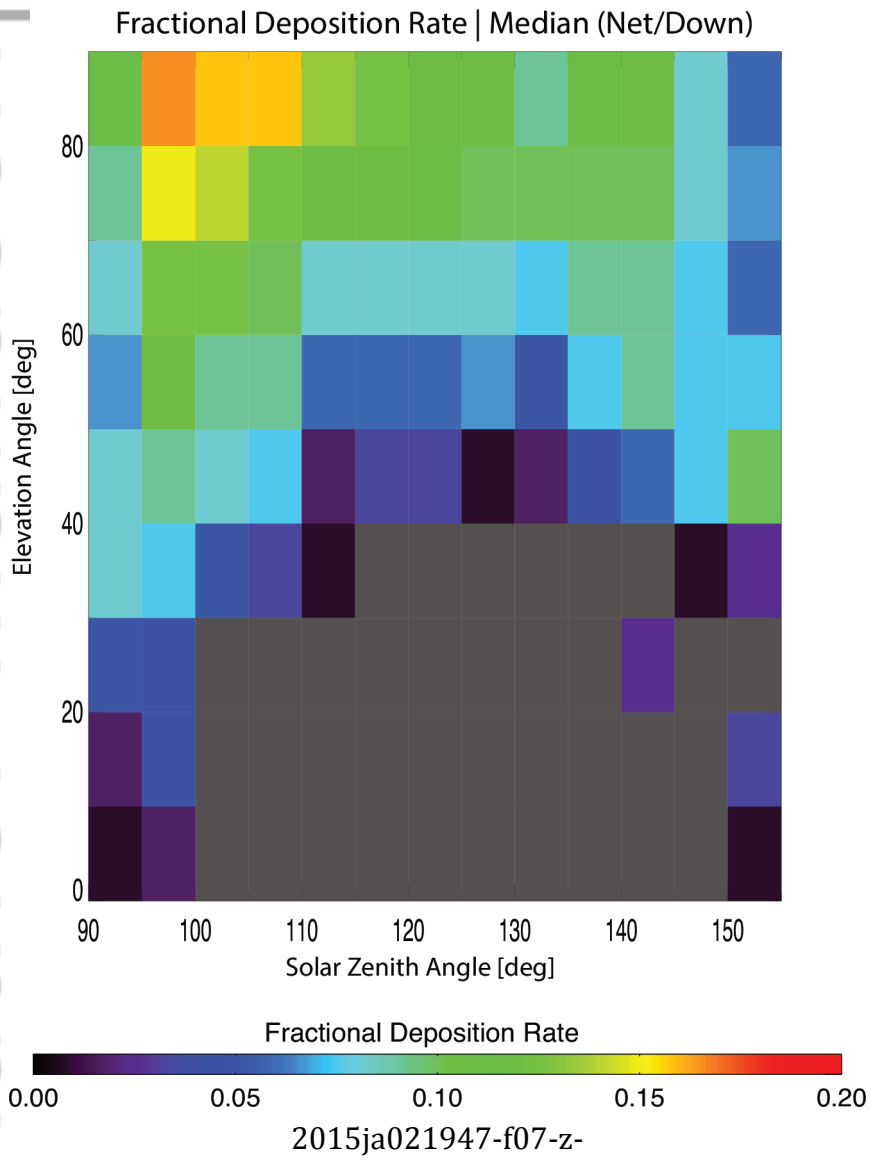
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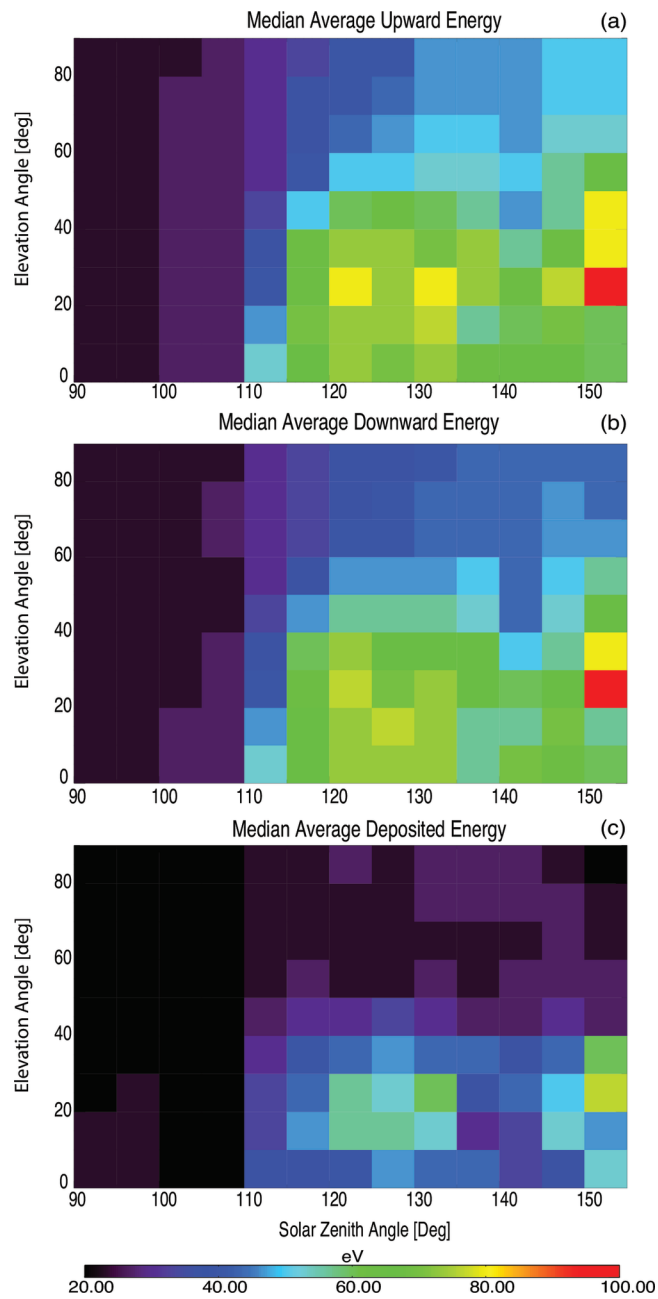


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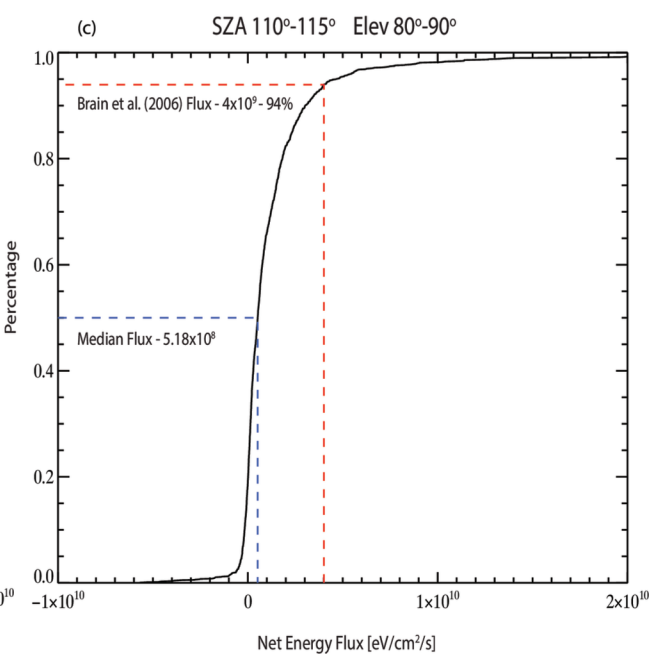
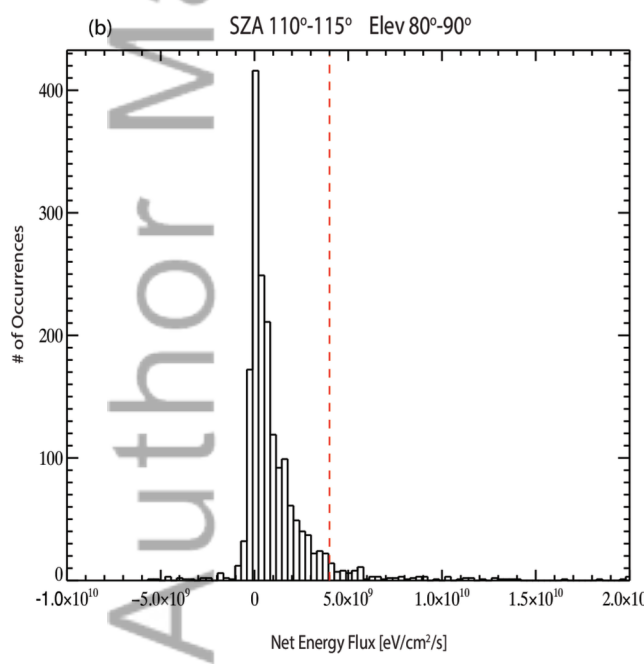
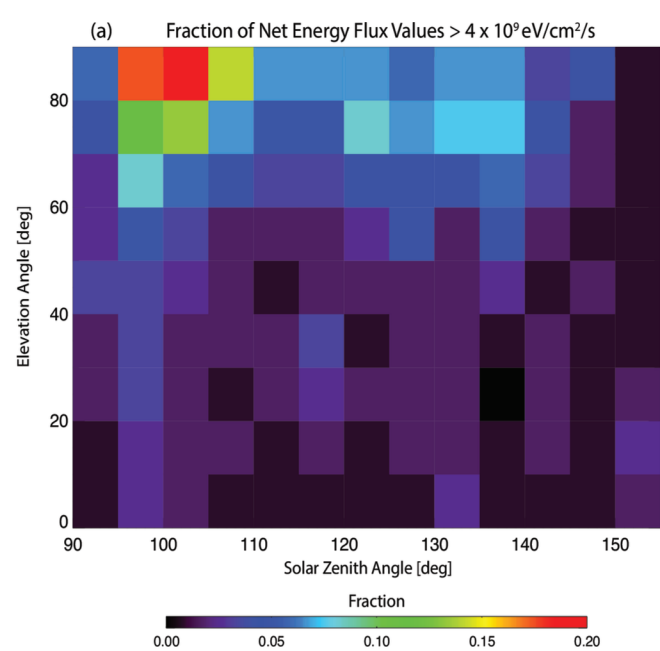


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