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MS MAURO BRUM MONTEIRO JUNIOR (Orcid ID : 0000-0002-9790-254X)
DR HEIDI ASBJORNSEN (Orcid ID : 0000-0001-8126-3328)

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Hydrological niche segregation defines forest structure and drought tolerance strategies in a seasonal Amazon forest

Mauro Brum^{*1}, Matthew A. Vadeboncoeur^{*2}, Valeriy Ivanov³, Heidi Asbjornsen^{2,4}, Scott Saleska⁵, Luciana F. Alves⁶, Deliane Penha⁷; Jadson D. Dias⁸, Luiz E.O.C. Aragão^{9,10}, Fernanda Barros¹, Paulo Bittencourt¹, Luciano Pereira¹, & Rafael S. Oliveira¹

** These authors contributed equally to this work*

¹.Department of Plant Biology, Institute of Biology, CP 6109, University of Campinas – UNICAMP, 13083-970, Campinas, SP, Brazil.

². Earth Systems Research Center, University of New Hampshire, 8 College Road, Durham, NH 03824 USA

³. Department of Civil and Environmental Engineering, University of Michigan, Ann Arbor, MI, 48109 USA.

⁴. Department of Natural Resources and the Environment, University of New Hampshire, 56 College Road, Durham, NH 03824 USA

⁵. Department of Ecology and Evolutionary Biology, University of Arizona, Tucson, AZ, 85721 USA

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31 ⁶. Center for Tropical Research, Institute of the Environment and Sustainability, University of California,
32 Los Angeles, CA 90095, USA

33

34 ⁷. Society, Nature and Development Department, Federal University of Western Pará - UFOPA, 68035-110,
35 Santarém, PA, Brazil.

36

37 ⁸. Laboratory of Isotopic Ecology, Center for Nuclear Energy in Agriculture - CENA University of São
38 Paulo -Piracicaba, Brazil 13416-000

39

40 ⁹. Remote Sensing Division, National Institute for Space Research, Av. dos Astronautas, 1.758, 12227-010,
41 São José dos Campos, SP, Brazil

42

43 ¹⁰. College of Life and Environmental Sciences, University of Exeter, EX4 4RJ, UK

44

45 Author for correspondence:

46 *Mauro Brum - maurobrumjr@gmail.com*

47 *Matthew Vadeboncoeur - matt.vad@unh.edu*

48 *Rafael Oliveira - rafaelsoliv@gmail.com*

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50

51 **Abstract**

52

53 1) The relationship between rooting depth and aboveground hydraulic traits can potentially define
54 drought-resistance strategies that are important in determining species distribution and
55 coexistence in seasonal tropical forests, and understanding this is important for predicting the
56 effects of future climate change in these ecosystems. .

57 2) We assessed the rooting depth of 12 dominant tree species (representing ~ 42% of the forest
58 basal area) in a seasonal Amazon forest using the stable isotope ratios ($\delta^{18}\text{O}$ and $\delta^2\text{H}$) of water
59 collected from tree xylem and soils from a range of depths. We took advantage of a major
60 ENSO-related drought in 2015/2016 that caused substantial evaporative isotope enrichment in
61 the soil and revealed water-use strategies of each species under extreme conditions. We
62 measured the minimum dry-season leaf water potential both in a normal year (2014; $\Psi_{\text{non-ENSO}}$)

63 and in an extreme drought year (2015; Ψ_{ENSO}). Furthermore, we measured xylem hydraulic traits
64 that indicate water potential thresholds trees tolerate without risking hydraulic failure (P_{50} and
65 P_{88}).

66 3) We demonstrate that coexisting trees are largely segregated along a single hydrological niche
67 axis defined by root depth differences, access to light, and tolerance of low water potential.
68 These differences in rooting depth were strongly related to tree size; diameter at breast height
69 (DBH) explained 72% of the variation in the $\delta^{18}\text{O}_{\text{xylem}}$. Additionally, $\delta^{18}\text{O}_{\text{xylem}}$ explained 49% of
70 the variation in P_{50} and 70% of P_{88} , with shallow-rooted species more tolerant of low water
71 potentials, while $\delta^{18}\text{O}$ of xylem water explained 47% and 77% of the variation of minimum $\Psi_{\text{non-}}$
72 ENSO and Ψ_{ENSO} .

73 4) We propose a new formulation to estimate an effective functional rooting depth, i.e., the likely
74 soil depth from which roots can sustain water uptake for physiological functions, using DBH as
75 predictor of root depth at this site. Based on these estimates, we conclude that rooting depth
76 varies systematically across the most abundant families, genera and species at the Tapajós
77 forest, and that understory species in particular are limited to shallow rooting depths.

78 5) Our results support the theory of hydrological niche segregation and its underlying trade-off
79 related to drought resistance, which also affect the dominance structure of trees in this seasonal
80 eastern Amazon forest.

81 6) **Synthesis:** Our results support the theory of hydrological niche segregation and demonstrate its
82 underlying trade-off related to drought resistance (access to deep water vs. tolerance of very low
83 water potentials). We found that the single hydrological axis defining water-use traits was
84 strongly related to tree size, and infer that periodic extreme droughts influence community
85 composition and the dominance structure of trees in this seasonal eastern Amazon forest

86
87 **Keywords:** *hydraulic traits; embolism resistance; Amazon functional diversity; stable isotopes;*
88 *cavitation; root depth; 2015 ENSO; water potential*

89 90 **Introduction**

91 Water availability is one of the most important factors influencing trait evolution and plant
92 species distribution across terrestrial ecosystems (Silvertown, Araya, & Gowing, 2015). Indeed,
93 drought tolerance is an important driver of species distribution across gradients of seasonality both

94 at the local and regional scale in the Amazon forest (Esquivel-Muelbert et al., 2016; Bonetti, Feng,
95 & Porporato, 2017; Cosme, Schiatti, Costa, & Oliveira, 2017). Nearly half the Amazon exhibits
96 marked seasonality in rainfall and is subject to additional high-magnitude water deficits caused by
97 positive phases of the El Niño - Southern Oscillation (ENSO) (Marengo, Tomasella, Alves, Soares,
98 & Rodriguez, 2011; Jiménez-Muñoz et al., 2016). Despite these periodically adverse conditions for
99 plant growth, trees can sustain transpiration, start new leaf flushing and maintain photosynthesis
100 during dry periods, though the mechanisms underlying this high drought resistance are still under
101 debate (Saleska et al., 2003; Oliveira, Dawson, Burgess, & Nepstad, 2005; Huete et al., 2006;
102 Malhi et al., 2009; Restrepo-Coupe et al., 2013, Wu et al., 2016, Giardina et al., 2018).

103 Deep rooting (Nepstad et al., 1994; Markewitz, Devine, Davidson, Brando, & Nepstad,
104 2010), root hydraulic redistribution (Oliveira, Dawson, Burgess, & Nepstad, 2005; Lee, Oliveira,
105 Dawson, & Fung, 2005), and root niche partitioning (Ivanov et al., 2012) are thought to be
106 important mechanisms explaining the sustained or increased photosynthetic productivity observed
107 during dry seasons of Amazon forests (Restrepo-Coupe et al., 2013; Wu et al., 2016).
108 Notwithstanding, empirical data on the depth distribution of roots of different species in seasonal
109 Amazon forests are scarce (but see Nepstad et al., 1994; Moreira, Sternberg, & Nepstad, 2000;
110 Romero-Saltos et al., 2005; Markewitz, Devine, Davidson, Brando, & Nepstad, 2010; Davidson et
111 al., 2011) relative to the high tree diversity in this ecosystem (Steege et al., 2013; Espírito-Santo,
112 Shimabukuro, Aragão, & Machado, 2005; Fauset et al., 2014; Bonetti, Feng, & Porporato, 2017).
113 During dry periods, root systems of different morphologies can facilitate the avoidance or
114 resistance to water stress. Deeply rooted trees can avoid the stress by accessing high water
115 potential water in deep soils, sustaining gas exchange over longer periods of water scarcity without
116 the need to adjust physiological regulation. Species with shallow roots are likely to become water
117 limited, especially under high VPD conditions common during droughts, leading to large declines
118 in plant water potential and thus implying the need for drought-tolerance strategies (Niinemets,
119 2010; Brum, Teodoro, Abrahão, & Oliveira, 2017).

120 Xylem embolism resistance, estimated as the water potential at which plants lose 50% or
121 88% of their hydraulic conductance (P_{50} and P_{88}), is one of the most important drought resistance
122 traits (Meinzer, Johnson, Lachenbruch, McCulloh, & Woodruff, 2009; Anderegg et al., 2016). This
123 structural trait determines the range of water potentials under which plants can safely transport
124 water without risking hydraulic failure via embolism (Hack et al., 2007; Bittencourt, Pereira, &

125 Oliveira, 2016; Pereira, Domingues-Junior, Jansen, Choat, & Mazzafera, 2017). Under water
126 stress, plants maintain water potential within a safe range by regulating stomatal conductance,
127 which also inevitably reduces carbon uptake (Sperry, Hacke, Oren, & Comstock, 2002; Choat et
128 al., 2012). A hydraulic safety margin can be estimated as the difference between the lowest water
129 potential observed under water-stressed conditions and P_{50} , and is a metric that is being used as a
130 proxy of drought vulnerability (Meinzer, Johnson, Lachenbruch, McCulloh, & Woodruff, 2009;
131 Anderegg et al., 2016).

132 Given the range of structural and physiological traits allowing different water use and
133 drought tolerance or avoidance strategies in different plant species, the hydrological niche
134 segregation (HNS) hypothesis (Silvertown, Araya, & Gowing, 2015) proposes that within a
135 community, plants may differ in hydraulic traits to avoid or tolerate drought along a water
136 availability gradient to avoid competition. These traits include water uptake capability (e.g.
137 different rooting depths or likely leaf water uptake capability), differences in stomatal control, and
138 differences in the xylem structure (Araya et al., 2011; Vinya et al., 2013; Oliveira et al., 2014;
139 Eller, Lima & Oliveira, 2016; Pina, Zandavalli, Oliveira, Martins, & Soares, 2016; Brum, Teodoro,
140 Abrahão, & Oliveira, 2017). In fact, these traits can exert a significant effect on hydrological
141 processes (García-Baquero, Silvertown, Gowing, & Valle, 2016) and determine differences in
142 drought resistance strategies in Amazonian tree species (Ivanov et al., 2012; Bonetti, Feng, &
143 Porporato, 2017). However, empirical data to test the HNS hypothesis in Amazonia are lacking.
144 Furthermore, whether and how different hydrological niches drive differences in drought
145 resistance traits and ecosystem processes are not fully understood (Ivanov et al., 2012;
146 Christoffersen et al., 2016).

147 Natural-abundance stable isotopes of water are a useful tool for determining the depths
148 from which plants acquire water in the soil (Dawson, Mambelli, Plamboeck, Templer, & Tu,
149 2002). However, this method is effective only when there is a gradient in water isotope ratios with
150 depth, caused by evaporative enrichment at the soil surface (Berry et al., 2017). As water isotopes
151 do not fractionate with root uptake, a tree's xylem water isotope ratio reflects the depths from
152 which it is drawing water (Dawson et al., 2002). However, substantial depth gradients in water
153 isotopes are not common in wet tropical forests because the rate of direct soil evaporation is slow
154 relative to the usually continual inputs of meteoric water (Moreira, Sternberg, & Nepstad, 2000;
155 Evaristo, McDonnell, Scholl, Bruijnzeel, & Chun, 2016). Only during extended very dry periods

156 can surface soils in wet tropical forests become significantly isotopically enriched, allowing the
157 use of soil water isotopes to estimate the effective rooting uptake depth at a time when water is an
158 important limiting resource.

159 Here, we took advantage of the severe drought recorded in the eastern Amazon basin
160 (Jiménez-Muños et al., 2016) during the 2015-16 El-Niño to investigate patterns of soil water use
161 among trees. We sampled soil and xylem water to determine the rooting depth of various tree
162 species (root niche partitioning), while also measuring embolism resistance and leaf water
163 potential. We tested two hypotheses: 1) the vertical canopy position of tree species (e.g. canopy,
164 subcanopy, and understory trees) relates directly to the vertical distribution of roots belowground
165 (Ivanov et al., 2012); and 2) shallow-rooted species are more drought-tolerant (i.e., feature traits
166 leading to greater xylem embolism resistance; lower P_{50}), as compared to deeply rooted species,
167 since they are more prone to seasonal water stress under long-term drought cycles. With our results
168 we propose a model to estimate functional rooting depth for the tropical seasonal Amazon forest.

170 **Methods**

171 *Study area*

172 This study was carried out in a lowland tropical rainforest in the Large-Scale Biosphere-
173 Atmosphere km-67 experimental site at Tapajós National Forest near Santarém, Pará, Brazil
174 (54°58'W, 3°51'S). The elevation is 185 m a.s.l., with topographic relief on the order of 10 m
175 (IBAMA, 2004). Soil depth is greater than 12 m and the water table is approximately 100 m deep
176 (Nepstad et al., 2002). Mean total annual precipitation (1998-2013) is 2,037 mm (Fig. S1). During
177 the prolonged dry season of 2015 (August-December), monthly precipitation averaged only 64 mm
178 (Restrepo-Coupe et al., 2016). Mean annual temperature and humidity are 25°C and 85%,
179 respectively (Rice et al., 2004).

181 *Species selection*

182 We studied 12 locally abundant tree species occupying a range of canopy positions along
183 the forest vertical profile (understory, subcanopy, and canopy); diameter at breast height (DBH)
184 ranged from 3 to 159 cm (Table S1). Ten canopy and subcanopy species were chosen based on a
185 long-term forest inventory database of 4 km of permanent transects. These ten species represent ~
186 41.5% of basal area of trees > 10 cm DBH (Pyle *et al.* 2008 updated by Longo, 2013; see S1). We

187 also sampled two very abundant understory species in which most individuals were < 10 cm DBH
188 (Table S1 and S2).

189

190 *Stable isotope analysis*

191 We analyzed water stable isotopes ($\delta^2\text{H}$) and ($\delta^{18}\text{O}$) from tree xylem water and from
192 different soil depths. To sample xylem water, we collected suberized twig segments (~5-10 mm
193 diameter) within reach or by ladder for small trees, and by climbing some larger trees. From other
194 large trees, we sampled sapwood using a 5-mm increment borer at 1.3 m height. For each of the 12
195 species, we sampled only mature individuals within a narrow DBH range to control for possible
196 ontogenetic effects (n= 3-5 trees per species, except for for *Endopleura uchi* where n=1; Table S2).

197 To sample soil water, we dug four pits along existing permanent survey transects and
198 collected soil samples from six different depths 0.15, 0.30, 0.60, 1.0, 1.5, and 2.0 meters. One
199 additional sample was collected at 0.02 m. We also sampled from two existing deeper pits at 4, 6,
200 8, and 10 m depths. Finally, we collected samples from two residential groundwater wells ~15 km
201 from the site, one 30 m deep and the other 60 m deep, to represent water deeper than 10 m. For
202 context, since we do not have rainfall isotope measurements, we used modeled precipitation
203 isotope data to represent the regional meteoric water line (RMWL) (<http://waterisotopes.org>
204 accessed 4/20/2017; Bowen, Wassenaar, & Hobson, 2005). We found a mismatch between the
205 ranges covered by plant and soil $\delta^2\text{H}$ (Fig. 1A), while the range of $\delta^{18}\text{O}$ observed in in plant and
206 soil samples was similar (Fig. S2). For this reason we used only $\delta^{18}\text{O}_{\text{xylem}}$ as a proxy for the depth
207 of water uptake.

208 All samples were quickly sealed in vials, wrapped tightly with parafilm (R) and kept frozen
209 in the laboratory. We extracted water from soil and plant samples at UNICAMP using a cryogenic
210 distillation method (Kryosis-HEKAtech; Ehleringer & Dawson, 1992). Stable isotope ratios of
211 extracted water were analysed at the Center for Stable Isotope Biogeochemistry - University of
212 California, Berkeley, using a hot chromium reactor unit (H/Device TM) interfaced with a Thermo
213 Delta Plus XL mass spectrometer. Data are expressed in delta (δ) notation relative to Vienna mean
214 ocean water standard (V-SMOW) (Coplen, 2011).

215 We collected all samples between 30 November and 3 December 2015, during the most
216 extreme drought on record in this part of the Amazon basin (Jiménez-Muñoz et al., 2016), with a
217 Palmer Drought Severity Index below -3 throughout much of the eastern Amazon. Temperatures

218 during the 2015/2016 ENSO reached a record of 1.5 °C higher than the maximum temperature
219 observed in ENSO October 1997 and 2 °C higher than the peak observed in ENSO January 1983
220 (Jiménez-Muñoz et al., 2016).

221

222 *Isotopic mixing model*

223 We used the 'simmr' package in R to solve mixing model equations for stable isotopic data
224 within a Bayesian framework (Parnell, 2016). The model is used to infer the proportion of water
225 taken up from the various depths of the soil profile based on stable isotope observations of xylem
226 water. The isotopic mixing model was run via the 'simmr_mcmc' function (Markov chain Monte
227 Carlo - MCMC) to produce 1,000 iterations over 4 MCMC chains. In the Bayesian context, the
228 MCMC repeatedly guesses the values of the water uptake proportion and finds those values that
229 best fit the data representing different source of water to plants defined by a soil depth range
230 (mean and standard deviation of $\delta^{18}\text{O}_{\text{soil}}$ in a given soil depth). The simulations thus produce
231 plausible contributions of each soil layer (in terms of proportion) to the xylem water isotope ratio
232 and return a posterior distribution representing a true probability density of data (Parnell et al.,
233 2013).

234 We fitted a segmented linear regression model to describe the relationship between $\delta^{18}\text{O}$
235 and soil depth (Fig. S3) using the 'segmented' package (Muggeo, 2008). The estimated breakpoint
236 was at 0.69 m depth (SD ± 0.17 m). Therefore, we split the soil data set into two distinct depth
237 ranges: 1) shallow soil above 1 m depth, which has a higher proportion of fine roots, larger
238 seasonal variation of water availability, and larger macropores (Nepstad et al., 1994; Broedel,
239 Tomasella, Cândido, & von Randow, 2017); and 2) deep soil at or below 1 m depth, where the soil
240 water content is greater, roots are less abundant, and percolation rates are lower (Broedel,
241 Tomasella, Cândido, & von Randow, 2017). The mixing model analysis was not conducted for
242 *Endopleura uchi* because we only sampled one individual.

243

244 *Embolism vulnerability measurements*

245 We measured xylem vulnerability to embolism as the relationship between the percentage
246 loss of xylem conductivity (PLC) and xylem water potential (Ψ_x in MPa). PLC was estimated from
247 percentage of air discharged (PAD) using the pneumatic method (Pereira et al., 2016). For the two
248 understory species, *Rinorea pubiflora* and *Amphirrhox longifolia*, we used the hydraulic bench

249 method to calculate the PLC (Sperry, Donnelly, & Tyree, 1988). Both methods provide similar
250 estimates of P_{50} and P_{88} values (i.e., water potentials at which the PLC is 50 and 88%; Pereira et
251 al., 2016; Zhang et al., 2018).

252 For both methods, we collected branches longer than 1 m from two to five individuals per
253 species, re-cut the ends under water, and let them rehydrate overnight keeping the leaves inside a
254 plastic bag. To induce embolism, we used the bench dehydration method (Sperry, Donnelly, &
255 Tyree, 1988). We measured Ψ_x as leaf water potential (Ψ_l), after equilibrating the branch inside a
256 black plastic bag for at least one hour prior to making the measurement, using a pressure chamber
257 (PMS 1000; PMS Instruments Co., Albany, OR, USA).

258 Air discharge was measured connecting the entire branch to a vacuum reservoir with 35-40
259 kPa absolute vacuum pressure and calculating the amount of air discharged from the plant to the
260 vacuum reservoir. Air discharge volume was calculated by measuring the pressure in the known
261 volume vacuum reservoir before and after connecting to the plant and using the ideal gas law. The
262 volume of air discharged from each branch was measured several times during branch dehydration
263 at different leaf Ψ_x values. PAD was calculated standardizing air discharge measurements for each
264 branch by minimum and maximum values. Then we calculated the P_{50} and P_{88} by fitting a Weibull
265 function to the data:

$$PAD = \frac{100}{1 + \exp\left[\frac{S_p}{25}(\Psi_x - \Psi_{P50})\right]} \quad (\text{eq.1})$$

266
267 where PAD is the percentage of the total air discharged, Ψ_{P50} is the Ψ_x when PAD equals to 50%,
268 and S_p is the slope of the curve (% PAD MPa^{-1}).

269 For the hydraulic method we used an ultra-low flow meter to measure PLC (Pereira &
270 Mazzafera 2012). Here, five segments of the base of dehydrated branches with ~ 4-6 cm length and
271 3-5 mm diameters length were cut under water, trimmed with a razor blade and attached to the
272 flow meter. Each segment was perfused with degasified and filtered 10 mmol KCl solution by a
273 gravity-induced pressure head (> 6 kPa) and the initial flow in each segment was measured. After
274 the initial measurements, segments were flushed at ~100 kPa to remove all bubbles based on
275 observation at the opposite side not attached to flow meter. The maximum flow, without
276 embolism, was then measured. This procedure was performed in several branches at different

277 dehydration stages. Using these measurements, we calculated the PLC curve by fitting against Ψ_x
278 using Eq. 1.

279

280 *Leaf water potential and hydraulic safety margin measurements*

281 We measured minimum leaf water potential at the peak of the dry season during a non-
282 ENSO year (December 2014; $\Psi_{\text{non-ENSO}}$) and during the ENSO drought year (December 2015,
283 Ψ_{ENSO}), using a pressure chamber (Table S2). We used fully expanded and exposed leaves,
284 collected on sunny, rain-free days between noon and 2:30 pm. For each species, we calculated the
285 hydraulic safety margin (HSM) as the difference between the min Ψ_{leaf} ($\Psi_{\text{non-ENSO}}$ or Ψ_{ENSO}) and
286 the P_{50} or P_{88} , to which we referred as HSM_{P50} and HSM_{P88} , respectively. The HSM was
287 calculated for both the ENSO and non-ENSO years (HSM_{ENSO} and $\text{HSM}_{\text{non-ENSO}}$, respectively).

288

289 *Statistical analysis*

290 We used R 3.3.3 to perform all statistical analyses (R Core Team, 2017). We tested
291 whether the effect of depth on $\delta^{18}\text{O}$ of soil water was strong enough to differentiate the various
292 parts of the soil profile using linear and multiple nonlinear models. We used the AIC criterion to
293 choose the best model to describe this relationship (Burnham & Anderson, 2003).

294 To test the hypothesis that the vertical canopy position of tree species relates directly to the
295 vertical pattern of water uptake by trees, we used linear regression to test for covariation between
296 DBH, $\delta^{18}\text{O}_{\text{xylem}}$, and the proportion of water uptake from the mixing model. We also used linear
297 regression to quantify the co-variation between the uptake depths and xylem resistance to
298 embolism traits. We performed a post-hoc multiple-comparisons Tukey-test to identify taxonomic
299 groups with similar estimated effective rooting depth. We also tested the validity of bivariate
300 models derived here, using observed vs predicted variables by linear models. In order to assess
301 error magnitude, we estimated the root mean squared deviation (RMSD) as

$$RMSD = \sqrt{\frac{1}{n-1} \sum_{i=1}^n (pred - obs)^2}$$

302 (eq.2)

303 which represents the mean deviation of the predicted value in relation to the observed value, in the
304 same units as the variable under evaluation (Piñeiro, Perelman, Guerschman, & Paruelo, 2008).

305 We used an ANCOVA to identify differences in slopes of the relationship between $\delta^{18}\text{O}$ and $\delta^2\text{H}$
306 for the regional meteoric water line and the water lines representing the xylem and soil water
307 samples.

308

309 *Effective functional rooting depth*

310 We used the observed relationships of oxygen isotopes with depth in soil water (Fig 1B),
311 and with tree size in xylem water (Fig 1C), to construct a model to infer the effective functional
312 rooting depth (EFRD), or depth of root water uptake, for any individual tree of known size.
313 Specifically, we fit models to describe the relationships between observed tree diameter, and
314 xylem $\delta^{18}\text{O}$, and between soil water $\delta^{18}\text{O}$ and depth. Firstly, we fit a model describing xylem $\delta^{18}\text{O}$
315 as a function of DBH (i.e., the inverse of relationship in Fig. 1C):

$$\delta^{18}\text{O} = k \cdot [1 - \exp(r \cdot \text{DBH})]$$

316 (eq. 3)

317 Secondly, we fit a model for soil depth predicted (Z_{soil}) as a function of soil water $\delta^{18}\text{O}$
318 (Fig. 1B):

$$Z_{\text{soil}} = \exp\left(\frac{\delta^{18}\text{O} - m}{n}\right)$$

319 (eq. 4)

320 Combining eqs. 3 and 4 yields (Fig. 1D; black line):

$$\text{EFRD} = \exp\left\{\frac{[k \cdot (1 - \exp(r \cdot \text{DBH})) - m]}{n}\right\}$$

321 (eq. 5)

322 EFRD is computed on a continuous basis, and its minimum and maximum values depend
323 on the DBH distribution of stems within the plot. To account for the effect of uncertainty of the
324 parameters fit in Eq. 3 and Eq. 4 on the predictions of Eq. 5, given the observed variations in
325 $\delta^{18}\text{O}_{\text{soil}}$ with soil depth and in $\delta^{18}\text{O}_{\text{xylem}}$ with tree size in larger trees (larger DBH), we performed a
326 bootstrapping analysis to propagate the uncertainty of the m , n , k and r parameters (Davison &
327 Hinkley, 1997). We started by creating a 1000 bootstrap replicate sets sampling from the data with
328 replacement combinations of $\delta^{18}\text{O}_{\text{soil}}$ and $\delta^{18}\text{O}_{\text{xylem}}$ (the *boot* (Canty & Ripley, 2016) and *nlstools*
329 (Baty et al., 2015) packages in R software were used). Using these replicate sets, we obtained the
330 m , n , k and r parameters by applying the fitting procedures to Eqns. 3 and 4. Each bootstrap

331 replicate set of parameters thus yielded a realization of the functional dependence of Eqn 5. The
332 median EFRD fit (i.e., based on 1,000 realizations for each DBH) was taken as the representative
333 form of this dependence (shown as the black line in Fig 1D), with the corresponding m , n , k and r
334 parameter set considered as ‘optimal’. We used this optimal set to evaluate EFRD for all trees in
335 the census inventory for the km-67 study site (4 ha area total, Pyle et al., 2008 updated by Longo,
336 2013).

337 To evaluate the EFRD at the ecosystem level for this site, we calculated the community-
338 weighted mean EFRD (following Muscarella & Uriarte, 2016 for community traits generally). For
339 that, we aggregated inventory data according to taxa (species, genera, and families): we averaged
340 EFRD (obtained for individual trees) within each taxa and weighted each taxon-specific EFRD by
341 the corresponding dominance, based on its fraction of total stem basal area at the site.

342

343 Results

344 *Water stable isotopes revealed species differences in rooting depth*

345 The extreme drought in 2015 caused substantial evaporative isotope enrichment in the first
346 meter of soil in the Tapajós forest, which allowed us to estimate a gradient of rooting depth of tree
347 species in Amazonia using natural-abundance stable isotopes of oxygen (Fig. 1-A-B). Overall, soil
348 depth explained 68% of the average $\delta^{18}\text{O}_{\text{soil}}$ variability (Table 1). Based on the regression model
349 with a segmented relationship between $\delta^{18}\text{O}_{\text{soil}}$ and soil depth, the breakpoint in the $\delta^{18}\text{O}_{\text{soil}}$ depth
350 profile was 0.69 m (Fig. S3, SE= ± 0.17 ; $r^2=0.82$; $p<0.01$). Splitting the soil data into two
351 categories based on this threshold as explained above, the average $\delta^{18}\text{O}_{\text{soil}(<1\text{m})}$ was -3.08 ‰ (SD=
352 ± 1.38), whereas the average $\delta^{18}\text{O}_{\text{soil}(>1)}$ of deeper soil layers were less enriched in $\delta^{18}\text{O}$ (-4.95 ‰
353 ± 0.95). Well water from 30-60 m depth was the most depleted, with $\delta^{18}\text{O}_{\text{soil}}$ varying from -5.85 to
354 -5.35 ‰ (Fig. 1B; blue rectangle), similar to that of the annual mean for precipitation ($\delta^{18}\text{O} = -5.7$
355 ‰; Fig. 1-A).

356 The slope of the regional meteoric water line (RMWL) (slope = 7.70; intercept = 10.99) is
357 0.30, which is lower than the global meteoric water line (GMWL) (ANCOVA; Table 2; $p<0.01$).
358 The soil evaporative line slope was 2.74, which is lower than the RMWL slope (Fig. 1A; Table 2).
359 The $\delta^{18}\text{O}_{\text{xylem}}$ ranged from -0.92 to -6.25 ‰ suggesting a large diversity of rooting depths for this
360 community (Fig. 1A). These data fall along a line with slope 1.43 lower than RMWL slope (slope

361 = 6.27; intercept = - 9.92; Fig. 1A; Table 2), but higher than the soil evaporative line (ANCOVA;
362 Table 2).

363 Species differed systematically in the contribution of water uptake from different depths.
364 The mixing model showed that four species used mostly shallow soil water (<1 m), while six
365 species derived at least 50% of their water from soils deeper than 1 m (Fig. 2A). Rooting depth
366 (using $\delta^{18}\text{O}$ as a proxy) was strongly related to the stem size differences across the species (Fig.
367 1C; $r^2 = 0.72$; $p < 0.01$). With the exception of the understory species *A. longifolia*, smaller trees
368 (DBH < 30 cm) mostly used water from the soil above 1 m depth ($r^2 = 0.41$; $p < 0.05$; Table 1),
369 while large trees (DBH > 40 cm) mostly used water from deeper soil ($r^2 = 0.41$; $p < 0.05$; Table 1;
370 Fig. 2B).

371

372 *Hydraulic traits and their relationships with rooting depth*

373 Among the species in our survey (Fig. 3A), P_{50} ranged from -5.01 to -1.52 MPa and P_{88}
374 ranges from -7.29 to -2.12 MPa. The variation of $\delta^{18}\text{O}_{\text{xylem}}$ explained 49% of the variation in P_{50}
375 and 70% of P_{88} (Table 1; $p < 0.05$), after removal of *Protium apiculatum*, which was an outlier as
376 evaluated by Cook's distance inspection (Aguinis et al., 2013). During a non-ENSO year, the
377 minimum $\Psi_{\text{non-ENSO}}$ ranged from -2.68 MPa to -1.10 MPa. The extreme ENSO drought increased
378 the variation of leaf water potential; Ψ_{ENSO} ranges from -4.43 to -1.06 MPa (Fig. 3B).

379 We observed a higher tolerance of low water potential in shallow-rooted species. Here, the
380 variation of $\delta^{18}\text{O}_{\text{xylem}}$ explained 47% and 77% of the variation of minimum $\Psi_{\text{non-ENSO}}$ and Ψ_{ENSO} ,
381 respectively (Table 1; $p = 0.06$ and $p < 0.01$). Trees with more enriched xylem water (i.e., taken up
382 from shallow soil) exhibited lower leaf water potentials in both non-ENSO and ENSO years (Fig.
383 3B). The HSMP₅₀ and HSMP₈₈ during normal and ENSO years were always positive or very
384 close to zero. The exception was the shallow-rooted species *Rinorea pubiflora*, which showed -
385 1.43 HSMP₅₀ during the ENSO year (Fig. 3-C). Furthermore, there was no relationship between
386 $\delta^{18}\text{O}_{\text{xylem}}$ and HSMP₅₀ and HSMP₈₈ during normal or ENSO years (Fig. 3C, D).

387

388 *Effective functional rooting depth model*

389 The models described in Eq. 3-5 were used to scale up the overall pattern of effective
390 rooting depth to the whole forest community at Tapajós (Fig. 1D). For Eq. 3, the parameters
391 derived from EFRD fit, based on 1,000 realizations for each DBH and considered as 'optimal'

392 representative form of this dependence are $k = -5.356$ and $r = -0.0516$. In this case, the $\delta^{18}\text{O}_{\text{predicted}}$
393 explained 52% of the variation of xylem $\delta^{18}\text{O}_{\text{observed}}$ ($r^2=0.52$; $p=0.002$; $\text{RMSD}=1.35\%$; Table S2).
394 The bootstrapping approach estimates the parameters as $k = -5.579$ ($\text{SE}=0.62$) and $r = -0.061$
395 ($\text{SE}=0.03$). For Eq. 4, using the same approach of Eq. 3, the parameters derived are $m = -3.829$
396 and $n = -0.588$. In this case, the predict Z_{soil} explained 84% of the variation of Z_{soil} observed
397 ($r^2=0.84$; $p<0.001$; $\text{RMSD}=12.9$ m; Fig. S4). Here, the results from bootstrapping were $m = -4.076$
398 ($\text{SE}=0.157$) and $n = -0.567$ ($\text{SE}=0.087$; Fig. 1-C). The model's deviation from the 1:1 line indicates
399 that this prediction was biased to shallow soil (modeled line below 1:1 line). Despite the larger
400 differences between predicted values and observed values ($\text{RMSD}=12.9$ m), we used our model
401 derived from Eq. 5 to extrapolate the EFRD to the larger km-67 study area at Tapajós Forest. The
402 estimated average maximum EFRD based on the maximum DBH measured within the plot was
403 13.33 m (25%-75% interquartile interval: 6.73 m to 30.14 m).

404

405 *Scaling rooting depth estimates to the ecosystem*

406 Our estimates of ecosystem-scale rooting depth distribution indicated a disproportionate
407 number of individuals drawing water from depths shallower than 1.32 meter (Median= 0.24 m;
408 Mean= 1.32 m). The dry season community-weighted mean of EFRD was 3.56 m. There were
409 specific families, genera, and species restricted to drawing water from shallow soil and others
410 exclusively drawing water from deep soil, as dictated by their DBH distribution in the studied
411 ecosystem. This was also confirmed when we contrasted the 10% most dominant taxa (higher
412 absolute dominance ($\text{ADo}_{(\text{DBH})}$: $\text{m}^2\cdot\text{ha}^{-1}$), as shown in Fig. S5 and S6 that illustrate differences in
413 EFRD based on multiple group comparison (Tukey HSD Test; $p<0.05$). The scatterplot showing
414 the variations in EFRD (hydrological axis) and DBH (light availability axis) across taxa suggests a
415 single predominant axis of variation representing a niche spectrum defined jointly by access to
416 water and light (Fig. 1c; Fig. 4). Differences in circle sizes given by post hoc Tukey test also show
417 that only a few taxonomic groups are more dominant than others (Fig. 4). The limited overlap
418 between circles demonstrates the segregation of taxa along a single niche axis of light-water
419 availability, from lower light availability and restricted water access, to higher light availability
420 and deep water access (Fig. 4; Fig. S6). By pooling individuals into DBH classes (every 10 cm) to
421 calculate the absolute dominance of each class ($\text{ADo}_{(\text{DBH})}$), we found a linear negative response of
422 $\text{ADo}_{(\text{DBH})}$ as a function of EFRD (Fig. S7; $r^2=0.63$; Table 1).

423

424 **Discussion**

425 Our results provide strong evidence for segregation of root water uptake in soil as an
426 important strategy allowing multi-species coexistence in a seasonal Amazon forest. Specifically,
427 using $\delta^{18}\text{O}$ to estimate the effective rooting depth, we empirically confirmed a modeling-based
428 hypothesis (Ivanov et al., 2012) that the above-ground vertical structure is related to rooting depth
429 distribution (Fig. 1-C). Furthermore, we showed that greater rooting depth (using $\delta^{18}\text{O}$ as proxy of
430 root depth) is associated with lower xylem resistance to embolism across taxa. Shallow-rooted
431 species, which dominate the understory, compensate for only having access to shallow soil with
432 lower water potentials by having greater xylem resistance to embolism (i.e. lower P_{50} and P_{88} ;
433 Fig. 2-A) and anisohydric stomatal control strategy (Tardieu, 1996), demonstrated by their weaker
434 year-to-year water potential regulation even during an extreme 2015/2016 ENSO drought (Fig. 3-
435 B). Our results also demonstrate, for the first time, the functional integration of below- and above-
436 ground hydraulic traits as drivers of drought-avoidance and drought-tolerance strategies for
437 Amazon tree species. These results support the niche theory in tropical forests where it is expected
438 that plant species strategies range from acquisitive with high growth rates at high resource levels
439 (e.g. light and water) to more conservative, slowly growing species that are tolerant of shade and
440 drought (Sterck et al., 2011). Additionally, our rooting depth results coupled with forest structure
441 parameters allowed the development of a new formulation to estimate the effective functional
442 rooting depth (EFRD), defined as the likely soil depth from which roots can sustain water uptake
443 for physiological functions (Fig. 1-D). The EFRD will be useful to modellers interested in
444 integrating the coordination between below- and above-ground plant functions into predictions of
445 forest productivity responses and forest resilience to climate change forecasts in seasonal
446 Amazonia (Meir et al., 2009; Markewitz, Devine, Davidson, Brando, & Nepstad, 2010; Ivanov et
447 al., 2012; Restrepo-Coupe et al., 2016; Christoffersen et al., 2016; Fun et al., 2017).

448

449 ***Ecological implications of soil vertical root partitioning***

450 The strong relationship between functional rooting depth and tree size (Fig. 1-B, Eq. 3)
451 that was used to scale our results to the ecosystem level suggests that soil water and light, which
452 both vary seasonally, are the resources for which trees must compete most strongly in this seasonal
453 Amazon forest, and that resource partitioning contributes to species coexistence (Sterck et al.,

2011). Indeed, the distributions of leaf area and light environments are strongly related to DBH distribution of trees in Tapajós, as individuals optimize their productivity over the vertical gradient to create consistent relationships between canopy light environments and biomass growth (Stark et al., 2012; 2015). Our results further suggest that rooting depth increases with tree height, compensating for the greater evaporative demand at the top of the canopy (McDowell & Allen, 2015) and allowing larger trees to be photosynthetically active during the dry season (Giardina et al., 2018). The greater light interception of taller trees may allow them to afford the carbon costs of growing deeper roots. Based on water isotopes in xylem and soil water, our estimates show that maximum effective root depth can be as deep as ~13 m (Fig. 1-D, black line; Table S2). This deeper root investment may be attributed to the temporal variability of water availability in Tapajós, inducing deep root investment to reduce water stress and competition during extreme dry conditions. These belowground allocation rules force a trade-off at the community level between light use (Stark et al., 2015; Wu et al., 2016) and water use strategies (Ivanov et al., 2012). Though whether rooting depth and drought tolerance correlate with phenological strategies remains to be tested, we would expect the leaf phenology of canopy trees to respond more to light availability than water, while the phenology of shallow-rooted smaller trees would be driven mostly by water availability.

The spatial variation in light and water (“eco-hydro-light” niche axes) along vertical profiles drives niche partitioning, and forest structural and taxonomic organization within the Tapajós community (Fig. 4). We found specific groups of families (five), genera (seven), and species (five) dominating a range of soil depths and canopy layers (given by DBH variance) (Fig. 4, Post-Hoc Tukey clustering, see legend). Among the 10% most dominant groups of species, genera and families, there is a relatively limited overlap in EFRD. These results illustrate above- and below-ground space partitioning at a fine scale by the most dominant groups and represent an average distance sufficiently small for species to minimize competition and dominate in each niche of light (Kohyama, 1993; Stark et al., 2012) and water availability (Araya et al., 2011). The strength of interspecific interaction among coexisting dominant species may reflect the distribution of long-term hydraulic traits selected within the community (Hillebrand, Bennet, & Cadotte, 2008). In fact, we showed a diversity of drought-related strategies dependent on rooting depth, xylem embolism resistance (Fig. 2), and aboveground forest vertical structure that might help explain the dominance structure of the Tapajós forest community.

485 We estimated an effective rooting depth for the whole community as ~3.6 m, considering
486 the community-weighted EFRD mean by species dominance. The weighted EFRD mean shows the
487 proportional influence of the dominant species (Muscarella & Uriarte, 2016) and determines the
488 depth with the greatest influence on water use in Tapajós, indicating the depth at which most of the
489 water extraction occurs during dry season. These results indicate the prevalence of a drought-
490 avoidance strategy defined by deeper water access to mitigate the seasonal and interannual drought
491 cycles. Indeed, quantification of fine-root vertical profiles showed only ~10% of the total root
492 mass occurred between 4 to 10 m in the eastern Amazon (Nepstad et al., 1994). A tracer-water
493 experiment suggested water access by dominant mid-canopy *Coussarea racemosa* (= *C. albicans*),
494 *Sclerolobium chrysophyllum* (= *Tachigali chrysophylla*) and *Eschweilera pedicellata* around 2
495 meter depth (Romero-Saltos et al., 2005). During a throughfall exclusion (artificial drought)
496 experiment conducted in Tapajós, based on measures of soil electrical resistivity, it was observed
497 that deep-root water uptake increased up to 18 m in the treatment, in contrast with the control plot
498 (Davidson et al., 2011). Further, root systems may allow the hydraulic redistribution of water from
499 shallow to deep soil (downward) for storage during the onset of the wet season; or the transfer of
500 stored water from deep to shallow soil (upward) to meet the demand by plants during the dry
501 season (Lee, Oliveira, Dawson, & Fung, 2005), as has been documented in *Manilkara elata*,
502 *Protium robustum* and *Coussarea racemosa* (= *C. albicans*), which are dominant species at our site
503 (Oliveira, Dawson, Burgess, & Nepstad, 2005).

504 Shallow-rooted small trees, between 20-30 cm diameter, (Fig. S4) are quite abundant,
505 representing 45% of the forest aboveground biomass in Tapajós (Vieira et al., 2004). This
506 structural pattern contributes to the median EFRD at a shallow soil depth (~0.37 m), indicating
507 high competition for water uptake in this layer. When water is abundant during the rainy season,
508 competition for water is not very important relative to other resources, such as nutrients and light.
509 However, the high climatic variability and the persistently high water demand during the dry
510 season forces a trade-off in terms of tree water use at the community level (Schwinning & Kelly,
511 2013). Moreover, considering the reduced light availability for understory trees during the rainy
512 season (Huete et al., 2006; Restrepo-Coupe et al., 2013), the relatively abundant light during the
513 dry season is a window of opportunity for these trees to maximize productivity, which must be
514 sustained by their strategy of drought-tolerance. The regular seasonal cycle of water scarcity and
515 the frequent droughts in the eastern Amazon appear to have selected for species that are successful

516 at either avoiding or tolerating water stress, as indicated by the great range of P_{50} and P_{88} we
517 observed (Fig. 3).

518 We showed that $\delta^{18}\text{O}$ variance (a proxy for rooting depth) explained 40% and 70% of
519 species-specific variation in P_{50} and P_{88} , respectively. Shallow rooted trees had greater xylem
520 resistance to embolism formation, i.e., the lower P_{50} and P_{88} values down to -5 MPa and -7.29
521 MPa, as compared to deeply rooted trees that do not require such high tolerance due to their access
522 to a larger reservoir of deep water (Fig. 3A). Consequently, shallow-rooted species showed greater
523 decreases in leaf water potential and hydraulic safety margin than deep-rooted species during the
524 extreme 2015 drought (Fig. 3D and Fig. S7). This suggests that shallow-rooted species can tolerate
525 a certain degree of embolism to maintain carbon uptake under dry conditions (Meinzer et al.,
526 2009), although most species presented positive safety margins. For instance, the higher seasonal
527 changes in hydraulic safety margins for shallow-rooted understory species (Fig. 3) reinforce the
528 notion that this group of species can decrease their water potential and operate under a narrow
529 safety margin in order to keep their stomata open and maintain gas exchange during the
530 opportunistic time of higher light level (with exception of *Protium apiculatum*, which was very
531 vulnerable to embolism and showed relatively little variation in leaf water potential). Indeed, the
532 shallow-rooted *R. pubiflora* experienced the most negative leaf water potential among species.
533 This behavior indicates a strategy consistent with ‘anisohydric’ pattern: reduced control over leaf
534 water potential in response to changing environmental conditions, therefore leading to a higher risk
535 of xylem embolism formation (Hacke *et al.* 2006; McDowell *et al.* 2008).

536 Such anisohydric behavior of the lower canopy is an important strategy to sustaining plant
537 productivity, considering that drought-induced mortality risk might be mitigated by some other
538 compensatory mechanism such as xylem structural reinforcement or plasticity (Fonti et al., 2010;
539 Markesteijn et al., 2011; Cosme, Schiatti, Costa, & Oliveira, 2017). In fact, our results help to
540 explain the low mortality rates observed in small trees (DBH < 20 cm) in throughfall exclusion
541 experiments in the Amazon (Nepstad, Tohver, David, Moutinho, & Cardinot, 2007; da Costa et al.,
542 2010), and even the increased growth rates of small trees following the substantial mortality of
543 larger trees during droughts at two Eastern Amazon forest sites (Brando et al., 2010; Rowland et
544 al., 2015).

545 We suggest that these patterns are likely also reflected along tree ontogeny, though our data
546 on mature trees of each species did not address this. Young trees of even dominant canopy species

547 must start out with relatively shallow roots, and during this phase of development they may need a
548 high degree of embolism resistance to survive in the dense understory environment (Fig. S7) where
549 competition for space and water among smaller trees can be intense (Rice et al., 2004; Starck et al.,
550 2015). Reduction in investment that confers greater embolism resistance would be consistent with
551 access to deeper soil water reservoirs with greater root depth (Fonti et al., 2010). Indeed, a
552 significant increase in P_{50} along with DBH during tree growth was observed across several species
553 in Caxiuana forest (Rowland et al., 2015).

554 Despite deep water access, larger trees are generally more vulnerable to xylem embolism
555 and appear to operate closer to their safety margins than do understory trees (Fig. 3-C-D),
556 particularly given the high-light and high-VPD conditions that they are frequently exposed to
557 during droughts. Indeed, drought-induced mortality of larger trees has been observed in many
558 areas of the Amazon forest (Nepstad, Tohver, David, Moutinho, & Cardinot, 2007; Phillips et al.,
559 2010; Rowland et al., 2015; Bennett, McDowell, Allen, & Anderson-Teixeira, 2015). Low
560 $HSMP_{50}$ may indicate that stomatal regulation takes the full advantage of the small safety range of
561 xylem pressures (Choat et al., 2012), while running the risk of severe hydraulic failure in drought
562 conditions extreme enough to deplete the deep soil water these trees use (Anderegg et al., 2016).
563 Deep water access represents a competitive advantage in terms of water use, making it possible for
564 trees to avoid drought and invest in growth, rather than investing in producing a very resistant
565 hydraulic transport system (Stark et al., 2015). However, basic principles of plant physiology
566 predict that vulnerability to drought stress increases with tree height; taller trees need to deal with
567 higher VPD, light interception, and hydraulic path length (McDowell & Allen 2015), which might
568 be compensated by other water regulation strategies, such as leaf turgor loss avoidance (Bartlett,
569 Scoffoni, & Sack, 2012; Skelton, West, & Dawson, 2015). Indeed, at the year-to-year scale (2014-
570 2015), deeply rooted larger tree species showed a more isohydric behavior (in terms of water
571 potential regulation; Fig. 3B).

572 Further studies need to investigate the minimum threshold of deep root water access
573 required for tree survival during prolonged drought. Taken together, our results suggest that long-
574 term droughts might increase embolism risk in species with higher P_{50} and P_{88} and, if associated
575 with the depletion of non-structural carbohydrates reserves beyond critical thresholds (McDowell
576 et al., 2008; Sala, Woodruff, & Meinzer, 2012), may lead to increased risk of large tree mortality.
577 Furthermore, additional studies are needed to elucidate the effect of hydrological drought on

578 decreasing deep water recharge and inducing canopy turnover and ecosystem changes in Amazon
579 (Taufik et al., 2017; Chitra-Tarak et al., 2018; Leitold et al., 2018). Greater mortality in deep-
580 rooted trees was observed during a drought in a dry tropical forest in India (Chitra-Tarak et al.,
581 2018), and attributed to the delayed recharge of deeper water following drought. Additionally,
582 periods with low groundwater recharge may amplify wildfire occurrence as observed in a tropical
583 forest in Borneo (Taufik et al., 2017), and may intensify the recently observed increase wildfire in
584 eastern Amazon (Aragão et al., 2018).

585

586 *Estimated Functional Rooting Depth*

587 Tree diameter is currently used in allometric models to estimate coarse root biomass (Tobin
588 et al., 2007; Gou et al., 2017). Here we suggest an empirically based allometric model to estimate
589 the effective rooting depth (EFRD) for a seasonal Amazon forest using DBH (Eq. 5). However,
590 some uncertainties should be considered.

591 When we examined EFRD responses using the bootstrapping estimates of parameters m , n ,
592 k and r , the results yielded an uncertainty of 6 to 30 m depth in soil profile (Fig. 1D and S4).
593 Despite this, the model was sufficiently sensitive to demonstrate a certain degree of vertical
594 rooting depth partitioning as a function of tree DBH in all simulations; thus, we are confident
595 about the existence of a rooting depth pattern in Tapajós forest that can be inferred by tree DBH.
596 The strong correlation between DBH and $\delta^{18}\text{O}_{\text{xylem}}$ and the isotopic mixing model results (Fig. 1C
597 and Fig. 2C) supports the idea of larger trees using relatively large amounts of water below 1 m
598 (and likely down to 13 m). While we are aware that our results only provide a picture of water use
599 dynamics during one dry period of a single year (the extreme drought of 2015), we also believe
600 there is a substantial plasticity of root water uptake to allow for shifts in effective rooting depth in
601 response to changes in soil dryness conditions (Doussan, Pierret, Garrigues, & Pagès, 2006;
602 Schröder, Javaux, Vanderborght, Körfgen, & Vereecken, 2008; Couvreur, Vanderborght, Draye, &
603 Javaux, 2014; Fun et al., 2017).

604 An additional source of uncertainty is that xylem water stable isotope composition reflects
605 a mixture of uptake from multiple depths (see Moreira, Sternberg, & Nepstad, 2000; Romero-
606 Saltos et al., 2004), considering that Eq. 5 requires the predicted $\delta^{18}\text{O}_{\text{xylem}}$ from Eq. 3 to be used
607 and this creates a noise in our estimates. However, the 2015/2016 ENSO induced both a strong
608 isotopic gradient in the soil water, mainly above 1 m depth (Fig. 1), as well as strong competition

609 for water, likely leading to a disproportionate uptake of water by each tree from the deepest soil
610 layers it had access to. Furthermore, the mixing model results highlight that even if xylem water
611 stable isotope composition reflects uptake from multiple depths, there was a clear distinction of
612 smaller trees using more shallow water (< 1 m) and larger trees using deeper water (2 -13 m; Fig.
613 2). The uptake of water from multiple depths may occur in the deeper soil layers, but we were not
614 able to clearly distinguish an isotopic signal in each layer below 1 m depth.

615 Our results are supported by the survey on root morphologies conducted at Tapajós
616 showing that the dimorphic rooting habit with tap roots growing vertically towards deeper soil
617 layers are strongly represented at this site (Oliveira et al., 2005). This contrasts with the pattern
618 observed in a hyper-humid aseasonal Panamanian forest, where larger trees (DBH > 50 cm)
619 acquired water from shallow soil layers (more $\delta^{18}\text{O}$ enriched water) and small trees from soil up to
620 2 m depth (Meinzer et al., 1999). These authors concluded that larger trees invest more in
621 widening their horizontal root distribution (Meinzer et al., 1999); however, we expect this shallow,
622 extensive strategy to be most common in environments without a long and intense dry season, or in
623 sites that lack a deep soil (Canadell et al., 1996; Fun et al., 2017). The strategy of competing for
624 shallow soil water would seem to be advantageous only when soil moisture at these depths is fairly
625 reliable, or if other drought-stress avoidance strategies such as deciduous dormancy are employed.
626 Intense competition for a limited amount of shallow soil water during an extended dry season
627 would likely require high embolism resistance, especially for evergreen canopy trees which must
628 endure high VPD and high radiation flux at the top of the canopy throughout the dry season.

629 Even with the aforementioned precautions, we highlight the model expressed in Eq. 5 as a
630 good approximation of the overall functional rooting depth in the Tapajós forest (Fig. 3C).
631 Evidence of root activity around 13 m depth in seasonal Amazon forest (Nepstad et al., 1994;
632 Markewitz, Devine, Davidson, Brando, & Nepstad, 2010; Davidson et al., 2011; Ivanov et al.,
633 2012), and the expected average maximum rooting depth for tropical evergreen forest of 15 (± 5.4)
634 m (Canadell et al., 1996), lend substantial support to scaling up the EFRD for a large area in
635 Tapajós (Rice et al., 2004). We can use simple-to-measure variables, such as DBH, to parameterize
636 model representations of difficult-to-measure underlying functions, such as rooting depth. It will
637 be relevant to determine whether the EFRD - DBH relationship holds in non-seasonal tropical
638 rainforests and whether we can better describe ecosystem processes by incorporating such
639 relationships into land-surface and ecohydrological models.

640

641 ***Conclusions***

642 The distribution of leaf area and light environments in seasonal Amazon forests is
643 integrally connected to the size (diameter) distribution of trees (Stark et al., 2015). We have
644 confirmed that tree size is also related to the effective rooting depth (Ivanov et al., 2012), and
645 demonstrated that such interdependence is associated with different hydraulic strategies.
646 Specifically, our data show coordination between the rooting depth and embolism resistance within
647 seasonal Amazon forests, suggesting a trade-off between drought avoidance (i.e. deep rooted) and
648 drought tolerance (i.e. embolism-resistance). Drought-avoiding species are characterized by deep
649 roots, which allow relatively little investment in mechanisms facilitating embolism resistance, as
650 well as operation near the limit of hydraulic safety margin. This is the predominant strategy in
651 canopy/overstory trees with greater canopy exposure to light. Drought tolerance, the predominant
652 strategy in shallow-rooted and light-limited understory trees, is demonstrated by the lower P_{50} and
653 P_{88} and high year-to-year variability in xylem water potential. These complementary strategies
654 allow niche partitioning within the ecosystem and affect the structure of dominant species in the
655 community, driven by both water and light availability. Further studies on these traits should be
656 conducted in other tropical forests to evaluate the generality of these findings.

657 Capturing species variation in structural and physiological traits in a hyper-diverse tropical
658 forest represents an enormous challenge for model parameterizations, but our results suggest that
659 much of the variation can be captured through the fairly straightforward relationships between tree
660 diameter, canopy position, rooting depth, and hydraulic trait parameters in seasonally dry tropical
661 forests growing on deep soils. Our results also help to explain the dry-season green-up and
662 increase in productivity observed in seasonal Amazon forests as dependent on canopy trees' access
663 to deep soil water (Saleska et al., 2003; Restrepo-Coupe et al., 2016). However, we stress that
664 further research is needed to better constrain the partitioning of water use and forest productivity
665 along vertical canopy gradient under a range of water-abundant and water-limiting conditions.

666

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678

679 **Authors contributions**

680 MB, MAV, and RSO conceived the project and wrote the first version of this manuscript, along
681 with HA, LFA and VI. MB, MAV, DP, and JDD collected field data. MB, MAV, PB, FB, SS, LP,
682 and VI contributed to data analyses. SS and LEOCA contributed with infrastructure at the field
683 site. All authors contributed to the final version of the manuscript.

684

685 **Data Accessibility**

686 Data available from the Dryad Digital Repository: <https://doi.org/10.5061/dryad.v704dj2> (Brum,
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688

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List of Tables

999 **Table 1** Results for linear models (response : predictive) performed.

Linear models	r ²	F	df	p<	a	b	c	Figure
log(Soil Depth) : $\delta^{18}\text{O}_{\text{soil}}$	0.68	8.77	12	0.001	-1.22	-4.5	-	1-B
DBH : $\delta^{18}\text{O}_{\text{xylem}}$ *	0.72	12.13	9	0.01	-1.95	10.32	-0.76	1-D
SWUP : log(DBH)	0.41	6.46	9	0.05	-0.12	0.81	-	2-B
DWUP : log(DBH)	0.41	6.46	9	0.05	0.18	0.12	-	2-B
P ₅₀ : $\delta^{18}\text{O}_{\text{xylem}}$ **	0.47	7.702	8	0.02	-3.10	-0.29	-	3
P ₈₈ : $\delta^{18}\text{O}_{\text{xylem}}$ **	0.70	19.32	8	0.001	-0.74	-8.7	-	3
MLWP _{normal} : $\delta^{18}\text{O}_{\text{xylem}}$	0.47	5.32	6	0.06	-0.29	-3.10	-	3
MLWP _{ENSO} : $\delta^{18}\text{O}_{\text{xylem}}$	0.77	20.98	6	0.01	-0.68	-5.19	-	3

1000 *The DBH : $\delta^{18}\text{O}_{\text{xylem}}$ is a polynomial relationship given by $\text{DBH} = b + a(\delta^{18}\text{O})^2 + c(\delta^{18}\text{O})^3$; ** After
 1001 removal of *Protium apiculatum*, which was an outlier; **SWUP**: shallow (<1m) water use
 1002 proportion; **DWUP**: deep (>1m) water use proportion; **MLWP**: midday leaf water potential;
 1003 **ADo_(DBH)**: absolute dominance of each DBH class. **P₅₀**: water potential at which plants lose 50%
 1004 of their hydraulic conductance; **P₈₈**: water potential at which plants lose 88% of their hydraulic
 1005 conductance; **DBH**: diameter at breast height; **EFRD**: estimated functional root depth.

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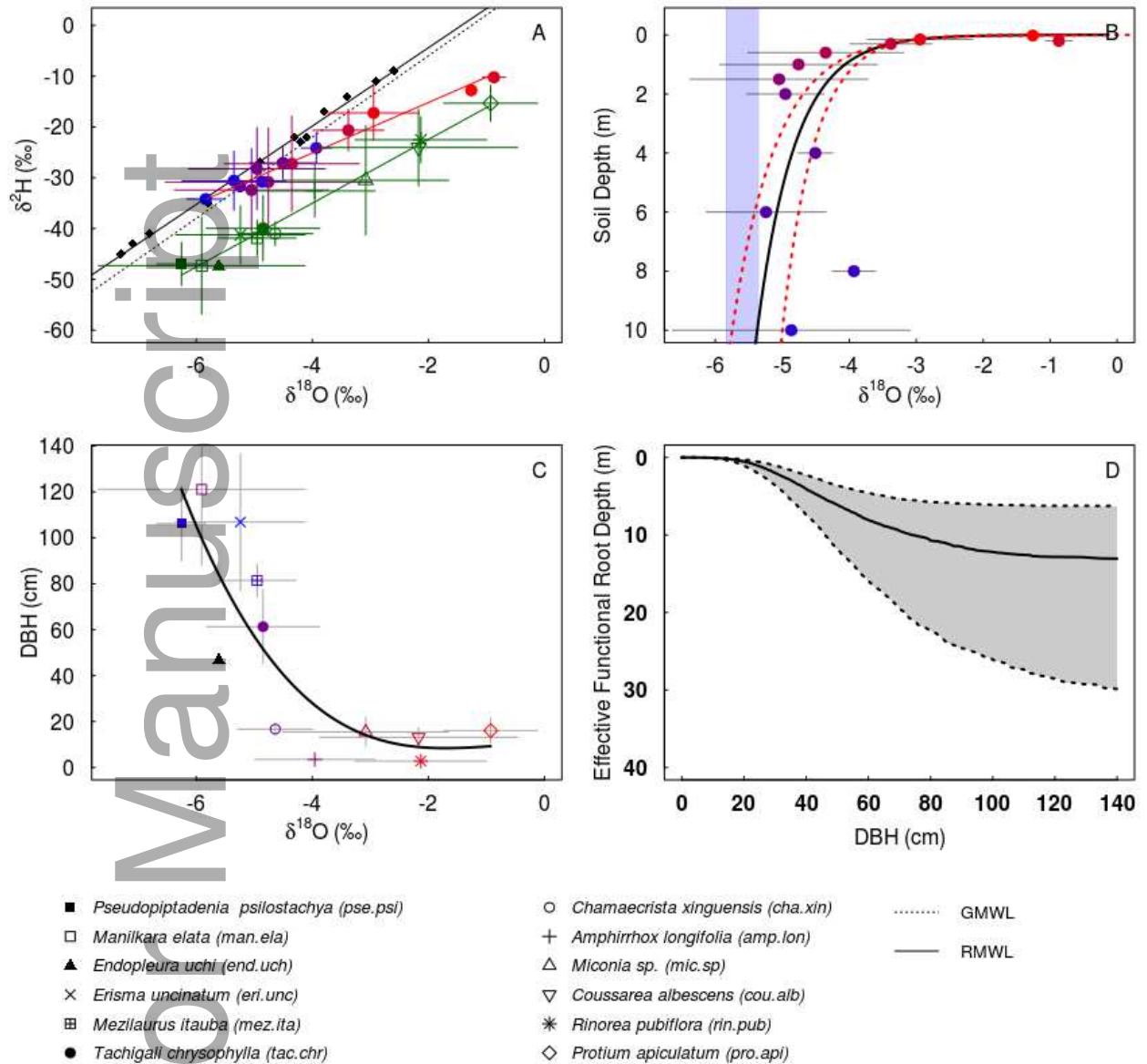
1008 **Table 2** ANCOVA statistics comparing the slopes of linear relationships between $\delta^{18}\text{O}$ and $\delta^2\text{H}$.

ANCOVA	F	df	AIC	p	Slope Difference
RMWL : SWEL	65.64	1	92.67	0.001	-2.74
GMWL : RMWL	6.127	1	64.12	0.01	-0.29
RMWL : PWEL	97.37	1	87.51	0.001	-1.43
SWEL : PWEL	17.60	1	101.50	0.003	-1.30

1009 **GMWL**: global meteoric water line; **RMWL**: regional meteoric water line; **SWEL**: soil water
 1010 evaporative line; **PWEL**: plant water evaporative line.

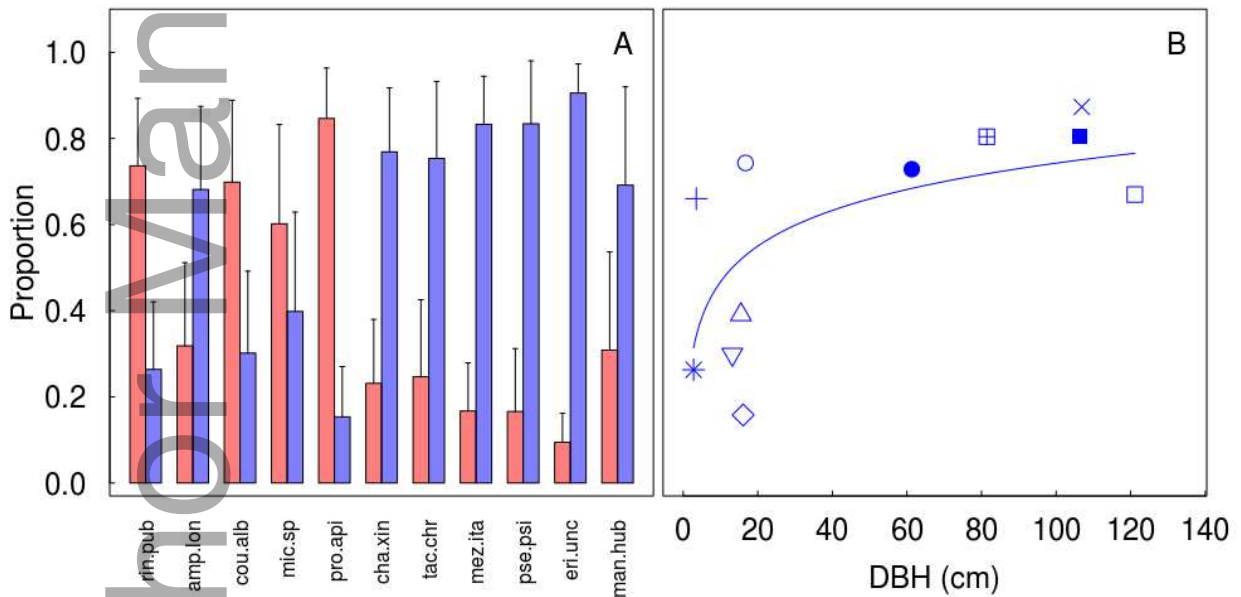
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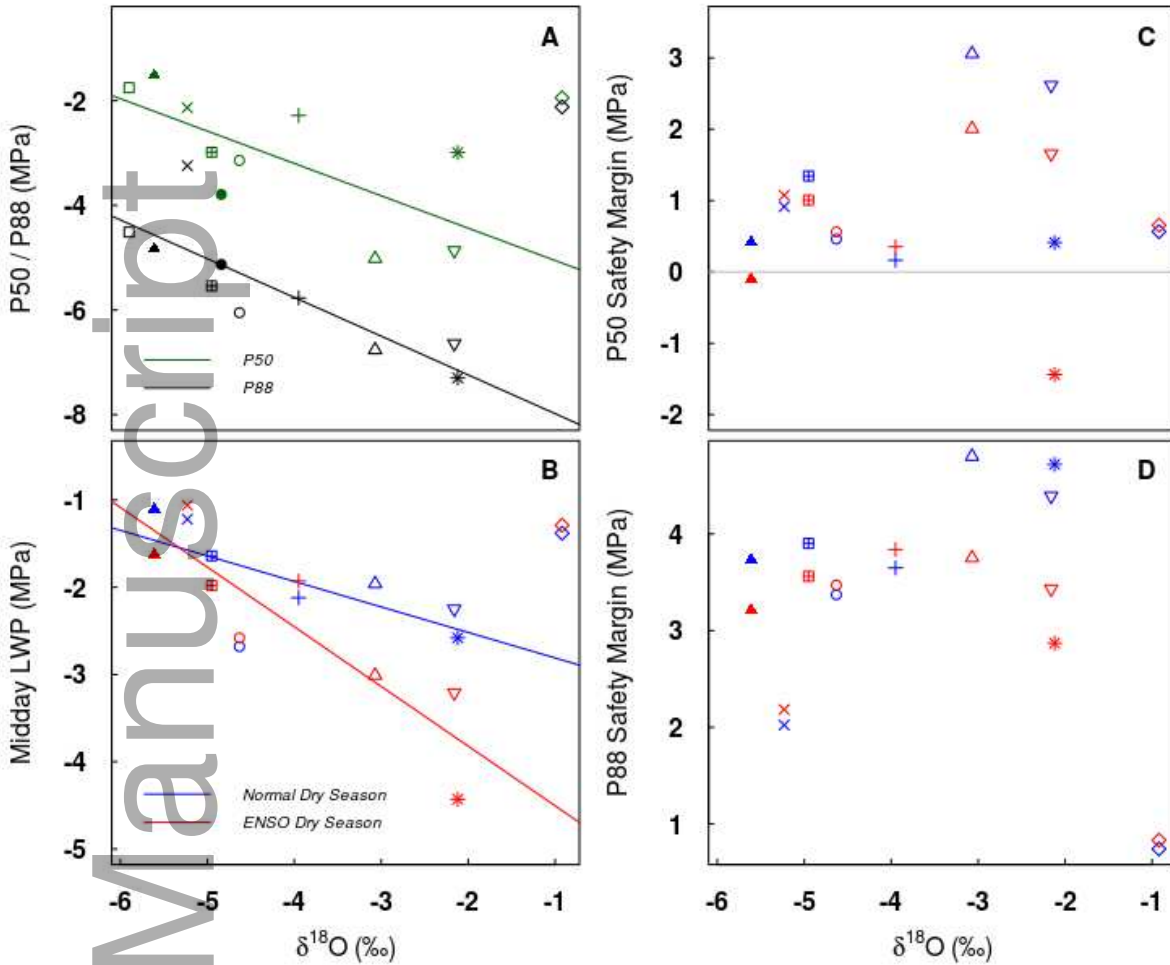


1015
 1016 **Figure 1** A- Dual $\delta^{18}\text{O}$ and $\delta^2\text{H}$ stable isotope plot showing the evaporative enrichment of soil
 1017 water (red to blue circles) and xylem water in trees (green line and symbols) collected during an
 1018 extreme dry season in an ENSO year (2015) at the Tapajós National Forest, Brazil. Tree species
 1019 symbols (green) are shown in the legend. Soil data are color coded by depth (see panel B for
 1020 depths). Black diamonds represent monthly mean meteoric water inputs. The dashed black line
 1021 represents the global meteoric water line (GMWL) and the continuous black line represents the
 1022 estimated regional meteoric water line (RMWL); B- The nonlinear relationship between soil depth
 1023 and $\delta^{18}\text{O}_{\text{soil}}$; horizontal bars represent the standard deviation around the mean for each soil depth.
 1024 The blue rectangle represents the range of mean values from samples below 10 m depth (30 and 60

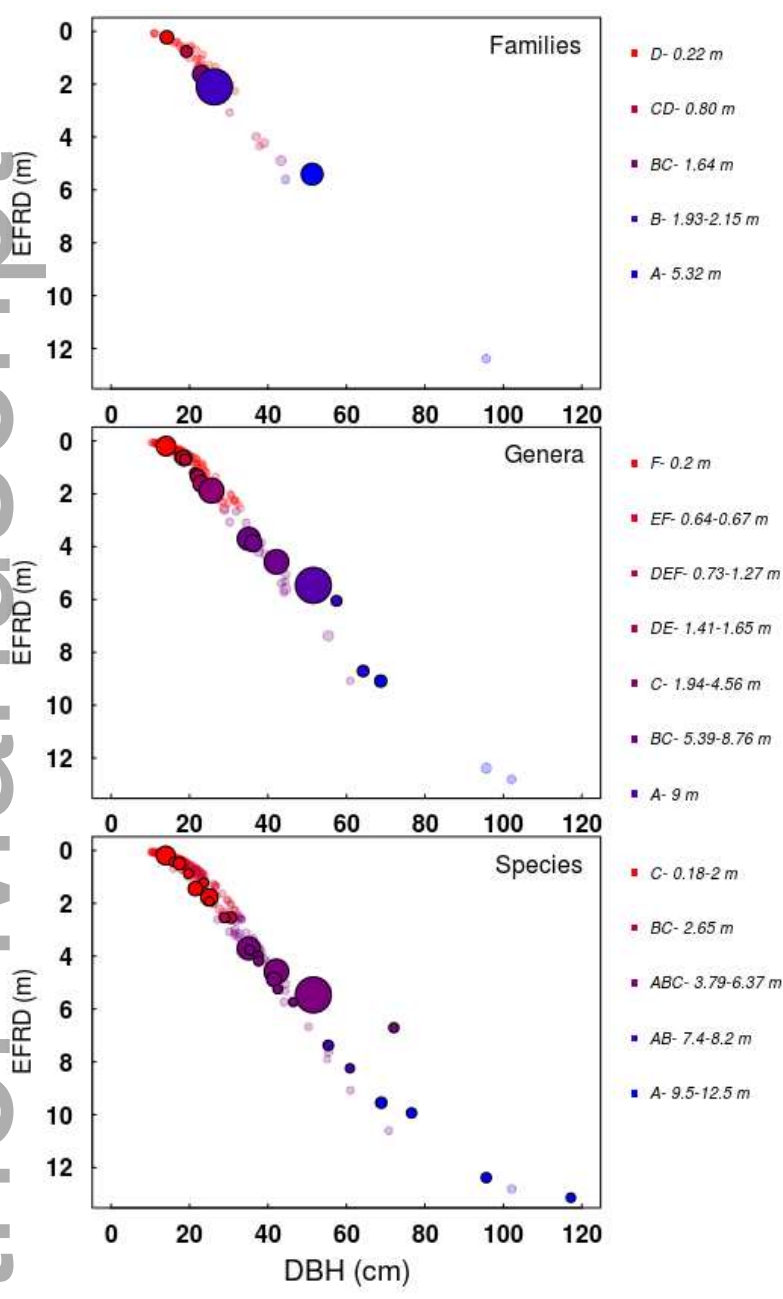
1025 m depths); the line represents the model of Eq. 3 and the dashed red lines show the standard error
 1026 of model parameters estimated with bootstrapping techniques (see Methods); **C-** Average diameter
 1027 at breast height (DBH) of each study species, plotted as a function of xylem water $\delta^{18}\text{O}$ (a proxy of
 1028 effective root depth). The line represents the non-linear regression model of Eq. 3. Error bars
 1029 show one standard deviation computed from data; **D-** Relationships between modelled functional
 1030 root depth (EFRD) and diameter at breast height (DBH) for each individual derived from equation
 1031 5. The continuous black line corresponds to median fitted values from bootstrapping of m, n, k and
 1032 r parameters and used on Eq. 5. The grey shaded area delimited by dashed lines show the 25% and
 1033 75% interquartile interval around the median fitted (see Methods).



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 1039 **Figure 2 A-** Proportion of water use from shallow (< 1 m; red) and deep (> 1 m; blue) derived
 1040 from the isotopic mixing model. Error bars show one standard deviation. Species are sorted from
 1041 left to right with increasing average DBH; species name abbreviations are described in Fig.1; **B-**
 1042 Nonlinear relationships between the mean DBH of each species and the proportion of water uptake
 1043 from deep soil. Species symbols follow the legend in Fig. 1.
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 1046 **Figure 3** Relationships between xylem $\delta^{18}\text{O}$ (used as proxy of functional root depth during an
 1047 extreme dry season in December 2015) and hydraulic traits of the study species. The lines
 1048 represent linear regressions when significant; $p < 0.05$). **A-** P₅₀ (green) and P₈₈ (black); **B-** Midday
 1049 leaf water potential (LWP) measured during peak of a normal dry season in December 2014 (blue)
 1050 and at the peak of an extreme ENSO dry season in December 2015 (red); **C and D-** Hydraulic
 1051 safety margins calculated relative to P₅₀ (HSMP₅₀) and P₈₈ (HSMP₈₈) safety margins calculated
 1052 for normal (blue) and ENSO (red) dry seasons. Species symbols follow the legend in Fig 1.
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Figure 4 Hydrological niche axis given by the average estimated functional root depth (EFRD) as a function of the average diameter at breast height (DBH), a proxy for light-availability. From top to bottom is the average value from both parameters for family, genus, and species level. The circle size is proportional to dominance ($m^2 ha^{-1}$) of the 10% most dominant taxa at each level (see Fig. S6). The lightly shaded circles represent all tree taxa at Tapajós (see Fig. S5). Colors represent the Post-Hoc Tukey comparison results of EFRD to each taxonomic level. The letters on right

1061 legend indicate differences between groups defined by Post-Hoc Tukey statistical differences on
1062 the EFRD range ($p < 0.05$) and the depth range of EFRD where each correspondent group belongs
1063 (see also Fig. S4).

1064 *Electronic Supplementary Information*

1065 **Hydrological niche segregation defines forest structure and drought tolerance** 1066 **strategies in a seasonal Amazon forest.**

1068 Mauro Brum, Matthew A. Vadeboncoeur, Valeriy Ivanov, Heidi Asbjornsen, Scott Saleska,
1069 Luciana F. Alves, Deliane Penha; Jadson D. Dias, Luiz E.O.C. Aragão, Fernanda V. Barros, Paulo
1070 R.L. Bittencourt, Luciano Pereira, & Rafael S. Oliveira

1071

1072 *Species abundance and basal area*

1073 We estimated basal area and local abundance for ten canopy and
1074 subcanopy tree species (Table S1) by extracting individual species data
1075 (diameter at breast height, cm for all live stems ≥ 10 cm, spatial
1076 position, species identity: year 2012) from a long-term forest tree
1077 inventory database ($n = 4$ transects of 1000×10 m located east of the
1078 TNF tower; Pyle et al., 2008, Longo, 2013). To estimate the local
1079 abundance (stems ha^{-1}) and basal area ($\text{m}^2 \text{ha}^{-1}$) for the two understory tree species,
1080 *Amphirrhox longifolia* and *Rinorea pubiflora* (Violaceae), we established five non-contiguous 500
1081 m^2 plots (10×50 m) within an area of 50×1000 m (Transect 1), and measured diameter (mm) at
1082 30 cm above ground and height (m) of all live individuals > 0.30 m tall.

1083

1084 *Water stable isotopes of $\delta^2\text{H}$*

1085 The average $\delta^2\text{H}_{\text{soil} (<1\text{m})}$ was -19‰ (± 8.07), while the average $\delta^2\text{H}_{\text{soil} (>1\text{m})}$ was -30‰
1086 (± 5.91). Well water from 30-60 m depth was the most depleted in heavier isotopes, with $\delta^2\text{H}_{\text{soil}}$
1087 varying from -34.3 to -30.6‰ . The $\delta^2\text{H}_{\text{xylem}}$ ranged from -15.33 to -47.50‰ . Due to the
1088 incomplete overlap between plant and soil $\delta^2\text{H}$ (in contrast to $\delta^{18}\text{O}$; Fig 1, Fig S2), we only used

1089 $\delta^{18}\text{O}$ to determine the water source of each tree species as a mixture of soil-water endmembers at
 1090 different depth intervals.

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1096 **Table S1** Biological and structural attributes of the species studied at Tapajós Forest km-67 LBA
 1097 study area, Brazil. Values are means across a 4-ha survey area of all trees larger than 10 cm.
 1098 Understory species were recorded in five 0.05-ha plots. Species are ordered from deepest rooting
 1099 to shallowest rooting as determined from xylem $\delta^{18}\text{O}$ (Fig. 1). DBH data are from the long-term
 1100 forest tree inventory database (Pyle et al., 2008, Longo, 2013). For two understory species
 1101 *Amphirrhox longifolia* and *Rinorea pubiflora*, too small to be included in the inventory, we report
 1102 the DBH of individuals sampled for stable isotope analysis.

Species	Family	Canopy Position	Basal Area (m ² .ha ⁻¹)	Abundance (ind.ha ⁻¹)	DBH (cm) Mean (SD)
<i>Manilkara elata</i> (Allemão ex. Miq.) Monach.	Sapotaceae	canopy	2.19	10.5	42 ±30
<i>Erismia uncinatum</i> Warm.	Vochysiaceae	canopy	3.64	11.0	51 ±39
<i>Pseudopiptadenia psilostachya</i> (DC.) G.P.Lewis & M.P.Lima	Fabaceae	canopy	0.23	2.0	30 ±21
<i>Endopleura uchi</i> (Huber) Cuatrec.	Humiraceae	canopy	0.33	2.3	37 ±22
<i>Mezilaurus itauba</i> (Meins.)Taub. ex Mez.	Lauraceae	canopy	0.40	1.5	55 ±19
<i>Tachigali chrysophylla</i> (Poepp.) Zarucchi & Herend.	Fabaceae	canopy	1.28	17.8	25 ±19
<i>Chamaecrista xinguensis</i> (Ducke) H.S.Irwin & Barneby	Fabaceae	subcanopy	2.01	15.5	35 ±19

<i>Protium apiculatum</i> Swart.	Burseraceae	subcanopy	0.65	24.3	17 ±6
<i>Coussarea albescens</i> (DC.) Müll.Arg.	Rubiaceae	subcanopy	1.48	92.5	13 ±3
<i>Miconia</i> sp.	Melastomataceae	subcanopy	0.08	2.5	18 ±5
<i>Amphirrhox longifolia</i> (A.St.-Hil.) Spreng	Violaceae	understory	0.35	908	3.5 ±0.92
<i>Rinorea pubiflora</i> (Benth.) Sprague & Sandwith	Violaceae	understory	2.45	3104	2.7 ±0.53

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Table S2 Sample size (n) for each parameter evaluated in this study.

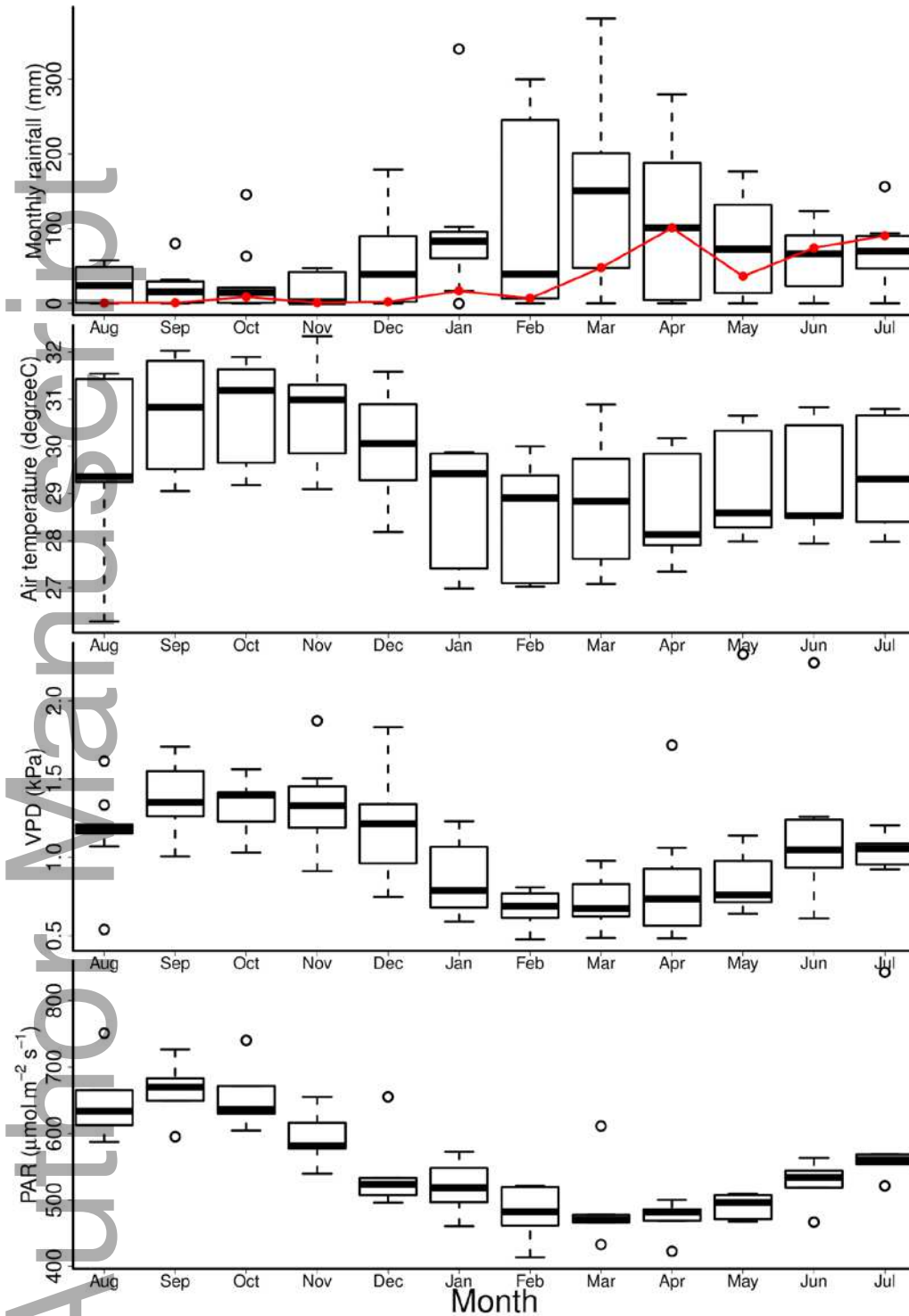
Species	Isotope ($\delta^{18}\text{O}$ and $\delta^2\text{H}$)	Hydraulic Traits (P_{50} and P_{88})	Leaf Water Potential (Ψ)	
			N° individual	N° leaf / individual
<i>Manilkara elata</i>	4	3	-	-
<i>Erismia uncinatum</i>	5	3	2	4-5
<i>Pseudopiptadenia psilostachya</i>	3	-	-	-
<i>Endopleura uchi</i>	1	2	1	5
<i>Mezilaurus itauba</i>	5	2	1	5
<i>Tachigali chrysophylla</i>	4	2	-	-

<i>Chamaecrista xinguensis</i>	5	3	2	4-5
<i>Protium apiculatum</i>	5	3	4	5
<i>Coussarea albescens</i>	5	3	4	4-5
<i>Miconia</i> sp.	5	3	4	5
<i>Amphirrhox longifolia</i>	5	5	4	3
<i>Rinorea pubiflora</i>	5	5	4	3

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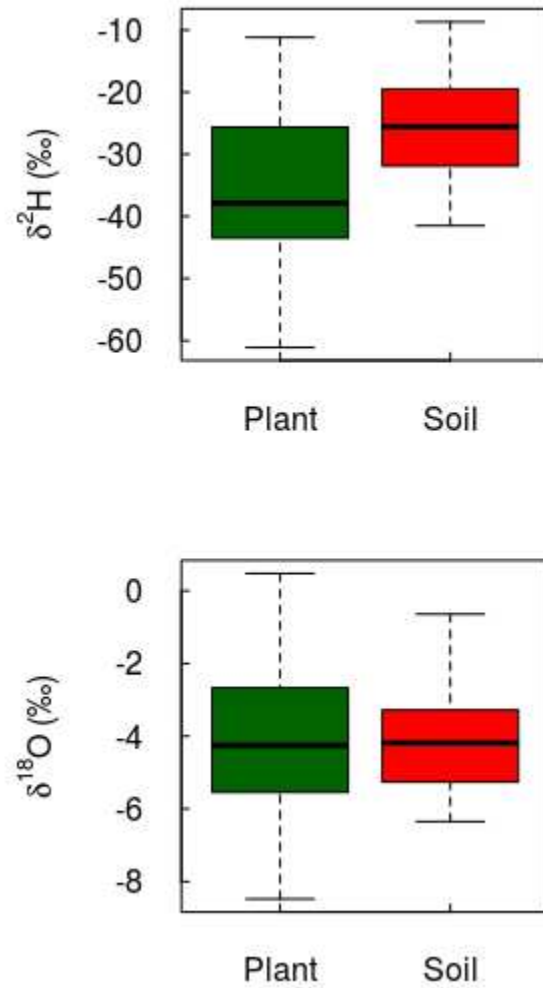
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 1114 **Figure S1** Box plot of monthly rainfall, air temperature, vapor pressure deficit (VPD) and
 1115 photosynthetic active radiation (PAR) from hourly measurements between 2006 to 2017 at Tapajós
 1116 National Forest, Brazil. (see Restrepo-Coupe et al., 2013; 2016). The red line shows the average
 1117 rainfall during 2015.

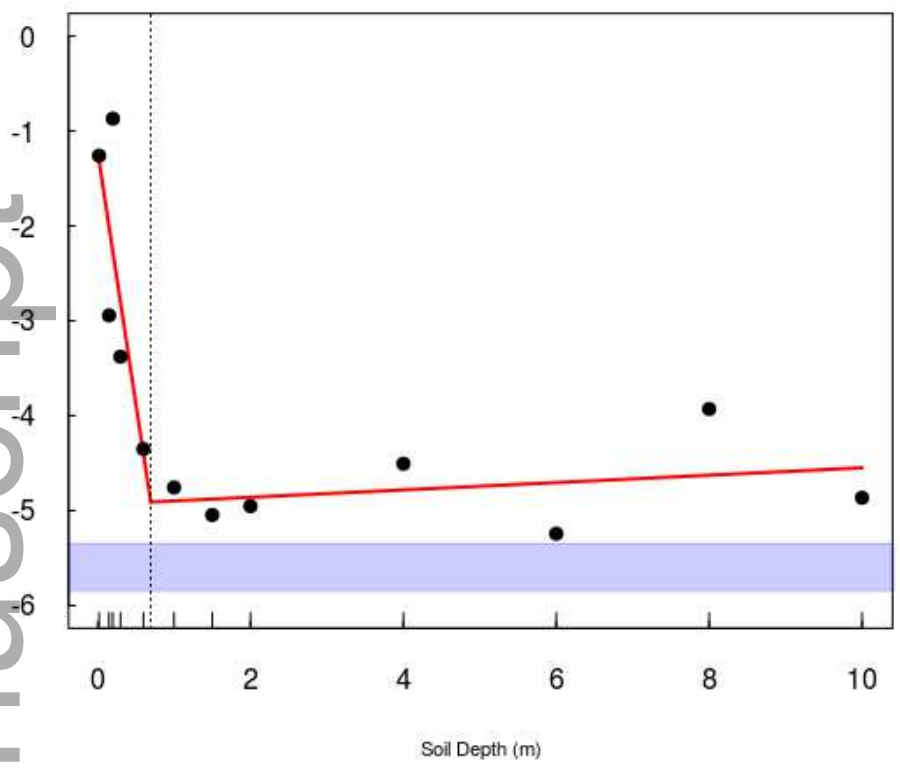
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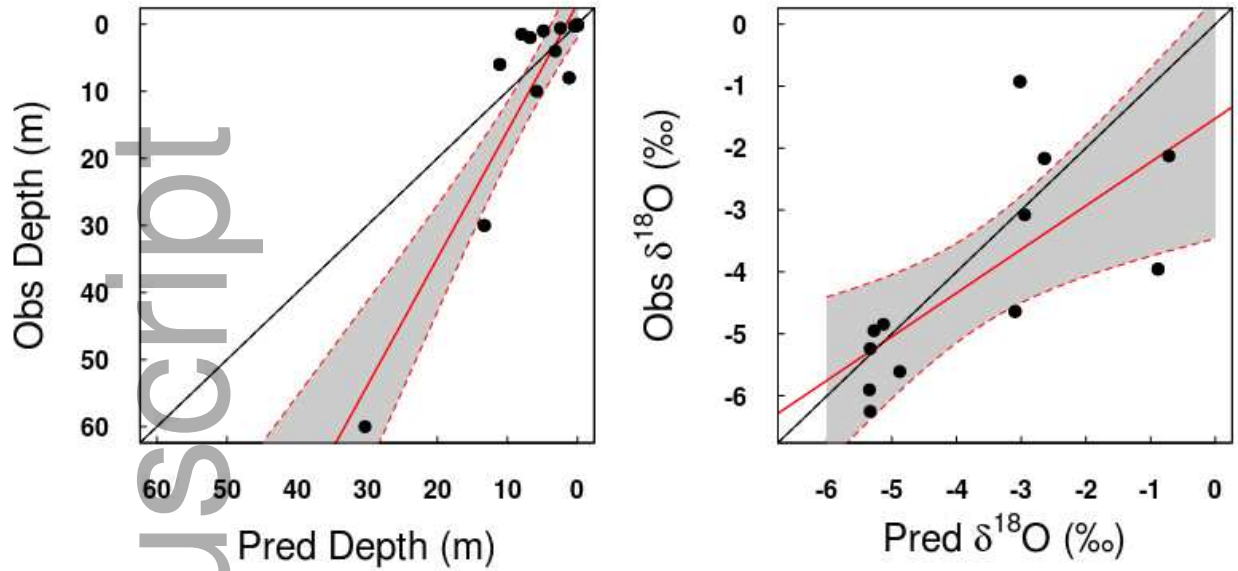
Figure S2 - Boxplot of $\delta^2\text{H}$ and $\delta^{18}\text{O}$ variation in all plants and all soil depths. On average, xylem water $\delta^2\text{H}$ (-34 ‰) was more depleted than soil water (-25 ‰) ($t=-4.83$; $p<0.001$), while $\delta^{18}\text{O}$ was more similar between xylem and soil water sampled ($t=0.28$; $p=0.77$).



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Figure S3 - Regression model with segmented relationship between average $\delta^{18}\text{O}$ and soil depth. The estimated break-point in the soil profile was 0.69 (± 0.17) m depth (vertical dashed line; see results for statistics). This threshold was used to define the depth categories above and below 1 m depth (see methods).

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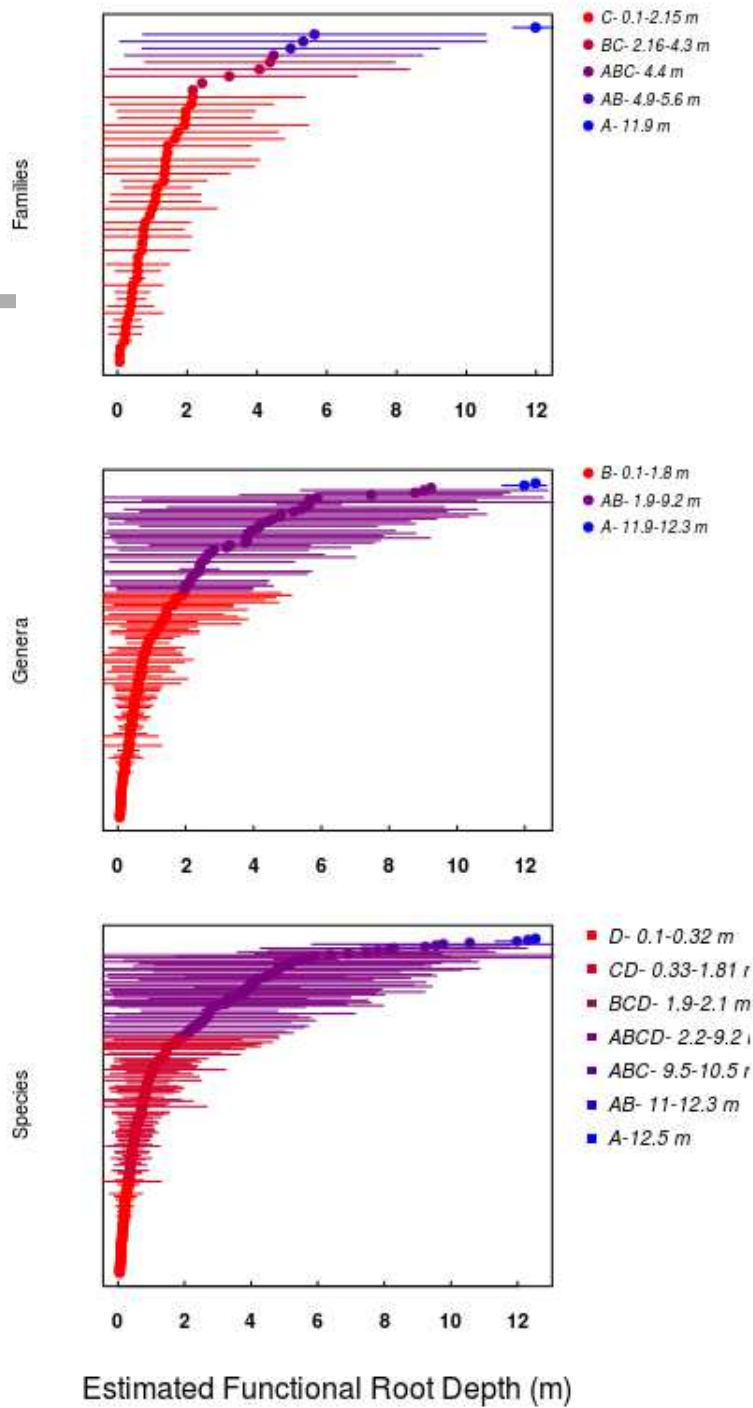
1138 **Figure S4** Relationship between observed and predicted soil depth and $\delta^{18}\text{O}_{\text{xylem}}$ models described
1139 in Eq. 3 (right plot) and Eq. 4 (left plot). The model parameters were derived from a bootstrap
1140 technique and the parameters were chosen by median fitted values from Eq. 5 (see methods). The
1141 black dashed line is the 1:1 line, the red continuous line is the linear model, and red dashed lines
1142 show the limit of the confidence interval (grey area) estimated by the model.

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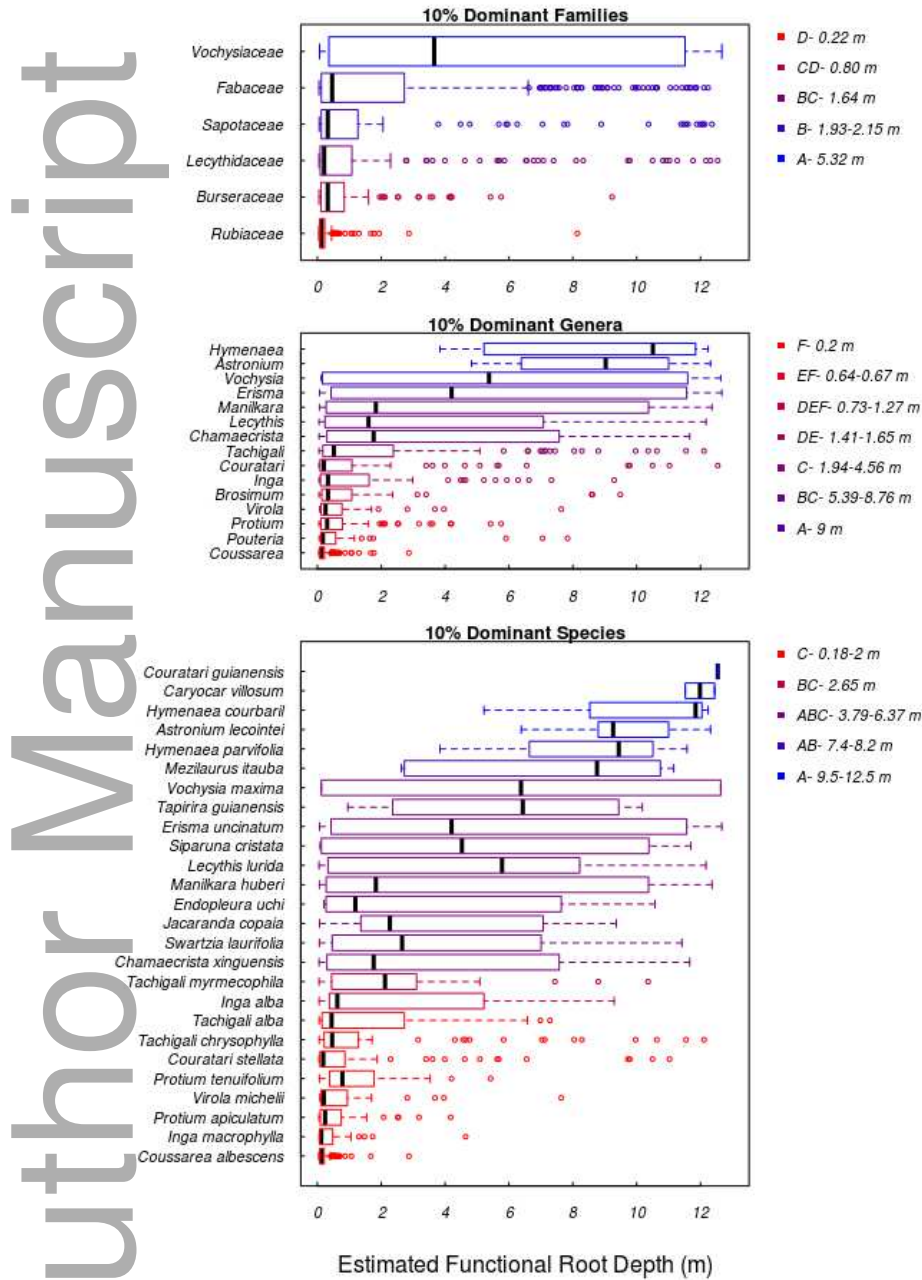
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 1148 **Figure S5** Extrapolated estimated functional root depth (EFRD) for Families, Genera and Species
 1149 (from top to bottom) using equation 5. We used coefficients derived using all soil $\delta^{18}\text{O}$ data (see
 1150 Fig. 3, Fig. S3 and Table S2) to scale across all individuals within the four 1-ha hectare study area
 1151 at Tapajós National Forest, Brazil (see S1). Colors represent the results of Post-Hoc Tukey

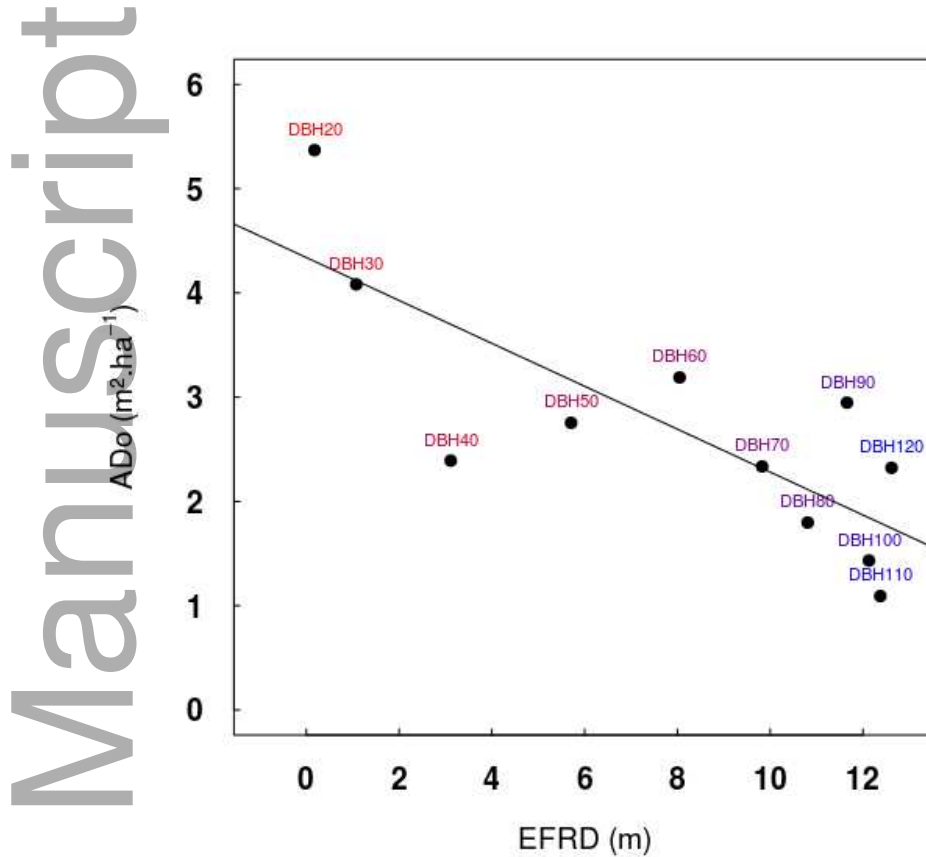
1152 comparison of EFRD at each taxonomic level. The letters on right legend indicate differences
 1153 between groups ($p < 0.05$) and the depth range of EFRD where each groups occurs.



1154
 1155 **Figure S6** Extrapolated estimated functional root depth (EFRD) for the most dominant 10% of
 1156 taxa at each level of organization (Family, Genera and Species), based on equation 5. We used
 1157 coefficients derived using all soil $\delta^{18}\text{O}$ data (see Fig. 3, Fig. S3 and Table S2) to scale up for these
 1158 dominant groups within the 4-ha study area in the Tapajós National Forest, Brazil (data from Pyle
 1159 et al., 2008; update by Longo, 2013; see S1). Colors represent the results of Post-Hoc Tukey

1160 comparison of EFRD for each taxonomic level. The letters in the legend on the right indicate the
1161 grouped taxa ($p < 0.05$) and the depth range of EFRD corresponding to each group. These data are
1162 the highlighted data points in Fig. 4.

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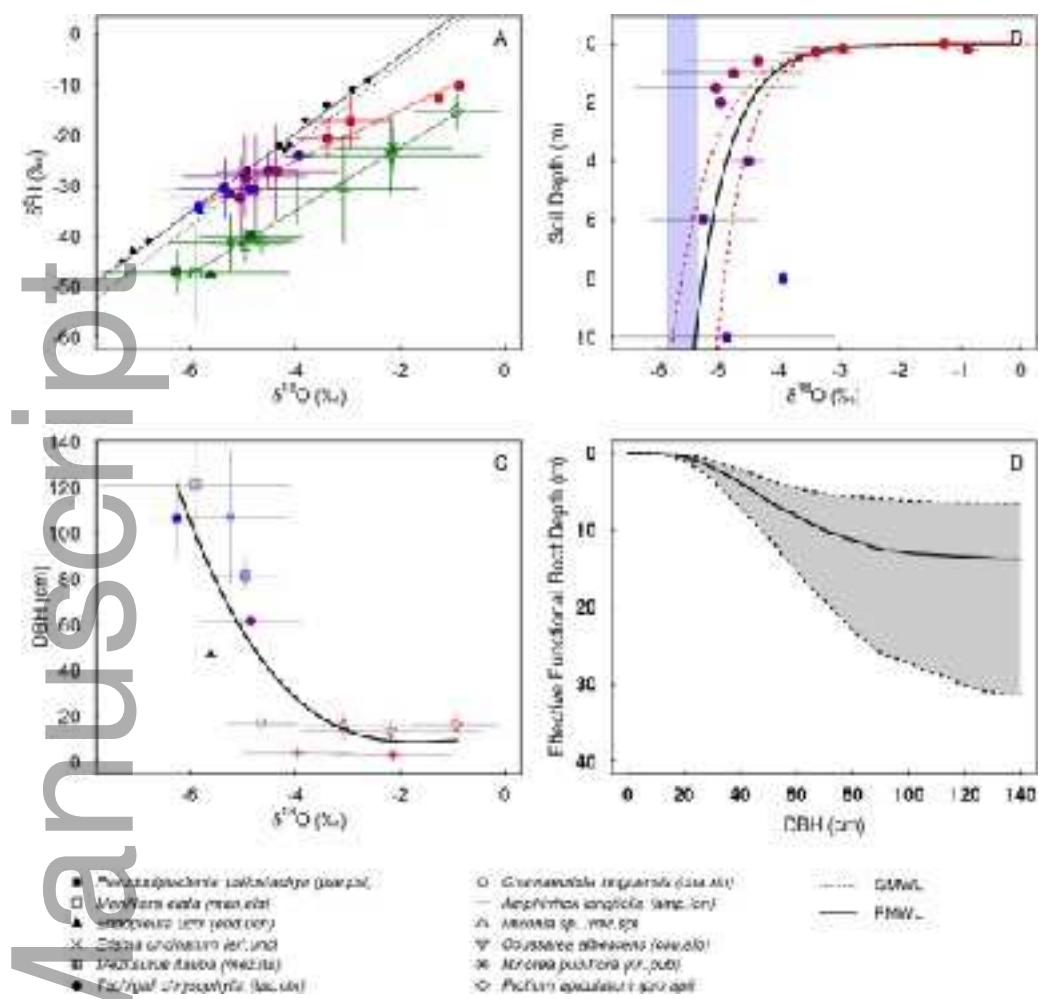
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1166 **Figure S7** Relationship between absolute dominance (ADo, $m^2 \cdot ha^{-1}$) and estimated functional
1167 root depth (EFRD) for each diameter class of trees (labels from 20 to 110; each 10 cm) from
1168 seasonal Amazon Forest - Tapajós National Forest, Brazil. The black line is a linear model
1169 ($r^2=0.63$; $p < 0.001$). Many studies ignore small plants or underrepresent the importance of this
1170 group to forest functioning.

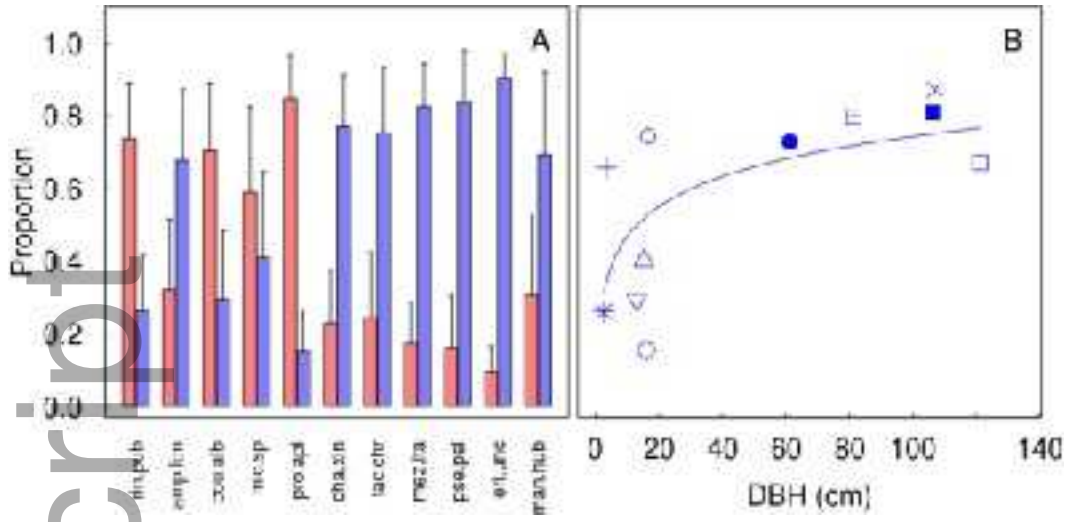
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1172 **Figure S8** Percentage of maximum air discharge (PAD) as a function of the xylem water
1173 potential (Ψ_x) for species described on figure 1. Also shown are percentage loss of conductance
1174 (PLC) data for the two species studied using only this method. Blue lines show Weibull functions

1175 fit using eq. 1. Red circles show the P_{50} . Embolism curves from *E. uncinatum* and *Miconia* sp.
1176 from Pereira et al. 2016.

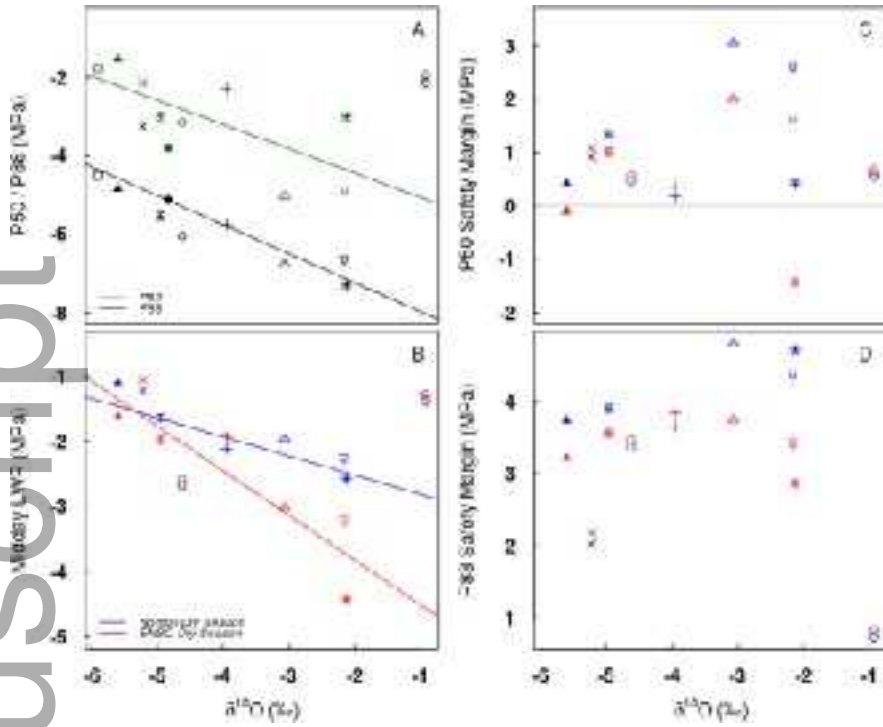
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