

Comparison of the Ionosphere during an SMC initiating substorm and an isolated substorm

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Key Points:

- Preconditioning of the ionosphere plays a role in which mode of transport the magnetosphere enters.
- No DP1 current system found during the initiating substorm of steady magnetospheric convection event (SMC).
- The Hall conductance is much stronger during the initiating substorm of the SMC and throughout the SMC compared to the isolated substorm.

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Abstract

In order to assess the effects of ionospheric feedback on different modes of energy transport in the magnetosphere, we investigate an isolated substorm and a steady magnetospheric convection (SMC) event with very similar solar wind drivers. The primary focus is on a comparison between the isolated substorm and the substorm that initiates the SMC. Auroral data from Polar UVI LBHI and LBHs, along with AMIE potential patterns are used as inputs to the Global Ionosphere-Thermosphere Model (GITM) to calculate conductances and Joule heating rates. Results from this study show that the conductances both before and during the events play a large role the ability of the magnetosphere to remain in steady driven state. The substorm that initiates the SMC event shows very different signatures in the ionosphere than isolated substorm, these signatures indicate that there is very weak substorm current wedge, or possibly a pseudo-breakup.

1 Introduction

When the interplanetary magnetic field (IMF) is southward and of moderate strength, energy and particles from the solar wind are loaded in the magnetotail. Eventually, the particles and energy need to be unloaded, this usually happens in the form of an isolated substorm. The term isolated indicates that the substorm is a single event and is not part of a larger geomagnetic storm or periodic substorm event. If the driving continues and is fairly steady, the magnetosphere may enter a mode of energy transport where there is no major loading and unloading; rather, the energy is continuously diverted and magnetospheric convection is steady on a large scale. This type of event is referred to as a steady magnetospheric convection event or SMC. Most SMCs start after an initiating substorm, however both *Kissinger et al.* [2011] and *DeJong et al.* [2008] found that about 10 percent of SMC events do not appear to have a preceding, or initiating, substorm. *Sergeev et al.* [1994] showed that the configuration of the magnetotail during an SMC lies between that of a substorm recovery phase and that of a substorm growth phase. These results indicate that a substorm may be necessary for the magnetosphere to enter into an active steady state such as an SMC. This raises questions about how important initiating substorms are to SMC events and how they differ from isolated substorms.

These two event types can have similar drivers [*DeJong et al.*, 2009; *Partamies et al.*, 2009]. While isolated substorms have been studied extensively, SMCs and their initiating substorms have not. *Juusola et al.* [2013] and *DeJong* [2014] found that SMCs are more

46 likely to occur during weak or non-storm times. *Kissinger et al.* [2010] show that there are
47 many fast flows in the tail during SMCs and that the flows occur along dawn and dusk as
48 opposed to the night side.

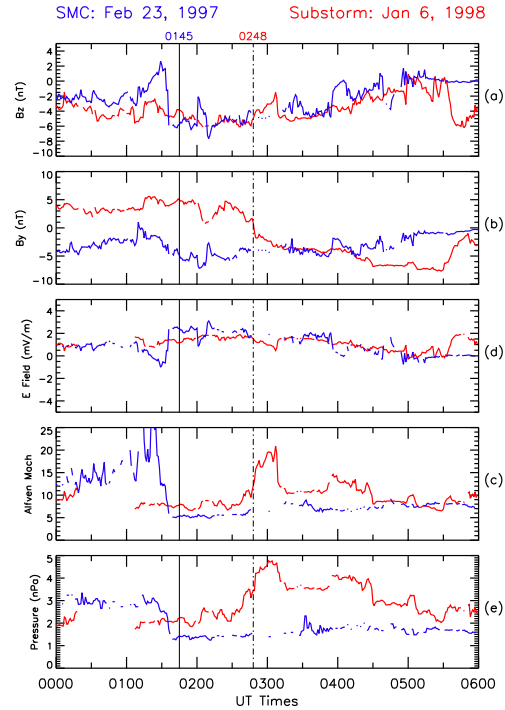
49 Many studies have investigated the solar wind drivers that lead to an active steady
50 state (SMC) as opposed to a loading and unloading state (substorm). While differences are
51 found in the drivers, there are also many times when the driving for these two event types
52 can be very similar [*DeJong et al.*, 2009; *Partamies et al.*, 2009]. Thus, drivers alone
53 cannot account for the differences in, the magnetospheric dynamics during these events, so
54 we will examine the role ionospheric feedback in determining why the magnetosphere
55 enters these different states. The modeling results of *Ridley et al.* [2004] show that
56 ionospheric conductance, both seasonal and auroral, can affect the state of the
57 magnetosphere. *Raeder et al.* [1996] found that specifying a higher conductance in their
58 simulations allows the magnetosphere to remain in a steady state. Along with these
59 findings, more recently *Welling and Liemohn* [2016] show that current models do not
60 accurately capture the storm time magnetospheric dynamics when the ionospheric
61 feedback is not included. This is most likely do to the fact that Pederson and Hall
62 conductances in the auroral zone are tied to field aligned currents [*Knight et al.*, 1972].

63 In order to isolate the effects of the ionosphere on the dynamics of the
64 magnetosphere, we selected two events with very similar drivers. We investigate an
65 isolated substorm that occurred on January 6, 1998 at 0248 UT and an SMC that occurred
66 February 23, 1997 with the onset of the initiating substorm at 0145 UT. Since these two
67 events occur in winter and at a very similar universal time, seasonal differences or
68 magnetometer locations should not affect the results. For the comparison of the
69 ionosphere during these events, we use the Global Ionosphere-Thermosphere Model
70 (GITM) to calculate both Pederson and Hall conductances along with Joule heating rates
71 [*Ridley et al.*, 2006]. Additionally, we use Polar UVI auroral energy flux, the average
72 auroral energy [*Germany et al.*, 1994], and Assimilative Mapping of the Ionosphere
73 Electrojet (AMIE) potential patterns as inputs for the model. Along with the modeling
74 result, we compare the residual ionospheric potential patterns after the onset of the
75 substorms, similar to the study by *Cai et al.* [2006].

2 Drivers and Indices

The solar wind drivers and IMF conditions for each event are plotted in Figure 1, where the blue line represents the SMC and the red line the isolated substorm. The solid and dotted vertical lines represent the onset times for the SMC initiating substorm and the isolated substorm respectively. The data for this plot comes from the OMNI data set. The IMF B_z (Figure 1a) and the electric field (Figure 1c) are very similar for both events, with the small exception of the increase and drop in B_z about 15 minutes before the onset of the SMC initiating substorm, indicating a possible trigger for the SMC. IMF B_y remains negative and steady during the SMC, but it changes direction from positive to negative at the onset of the isolated substorm. Both the dynamic pressure and the Alfvénic Mach number drop just before the start of the SMC event, but they increase at the onset of the isolated substorm. These figures indicate that the isolated substorm is most likely triggered by the change in pressure, Mach number, and B_y . Despite these small differences, these two events share very similar drivers and we should expect a similar response in the magnetosphere. However, this is not what occurs.

Figure 2 shows the magnetospheric response to these similar drivers; AU, AL, AE, and Sym-H are all from the OMNI data set. The F_{PC} , or open magnetic flux, was calculated for *DeJong et al.* [2007] and is once again plotted here to show the approximate change in the polar cap. The cross polar cap potential or CPCP is calculated from AMIE [*Kihn and Ridley*, 2005]. The B_z component of the magnetic field from the GOES 8 satellite in GSM coordinate, is also plotted. During this time GOES 8 was on the nightside in the pre-midnight sector when both substorms occurred. The AU and the Sym-H show very little activity for both events. The Sym-H is close to zero for the isolated substorm and it only decreases to about -30 nT for the SMC. Thus, neither of these events occur during storm time. While the IMF B_z and the electric fields are very similar for these events (Figure 1), their AL response is very different (Figure 2b). At the onset of the SMC initiating substorm (solid line) there is a decrease in AL from 0 nT to about -300 nT over an hour, indicating a weak substorm that then leads into the SMC. This is also roughly the time of the CPCP increase, confirming an onset time. The GOES B_z shows a positive change during this time, possibly indicating a dipolarization event occurred. This SMC was identified using the method in *DeJong and Clauer* [2005], which requires a steady F_{PC} for at least 3 hours, and this can be seen in Figure 2e. The onset of the isolated substorm is much more obvious in AL, the CPCP, and the F_{PC} as all changes

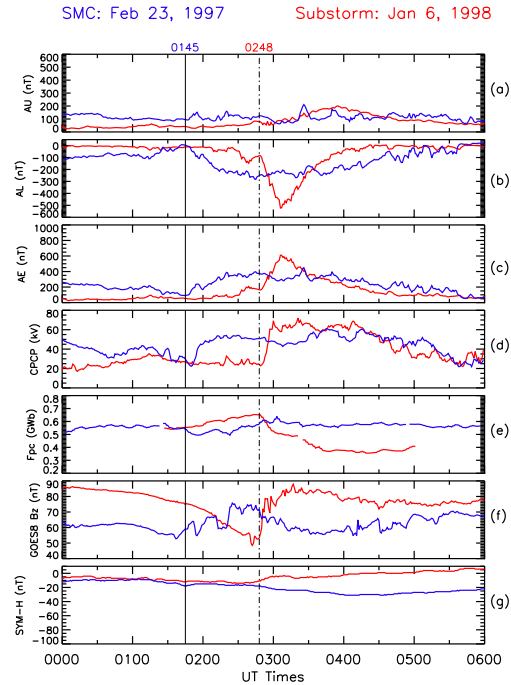


91 **Figure 1.** Stack plot of the drivers for the substorm on Jan. 6 1998 in red and the SMC on Feb. 23, 1997 in
 92 blue. The panels from top to bottom are as follows: IMF B_z , IMF B_y , Electric Field, Alfvén Mach number,
 93 and Dynamic Pressure. The solid vertical line represents the onset of the substorm that precedes the SMC and
 94 the dotted vertical line is the onset of the isolated substorm.

113 line up with the onset time of 0248 UT (dotted line). The AL drops to about -500 nT 20
 114 minutes after the onset, which is then followed by a recovery phase that lasts about 1 hour.
 115 At the same time, the polar cap potential difference changes from 20 to 70 kV and the
 116 polar cap flux shrinks from 0.65 to 0.35 GWb. The GOES B_z drops from almost 90 nT
 117 down about 50 nT before the onset as the magnetic fields at geosynchronous orbit are
 118 stretched, then a rapid increase at the substorm onset indicates the dipolarization that
 119 occurs at the onset of the expansion phase.

120 The precondition for these two events can be seen from the 0000 UT point on both
 121 days. While the IMF B_z is negative and moderate (-3 nT) the magnetosphere shows little
 122 to no response on January 6th 1998 with AL very close to zero until the onset of the
 123 substorm. During same IMF B_z conditions the magnetosphere has a response of about
 124 -100 nT in AL before the SMC, the AL then quiets down before the onset of the SMC.

125 The differences the magnetospheric response during the time leading up to these events
 126 could very well have an impact on the final response mode of the magnetosphere.



127 **Figure 2.** A stack plot of the magnetospheric responses for the substorm on Jan. 6 1998 in red and the
 128 SMC on Feb. 23, 1997 in blue. The panels from top to bottom are as follows: AU, AL, AE, Cross polar cap
 129 potential, open magnetic flux, GOES 8 B_z , and Sym-H. The solid vertical line represents the onset of the
 130 substorm that precedes the SMC and the dotted vertical line is the onset of the isolated substorm.

131 3 Modeling and Methods

132 We utilize the Global Ionosphere-Thermosphere Model (GITM) to simulate auroral
 133 conductances and Joule heating [Ridley *et al.*, 2006]. In order to simulate the state of the
 134 ionosphere-thermosphere during the SMC and the isolated substorm events, we input the
 135 solar wind drivers, the AMIE potential patterns, the calculated average auroral energy
 136 flux [Germany *et al.*, 1994], and the auroral energy flux from Polar UVI. The auroral energy
 137 flux is taken from the Polar UVI LBHI filter, and the average auroral energy is
 138 calculated from the UVI images using the ratio of the measured LBHI to the measured
 139 LBHs and the atmospheric model of Germany *et al.* [1994]. Because the LBHI and LBHs

140 images are not taken at the same time, as UVI switches through the filters, the images
141 must be linearly interpolated to match the time stamp of the other image. This allows us
142 to generate an average energy and energy flux input approximately every 1 to 2 minutes.
143 These time-dependent auroral drivers are then input into GITM as the simulation evolves,
144 allowing us to investigate how the ionosphere responds to these time-varying high latitude
145 inputs.

146 We employ GITM, because it is the only coupled ionosphere-thermosphere model to
147 relax the hydrostatic constraint, which allows it to self-consistently solve the momentum
148 equation in the vertical direction [Ridley *et al.*, 2006]. This allows GITM to accurately
149 simulate the ionospheric response to intense, localized heating in the aurora [Deng *et al.*,
150 2008a]. GITM also includes ionospheric chemistry coupled with neutral chemistry,
151 self-consistently. Moreover the ionization and dissociation related to auroral precipitation
152 are included as part of the chemical drivers. This allows us to calculate the net ionization
153 in the ionosphere during these events, which further allows us to calculate the Pedersen
154 and Hall conductivities in the ionosphere according to the well-known formulas involving
155 the ion-neutral collision frequencies and the local ion gyro frequency [Schunk and Nagy,
156 2004]. GITM also accounts for both Joule Heating and direct heating by auroral
157 precipitation, allowing us to combine thermal structure, dynamical, and chemical
158 responses of the upper atmosphere to the high-latitude driving specified by AMIE and
159 Polar UVI. We use the model-simulated conductances to examine the state of the
160 ionosphere during the SMC and the isolated substorm.

161 Each GITM Simulation was run for an initial twenty-four hours to remove any
162 impacts from start-up. We include date-appropriate F10.7-cm values, the waves and tides
163 provided by the Global Scale Wave Model (GSWM) Hagan *et al.* [1999], and lower
164 boundary temperatures and winds are provided by the NRL MSIS-00 Picone *et al.* [2002].
165 The AMIE potential patterns were incorporated into GITM at each time-step to drive the
166 upper atmosphere through an imposed magnetospheric electric field and through particle
167 precipitation into the auroral zones. We column-integrate the conductances and the joule
168 heating rates within GITM itself during run-time to ensure the self-consistency of the
169 outputs. These methods are consistent with numerous studies of the
170 thermosphere-ionosphere (Ridley *et al.* [2006], Bougher *et al.* [2015], Bell *et al.* [2014],
171 Deng *et al.* [2008b]).

4 Results

The potential patterns for the isolated substorm and the SMC are created from the Assimilative Mapping of the Ionosphere Electrojet (AMIE) with inputs of only magnetometers as in *Ridley et al. [2004]* and *Cai et al. [2006]*. The focus of this study is the development of the DP1 pattern during the expansion phase of the isolated substorm and the substorm preceding the SMC. The DP1 pattern is associated with the substorm current wedge, and it begins to develop right after the onset of the expansion phase and continues to strengthen during this phase. A similar pattern was found for global sawtooth injections, with a larger potential difference than the substorms [*Cai et al., 2006*]. During the substorm expansion phase, the DP1 potential pattern dominates, but it can be difficult to observe. Thus, to isolate the pattern that arises during the events, the onset time must first be removed from all succeeding times *Cai et al. [2006]*, creating a residual pattern.

Figure 3 shows the AMIE potential patterns for the isolated substorm on Jan. 8th, 1998. The plots are in magnetic local time (MLT) coordinates with noon at the top, and each dotted circle is 10 degrees magnetic latitude. All of the plots have had the onset of the expansion phase, 0248 UT, removed, leaving the first plot blank after the subtraction. The plots that follow are snap shots at 5 minute intervals for the 2 hours after the start of the isolated substorm. The DP1 pattern begins to appear as early as 5 minutes after the onset, with a peak in the pattern at 40 minutes, as seen by the blue contours. By minute 60, the pattern has changed back to an enhanced DP2 pattern during the substorm's recovery phase. After 2 hours the potential patterns have weakened to their initial levels. These results are consistent with those shown for isolated substorms and global sawtooth oscillations by *Cai et al. [2006]*.

The initiating substorm of the SMC shows very different residual potential patterns than those of the isolated substorm. Figure 4 shows the SMC initiating substorm on Feb. 23rd, 1997, in the same format as Figure 3. The residual potential patterns shows an enhanced two cell convection, DP2, with no sign of a DP1 pattern. Since there is a substorm before the magnetosphere enters the steady state, a DP1 pattern is to be expected. There is a small positive cell that develops around 30 minutes after the onset, and this could reveal a very weak DP1 pattern overlapping the DP2. However, when the 0215 UT pattern (30 minutes after onset) was removed as the background, there was no

203 indication of a DP1 pattern. This lack of a DP1 pattern shows that the substorms that
204 precede SMCs may in fact be different than isolated substorms.

209 Figures 5 and 6 are magnetic local time (MLT) plots of the auroral energy flux,
210 average auroral energy, integrated Hall conductance, integrated Pederson conductance, and
211 Joule heating for the isolated substorm and the SMC initiating substorm respectively. The
212 auroral energy fluxes are from Polar UVI LBHI images, the average auroral energy is
213 calculated from *Germany et al.* [1994], and the conductances and Joule heating are
214 calculated from GITM. The first column of images are from 15 minutes before the onset
215 of the expansion phase and each succeeding column snap shots at 15 minute increments
216 up to 1 hour after the onset. The plots have magnetic north at the center of the image, and
217 the circles represent 10 degrees magnetic latitude down to 50 degrees. The top of each
218 circle is magnetic local noon and the bottom is magnetic local midnight and the sides are
219 dawn and dusk. The number on the bottom right of each plot is the maximum value at
220 that time stamp. As expected, the integrated Hall conductances are higher than the
221 Pederson conductances for both the isolated substorm and the initiating substorm of the
222 SMC. Both patterns follow the auroral images (top row) fairly closely. The same data is
223 shown in Figures 7 and 8 in keogram format for 22 MLT; this magnetic local time was
224 chosen since both onsets occur near this location. The data and color bars for each row
225 are the same for all four figures. The keograms begin one hour before onset and continue
226 to 2 hours after onset.

227 The isolated substorm on January 6th, 1998 in Figures 5 and 7 proceeds as
228 anticipated in the auroral images. The substorm enters into the recovery phase about 1
229 hour after the onset, even though the solar wind B_z is still slowly approaching zero and
230 will not become positive for another hour (Figure 1a). The Joule heating begins about 5
231 minutes after onset and ends when once the recovery phase begins. *Palmroth et al.* [2004]
232 found that Joule heating increases about 5 minutes after a pressure pulse during steady
233 IMF conditions. Thus, the Joule heating is most likely a combination of the onset of the
234 substorm and the pressure pulse seen in Figure 1. The Hall conductance continues to
235 increase even as the auroral energy flux starts to diminish. By the end of the time frame
236 in Figure 7, the substorm has ended and the aurora is back to pre-substorm levels.

237 The SMC and its initiating substorm on February 23rd is shown in Figures 6 and 8.
238 The auroral energy flux during the SMC does not change significantly and the initiating

239 substorm does not show a strong poleward movement, as seen in both Figure 6 and the
240 F_{PC} in Figure 2. The auroral energy flux and conductances do intensify and spread
241 toward dawn and dusk, as seen in Figure 6. The Joule heating does not increase
242 significantly until 40 minutes after the onset of the initiating substorm. We note in Figure
243 8 that the aurora becomes active about 10 minutes before the onset on the substorm,
244 which could be due to a pre-onset pseudo-breakup. This could not be considered the
245 onset, since the second brightening occurs at the same time as the decrease in AL and the
246 increase in cross polar cap potential (Figure 2). The SMC event continues past the time
247 frame shown in Figure 8. By definition from *DeJong and Clauer* [2005], the event must
248 last at least one more hour for a total of 3 hours to be classified as an SMC.

259 5 Discussion

260 The results in the previous sections show that these two events types can have
261 distinct ionospheric signatures. However, one of the most significant differences is the
262 preconditioning before the SMC event. Not only can it be seen in the AL and AE data,
263 but also in the keograms in Figures 7 and 8 that show the aurora and integrated
264 conductances are more active before the onset of the SMC. It is unusual that the solar
265 wind IMF B_z , which is at -4 nT for 3 hours, does appear to impact the magnetosphere and
266 ionosphere before the isolated substorm. The interval before the SMC on the other hand
267 has activity that most likely leads to preconditioning of the ionosphere and magnetosphere
268 allowing the SMC to occur. The higher conductance before the SMC event most likely
269 plays a role in allowing the magnetosphere to unload energy in smaller amounts such as
270 fast flow rather than one large unloading. If the conductance is lower, then the energy
271 can't flow as well, forcing the magnetotail to store the energy until the system is
272 overloaded and the energy is unloaded all at once.

273 During the events there are also differences in the ionosphere both the integrated
274 Pederson and Hall conductances are higher during the SMC, however it takes longer for
275 these conductances to build up during this event. For the SMC initiating substorm, the
276 peaks in the conductances occur an hour after onset and they remain strong for the
277 duration of the time shown. The peaks are also much more spread out in the auroral zone
278 during this event, with the largest Hall conductances at 49 mhos and the Pederson at 20
279 mhos, as seen in Figure 6. By contrast, the peak conductances for the isolated substorm
280 occur approximately 45 minutes after the onset, with the Hall at 37 mhos and Pederson at

281 12 mhos. The auroral energy flux and average energy however peak at the onset of the
282 isolated substorm. Thus, the auroral conductance is much higher when the magnetosphere
283 is in a steady state. This agrees with *Raeder et al.* [1996] who found that higher
284 conductances in MHD simulations lead to a steadier magnetosphere. This suggests that
285 higher conductance allow the magnetosphere to slowly and steadily unload its energy as
286 opposed to having a large reconfiguration event like a strong substorm.

287 Another interesting observation is that neither AL nor the solar wind parameters
288 during these events are well correlated with the Hall conductance as found by *Aksnes et al.*
289 [2002]. There is a small amount of activity in AL before the onset of both events (Figure
290 2b), but during the SMC it slowly decreases over the first hour and never gets below -300
291 nT. On the other hand, the isolated substorm AL decreases sharply over the first 30
292 minutes and extends beyond -500 nT before entering the recovery phase. The solar wind
293 conditions are almost exactly the same for these two events, yet the Hall conductance is
294 quite different. Thus, if either AL or the solar word is related to the Hall conductance,
295 then it is not apparent in our results, since the event with the weaker AL has a much
296 stronger Hall conductance.

297 The Joule heating rates for these events do not show the same patterns as the
298 conductances. Both events reach a maximum heating rate of close to 15 mW/m^2 . During
299 the isolated substorm the heating starts almost immediately after the onset, whereas the
300 heating during the SMC does not initiate until 40 minutes into the event. The Joule
301 heating also takes place a much higher latitude (70-80 degrees) during the isolated
302 substorm than during the SMC (60-70 degrees). *Bjoland et al.* [2015] found enhancements
303 in the Joule Heating at 70 degrees latitude at 15 MLT and 2 MLT during all IMF
304 conditions. We only see the dayside (15 MLT) enhancement during the SMC starting at
305 15 minutes after onset, Figure 6. We also have a slight increase in Joule heating that
306 occurs at 2 MLT at 15 minutes before the onset of the isolated substorm and 45 minutes
307 after the onset in Figure 5. Most of the simulated Joule heating during these events occurs
308 around at dusk, where the auroral inputs are the most intense.

309 *Zhou et al.* [2011] showed that nightside enhancements in Joule heating during
310 substorms is associated with unloading, or a substorm current wedge. Thus, the isolated
311 substorm shows a strong current wedge signature in both the Joule heating (Figure 5) and
312 the potential patterns (Figure 3). These potential patterns during the isolated substorm will

313 produce a strong electric field that subsequently induces stronger ion velocities. These
314 enhanced ion winds, since they are flowing against the background neutral winds, naturally
315 produce stronger Joule Heating in those regions.

316 The lack of a DPI pattern during the SMC initiating substorm along with the weak
317 substorm signatures seen in the GOES8 data and auroral indices indicates the initiating
318 substorm of the SMC is very weak or even a pseudo-breakup. Thus, while a substorm
319 current wedge is initially created the energy from tail soon gets diverted through
320 enhanced convection allowing the magnetosphere to enter a steady state. This is consistent
321 with the auroral images shown in Figure 6 and the F_{PC} in Figure 2, where the aurora
322 does not show a large poleward movement but rather spreads out along the oval. *Kissinger*
323 *et al.* [2010] showed that fast flows, carrying energy from the magnetosphere to the
324 ionosphere, occur frequently during SMCs. If these flows can carry enough energy then a
325 substorm current wedge, or at least not a strong one, may not need to fully develop during
326 these substorms that initiate SMCs. This indicates that the magnetosphere does not have a
327 strong tail reconfiguration during the expansion phase of the SMC initiating substorm.

328 The Joule heating also supports the idea that there is only a weak substorm current
329 wedge, or unloading, associated with the onset of the substorm that initiates the SMC. The
330 Joule heating does not increase until 35 minutes into the event (Figures 6 and 8), about
331 the same time a small positive cell appears to overlap with the enhanced convection
332 pattern of the SMC (Figure 4). This could indicate that a weak current wedge is
333 established during the SMC.

334 *Liou et al.* [2011] also studied this substorm but identified the 0215 UT auroral
335 brightening as the substorm onset instead of our 0145 UT onset. However, since there is
336 no real change in AL, CPCP, or F_{CP} at this time, it is more likely just an enhancement in
337 activity. Thus, the SMC starts off without a current wedge, but enhances the aurora and
338 AL by unloading in small amounts as opposed to one large reconfiguration.

339 Overall, the initiating substorm of this SMC is very different from an isolated
340 substorm. In order to more fully investigate whether a weak substorm current wedge is
341 normal for initiating substorms of SMC and if most SMCs have a higher conductance,
342 more events must be studied.

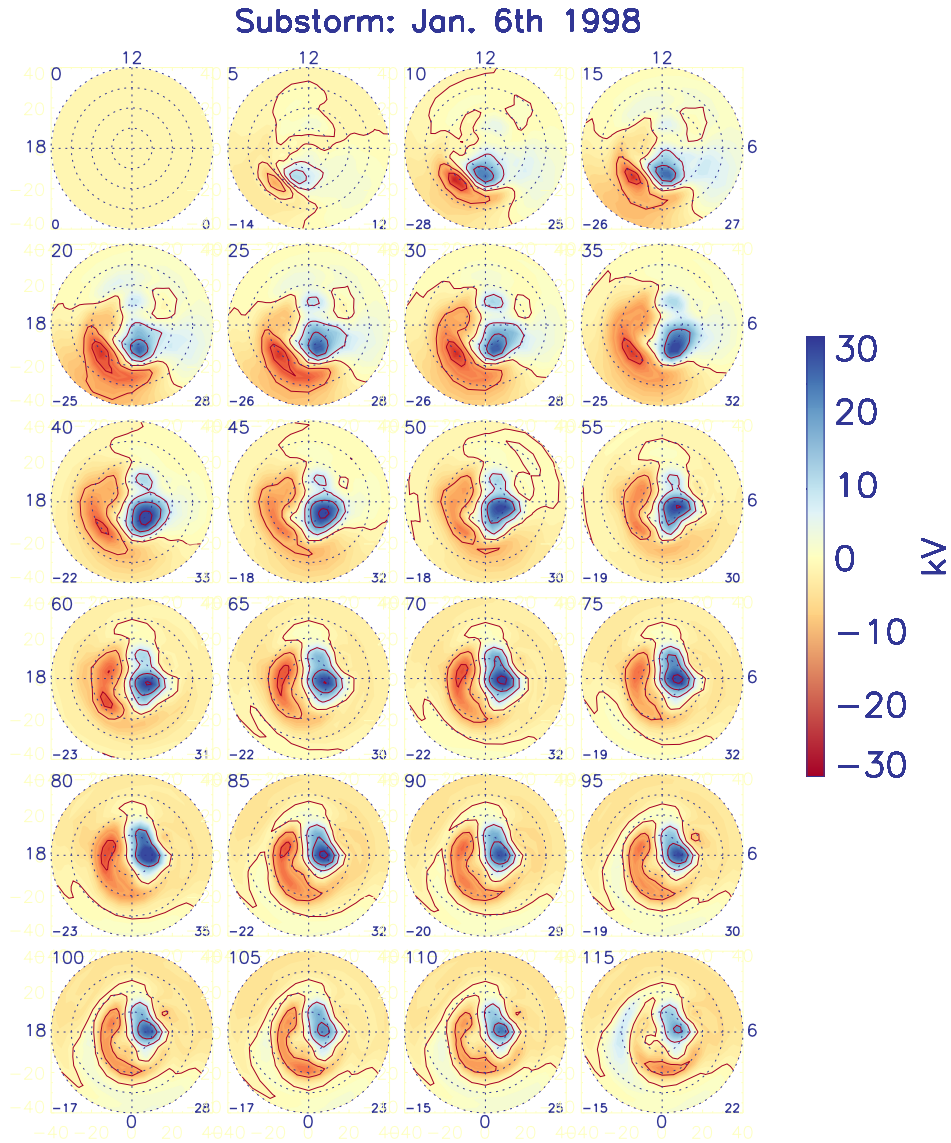
343 This study has shown that there more questions to investigate:

- 344 • How similar are other SMC initiating substorms to this one, including
345 preconditioning?
- 346 • How important is the larger Hall conductance in allowing the magnetosphere to
347 remain steady? Does the higher conductivity allow the magnetosphere to stay
348 steady or are the small fast flows creating a higher conductance?

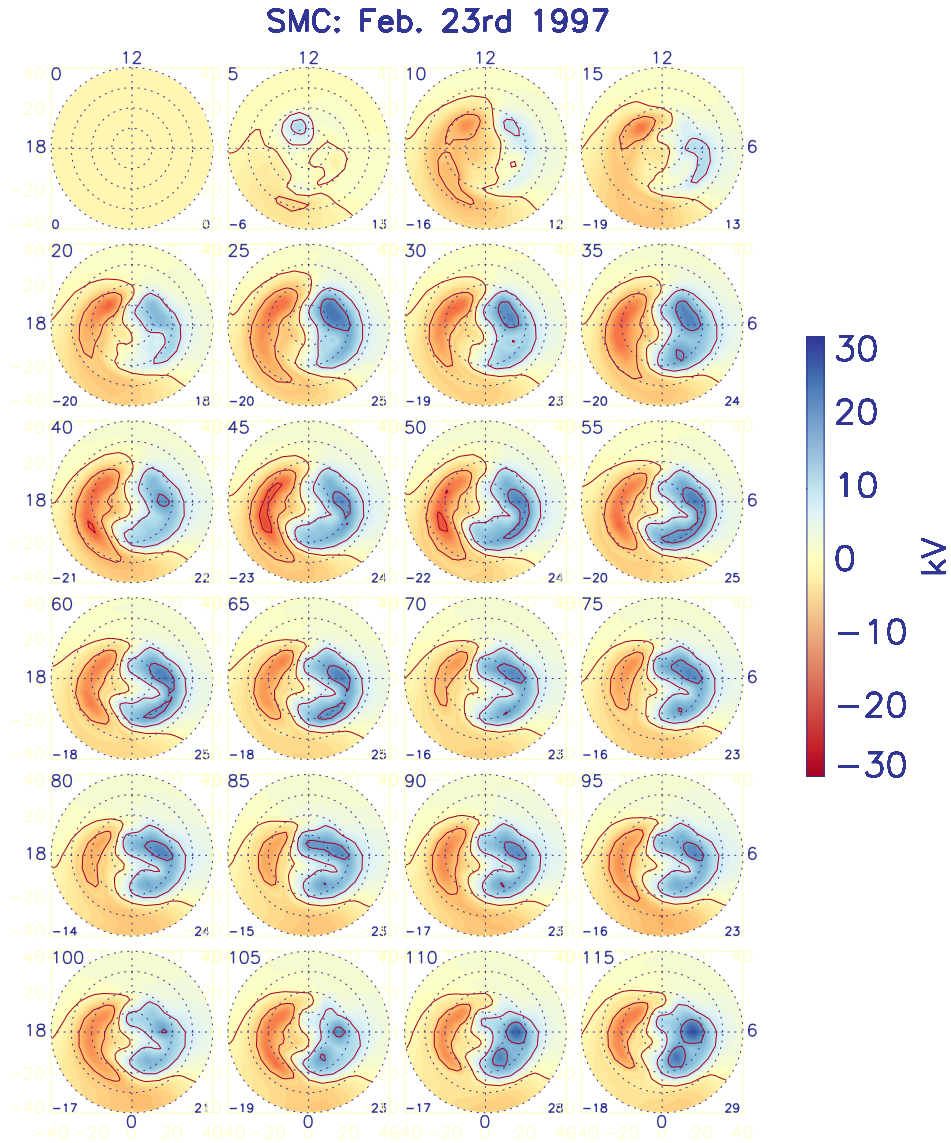
349 **6 Summary**

350 The ionosphere during two substorms with very similar solar wind drivers are
351 investigated in this study, one substorm is isolated while the other initiates a steady
352 magnetospheric convection event. Both events occur during a similar universal time so
353 that our differences in the data are not due to instrument location and both take place
354 during the winter months to help account for any difference in the conductivity of the
355 ionosphere. We found that the ionosphere plays a large role in these two events, allowing
356 one to remain in an enhanced steady state, while the other unloads the magnetosphere and
357 goes back to a quiet state. The major findings are listed below.

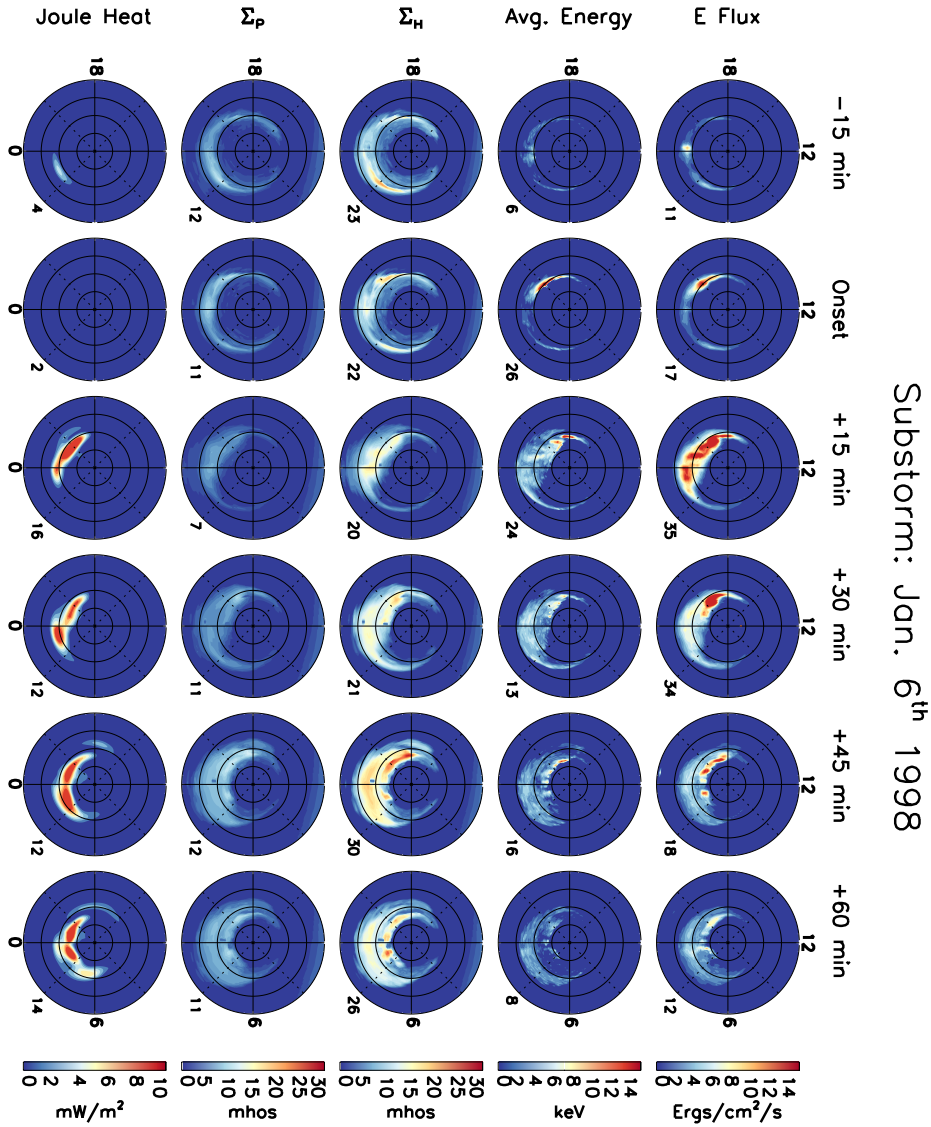
- 358 1. Preconditioning appears to play a large role in the type of event the magnetosphere
359 enters.
- 360 2. There is no DP1 current development after the onset of the SMC initiating
361 substorm, instead there is an enhanced DP2 convection pattern.
- 362 3. Integrated Hall conductance is much greater during the SMC initiating substorm
363 and continues to increase throughout the SMC.
- 364 4. Joule heating is stronger during the isolated substorm and there little to no Joule
365 heating until 30 minutes into the SMC.
- 366 5. The weaker substorm or pseudo-breakup that that initiates SMC is weak enough
367 that the energy from the magnetosphere can quickly be deposited by smaller fast
368 flows that reach the ionosphere.



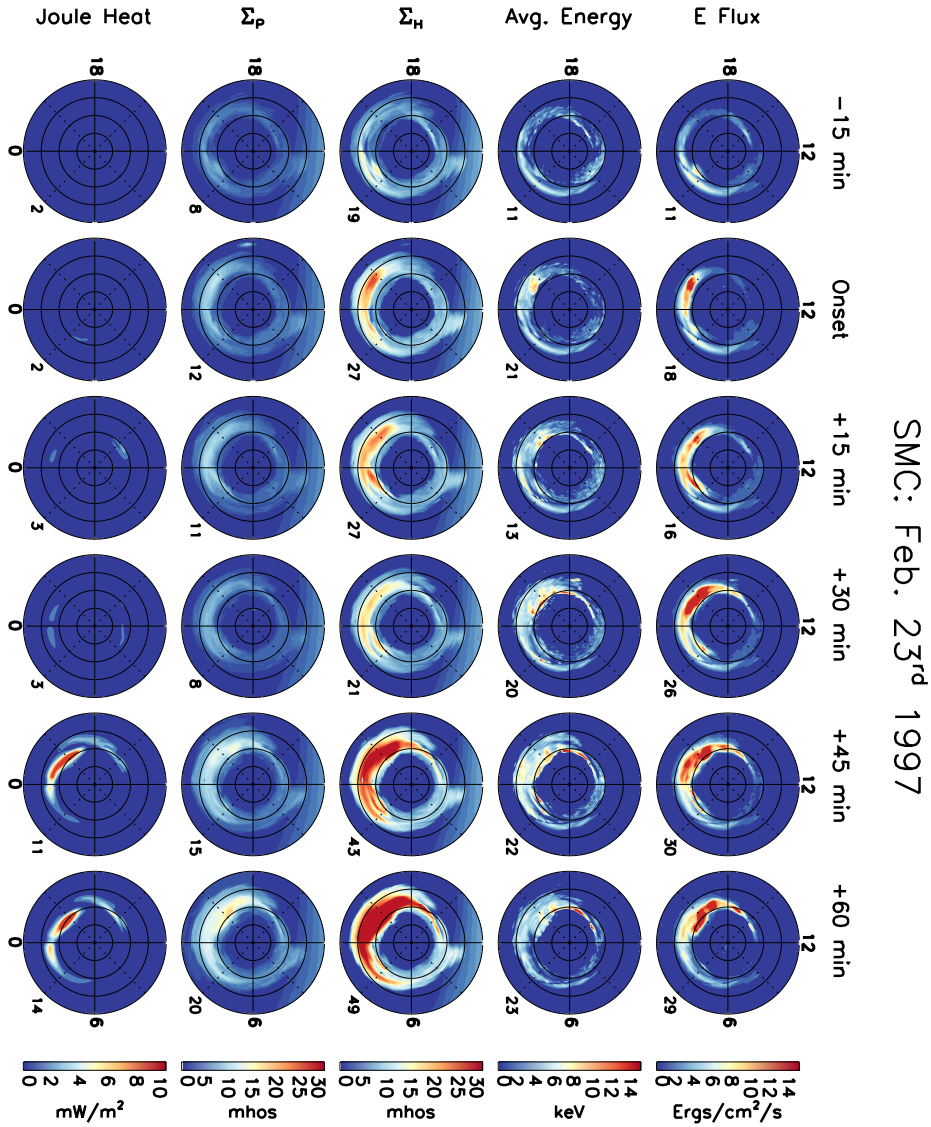
205 **Figure 3.** Potential patterns from AMIE, the onset of the Substorm (0248) has been removed from all times
 206 post onset. Thus, the patterns shown are the residual patterns from the event.



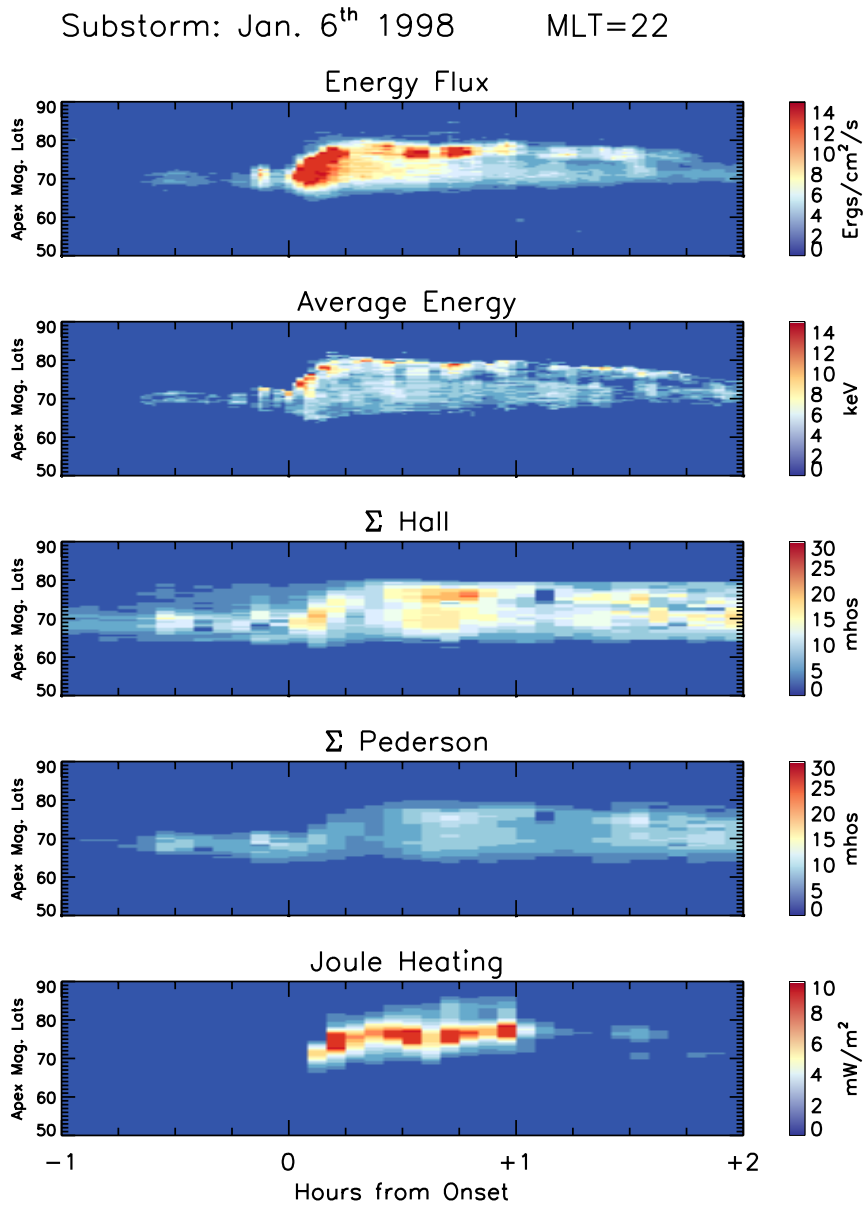
207 **Figure 4.** Potential patterns from AMIE, the onset of the SMC (0145) has been removed from all times post
 208 onset. Thus, the patterns shown are the residual patterns from the event.



249 **Figure 5.** All plots are in MLT coordinates with noon at the top and midnight at the bottom. The first
 250 column is 15 minutes before the onset of the isolated substorm, the second is at the onset time (0248 UT) and
 251 the next four columns are in 15 minute increments after the onset up to 1 hour.

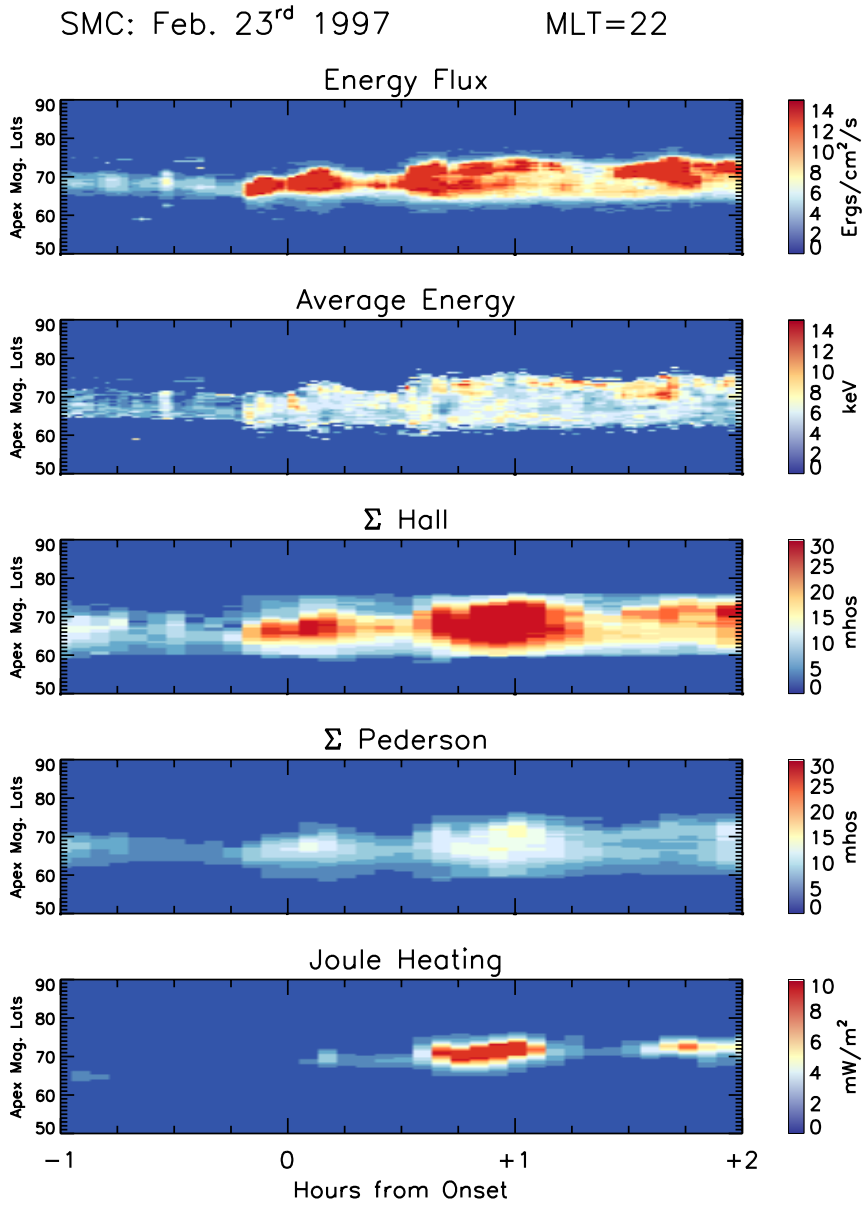


252 **Figure 6.** All plots are in MLT coordinates with noon at the top and midnight at the bottom. The first
 253 column is 15 minutes before the onset of the initiating substorm, the second is at the onset time (0145 UT)
 254 and the next four columns are in 15 minute increments after the onset up to 1 hour.



255 **Figure 7.** Keograms from an MLT of 22, with the same data from Figure 5 starting from 1 hour before

256 onset to 2 hours after the onset.



257 **Figure 8.** Keograms from an MLT of 22, with the same data from Figure 6 starting from 1 hour before

258 onset to 2 hours after the onset

Acknowledgments

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References

- Aksnes, A., J. Stadsnes, J. Bjordal, N. Østgaard, R. R. Vondrak, D. L. Detrick, T. J. Rosenberg, G. A. Germany, and D. Chenette (2002), Instantaneous ionospheric global conductance maps during an isolated substorm, *Annales Geophysicae*, *20*, 1181–1191, doi:10.5194/angeo-20-1181-2002.
- Bell, J. M., J. Hunter Waite, J. H. Westlake, S. W. Bougher, A. J. Ridley, R. Perryman, and K. Mandt (2014), Developing a self-consistent description of Titan’s upper atmosphere without hydrodynamic escape, *Journal of Geophysical Research (Space Physics)*, *119*, 4957–4972, doi:10.1002/2014JA019781.
- Bjoland, L. M., X. Chen, Y. Jin, A. S. Reimer, Å. Skjæveland, M. R. Wessel, J. K. Burchill, L. B. N. Clausen, S. E. Haaland, and K. A. McWilliams (2015), Interplanetary magnetic field and solar cycle dependence of Northern Hemisphere F region joule heating, *Journal of Geophysical Research (Space Physics)*, *120*, 1478–1487, doi: 10.1002/2014JA020586.
- Bougher, S. W., D. Pawlowski, J. M. Bell, S. Nelli, T. McDunn, J. R. Murphy, M. Chizek, and A. Ridley (2015), Mars Global Ionosphere-Thermosphere Model: Solar cycle, seasonal, and diurnal variations of the Mars upper atmosphere, *Journal of Geophysical Research (Planets)*, *120*, 311–342, doi:10.1002/2014JE004715.
- Cai, X., C. Clauer, and A. Ridley (2006), Statistical analysis of ionospheric potential patterns for isolated substorms and sawtooth events, *Ann. Geophysicae*, *24*, 1977–1991.
- DeJong, A., and C. Clauer (2005), Polar uvi images to study steady magnetospheric convection events: Initial results, *Geophys. Res. Lett.*, *32*, L24,101, doi:10.1029/2005GL024498.
- DeJong, A. D. (2014), Steady magnetospheric convection events: How much does steadiness matter?, *Journal of Geophysical Research (Space Physics)*, *119*, 4389–4399, doi:10.1002/2013JA019220.

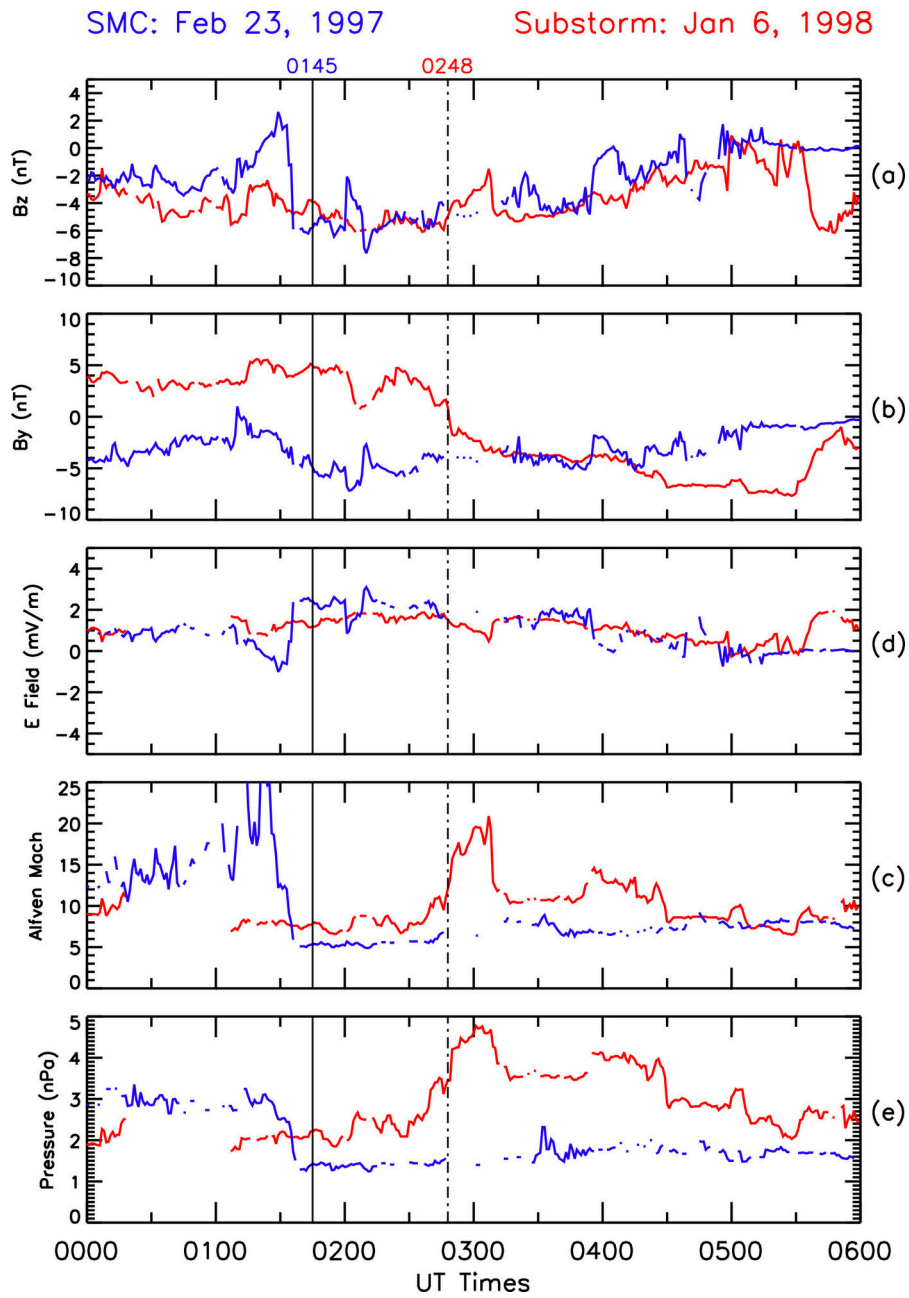
- 401 DeJong, A. D., X. Cai, C. R. Clauer, and J. F. Spann (2007), Aurora and open magnetic
402 flux during isolated substorms, sawteeth, and SMC events, *Annales Geophysicae*, *25*,
403 1865–1876, doi:10.5194/angeo-25-1865-2007.
- 404 DeJong, A. D., A. J. Ridley, and C. R. Clauer (2008), Balanced reconnection
405 intervals: four case studies, *Annales Geophysicae*, *26*, 3897–3912, doi:10.5194/
406 angeo-26-3897-2008.
- 407 DeJong, A. D., A. J. Ridley, X. Cai, and C. R. Clauer (2009), A statistical study of BRIs
408 (SMCs), isolated substorms, and individual sawtooth injections, *J. Geophys. Res.*, *114*,
409 A08215, doi:10.1029/2008JA013870.
- 410 Deng, Y., A. D. Richmond, A. J. Ridley, and H.-L. Liu (2008a), Assessment of the non-
411 hydrostatic effect on the upper atmosphere using a general circulation model (GCM),
412 *Geophysical Research Letters*, *35*, L01104, doi:10.1029/2007GL032182.
- 413 Deng, Y., A. Maute, A. D. Richmond, and R. G. Roble (2008b), Analysis of
414 thermospheric response to magnetospheric inputs, *Journal of Geophysical Research*
415 (*Space Physics*), *113*, A04301, doi:10.1029/2007JA012840.
- 416 Germany, G., D. G. Torr, P. G. Richards, M. R. Torr, and S. Jong (1994), Determination
417 of ionospheric conductivities from fuv auroral emission, *J. Geophys. Res.*, *99*, 23,297–
418 23,305.
- 419 Hagan, M. E., M. D. Burrage, J. M. Forbes, J. Hackney, W. J. Randel, and X. Zhang
420 (1999), GSWM-98: Results for migrating solar tides, *Journal of Geophysical Research*,
421 *104*, 6813–6828, doi:10.1029/1998JA900125.
- 422 Juusola, L., N. Partamies, and Tanskanen (2013), Effect of the ring current on
423 preconditioning the magnetosphere for steady magnetospheric convection, *Geophys. Res.*
424 *Lett.*, *40*, doi:10.1002/grl.50405.
- 425 Kihn, E. A., and A. J. Ridley (2005), A statistical analysis of the assimilative mapping
426 of ionospheric electrodynamics auroral specification, *Journal of Geophysical Research*
427 (*Space Physics*), *110*, A07305, doi:10.1029/2003JA010371.
- 428 Kissinger, J., R. L. McPherron, V. Angelopoulos, T.-S. Hsu, and J. P. McFadden (2010),
429 An investigation of the association between steady magnetospheric convection and CIR
430 stream interfaces, *Geophys. Res. Lett.*, *37*, L04105, doi:10.1029/2009GL041541.
- 431 Kissinger, J., R. L. McPherron, T.-S. Hsu, and V. Angelopoulos (2011), Steady
432 magnetospheric convection and stream interfaces: Relationship over a solar cycle, *J.*
433 *Geophys. Res.*, *116*, A00I19, doi:10.1029/2010JA015763.

- 434 Knight, D. E., R. Uribe, and B. E. Woodgate (1972), Extreme ultra-violet absorption
435 cross-sections in the Earth's upper atmosphere, *Planetary and Space Science*, *20*, 161–
436 164, doi:10.1016/0032-0633(72)90098-0.
- 437 Liou, K., Y.-L. Zhang, P. T. Newell, L. J. Paxton, and J. F. Carbary (2011), TIMED/GUVI
438 observation of solar illumination effect on auroral energy deposition, *Journal of*
439 *Geophysical Research (Space Physics)*, *116*, A09305, doi:10.1029/2010JA016402.
- 440 Palmroth, M., T. I. Pulkkinen, P. Janhunen, D. J. McComas, C. W. Smith, and H. E. J.
441 Koskinen (2004), Role of solar wind dynamic pressure in driving ionospheric Joule
442 heating, *Journal of Geophysical Research (Space Physics)*, *109*, A11302, doi:10.1029/
443 2004JA010529.
- 444 Partamies, N., T. I. Pulkkinen, R. L. McPherron, K. McWilliams, C. R. Bryant,
445 E. Tanskanen, H. J. Singer, G. D. Reeves, and M. F. Thomsen (2009), Statistical survey
446 on sawtooth events, SMCs and isolated substorms, *Advances in Space Research*, *44*,
447 376–384, doi:10.1016/j.asr.2009.03.013.
- 448 Picone, J. M., A. E. Hedin, D. P. Drob, and A. C. Aikin (2002), NRLMSISE-00 empirical
449 model of the atmosphere: Statistical comparisons and scientific issues, *Journal of*
450 *Geophysical Research (Space Physics)*, *107*, 1468, doi:10.1029/2002JA009430.
- 451 Raeder, J., J. Berchem, and M. Ashour-Abdalla (1996), The importance of small scale
452 processes in global MHD simulations: Some numerical experiments, *The Physics of*
453 *Space Plasmas*, edited by *T. Chang and J. R. Jasperse*, vol. 14, 403 pp., MIT Center for
454 Theoretical Geo/Cosmo Plasma Physics, Cambridge, Mass.
- 455 Ridley, A., T. Gombosi, and D. Dezeewu (2004), Ionospheric control of the
456 magnetosphere: conductance, *Annales Geophysicae*, *22*, 567–584, doi:10.5194/
457 angeo-22-567-2004.
- 458 Ridley, A. J., Y. Deng, and G. Tóth (2006), The global ionosphere thermosphere model,
459 *Journal of Atmospheric and Solar-Terrestrial Physics*, *68*, 839–864, doi:10.1016/j.jastp.
460 2006.01.008.
- 461 Schunk, R. W., and A. F. Nagy (2004), *Ionospheres*, 570 pp., Cambridge University Press,
462 Cambridge, Mass.
- 463 Sergeev, V. A., T. I. Pulkkinen, T. I. Pellinen, and N. A. Tsyganenko (1994), Hybrid state
464 of the tail magnetic configuration during steady convection events, *J. Geophys. Res.*, *99*,
465 23,571–+, doi:10.1029/94JA01980.

466 Welling, D. T., and M. W. Liemohn (2016), The ionospheric source of magnetospheric
467 plasma is not a black box input for global models, *Journal of Geophysical Research*
468 (*Space Physics*), *121*, 5559–5565, doi:10.1002/2016JA022646.

469 Zhou, X.-Y., W. Sun, A. J. Ridley, and S. B. Mende (2011), Joule heating associated with
470 auroral electrojets during magnetospheric substorms, *Journal of Geophysical Research:*
471 *Space Physics*, *116*(A5), n/a–n/a, doi:10.1029/2010JA015804, a00I28.

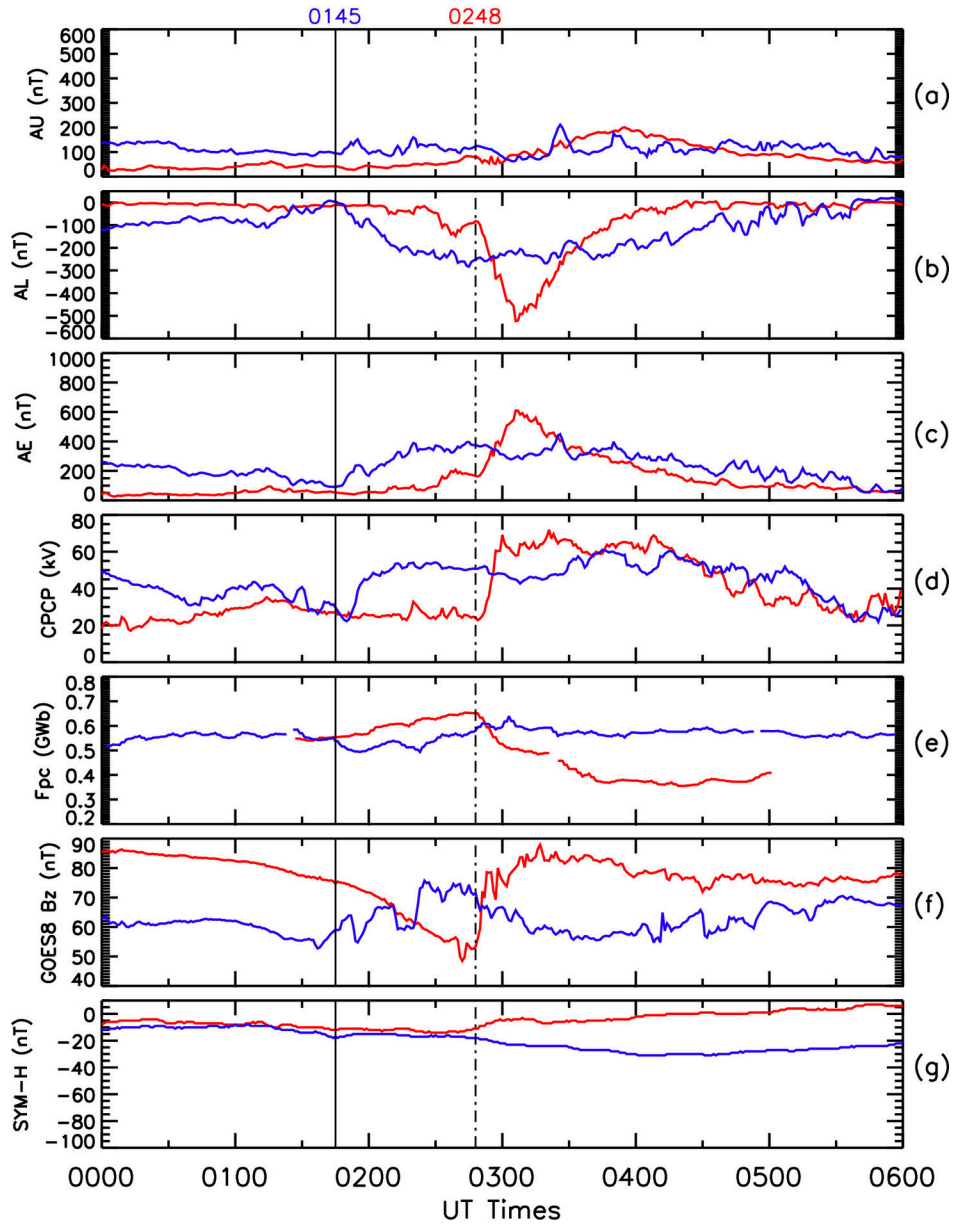
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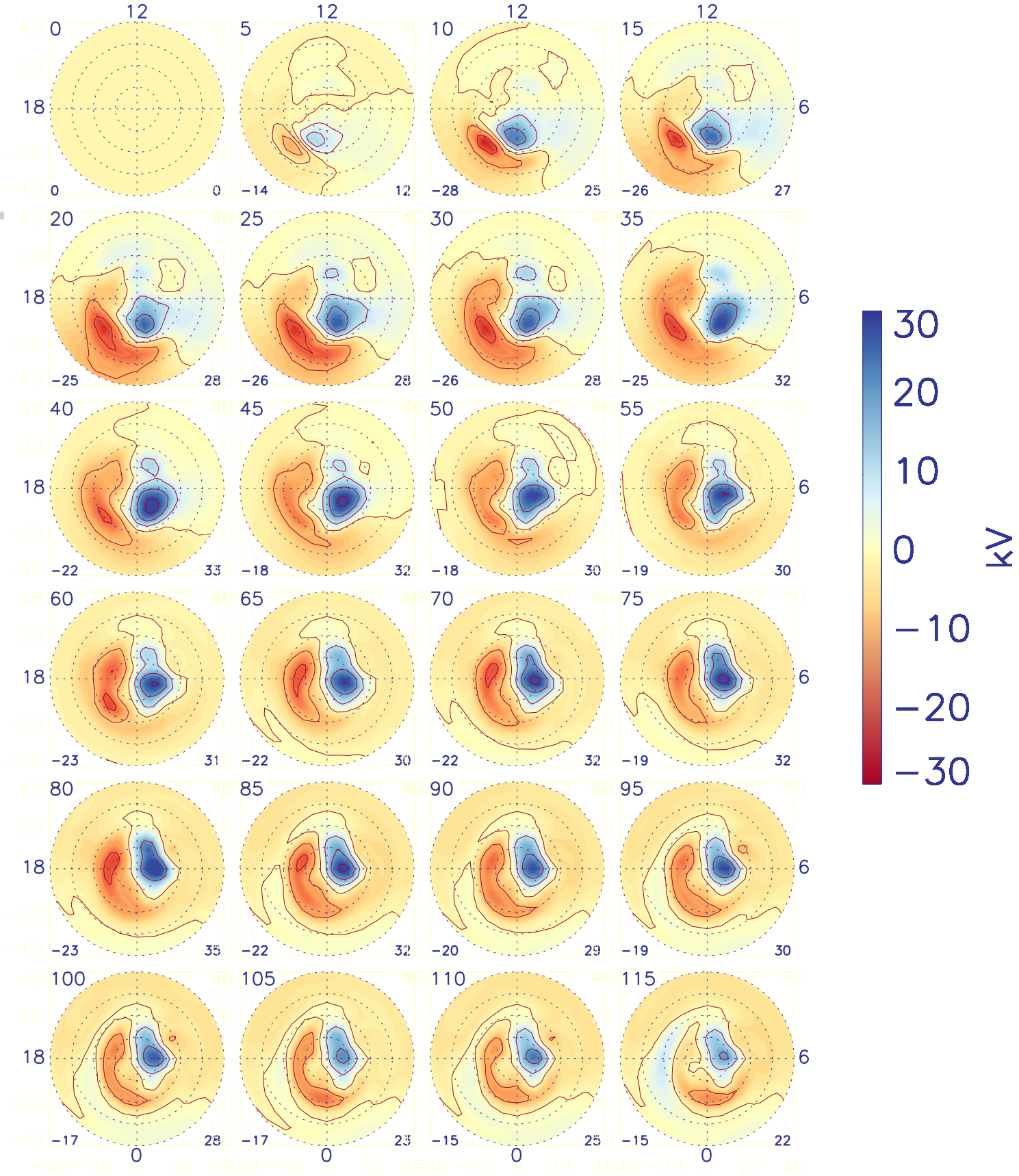
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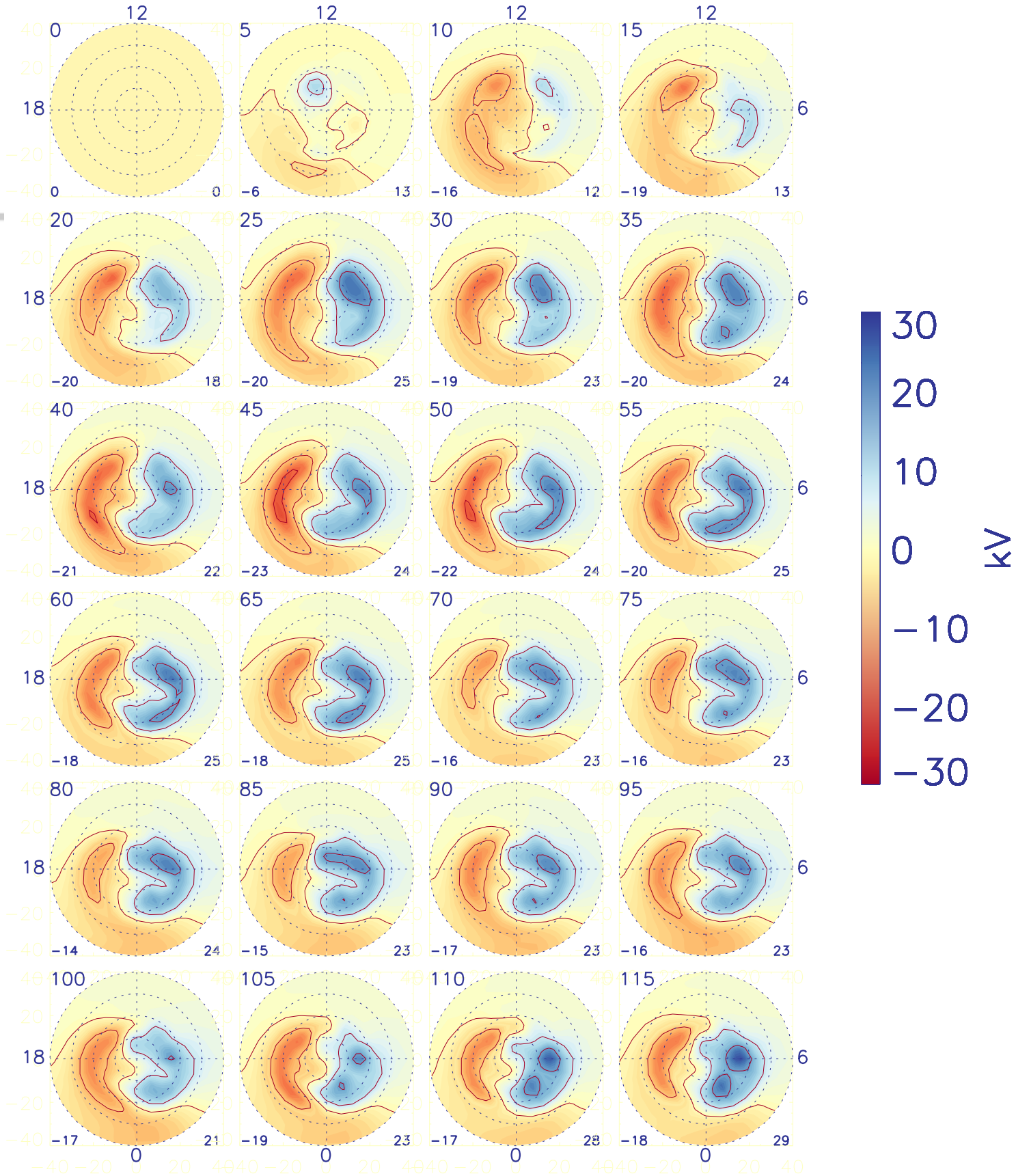


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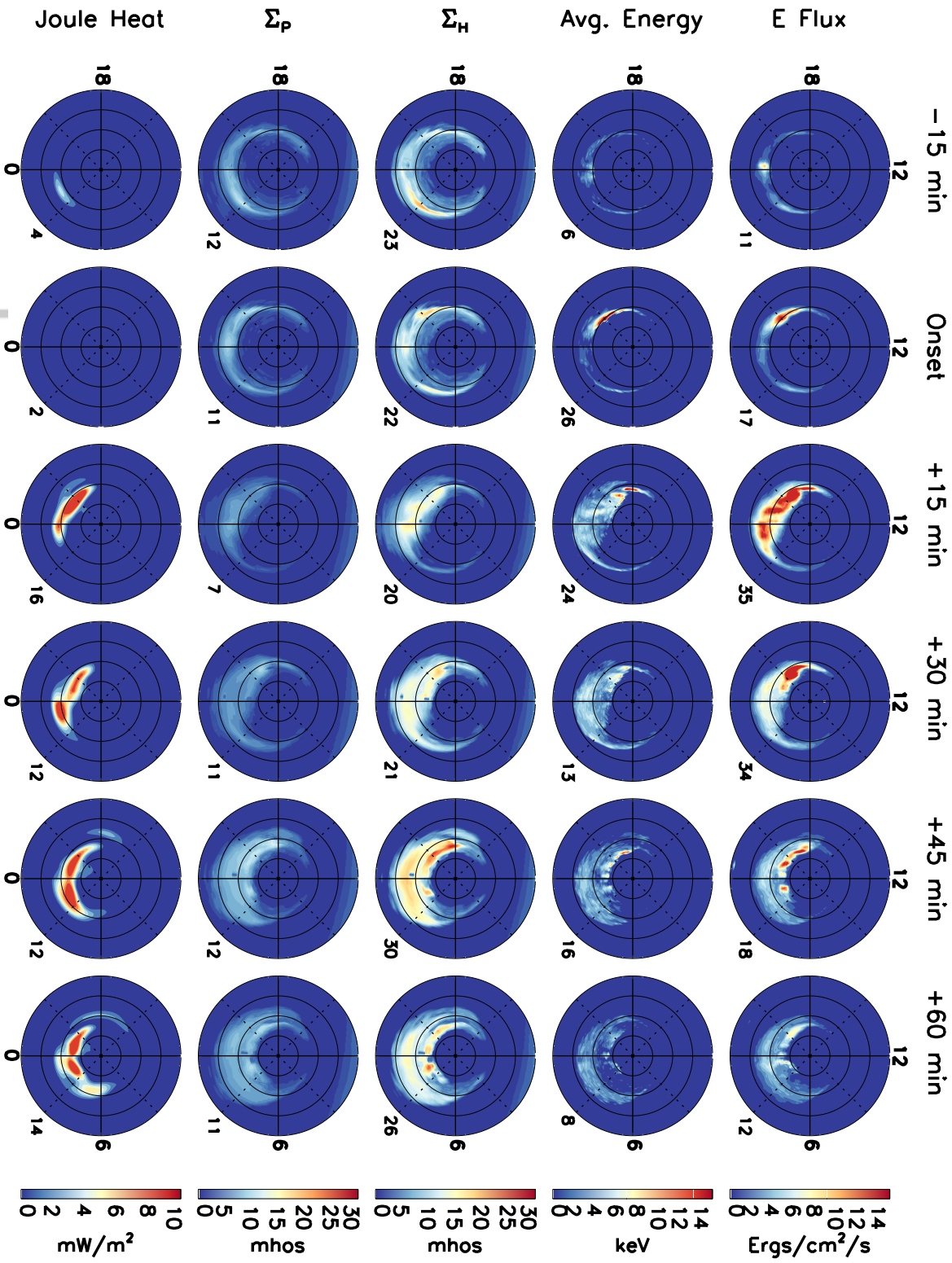
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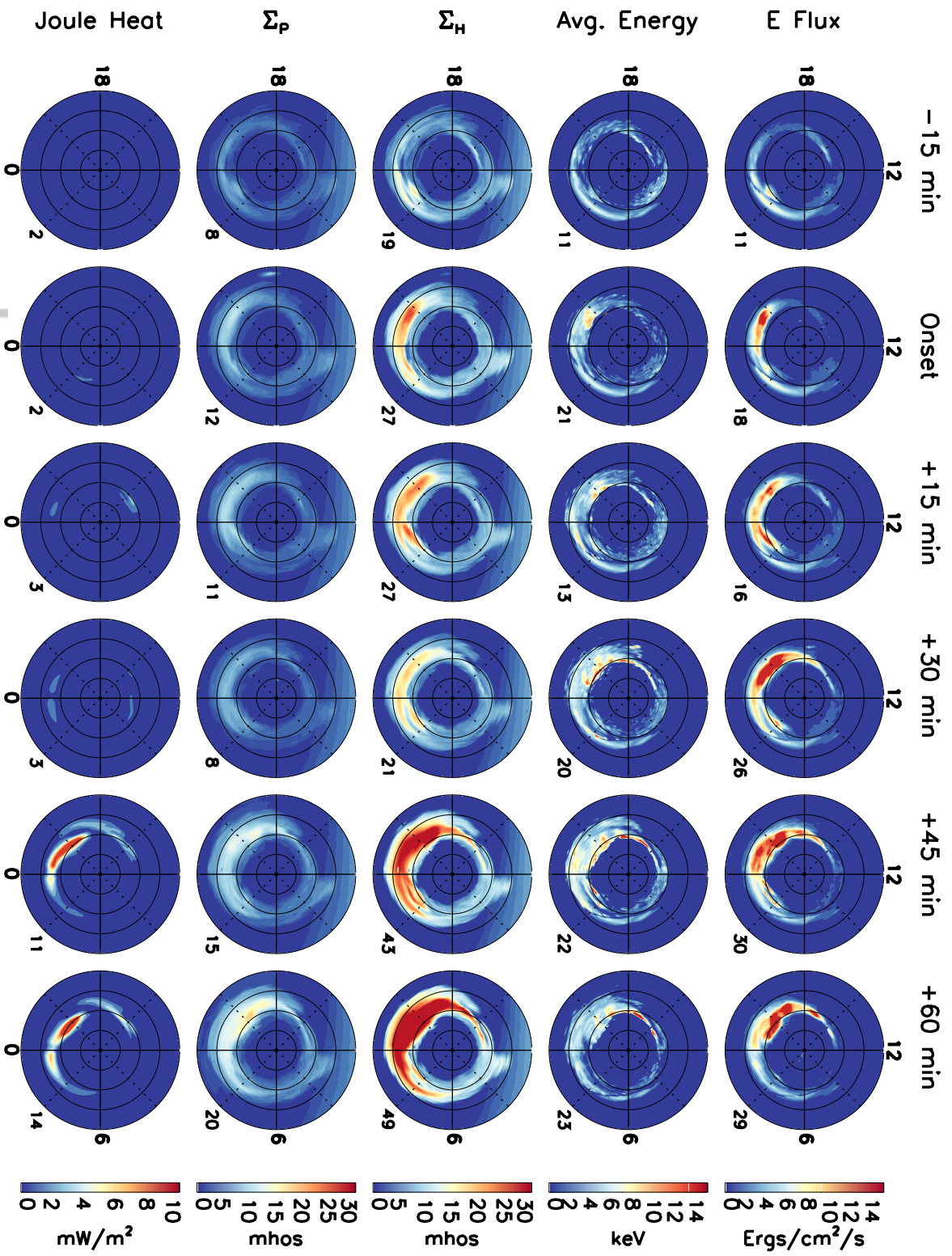
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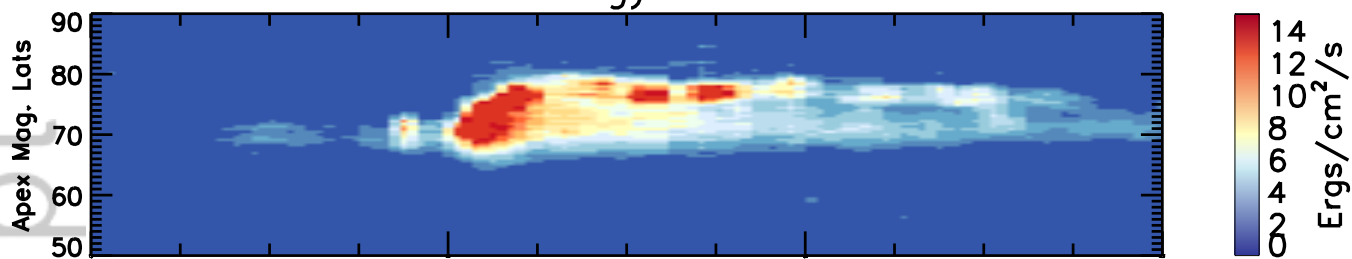
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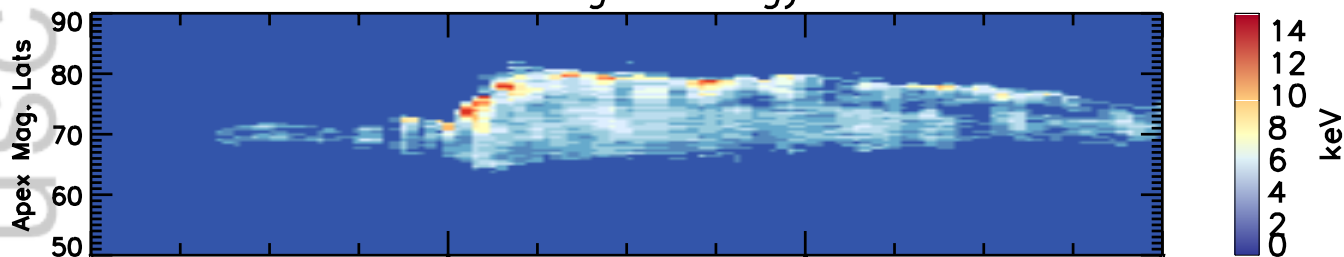
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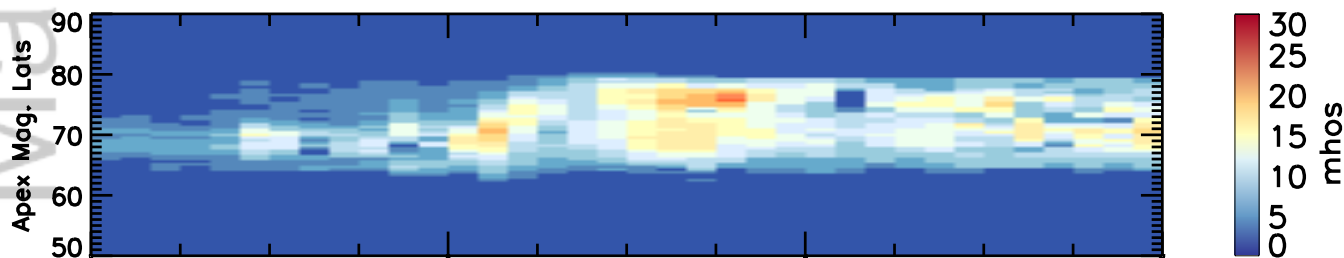
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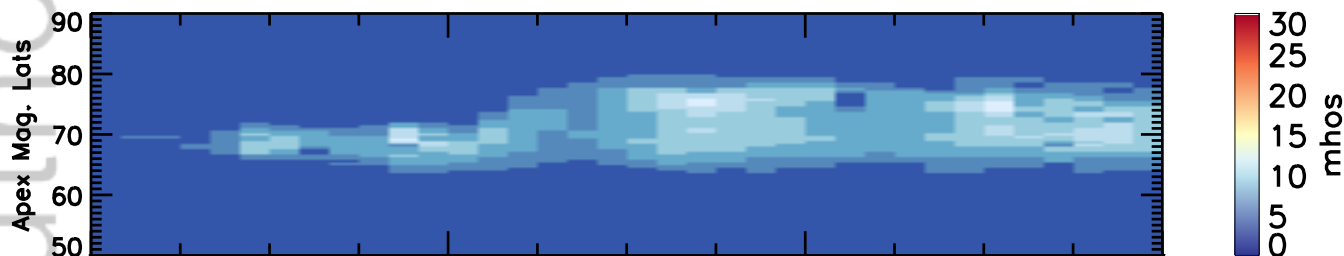
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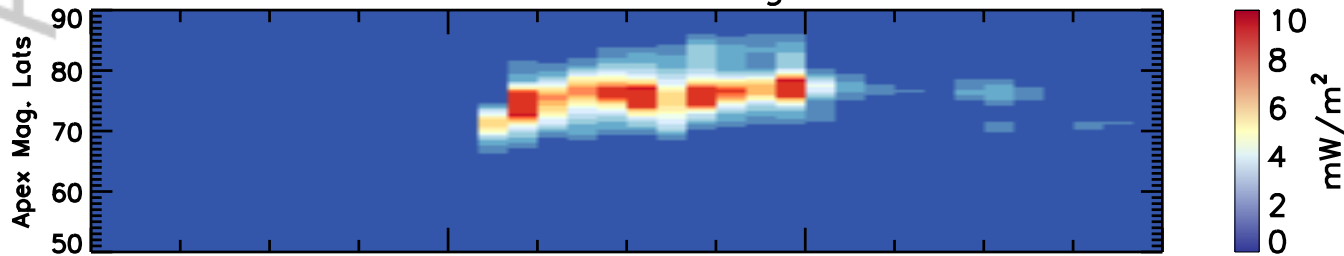
Σ Hall



Σ Pederson



Joule Heating



-1 0 +1 +2

Hours from Onset

