1 Title:

Range position and climate sensitivity: the structure of among-population demographic responses to
 climatic variation

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- 39 Demographic processes and climate interact and vary across a species' range to determine how species'
- 40 distributions will respond to climate change. We predicted that populations at the extremes of a species'
- 41 climate envelope are most sensitive to climate shifts. We tested this using a dynamic species distribution
- 42 model linking demographic rates to variation in climate for wood frogs (*Lithobates sylvaticus*) in North
- 43 America. Sensitivity to changes in climate cannot be predicted simply by knowing locations within the
- 44 species' climate envelope.
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66 relationships varied with respect to long-term climate. Some models supported the 67 predicted pattern, with negative effects of extreme summer temperatures in hotter areas 68 and positive effects on recruitment for summer water availability in drier areas. We also 69 found evidence of interacting temperature and precipitation influencing population size, 70 such as extreme heat having less of a negative effect in wetter areas. Other results were 71 contrary to predictions, such as positive effects of summer water availability in wetter 72 parts of the range and positive responses to winter warming especially in milder areas. In 73 general we found wood frogs were more sensitive to changes in temperature or 74 temperature interacting with precipitation than to changes in precipitation alone. Our 75 results suggest that sensitivity to changes in climate cannot be predicted simply by 76 knowing locations within the species' climate envelope. Many climate processes did not 77 affect population growth rates in the predicted direction based on range position. 78 Processes such as species-interactions, local adaptation, and interactions with the physical 79 landscape likely affect the responses we observed. Our work highlights the need to 80 measure demographic responses to changing climate.

81

82 Introduction

83 A persistent theme in ecology is the need to understand the factors that shape and 84 describe species distributions (Grinnell, 1917; MacArthur, 1972; Gaston, 2009; Sexton et 85 al., 2009). These factors have been touted as a means to understand the conditions that 86 facilitate sustainable populations currently and in the future (Sexton et al., 2009). Species 87 distributions are determined by a spectrum of biotic and abiotic factors that act across 88 varying spatial and temporal scales (Anders & Post, 2006; Sexton et al., 2009). Among 89 abiotic factors, climate is thought to be one of the most important determinants of species 90 occurrence and key to the formation, maintenance and evolution of species distributions 91 (Darwin, 1859; Sexton et al., 2009; Araújo & Peterson, 2012). Climate may affect 92 species directly via constraints in physiological tolerances, indirectly via its influence on 93 community assemblages and habitats, or by complex interactions of both (Menge & 94 Olson, 1990). Understanding when and where climate constrains species' occurrence is 95 useful in predicting future responses, conserving and managing species in the face of 96 ongoing global climate change (Pearson & Dawson, 2003; Araújo & Peterson, 2012) and

97 identifying areas where other factors aside from climate are more strongly influencing
98 distributions (e.g., biotic interactions; Urban et al., 2013).

99 Attempts to quantify the role climate plays in shaping species distributions 100 frequently rely on the correlation between species occurrence and climate (i.e., a species' 101 bioclimatic envelope; Araújo & Peterson, 2012) to characterize current and to predict 102 future range dynamics. These static distribution modeling approaches are used to identify 103 broad-scale patterns contributing to range limits (Pearson & Dawson, 2003; Hijmans & 104 Graham, 2006) and to predict range-wide effects of climate change on species 105 distributions (Pearson & Dawson, 2003; Araújo et al., 2005; Thuiller et al., 2005). Under 106 bioclimatic envelope models, climatic conditions where a species is not observed are 107 assumed to prevent establishment of viable populations and thus are the environmental 108 conditions that set range limits (Araújo & Peterson, 2012). However, these 109 phenomenological models assume (1) species' ranges are in equilibrium with climate 110 conditions and (2) species responses are static across the range (Hijmans & Graham, 111 2006; Franklin, 2010). These assumptions do not realistically represent the dynamic 112 nature of the physical environment and the species themselves, especially for broadly 113 distributed species (Zurell et al., 2009). Static models of species responses to climate are 114 insufficient to understand the effect annual climate variation can have on population 115 persistence (Zurell et al., 2009; Franklin, 2010). Furthermore, the focus on species 116 occurrence data ignores the temporal variation in species responses and the demographic 117 processes that determine how a species will respond to climatic shifts (Merow et al., 118 2014; Thuiller et al., 2014).

119 If climate shapes species distributions, changes in climate should have the greatest 120 effect on populations occurring near the climatic extremes (e.g., increased temperature 121 will have the greatest effect on populations in the warmest part of the range; MacArthur, 122 1972; Hoffman & Parsons, 1997; Parmesan et al., 2000). We test this by measuring 123 sensitivity of demographic responses to climatic variation across the range of the wood 124 frog (Lithobates sylvaticus). Specifically, we define sensitivity as the expected change in 125 annual population growth rate (r) with respect to change in an annual climatic measure 126 (e.g., summer extreme heat). We expect range contractions to occur when populations are 127 lost because growth rate is negative for an extended period of time. Similarly, range

128 expansions may occur when populations are gained because growth rate is positive for an 129 extended period of time. Measuring sensitivity tells us how much growth rate is expected 130 to change with a change in average annual conditions and thus how likely population 131 declines (or expansions) are to occur. We test whether sensitivity of population growth 132 rates to year-to-year variation in climate is stronger at the climatic extremes than at the 133 climatic center of a species range (i.e., the bioclimatic envelope prediction; MacArthur, 134 1972; Hoffman & Parsons, 1997; Parmesan et al., 2000). Failure to find evidence to 135 support this hypothesis could result from processes such as local adaptation, biotic 136 interactions, and other abiotic variables leading to different patterns in sensitivity to 137 change. Testing this hypothesis requires an understanding of how life history is impacted 138 and thus how demographic rates respond to climatic variation (e.g., the relationship 139 between population growth rate and temporal variation in environmental conditions; 140 Normand et al., 2014; Ross et al., 2015). This approach captures more of the process 141 underlying range shifts rather than simply the observed pattern that previous correlative 142 approaches have used to predict range shifts.

143 Amphibians make an interesting focal taxon to test the importance of population-144 level sensitivity to climate variation in range dynamics. Amphibians are expected to be 145 particularly sensitive to the effects of climate due to their physiology and life history 146 (Hutchinson & Dupré, 1992; Duellman, 1999), generally limited dispersal abilities 147 (Beebee, 1996, but see Smith & Green, 2005), and reliance on seasonal precipitation and 148 temperature patterns to create breeding habitats and facilitate movement (Pechmann et 149 al., 1989; Rittenhouse & Semlitsch, 2007; Urban et al., 2014). Their ecological 150 importance as a link between terrestrial and aquatic systems (Ranvestel et al., 2004; Earl 151 & Semlitsch, 2012) and the decline of even common species (Stuart et al., 2004; Adams 152 et al., 2013; Grant et al., 2016) make understanding the importance of climate in 153 influencing population level dynamics important for forecasting future extinction risk. 154 Here we focus on the demographic responses to climate for a species of pond-breeding 155 frog, the wood frog, whose range extends across much of northern North America (Fig. 156 1). 157

Using a spatially and temporally rich dataset, we tested the prediction that woodfrog populations are most sensitive to annual climatic variation at sites near the climatic

159 extremes of their distribution (Fig. 2) and that the species distribution is shaped by the 160 interaction of long-term and annual climate conditions on population growth rates. 161 Population growth rates at sites may have three predicted responses based on their range 162 position. For example, if sensitivity of wood frog populations to variation in temperature 163 differs across the range, we predicted that 1) populations in the colder portion of the 164 range (blue; Fig. 2a) would be positively affected by warmer than average annual 165 temperatures, meaning that if warming occurred this could lead to more frequent years of 166 high population growth rates and potential range expansion; 2) populations in the warmer 167 portion of the range (red; Fig. 2a) would be negatively affected by warmer than average 168 annual temperatures, meaning that if warming occurred this would lead to more frequent 169 years of low population growth rates and potential range contraction; and 3) populations 170 in the middle of the range (black; Fig. 2a) are far from climate extremes (Fig. 2b) and 171 annual temperatures would not strongly affect population growth rates. By fitting 172 dynamic models that estimate annual changes in abundance in relation to long-term 173 climate, we can better understand which populations within a species' range are most 174 likely to respond to changing climate.

175 Study System and Methods

176 We build on previous static approaches to model bioclimatic determinants of 177 species distributions (e.g., Guisan & Zimmerman, 2000; Hijmans & Graham, 2006) by 178 measuring local demographic responses of populations using a dynamic species 179 distribution model (DSDM). The DSDM approach allowed us to test the importance of 180 range position in determining responses to climate by measuring local sensitivity of 181 population growth rate to annual variation in climate covariates. Our model takes the 182 form of a hierarchical state-space model (SSM; De Valpine & Hastings, 2002; Buckland 183 et al., 2004; Kéry & Schaub, 2012; Ross et al. 2015), allowing us to link annual 184 population dynamics across different sites and study areas to annual variation in climatic 185 variables. The results provide a measure of climate sensitivity (i.e., the expected change 186 in mean population growth rates in response to changes in mean climate; Thuiller et al., 187 2005; Thomas, 2010; Burrows et al., 2014).

188 Study System and Life History

189 Wood frogs occupy an extensive range, occurring from northern Alaska to 190 Canada and south to the south central United States (USGS National Amphibian Atlas, 191 2014; Fig. 1), spanning a large gradient of climatic conditions (Fig. 1). Specific elements 192 of the wood frog life history potentially make them sensitive to changes in local climate. 193 Breeding normally occurs in early spring when rising temperatures rouse animals and 194 warm spring rains facilitate movement into breeding ponds. Adult frogs show high 195 fidelity to breeding sites (Berven & Grudzien, 1990; Green & Bailey 2015). Breeding 196 generally occurs in a short window of time, anywhere from a few consecutive evenings to 197 a few weeks in length depending on location (Authors, pers. observations; Crouch & 198 Paton, 2000). Female wood frogs become sexually mature between two and four years of 199 age and males between one and three years of age (Berven, 1982a; Berven, 2009; Green 200 & Bailey 2015), and both can live up to six years (Redmer & Trauth, 2005). Females 201 typically lay one egg mass during each breeding season, and these egg masses are 202 visually distinct and easy to locate and count (Crouch & Paton, 2000; Grant et al., 2005; 203 Green et al., 2013). Comparison of census methods show that counts of total egg masses 204 seen per season serves as a suitable proxy for total breeding females per season in a pond 205 (Crouch & Paton, 2000).

206 Field sampling

207 We used egg mass counts from 746 sites within 27 study areas across the wood 208 frog range (Fig. 1; Table S1). A site consisted of a pond or wetland (area ≤ 0.10 ha to 209 5.24 ha) that was visually sampled for wood frog egg masses during the peak of each 210 breeding season and where wood frog egg masses were observed at least once during 211 years when surveys occurred. Study areas designate geographic clusters of sites that 212 occurred within relatively close proximity (e.g., within a single national park). Sites were 213 surveyed in multiple years (range = 3-22 years, mean = 10 years; Table S1) with most, 214 but not all, sites being surveyed multiple times within each year. Surveys occurred during 215 or right after peak breeding based on the lack of calling adults and/or no additional egg 216 masses during subsequent surveys, and a maximum count at a site was recorded each 217 breeding season and used as the response variable in analyses. Wood frog egg masses are 218 conspicuous and detection probability is high ($p = 0.96 \pm 0.02$ to 0.95 ± 0.01 ; Grant et al., 219 2005).

220 Climate covariates

We tested specific predictions with each model about the variation in sensitivity of population growth rates to four climate covariates (Table 1, 2; Fig. S1): 1) spring precipitation (Precip), 2) summer water availability (Hydro), 3) summer extreme heat (Heat), and 4) winter severity (Cold; Table 1). As our sites cover a broad geographic space, wood frog breeding was not synchronous across all study areas. Months used to calculate Precip and Hydro were benchmarked to average breeding dates in each study area, reflecting differences in seasonality across the wood frog range (Table S1).

We obtained global climate normal (~1960-1990, 2.5 arc-minutes resolution) 228 229 rasters from WorldClim (Hijmans et al., 2005) and created 30-year climate normal maps 230 of North America in program R (R Core Team, 2016). We determined 30-year mean 231 annual temperature and precipitation values across North America and within the 232 recorded range of wood frog occurrence (IUCN, 2015) to determine where the species 233 occurs within the broader North American climate space (Fig. 3). These values were used 234 to depict the climate space of wood frogs and our sampled populations in Figure 3 but 235 were not used in SSMs. Using PRISM (Daly et al., 2002) model output for the US and 236 weather station data for Canada (Environment Canada, 2015), we calculated annual 237 climate values for Precip, Hydro, Heat, and Cold at every site every year for SSMs 238 (Table 1). To model differences in long-term climate, we determined 30-year climate 239 normal (average) values (Hijmans et al., 2005) at every site over the same seasonal 240 periods as our annual climate covariates for SSMs (Table 2; nmPrecip, nmHydro, 241 nmHeat, and nmCold). For example, at northern sites we calculated total precipitation 242 values each year for February, March, and April, due to their importance in timing wood 243 frog migrations and pond filling, and averaged them for an annual spring precipitation 244 value (Precip). We then averaged total precipitation values over the same months across 245 30 years to get a long-term climate normal value (nmPrecip) that varied across but not 246 within sites. Annual climate values were standardized by 33-year (1981-2013) mean and 247 standard deviations at a site. Climate normals were standardized using the mean and 248 standard deviation from the entire extent of the wood frog range.

249 Data analysis

250 We used SSMs to estimate the effect of annual variation on population growth 251 rate (De Valpine & Hastings, 2002; Buckland et al., 2004; Kéry & Schaub, 2011; Ross et 252 al. 2015). Models were fit in JAGS (Plummer 2003) and implemented in program R via 253 the R2jags package (Su & Yajima, 2012; see Appendix S1 for JAGS code). The 254 hierarchical model allows for estimation of latent state and observation processes 255 characterizing sampled populations while simultaneously accounting for process 256 variation and observation error (Buckland et al., 2004; Kéry & Schaub, 2011). We were 257 interested in understanding how these latent processes were affected by annual climate 258 variation across the range. At the same time, the modeling framework allowed us to 259 account for observation error in counts (e.g., through variable detection, field conditions, 260 variable observer expertise) that was unrelated to the underlying population processes 261 (MacKenzie et al., 2006).

We described changes in wood frog population size (as based on egg mass counts that serve as a proxy for number of breeding females in a season) using an exponential population growth model

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$$\mathbf{N}_{t+1} = \mathbf{N}_t * \mathbf{e}^r \tag{1}$$

where population size N_{t+1} is a function of the previous population size N_t (from the previous year) and the per capita annual growth rate (r, the exponent of the instantaneous growth rate). Using this as a starting point, we estimated regression coefficients characterizing the relationship between annual weather and the realized growth rate (r_{ti}) for a given year (t) and a given site (i) for each climate hypothesis.

To fit the model, we reformulated Eq. 1 by taking the natural logarithm of each side of the equation and indexing all parameters by year (t) and site (i) to capture annual and site-specific variation in the climate covariates and population responses. We added one to all observations to accommodate zeros in the data prior to log transformation.

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$$\log(N_{t+1, i}) = \log(N_{ti}) + r_{ti}$$
(2)

The now additive growth rate r_{ti} was modified to include the effects of climate covariates and unexplained annual variation captured using random-error terms. Our goal was to estimate the effect of annual variation in each of our four climate covariates and how those effects differed across the range. We estimated these relationships using a linear model that included the main effects of annual climate values and the climatenormal along with the interaction of the two (Table 3). The model took the form of:

282 283 $\mathbf{r}_{ti} = \beta_1 * \text{Annual Climate}_{ti} + \beta_2 * \text{Climate Normal}_i +$ (3)

 β_3 * Annual Climate_{ti} * Climate Normal_i + δ_i + ε_{ti}

284 The model allowed us to determine sensitivity, defined as the expected change in annual 285 population growth rate (r) for a one standard deviation increase in annual conditions, to 286 long-term climate normal conditions. Specifically, the interaction term allowed us to 287 quantify the amount of change in population growth rates given an annual shift in climate in respect to the climatic range position in which a population exists. We included a 288 random effect for site level differences, $\delta_i \sim \text{Normal}(0, \sigma^2_{\text{Site}})$. This effect served as the 289 290 local site-level intercept for growth rate, which we expected to vary around a mean value 291 of 0. We included a second random error component for additional annual variation in growth rate not explained by the climate covariates, $\varepsilon_{ti} \sim \text{Normal}(0, \sigma^2_{\text{proc}})$. To account 292 293 for observation error in counts that was not explained by the population level state 294 processes, we assumed that the log observed count of egg masses for that site and year, y_{ti} , is given by $y_{ti} \sim Normal (log[N_{ti}], \sigma^2_{obs})$. It was also necessary to estimate a starting 295 296 population size for each site. We used a prior value of $\log (N_{1i}) \sim Normal (0, 100)$. We used vague priors for random effect variance components ($\sigma^2_{obs}, \sigma^2_{proc}$) with 297 298 uniform distributions bounded between 0 and 5. For σ^2_{Site} we used a uniform prior 299 bounded between 0 and 0.2 to facilitate convergence. Priors for all regression coefficients 300 were $\beta_k \sim \text{Normal}(0, 100)$. We ran three parallel chains for 50,000 iterations each and

301 discarded the first 1000 iterations as burn-in to allow for model convergence. Model 302 convergence was determined visually from traceplots and Gelman Rubin statistics ($\hat{R} <$

302 convergence was determined visually from traceplots and German Rubin statistics (*R*303 1.05; Gelman & Rubin 1992).

We predicted that climate covariates could have both immediate and lagged effects on annual growth rate (r_{ti} ; Fig. S1). We predicted that covariates that disproportionately impact adult survival and season-to-season variation in breeding would lead to changes in growth rate in the same year. In the case where we expected a covariate to impact the survival of eggs and tadpoles in a wetland and thus the number of potential recruits from a cohort, these were predicted to lead to changes in growth rates after a 2-year lag. Female wood frogs take approximately two years to reach sexual

311 maturity in our study sites (e.g., lowland populations; Berven, 1982a, 2009; Green & 312 Bailey 2015). Therefore the effects of reproductive failure (e.g., desiccation of tadpoles in 313 a dry year) on growth rates would not be evident in counts of egg masses in the year 314 immediately following these suboptimal conditions. We hypothesized that annual Precip 315 values affect movement of adult animals and the opportunity for successful oviposition 316 (i.e., pond filling), with low Precip values resulting in fewer egg masses laid and thus 317 reduced recruitment two years later. Hydro values reflect desiccation risk for developing 318 tadpoles (realized as altered recruitment two years later) and also drier summer 319 conditions that can decrease adult survival during foraging or return to overwintering 320 sites. Heat and Cold values reflect late summer dryness and overwintering cold stress 321 expected to impact adults. While any number of time lag combinations and effects are 322 possible, we fit the model (Eq. 3) focusing on these key periods due to their biological 323 importance and support in the literature (Table 1).

324 We were also interested in how water availability and temperature may interact to 325 explain variation in climate sensitivity. We expected that years of low precipitation 326 (Precip) would have a greater negative effect in sites with higher mean annual summer 327 temperatures (e.g., hotter areas; nmHeat) as increased water on the landscape may help 328 keep permeable amphibian skin moist and lessen desiccation risk (Rittenhouse et al., 329 2009; Köhler et al., 2011). Similarly, we expected reduced winter severity (Cold) and its 330 indirect effect on water availability and pond filling in the spring to be greater in areas 331 that receive less spring precipitation (e.g., drier areas; nmPrecip). We tested for these 332 effects by including the interaction of different annual and long-term climate covariates 333 (e.g., Precip*nmCold, Hydro*nmHeat; Table 3). Annual covariates included the same 334 time lags as previously discussed. Models with both temperature and precipitation 335 included all annual and long-term covariates for each climate measure and an additional 336 two interaction terms allowing annual and long-term covariates to interact (Table 3). This 337 means a total of eight models testing climate hypotheses (Table 1 and 2) were run. None 338 of the selected climate covariates were strongly correlated ($|\mathbf{r}| < 0.4$). 339 When fitting models, we tested for goodness of fit using a posterior predictive

340 check to test whether observed variability in counts was consistent with expected

341 variation. We calculated observed variance in our data for each of the sites and

342 determined if on average variance was less than or greater than the predicted variance of

343 simulated data based on our model. We report the proportion of the time that the

344 observed variance was greater than the predicted variance, with the expectation that if the

model fits the data well we expect this proportion to be 0.5.

Additionally, we were interested in estimating the overall expected rate of change in wood frog population growth rates, dr/dt, based on our estimated climate relationships. Expected change is a function of the local sensitivity to each of our climate covariates, dr/dX, as measured in our models as well as the rate of change in mean climate over that time period, dX/dt, where:

$$\frac{dr}{dt} = \frac{dr}{dX}\frac{dX}{dt}$$

We calculated rate of change in each of our climate variables at each of our sites using linear regression where year was the predictor variable and annual values of each of our climate variables over a 30-year period from 1984 to 2013 were our response variables. We mapped these to geographic and climate space to highlight areas where climate may currently be altering the wood frog distribution.

356 Results

357 Sites spanned a >23 degree range in latitude and >50 degree range in longitude from 358 North Carolina to Jasper National Park, Alberta, Canada. Study areas fell into 16 359 different states, one administrative subdivision (Washington D.C.) and one Canadian 360 province (Alberta) (Fig.1). Our data show good geographic coverage along the wood 361 frog's southern and easternmost range limit but are restricted in geographic coverage in 362 the northern and westernmost portions of the wood frog range. This was reflected in our 363 coverage in climate space (Fig. 3), with best coverage in the portion of the range with 364 warmer temperatures and higher precipitation. Therefore, we limit the presentation of 365 results and their interpretation to only the sampled portion of the wood frog range. 366 Additionally, support for models was judged by whether or not credible intervals of 367 parameter estimates overlapped zero, and we have limited our presentation of results to 368 those models with the strongest support and thus credible intervals for interaction terms 369 that did not overlap. Our posterior-predictive check values for each of our models were

between 0.493 and 0.541, indicating that our models did a good job of capturing actualvariation in growth rates.

372 Our first three models tested the effect of moisture on population growth rates, 373 with the first focused on spring precipitation and the second on late summer water 374 availability. Contrary to our predictions, we found a negative relationship between Precip 375 and wood frog population growth rates across all areas two years later (Table 4; Fig. 4a; 376 Fig. S2). The relationship of increased annual Hydro values to wood frog population 377 growth rates differed depending on if a time lag was incorporated (Table 4; Fig. 4b,c; Fig. S2). The same-year effect of Hydro was dependent on long-term climate, with 378 379 populations in wetter areas responding most positively to wetter annual conditions as 380 compared to those in drier areas (Fig. 4b). When incorporating a two-year time lag, 381 increased values of Hydro were positively associated with growth rates only in drier areas 382 (Fig. 4c; Table 4), agreeing with our bioclimatic envelope predictions of increased 383 sensitivity to water availability in drier portions of the range.

384 Our next two models focused on the effect of extreme heat and cold severity on 385 population growth rates. The relationship between increased values of Heat and wood 386 frog population growth rates (Table 4) depended upon long-term climate. Years with 387 hotter summer temperatures had higher population growth rates in areas with cooler 388 summer climates. However, there was a negative association between warmer summers 389 and population growth in areas with hotter summer climates (Fig. 4d; Fig. S2). This 390 agrees with our bioclimatic envelope prediction, where we expect population growth rate 391 to be most sensitive to warming in the warmest portion of the range. The relationship of 392 Cold to population growth rates showed increased growth rates associated with milder 393 winters across all areas (Table 4) with the most positive association in areas with milder 394 winter climates (Fig 4e; Fig. S2).

Finally we examined how precipitation and temperature interacted to affect
population growth rates. We found that the two-year lag effect of annual variation in
spring precipitation did not depend on long-term winter climate (Precip*nmCold; Table
5; Fig. 5a; Fig. S2), and the effect of annual variation in winter severity did not vary
significantly by long-term spring precipitation (Cold*nmPrecip; Table 5; Fig. S2).

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402 the effect of warmer summers differed between drier and wetter areas. Hotter summer 403 temperatures had a positive relationship to population growth rates in wetter areas but 404 negatively impacted growth rates in drier areas (Heat*nmHydro; Fig. 5d; Fig. S2). We 405 found a positive relationship between increased summer water availability and wood frog 406 growth rates two years later in areas with cooler summer temperatures but a negative 407 relationship in areas with hotter summer temperatures (Table 6; Fig. 5e; Fig. S2). The 408 interaction of increased summer temperatures had a similar impact on population growth 409 rates as the summer water availability model with no time lag, with a positive effect of 410 increased summer heat in wetter versus drier areas (Table 6).

Expected rate of change in population growth rates over the previous 30 years that could be attributed to changes in climate showed few major increases or decreases across the wood frog range (Fig. S3, S4). The biggest changes in population growth rates were estimated to have occurred for variables related to temperature. These suggest some reductions in growth rates in the southern portion of the wood frog range due to changes in heat and cold (Fig. S3d,e; S4d,e).

417 Discussion

418 We tested the prediction that the effect of climate on population growth rates 419 varies in a predictable pattern based on local, long-term climate (i.e., bioclimatic 420 envelope prediction; MacArthur, 1972; Hoffman & Parsons, 1997; Parmesan et al., 421 2000). Populations near the climatic extremes of the species range were predicted to be 422 the most sensitive to annual variation in climate. Our use of hierarchical SSMs (De 423 Valpine & Hastings, 2002; Kéry & Schaub, 2012) allowed us test this broad-scale 424 prediction by simultaneously linking climate directly to demographic rates at the 425 temporal- (short-term variation in weather) and spatial- (individual populations) scales at 426 which climate acts to affect species distributions. We acknowledge that our sampled sites 427 are only a portion of the wood frog range and thus limit the interpretation of our results to 428 conditions represented in this study. Our results provided mixed evidence to support this 429 prediction, with differences in climate sensitivity often occurring in the opposite direction 430 of this prediction. For example, the effect of summer temperature was consistent with our

prediction – warmer summers had a more detrimental effect in the warmest part of the
range. The effect of summer water availability was also consistent with this prediction,
where increased moisture had a positive effect two years later in drier areas. On the other
hand, variation in spring precipitation, summer water availability in the current year, and
winter severity did not conform to predictions based on position within the range.

436 We also tested the climate sensitivity of populations to interactions of temperature 437 and precipitation. We again predicted that population growth rates would be most 438 sensitive to annual variation in one factor (e.g., increased summer heat) as they 439 approached climate extremes of the other (e.g., drier areas). Again, we found mixed 440 support for this prediction. Hotter summers had a positive effect on wood frog growth 441 rates in wetter areas but a negative effect in drier areas as predicted. However, we found a 442 contradictory positive effect of increased summer water availability two years later in 443 cooler areas and no significant association of spring precipitation and winter severity to wood frog population growth rates. This suggests that expected shifts due to changing 444 445 climate for wood frogs may not be strongest at the climatic extremes of the range or 446 easily predicted solely by climate, which is surprising given the expected sensitivity of 447 amphibians to abiotic conditions.

448 Many of the metabolic, reproductive and phenological processes in amphibians 449 are strongly linked to temperature (Berven, 1982a,b; Beebee, 1996; Gibbs & Breisch, 450 2001) and can be of key importance in structuring species distributions (Tingley et al., 451 2009; Cahill et al., 2014). This may explain why bioclimatic envelope model predictions 452 regarding temperature, specifically heat, were better supported in our models. 453 Temperature may have a more uniform effect across the landscape and may be better 454 represented by coarse measures. Alternatively, precipitation largely acts through its effect 455 on hydrological processes during the reproductive phase and interactions between water, 456 soil, and vegetation during non-breeding periods (Drexler et al., 2004; Bauder, 2005; 457 Davis et al., in prep). Hydrologic deficits (Brooks, 2004), landscape topography (Boswell 458 & Olyphant, 2007), pond-selection by breeding animals (Pechmann et al., 1989; Skidds 459 et al., 2007; Amburgey et al., 2014), and plasticity in development (Relyea, 2002; 460 Amburgey et al., 2012) are among the many factors that may attenuate the relationships 461 between water availability and amphibian population growth rates. Our inferences are

also limited to the study area that we were able to sample. Limited sampling of the colder
and drier edge of climate space (Fig. 3) may restrict our ability to detect relationships
occurring at those extremes. Our study did, however, provide good coverage at the warm
and wet edge of the wood frog range, which is most susceptible to the effects of climate
change (Corn 2005; Meehl et al., 2007).

467 A multitude of other factors (e.g., local adaptation, biotic interactions, and other 468 abiotic variables) can affect populations and lead to patterns contradictory to bioclimatic 469 envelope predictions of climate sensitivity (HilleRisLambers et al., 2013; Urban et al., 470 2013). The effect of moisture on the landscape likely depends on the form and the timing 471 of precipitation and can also impact biotic factors that likewise contribute to 472 heterogeneity in population growth rates. Increased spring precipitation may come as 473 early spring snow and ice storms that can increase adult mortality through reduced freeze 474 tolerance (Costanzo & Lee Jr., 1992) or truncate the breeding season (Berven, 1982b). 475 Increased moisture on the landscape may increase the probability of egg mass or tadpole 476 stranding in temporary flooded areas or facilitate colonization or persistence of predators 477 in ponds (Werner et al., 2009). Local adaptation to annual climate variation may alter 478 climate sensitivity, with populations nearer to climate extremes accustomed to increased 479 annual variation while populations farther away from extremes are not (e.g., Berven, 480 1982a, Laugen et al., 2003; Amburgey et al., 2012), though we cannot test this directly 481 with our approach. Local dynamics may also vary spatially, where populations near 482 climate extremes are at low enough densities that they are unable to respond to the 483 benefits of years with more suitable climate conditions.

484 Species biology may additionally structure population responses to climate and 485 result in deviations from bioclimatic envelope predictions. Wood frogs are freeze tolerant 486 (Storey & Storey, 1986; Costanzo & Lee Jr., 1992) though extended or extreme periods 487 of freezing temperatures can impact overwintering survival (Costanzo et al., 1991; 488 O'Connor & Rittenhouse, 2016). In a portion of the range that encompassed our study 489 areas, no differentiation in wood frog thermal tolerance was found (Manis & Claussen, 490 1986); however, far northern populations in Alaska have shown increased cold tolerance 491 (Larson et al., 2014). However, mild winters in colder areas may result in freeze-thaw 492 cycles that rouse animals from torpor, resulting in increased energetic demands (Storey,

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1987), mating behavior impairment (Costanzo et al., 1997), and reduced fecundity
(Benard, 2015). Additionally, the life stage on which climate most strongly acts may
influence the population response. In amphibians, the aquatic larval stage already
experiences heightened mortality, and climate conditions that affects tadpole survival
may not lead to differential climate sensitivity at the population level as much as those
factors that influence terrestrial juvenile and adult survival (Biek et al., 2002; Harper et
al., 2008).

500 Currently, species distributions and range dynamics are frequently modeled using 501 static approaches that treat climate and species responses as fixed across space and time 502 (Hijmans & Graham, 2006; Franklin, 2010). However, species responses to climate are 503 spatially complex, especially for those with multistage life histories. Climate shifts will 504 likely alter species distributions by acting on demographic processes where sensitivity to 505 change is greatest. Combining estimates of climate sensitivity with data about observed 506 or predicted changes in climate allows for predictions about local changes in population 507 growth rate to be made. We did this for the last 30-year period, highlighting the 508 variability in population response across the range (Fig. S3, S4). Demographic response 509 for some climate variables fit predictions (e.g., negative responses to warming in the 510 warmest regions). However, estimated demographic changes related to water availability 511 and interactions with temperature follow much less clear patterns, which would not easily 512 be predicted using static modeling approaches. Our results demonstrate that focusing on 513 demographic processes provides insights for understanding how species distributions 514 may respond to change not possible with presence-absence correlative models focused on 515 pattern (Normand et al., 2014; Ross et al., 2015). Correlative approaches based on a 516 static snapshot of species distribution do not measure the actual mechanistic processes 517 impacting populations (Dormann et al., 2012; Cahill et al., 2014) and do not estimate 518 rates of change that demographic models can incorporate (Normand et al., 2014). Thus, 519 correlations may break down with no-analogue climates (Williams & Jackson, 2007) and 520 lack the predictive power explicit estimates of climate-demography relationships can 521 offer (Normand et al., 2014). While our model is still correlative in relating demographic 522 rates to climatic variation, it provides a finer scale approach that provides insights to 523 potential mechanisms while also explaining broader patterns. Bioclimatic envelope

524 modeling does not include other potentially important factors (e.g., biotic interactions,

525 genetic differentiation, and geographical barriers) that may set species range limits alone

or in concert with climate (HilleRisLambers et al., 2013; Urban et al., 2013). However,
such demographic models can be modified to include such information and better inform
our understanding of species range dynamics.

529 A demographic understanding of species distributions is essential to evaluating 530 and understanding range limits, forecasting range shifts and stability, and managing 531 species and conserving habitats. These aims will be critical in the context of changing climate. By pairing large-scale modeling studies with targeted experimental or 532 533 demographic studies, we can better understand the way these broad-scale measures are 534 realized on the landscape and influence local populations (Merow et al., 2014; Normand 535 et al., 2014). In the future, all species are likely to experience some change to their 536 current distributions, whether through range contractions (via altered habitat suitability 537 through changing climate) or expansions (via altered climate facilitating colonization of 538 new habitats; Thuiller et al., 2008). With increasingly limited conservation resources, 539 identification and prioritization of critical areas where species are most sensitive to 540 changing climate (Beissinger & Westphal, 1998; Keith et al., 2008) and where range 541 shifts may occur (Thuiller et al., 2008) will allow for more efficient and effective 542 conservation management.

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Figure 1. The wood frog is a broadly distributed species that spans most of the
northeastern United States into Canada and Alaska. Red dots indicate sites where egg
mass counts were obtained. Thirty-year annual (a) precipitation and (b) temperature
(Hijmans et al. 2005) maps show the broad range of climate conditions this species
experiences across its range (IUCN 2014).

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807 Figure 2. (a) The wood frog range (light grey) with an example of a northern (blue), 808 central (black), and southern (red) population. (b) These populations come from different 809 long-term climate normals (e.g., colder to warmer represented by mean 30-year 810 temperature). If wood frog responses are consistent with bioclimatic envelope 811 predictions, the probability of occurrence of wood frogs peaks at some optimal 812 temperature and declines in more extreme conditions. (c) Sensitivity of wood frog 813 population growth rates to annual climate variation is predicted to vary by long-term 814 climate (shaded regions are 95% credible intervals). Sensitivity is the expected change in 815 annual population growth rate (r) for a one standard deviation increase in annual 816 conditions. We predict 1) populations in colder areas (blue) will be sensitive to warmer 817 than average annual temperatures, leading to higher population growth rates (positive 818 values); 2) populations in hotter areas (red) will be sensitive to warmer than average 819 annual temperatures, leading to lower population growth rates (negative values); and 3) 820 populations in areas far from climate extremes (black) will not be strongly affected by 821 year to year deviations in temperature, leading to fairly consistent population growth 822 rates (values around zero).

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Figure 3. The climate space [based on 30-year mean annual temperature (°C*10) and precipitation values (mm)] that encompasses North America (dark blue), the wood frog

range (light green), and our sites (red). Points on the scatterplot represent all temperature
by precipitation raster cell values where wood frogs occur (light green) and do not occur
(dark blue), with our sites in red. Precipitation values were truncated at 3500 mm for
visualization purposes. Histograms represent frequencies of these same 30-year annual
precipitation (top) and temperature (right) values in just the wood frog range. Boxplots of
precipitation and temperature values from our sites show the minimum, median,
maximum and 25th and 75th quartiles (box).

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834 Figure 4. We estimated how sensitivity of wood frog population growth rates to annual 835 climate variation changed with respect to long-term climate differences (shaded regions 836 are 95% credible intervals). Sensitivity is the expected change in annual population 837 growth rate (r) for a one standard deviation increase in annual conditions (y-axis). Long-838 term differences in mean climate are calculated using 30-year climate normals for 839 conditions during the same portion of the year that annual covariates are measured (x-840 axis; see Tables 1 and 2) at our sampled sites. (a) Annual wood frog population growth 841 rate two years later responded negatively to spring precipitation (PRECIP lag) across all 842 areas, (b) annual wood frog population growth rate responded positively to years with 843 more summer water availability (HYDRO) in areas where long-term average summer 844 precipitation was higher (>50 mm), (c) annual wood frog population growth rate two 845 years later responded negatively to years with more summer water availability (HYDRO 846 lag) in areas where long-term average summer precipitation was higher (>105 mm) and positively in years where long-term averages were lower (<105 mm), (d) annual wood 847 848 frog population growth rate responded negatively to extreme summer temperatures 849 (HEAT) in areas where long-term average extreme temperature was higher (>24°C) and 850 positively where long-term averages were lower ($<24^{\circ}$ C), (e) annual wood frog 851 population growth rate responded positively to increased winter severity (COLD) in areas 852 where long-term average minimum temperature was milder (>-6.25°C) and negatively 853 where long-term averages were colder (<-6.25°C). 854

Figure 5. We estimated how sensitivity of wood frog population growth rates to annual
climate variation changed with respect to long-term climate differences (shaded regions)

857 are 95% credible intervals). Sensitivity is the expected change in annual population 858 growth rate (r) for a one standard deviation increase in annual conditions (y-axis). Long-859 term differences in mean climate are calculated using 30-year climate normals for 860 conditions during the same portion of the year that annual covariates are measured (x-861 axis; see Tables 1 and 2) at our sampled sites. (a) Annual wood frog population growth 862 rate two years later did not significantly respond to spring precipitation (PRECIP lag) 863 regardless of long-term winter severity, (b) annual wood frog population growth rate did 864 not significantly respond to winter severity (COLD) regardless of long-term spring 865 precipitation, (c) annual wood frog population growth rate did not significantly respond 866 to summer water availability (HYDRO) regardless of long-term extreme summer heat, 867 (d) annual wood frog population growth rate responded positively to years with more 868 extreme summer temperatures (HYDRO lag) in areas where long-term average summer 869 precipitation was higher (>20 mm), (e) annual wood frog population growth rate two 870 years later responded negatively to increased summer water availability (HYDRO lag) in 871 areas where long-term average extreme temperature was higher (>26.25°C) and 872 positively where long-term averages were lower (<26.25°C).

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1 **Table 1.** Annual climate covariates selected for state space models based on their potential importance in wood frog breeding and survival. Annual

2 values at each site were used in modeling the effect of annual climate variation on wood frog population growth rates.

Covariate	Definition	Ecological Importance
Precip =	Deviation of the observed precipitation value	Values represent the wetness of an area during the start of spring
Standardized Precipitation	from the estimated median for an area	breeding (e.g., Feb-Apr) such that a more positive value indicates more
Index 3-month (SPI3) ¹	calculated over a 3-month period, uses only	precipitation than predicted. Spring precipitation is important as a cue for
0	precipitation values (only inputs to the system)	breeding adults to migrate to ponds and for filling ephemeral ponds ^{2,3}
Hydro =	Deviation of the observed precipitation value	Similar to SPI3 but includes the effect of temperature on
Standardized Precipitation	from the estimated median for an area	evapotranspiration rates, considers the way these rates will influence
Evapotranspiration Index 3-	calculated over a 3-month period, uses	drought severity and can be used as a measure of water available on the
month (SPEI3) ⁴	precipitation and evapotranspiration values	landscape, calculated during summer (e.g., May-July) to get at pond
m	(inputs and outputs to the system)	drying. Hydroperiod impacts desiccation risk of tadpoles and can
		approximate dry summers that increase desiccation risk of adults ⁵
Heat =	Hottest 10-day average temperature, falls in the	Periods of intense heat increase the risk of heat stress and desiccation
Extreme Heat Index (EHI) ⁴	late summer for North America	while moving between sites ^{3,5}
Cold =	Cumulative index of freeze severity and frost	Though freeze tolerance has been demonstrated in this species ⁷ , extreme
Air Freezing Index (AFI) ⁴	depth that factors in magnitude and duration of	cold temperatures and long durations of cold temperatures may reduce
0	below freezing air temperatures ⁶	overwinter survival of juveniles and adults ⁸
¹ National Climatic Data Center, N	NOAA (2015)	⁵ Brooks (2004)
² Rittenhouse et al. (2009)		⁶ Bilotta et al. (2015)
³ Davis et al. (in prep)		⁷ Storey and Storey (1986)
⁴ Daly et al. (2002)		⁸ O'Connor and Rittenhouse (2016)
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1 Table 2. Thirty-year normal climate covariates selected for state space models to account for long-term effects of climate at a site (i.e.,

2 values are constant over time). Their interaction with annual climate covariate values (Table 1) indicated if population growth rates differ

3 in sensitivity across the range. The predicted relationship of the interaction between annual and long-term climate covariates to population

4 growth rates across the wood frog range represent hypotheses from the bioclimatic envelope model.

Comminto	Definition	Factorial Importance	Predicted Annual and Long-term
Covariate	Definition	Ecological Importance	Interaction (Bioclimatic Envelope Model)
nmPrecip =	30-year mean monthly	Measure of precipitation and water availability	Precip * nmPrecip
Precip Normal ⁸	precipitation over same 3-	during spring breeding, long term moisture	Negative impact of drier years in drier areas
<u> </u>	month period as SPI3	dynamics of areas	
nmHydro =	30-year mean monthly	Measure of precipitation and water availability	Hydro * nmHydro
Hydroperiod Normal ⁸	precipitation over same 3-	during tadpole development, long term moisture	Negative impact of drier years in drier areas
	month period as SPEI3	dynamics of areas	
nmHeat =	30-year maximum monthly	Measure of extreme heat patterns occurring during	Heat * nmHeat
Heat Normal ⁸	temperature over similar late	the late summer, long term heat regime	Negative impact of hotter years in hotter
	summer period as EHI		areas
nmCold =	30-year minimum monthly	Measure of winter severity patterns, long term cold	Cold * nmCold
Cold Normal ⁸	temperature over similar mid-	regime	Negative impact of colder years in colder
5	winter period as AFI		areas

5 ⁸ WorldClim; Hijmans et al. (2005)



Table 3. All candidate state space models investigated for modeling wood frog egg mass counts. Each main model consists of the annual climate covariate (Precip, Hydro, Heat, Cold), the respective long-term climate normal (nmPrecip, nmHydro, nmHeat, nmCold), and the interaction between each annual and long-term covariate. Combination models are those with additional crossed interactions between annual climate covariates and long-term climate normals representing a different climate component (e.g., Hydro*nmHeat investigates the interaction between annual summer precipitation by long-term late summer maximum temperatures). The random effects of site (δ_i) and observation error (ε_{ti}) were included in all models.

Model Name	Parameters
Precip (2-yr lag)	$\beta_1(\operatorname{Precip}_{2yr}) + \beta_2(\operatorname{nmPrecip}) + \beta_3(\operatorname{Precip}_{2yr} * \operatorname{nmPrecip}) + \delta_i + \varepsilon_{ti}$
Hydro	β_1 (Hydro) + β_2 (nmHydro) + β_3 (Hydro * nmHydro) + δ_i + ε_{ti}
Hydro (2-yr lag)	$\beta_1(\text{Hydro}_{2yr}) + \beta_2(\text{nmHydro}) + \beta_3(\text{Hydro}_{2yr} * \text{nmHydro}) + \delta_i + \varepsilon_{ti}$
Heat	$\beta_1(\text{Heat}) + \beta_2(\text{nmHeat}) + \beta_3(\text{Heat} * \text{nmHeat}) + \delta_i + \varepsilon_{ti}$
Cold	$\beta_1(\text{Cold}) + \beta_2(\text{nmCold}) + \beta_3(\text{Cold} * \text{nmCold}) + \delta_i + \varepsilon_{ti}$
Precip (2-yr lag) and Cold	$\beta_1(\operatorname{Precip}_{2yr}) + \beta_2(\operatorname{nmPrecip}) + \beta_3(\operatorname{Precip}_{2yr} * \operatorname{nmPrecip}) + \beta_4(\operatorname{Cold}) + \beta_5(\operatorname{nmCold}) + \beta_6(\operatorname{Cold} * \operatorname{nmCold}) + \beta_6(\operatorname{Cold} * $
by long-term climate	$\beta_7(\text{Precip}_{2yr}* \text{nmCold}) + \beta_8(\text{Cold}* \text{nmPrecip}) + \delta_i + \varepsilon_{ti}$
Hydro and Heat by long-	$\beta_1(\text{Hydro}) + \beta_2(\text{nmHydro}) + \beta_3(\text{Hydro} * \text{nmHydro}) + \beta_4(\text{Heat}) + \beta_5(\text{nmHeat}) + \beta_6(\text{Heat} * \text{nmHeat}) + \beta_6(\text$
term climate	β_7 (Hydro * nmHeat) + β_8 (Heat * nmHydro) + δ_i + ε_{ii}
Hydro (2-yr lag) and Heat	$\beta_1(\text{Hydro}_{2yr}) + \beta_2(\text{nmHydro}) + \beta_3(\text{Hydro}_{2yr} * \text{nmHydro}) + \beta_4(\text{Heat}) + \beta_5(\text{nmHeat}) + \beta_6(\text{Heat} * \text{nmHeat}) + \beta_6(\text{Heat} * nmHe$
by long-term climate	β_7 (Hydro _{2yr} * nmHeat) + β_8 (Heat * nmHydro) + δ_i + ε_{ti}

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Table 4. Parameter estimates from the four main climate covariate models. Precip, Hydro, Heat,
and Cold represent annual climate values. nmPrecip, nmHydro, nmHeat, and nmCold are the
long-term (~30 year) climate normal values. Interaction terms of annual and normal values (e.g.,
Precip*nmPrecip) represent the effect of an annual climate value by different long-term climate.
SD is the standard deviation of a parameter estimate, and q_{0.025-0.975} represent 2.5th, 50th, and
97.5th quartile values.

Model: Precip					
Parameter	Mean	SD	q _{0.025}	Q 0.500	Q 0.975
Precip (2-yr lag)	-0.0840	0.0370	-0.155	-0.0840	-0.0120
nmPrecip	4.00e-03	4.00e-03	-4.00e-03	4.00e-03	0.0130
Precip (2-yr lag)*	4.00e-03	0.0140	-0.0240	4.00e-04	0.0320
nmPrecip					
Model: Hydro					
Hydro	0.0533	0.0387	-0.0229	0.0534	0.129
nmHydro	-0.0137	5.44e-03	-0.0243	-0.0137	-2.98e-03
Hydro* nmHydro	0.0732	0.0244	0.0255	0.0732	0.121
Hydro (2-yr lag)	0.0925	0.0415	0.0111	0.0924	0.174
nmHydro	-7.21e-03	5.47e-03	-0.0179	-7.22e-03	3.57e-03
Hydro (2-yr lag)*	-0.0605	0.0253	-0.110	-0.0605	-0.0109
nmHydro					
Model: Heat					
Heat	0.340	0.0670	0.208	0.340	0.471
nmHeat	7.50e-03	5.46e-03	-3.12e-03	7.51e-03	0.0183
Heat* nmHeat	-0.266	0.0383	-0.341	-0.266	-0.191
Model: Cold					
Cold	-0.438	0.117	-0.666	-0.438	-0.209
nmCold	0.0115	4.81e-03	2.09e-03	0.0115	0.0210
Cold* nmCold	0.258	0.0580	0.145	0.259	0.372

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1 Table 5. Parameter estimates from the interaction model of spring precipitation and winter

- 2 severity. Precip and Cold represent annual climate values. nmPrecip and nmCold are the long-
- 3 term (~30 year) climate normal values. Interaction terms of annual and normal values (e.g.,
- Precip*nmPrecip) represent the effect of an annual climate value by different long-term climate. 4

SD is the standard deviation of a parameter estimate, and $q_{0.025-0.975}$ represent 2.5th, 50th, and 5

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97.5th quartile values. 6

Model: Precip (2-yr lag) and Cold by long-term climate conditions							
Parameter	Mean	SD	q _{0.025}	Q 0.500	q 0.975		
Precip (2-yr lag)	-0.192	0.126	-0.439	-0.192	0.055		
nmPrecip	-8.00e-03	0.0110	-0.0300	-8.00e-03	0.0140		
Precip (2-yr lag)*	9.00e-03	0.0170	-0.0240	9.00e-03	0.0420		
nmPrecip							
Cold	-0.398	0.125	-0.644	-0.398	-0.152		
nmCold	0.0210	0.0130	-4.00e-03	0.0210	0.0460		
Cold* nmCold	0.220	0.0630	0.0970	0.220	0.344		
Precip (2-yr lag)*	0.0620	0.0700	-0.0760	0.0620	0.200		
nmCold							
Cold*nmPrecip	0.0150	0.0240	-0.0320	0.0150	0.0630		

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Table 6. Parameter estimates from the interaction model of summer water availability and
 extreme heat. Hydro and Heat represent annual climate values. nmHydro and nmHeat are the
 long-term (~30 year) climate normal values. Interaction terms of annual and normal values (e.g.,
 Hydro*nmHydro) represent the effect of an annual climate value by different long-term climate.

5 SD is the standard deviation of a parameter estimate, and $q_{0.025-0.975}$ represent 2.5th, 50th, and

6 97.5th quartile values.

Nodel: Hydro and Heat by long-term climate conditions							
Parameter	Mean	SD	q _{0.025}	Q 0.500	Q 0.975		
Hydro	-0.0153	0.0701	-0.152	-0.0155	1.22e-01		
nmHydro	-0.0404	0.0117	-0.0634	-0.0405	-1.74e-02		
Hydro*nmHydro	0.0736	0.0267	0.0210	0.0736	1.26e-01		
Heat	0.202	0.0741	0.0562	0.202	3.47e-01		
nmHeat	0.0383	0.0117	0.0153	0.0383	6.12e-02		
Heat*nmHeat	-0.259	0.0415	-0.340	-0.259	-1.77e-01		
Hydro*nmHeat	0.0191	0.0474	-0.0742	0.0191	1.11e-01		
Heat*nmHydro	0.0913	0.0269	0.0384	0.0914	1.44e-01		
Model: Hydro (2-yr lag) and Heat by long-term climate conditions							
Hydro (2-yr lag)	0.333	0.0787	0.179	0.333	0.487		
nmHydro	-0.0354	0.0118	-0.0585	-0.0354	-0.0122		
Hydro (2-yr lag)*	-0.0120	0.0263	-0.0635	-0.0120	0.0393		
nmHydro							
Heat	0.252	0.0722	0.111	0.252	0.393		
nmHeat	0.0399	0.0117	0.0169	0.0399	0.0628		
Heat*nmHeat	-0.300	0.0390	-0.377	-0.300	-0.224		
Hydro (2-yr lag)*	-0.198	0.0475	-0.291	-0.198	-0.106		
nmHeat							
Heat*nmHydro	0.0913	0.0267	0.0389	0.0914	0.144		

Model: Hydro and Heat by long-term climate condition

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