CONTRIBUTIONS FROM THE MUSEUM OF PALEONTOLOGY

THE UNIVERSITY OF MICHIGAN

VOL. 29, NO. 7, PP. 135-184

<u>November 30, 1995</u>

MORPHOLOGY OF CARBONIFEROUS AND PERMIAN CRINOIDS

BY

MIKE FOOTE



MUSEUM OF PALEONTOLOGY THE UNIVERSITY OF MICHIGAN ANN ARBOR

CONTRIBUTIONS FROM THE MUSEUM OF PALEONTOLOGY

Philip D. Gingerich, Director

This series of contributions from the Museum of Paleontology is a medium for publication of papers based chiefly on collections in the Museum. When the number of pages issued is sufficient to make a volume, a title page and a table of contents will be sent to libraries on the mailing list, and to individuals on request. A list of the separate issues may also be obtained by request. Correspondence should be directed to the Museum of Paleontology, The University of Michigan, Ann Arbor, Michigan 48109-1079.

VOLS. 2-29. Parts of volumes may be obtained if available. Price lists are available upon inquiry.

MORPHOLOGY OF CARBONIFEROUS AND PERMIAN CRINOIDS

By

MIKE FOOTE¹

Abstract—To allow further investigation of morphological diversification in Paleozoic crinoids, this paper supplements discrete-character data on Ordovician-Devonian crinoids and presents data for Carboniferous and Permian species. In addition to 75 characters used previously, three new characters are incorporated into the data. Analysis complements previous work in suggesting that most of the spectrum of crinoid form was attained early. Morphological diversity did not continue to increase through most of the Paleozoic, even in the face of several phases of taxonomic diversification. This result further supports previous arguments for the importance of basic morphological constraints in crinoid evolution. Shifts in the distribution of morphology largely corresponded with the waning or diversification of particular higher taxa.

INTRODUCTION

In documenting the evolutionary history of major biologic groups, data on morphological disparity are important in complementing other aspects of diversity and in leading to inferences about evolutionary mechanisms (e.g., Ward, 1980; Derstler, 1981, 1982; Saunders and Swan, 1984; Campbell and Marshall, 1987; Runnegar, 1987; Jacobs, 1990; Anstey and Pachut, 1992; Briggs et al., 1992; Foote, 1992b, 1993, 1994a,b, 1995; Sprinkle, 1992; Roy, 1994; Wagner, 1993; Smith, 1994; Wills et al., 1994). Because they have a long history marked by substantial taxonomic richness and morphological variety, a good fossil record, and a skeleton that reflects soft anatomy and function (Lane, 1963b; Brower, 1966, 1988b; Breimer, 1969; Meyer, 1973, 1979, 1983; Breimer and Webster, 1975; Breimer and Lane, 1978; Ubaghs, 1978a; Ausich, 1980, 1983, 1986, 1988; Brett, 1981; Kammer, 1985; Kammer and Ausich, 1987; Broadhead 1988a; Donovan, 1988, 1990; Riddle et al., 1988; Baumiller and Plotnick, 1989; Kendrick, 1992; Baumiller, 1990, 1993), crinoids represent an ideal clade in which to study morphologi-

¹Museum of Paleontology and Department of Geological Sciences, The University of Michigan, Ann Arbor, Michigan 48109-1079. Present address: Department of the Geophysical Sciences, University of Chicago, 5734 S. Ellis Avenue, Chicago, Illinois 60637.

cal diversification. Although crinoids and other echinoderms may be regarded as structurally complex (Moore, 1952), this feature is no prerequisite to the investigation of disparity, which concerns variation among forms regardless of their complexity. Previous work on early and middle Paleozoic Crinoidea suggested that this clade demonstrates early morphological diversification at low taxonomic diversity (e.g., Kolata, 1975; Brower and Veinus, 1974, 1978; Sprinkle, 1982, 1990; Guensburg, 1984; Guensburg and Sprinkle, 1990, 1992; Sprinkle and Guensburg, 1991; Brower, 1992a,b, 1994; Foote, 1994a,b, 1995). Whether or not the evolutionary mechanisms underlying this pattern need be considered "special" (Gould, 1989, 1991, 1993; Briggs and Fortey, 1989; Smith 1990; Briggs et al., 1992; McShea, 1993b; Ridley, 1993; Wills et al., 1994), such early bursts of morphological diversification point toward initially accelerated rates of morphological change, or what is essentially equivalent, subsequently decelerated rates (Paul, 1977, 1979; Sprinkle, 1980, 1983, 1992; Valentine, 1986, 1991, 1992; Campbell and Marshall, 1987; Jacobs, 1990; Anstey and Pachut, 1992; Foote and Gould, 1992; Foote, 1993; Wagner, 1993; McShea, 1993b; Wills et al., 1994; but see Smith, 1990). That the level of disparity attained by crinoids in the Ordovician was not exceeded at any time through the end of the Devonian, despite striking taxonomic diversification, suggests that some basic constraints to form may have been reached early (Ausich, 1988; Foote, 1994a,b, 1995). However, the possibility remains that the apparent limits which bounded the spectrum of crinoid form were transient.

To allow analysis of morphological diversification in the later Paleozoic, this paper will supplement the set of discrete characters used previously (Foote, 1994a,b), and present morphological data for a number of Carboniferous and Permian crinoid species. Some additions and small corrections to data on Ordovician-Devonian crinoids (Foote, 1994a) will also be given. Most of these changes reflect standardization of characters following the incorporation of data on additional species. Analysis of the data will show that some morphological extremes reached before the Carboniferous were eventually surpassed to some extent. However, the evacuation of morphological space that had been occupied before the Carboniferous was at least as extensive as the colonization of new morphospace. Thus, despite renewed taxonomic and morphological diversification, the level of overall morphological diversity was no greater in the Carboniferous and Permian than it had been during the Ordovician through Devonian. Much of the evacuation and colonization of regions in morphospace reflects changes in taxonomic diversity within single higher taxa. Therefore, although homeomorphy in crinoids is common (e.g., Moore and Laudon, 1943; Sprinkle, 1980, 1983; Ausich, 1988; Broadhead, 1988a,b; Kendrick, 1992), there seems to be a fair correspondence between certain large-scale patterns of taxonomic and morphological evolution in this clade.

MATERIALS AND METHODS

Quantification of crinoid form, sampling of species, and measurement of morphological disparity follow the methods described previously (Foote 1994a,b). A brief summary is given here.

Crinoid Morphology.—Because crinoids as a whole show enormous variability in the number and arrangement of skeletal plates, this paper follows previous work in using discrete (mostly binary) characters to quantify form (e.g., Lane, 1963a; Brower, 1973, 1982, 1988a; Macurda, 1974; Webster, 1981; Kammer and Ausich, 1992, 1993, 1994). The characters, chosen to be applicable to the majority of crinoid species, broadly cover four principal regions of the crinoid skeleton: the pelma (14 characters), the dorsal cup (28 characters), the arms, including fixed brachials and interbrachials (28 characters), and the anus and tegmen (8 characters). These 78 characters comprise the same 75 characters used to document morphological diversity in Ordovician-Devonian crinoids, plus three additional characters: 40A, presence of acrocrinid-

CRINOID MORPHOLOGY

TABLE 1—Additional characters used to quantify crinoid form. Only modifications to characters in Foote (1994a: table 1) are presented here. Characters are denoted as binary (B), ordered multistate (O), or unordered multistate (U).

Character	Character description and states
Pelma 1(U)	Pelma: 0, absent or directly attached. 1, multiplated holdfast. 2, column. Note: State 0 incorporates state 3 of Foote (1994a).
Dorsal cup	Note: This work follows Lane (1967) in considering the plates of Codiacrinacea having only two circlets in the dorsal cup to be basals and infrabasals.
33(U)	Shape of cup (sag.): 0, cylinder. 1, cone. 2, bowl. 3, globe. 4, inverted cone. 5, inverted bowl. 6, splayed bowl or cone (as in Uperocrinus). 7, irregular. Note: State 7 corresponds to state 4 of Foote (1994a).
36(U)	Symmetry of cup (trans.): 0, asymmetric. 1, strongly bilateral. 2, triradial. 3, tetraradial. 4, pentaradial with strong bilateral overprint. 5, pentaradial (with or without weak bilateral overprint). Note: In Foote (1994a,b) states 3, 4 and 5 were presented as states 4, 5, and 6, and there was no state 3.
40A(B)	Intercalary plates (as in Acrocrinus): 0, absent. 1, present.
40B(O)	Number of circlets of intercalaries: state = count. Note: Characters 40A and 40B were not listed in Foote (1994a). No species in Foote (1994a) is coded as having intercalary plates.
Arms	
67A(B)	Torted arms (as in Mespilocrinus): 0, absent. 1, present. Note: Character 67A was not listed in Foote (1994a). No species in Foote (1994a) is coded as having torted arms.

type intercalary cup plates; 40B, number of intercalary cup plates; and 67B, presence of torted arms (Table 1). These three characters were not observed to vary among the Ordovician-Devonian species initially studied (Foote, 1994a,b). Characters were coded mainly from illustrated descriptions of well preserved crinoid species in the primary literature. Species in which the arms are not preserved were omitted, but species considered to lack arms were used (Lane, 1967; see discussion below). Columnal species and other form taxa were not used. Morphological data and sources for these data are given in Appendices 1 and 2.

The homologies among crinoid plates are uncertain (e.g., Moore, 1962; Warn, 1975; Kelly, 1982; Broadhead, 1988a,b; Simms, 1993). However, because this research seeks to investigate the overall spectrum of realized forms, it is appropriate to code characters based on the topological arrangement of plates. For example, in a recent reinterpretation of crinoid plate homologies, Simms (1993; see also Kelly, 1982; Broadhead, 1988b) suggests that the arms, which are conventionally regarded as supported by radial plates (which by implication are homologous) in all crinoids (Moore, 1962), are sometimes supported by the radials, sometimes by basals, and sometimes by both on the same individual. As far as understanding how the major groups of crinoids are related, it is essential that we know whether Simms' view, the conventionally call *radials* share a common evolutionary origin, they play essentially the same role structurally in the vast majority of crinoids, and for analysis of overall crinoid form it is reasonable to consider them the same element in the various crinoid taxa (see Foote 1995).

Time Scale and Sampling.—Stratigraphic intervals, based mainly on Harland et al. (1990), were chosen to be fine enough to give resolution of evolutionary sequences, but coarse enough to allow reasonably large sample sizes (Table 2). The Carboniferous and Permian intervals

TABLE 2—Carboniferous-Permian time scale, generic richness, and sample sizes. Totals for generic richness include genera of uncertain affinities. Ages (Ma) and durations (m.y.) from Harland et al. (1990). Corresponding information for Ordovician-Devonian intervals can be found in Foote (1994a,b). Note that most higher taxa during most intervals are sampled in proportion to their known generic diversity; exceptions are discussed in the text.

Stratigraphic interval	Age (base)	Dura- tion	Mono- bathrida	Diplo- bathrida	Dispar- ida	Clad- ida	Flexi- bilia	Total
					Number of	of genera		
Tournaisian (T)	363	13	51	3	9	63	13	141
Visean (V)	350	17	43	3	8	88	15	160
Serpukhovian (Sr)	333	10	15		4	58	4	81
Bashkirian (B)	323	12	13		6	62	6	87
Moscovian (M)	311	8	14		6	103	6	129
Stephanian (St)	303	13	6		4	99	6	117
Permian-1 (P1)	290	21	5		5	57	4	73
Permian-2 (P2)	269	13	11		11	86	12	121
Permian-3 (P3)	256	11	4		1	13	3	21
				Nun	nber of sp	ecies san	npled	
Tournaisian (T)	363	13	30	3	5	24	9	71
Visean (V)	350	17	17	1	4	49	10	81
Serpukhovian (Sr)	333	10	3		0	25	3	31
Bashkirian (B)	323	12	0		0	9	0	9
Moscovian (M)	311	8	0		0	31	2	33
Stephanian (St)	303	13	1		1	31	1	34
Permian-1 (P1)	290	21	1		1	15	3	20
Permian-2 (P2)	269	13	3		4	28	8	43
Permian-3 (P3)	256	11	1		0	3	1	5

used are: (1) Tournaisian, including the North American sequence through the Burlington Limestone and equivalents; (2) Visean, including the North American sequence from the Keokuk and equivalents roughly through the Gasperian; (3) Serpukhovian, including roughly Hombergian through Elvirian strata; (4) Bashkirian, assumed roughly equivalent to Morrowan; (5) Moscovian, assumed roughly equivalent to Atokan and Desmoinesian; (6) Stephanian. assumed roughly equivalent to Missourian and Virgilian; (7) Permian-1, including Asselian, Sakmarian, and Wolfcampian; (8) Permian-2, including Artinskian and Kungurian; and (9) Permian-3, including the remainder of the Permian. The mean interval length for the Carboniferous and Permian is 13.1 m.y. (SD = 3.8 m.y.) (cf. subdivision of pre-Carboniferous, mean = 16.4 m.y., SD = 8.5 m.y. [Foote, 1994a,b]). Although the rich faunas from Timor (e.g., Wanner, 1916, 1924; Springer, 1924) have traditionally been considered as Upper Permian, I have followed recent evidence (e.g., Webster, 1987, 1990; Webster and Jell, 1992) that they may be closer to Artinskian (Permian-2) in age. Peter Holterhoff (personal communication, 1994) has kindly pointed out that, contrary to my assumption, the Wolfcampian-Leonardian boundary may in fact lie above the Sakmarian-Artinskian boundary. However, most (15 of 17) of the North American Wolfcampian species in the data come from the middle Wolfcampian, and therefore should be within the Permian-1 interval. The effect of uncertainties in the stratigraphic placement of Permian species will be discussed below.

Publications describing crinoid species were identified using several bibliographies, especially Webster (1969, 1977, 1986, 1988, 1993), Moore and Teichert (1978), and the Zoological Record, and species were chosen from these publications. In effect, all genera found by the sampling procedure and represented by relatively complete crowns were used. No more than one species per genus was sampled in any time interval. When numerous species were available, an effort was made to sample the most completely preserved representative of the genus. Restricting sampling in this way is meant to keep the study tractable, to emphasize disparity at the large scale of crinoids as a whole, and to reduce the effects of variable species concepts (Foote, 1994a,b). Although discrete characters were coded for the representative species, not the genus as a whole, the method of sampling effectively uses the genus as the basic sampling unit. This is reasonable in light of the observation that crinoid genera seem to be rather coherent units of morphology (Kammer and Ausich, 1992, 1993, 1994; see Allmon, 1992). Previously, there were two instances in which I inadvertently included two species from a single genus within a stratigraphic interval (Foote, 1994a). Although this oversight had little effect on patterns of morphological diversity, these extra species have since been removed from the Ordovician-Devonian data. In addition, eight more Ordovician-Devonian species were coded (Appendix 1). Sample sizes for the nine Ordovician-Devonian intervals are now 5, 34, 31, 19, 42, 27, 43, 42, and 18. Although there is now one more species (Sprinkle and Wahlman, 1994) representing the Lower Ordovician, many more newly discovered species remain to be described (Guensburg and Sprinkle, 1990; Sprinkle and Guensburg, 1991; T. E. Guensburg, personal communication, 1994). Therefore, future measures of disparity for this interval may be higher than current estimates.

The sample of species is large enough and sufficiently representative of the diversity of Paleozoic crinoids that the history of morphological disparity can be reasonably documented, although not without some limitations. On the whole, 50% of all recognized Paleozoic genera are represented by morphological data in this study, and the vast majority of those genera not included are known from relatively incomplete material. Moreover, the major higher taxa of crinoids are generally represented in the morphological sample in proportion to their known diversity (Table 2). However, there are some noteworthy exceptions: (1) The Serpukhovian sample contains no disparids; (2) the Bashkirian sample contains only cladids; and (3) the Moscovian sample contains no monobathrids or disparids. These shortcomings reflect the dearth of relatively complete crowns from intervals in which isolated cups are quite common (e.g., Moore and Plummer, 1938, 1940; Knapp, 1969; Moore and Strimple, 1973). Because morphological diversity reflects differences among and within groups, the absence of some groups, which occupy largely different regions in morphological space (see below), is likely to bias estimates of disparity downward. This bias is explored more elsewhere (Foote 1995), and will be touched upon below.

Morphological Disparity.—Metrics of disparity convey the magnitude of morphological differences among species (Wills et al., 1994). Morphological disparity within a stratigraphic interval is measured here as the mean pairwise character difference between species (normalized by the number of characters compared) (Sneath and Sokal, 1973; Cherry et al., 1982; Foote, 1992, 1994a,b; Wills et al., 1994). Calculation of character differences follows Foote (1994a: Appendix 1). Without a reasonable way to assign character weights differentially, I have adopted the standard practice of giving all characters equal weight (Sneath and Sokal, 1973). Although equal weighting is in a sense as arbitrary as any other scheme of weighting, it was shown previously (Foote, 1994a) that large-scale patterns of morphological diversity such as those documented here are unlikely to reflect the peculiarities of character weighting. This point will be explored further below.

It has been suggested (e.g., Smith, 1994) that the *total* amount of evolutionary change (including reversal), rather than *net* differences among lineages, should be the focus of studies of morphological diversity. When Smith (1994) states that phenetic distance confounds convergence with homology and underestimates the amount of morphological change that has



FIG. 1-Morphological disparity and taxonomic diversity in Paleozoic crinoids. A, Mean pairwise distance between species. Error bars in this and subsequent figures show 1 SE on either side of disparity estimate and are based on bootstrap resampling of data (Efron, 1982; see Foote, 1994b for interpretation of error bars). B, Generic richness (total number of described genera, not the number sampled for this study). Error bars estimated as $\pm \sqrt{\text{number of genera}}$ (see Sepkoski and Raup, 1986; Abbreviations for Ordovician-Devonian intervals: LO, Lower Ordovician; O2, Foote, 1994b). Ordovician-2; O3, Ordovician-3; Ld, Llandoverian; W, Wenlockian; US, Upper Silurian; LD, Lower Devonian; MD, Middle Devonian; UD, Upper Devonian. Abbreviations for Carboniferous and Permian intervals as in Table 2. Xs in Permian-2 and Permian-3 show the values disparity and diversity would have if species from Timor were assigned to Permian-3 rather than Permian-2. Os in Permian-1 and Permian-2 show the values disparity and diversity would have if species from Timor and from the Australian Callytharra Formation were assigned to Permian-1 rather than Permian-2. Despite expansions of disparity in Lower Carboniferous and Permian, maximal disparity in Paleozoic crinoids had been reached long before the time of maximal diversity. It is suggested in the text that the low disparity in the Serpukhovian through Permian-1 may be exaggerated by poor sampling (see Table 2, Fig. 2).

occurred, he seems to imply, incorrectly, that distance-based measures of disparity are intended to measure total character change. Just as taxonomic turnover (analogous to total character change) and standing diversity (analogous to disparity) reflect different aspects of taxonomic diversification, so do total change and net change reflect different aspects of morphological diversification. Neither is more fundamental than the other; they are complementary. For example, the discordance between abundant total character change (documented, even if not



FIG. 2—Total variance (sum of univariate variances; Van Valen 1974) of first twenty principalcoordinates, with average generic morphology ranging through all stratigraphic intervals between first and last occurrence of a genus. Curve shows much less variable disparity than in raw data. In particular, comparison with Figure 1 suggests that low disparity in the raw data for the Serpukhovian through Permian-1 may be exaggerated by poor sampling (see Table 2). Sum of first twenty eigenvalues is equal to 86% of trace of distance matrix.



FIG. 3—Morphological disparity with armless codiacrinacean species omitted. Note drop in disparity, relative to Figure 1, in Permian-2 and Permian-3.

fully quantitatively, by over a century of studies revealing abundant morphological evolution and the origins of hundreds of new genera over the Paleozoic) and limited net change (supported by the failure of overall disparity to increase over most of the Paleozoic) provides one line of evidence for the importance of morphological constraints in crinoid evolution (Ausich, 1988; Lee 1992; Foote, 1994a,b, 1995). Insistence on the exclusive value of one aspect of diversity or diversification can hinder the insight that comes from considering many perspectives (Gould, 1991; Hickman, 1993a,b).

ANALYSES AND RESULTS

Diversity and Disparity.—To compare disparity with taxonomic diversity, stratigraphic ranges of crinoid genera were compiled, based primarily on the bibliographies of Webster (1969, 1977, 1986, 1988, 1993). The compilation is described in more detail elsewhere (Foote, 1994b). Figure 1 compares generic richness with morphological disparity over the Paleozoic. The temporal patterns for the Ordovician-Devonian have not been substantially altered by the minor changes made to the data for these intervals, and therefore will not be discussed at length here (see Foote, 1994a,b). Extensive taxonomic diversification in the Lower Carboniferous corresponds with an increase in morphological diversity above the level of the Upper Devonian, but Carboniferous diversification did not result in the high level of disparity that had been achieved in the mid Ordovician. Low disparity immediately after the Visean is perhaps to be expected, following the final extinction of diplobathrid camerates. However, it is suggested below that this low level of disparity is exaggerated by biased sampling. Morphological diversity appears to increase between the Upper Carboniferous and the Upper Permian.

By allowing genera that were extant but not sampled for a stratigraphic interval to be represented in the morphological data for that interval, it is possible to take some steps toward compensating for poor sampling in the Serpukhovian through Moscovian. All species were first ordinated using principal coordinates (Gower, 1966). Treating these coordinates as continuous variables, the average form of each genus (mean form of all species in the genus) was computed, and a generic morphology was allowed to range through intervals during which the genus was extant but not sampled (see Foote, 1991, 1993, 1994b, 1995). The result of this procedure is a much flatter morphological diversity curve (Fig. 2). This result may be partly a taxonomic artifact, if unrelated homeomorphs separated by great spans of time are assigned to the same genus. However, it at least suggests that morphological disparity in the Serpukhovian through Permian-1, while it was probably higher than shown in Figure 1, was not higher than in the Tournaisian, Visean, or Permian-2.

The maximal level of morphological diversity after the Devonian is shown by the Permian-2 sample. This interval includes the diverse faunas from Timor (e.g., Wanner, 1916, 1924), which contain many forms one might be tempted to regard as unusual or "bizarre" (e.g., Springer, 1926: p. 93). Of course, there is no necessary correspondence between what we perceive as strange and what is morphologically extreme in some quantitative scheme. Nonetheless, it is worth asking to what extent the high disparity in the Permian-2 sample reflects only the 23 Timor species. If these species were assigned to Permian-3 rather than Permian-2, there would be little difference in the estimates of disparity for either of these intervals (Fig. 1). This stands in contrast to the blastoid record, where estimates of disparity in the Permian rest heavily on the Timor faunas (Foote, 1991, 1993). At the same time, it is possible that the Timor species, as well as the three species from Callytharra Formation of Australia, should be assigned a Permian-1 age (P. Holterhoff, personal communication, 1994). Accepting such an assignment would yield a substantially higher estimate of Early Permian morphological disparity (Fig. 1), implying a more rapid diversification of morphology in the Permian.

This study includes armless forms of the cyathocrine cladid superfamily Codiacrinacea. In contrast to many armless disparids that appear to represent early growth stages of arm-bearing adults (e.g., Koenig, 1965; Lane, 1967; Lane and Sevastopulo, 1981, 1982a,b, 1985), the armless forms included here have not been traced to brachiate adult forms (Lane, 1967). Nevertheless, the possibility remains (Broadhead, 1988a) that at least some of these forms are juvenile stages. The effect of including armless codiacrinaceans can be assessed if we take the extreme position that all abrachiate species should be omitted as juvenile stages. Doing so

yields lower disparity in Permian-2 and especially Permian-3 (Fig. 3). Therefore, part of the temporal pattern of disparity is rather sensitive to the status of these unusual forms.

Morphological Distribution of Higher Taxa.—Based on principal-coordinates analysis, the major higher taxa of crinoids occupy reasonably distinct regions in morphological space (Fig. 4) (see also Lane, 1963b; Ausich, 1980). As found previously (Foote, 1994a), the Flexibilia overlap with the other higher taxa. Within Monobathrida (Fig. 5), Compsocrinina and Glyptocrinina are somewhat distinct; this is implicit in Ubaghs' (1978b) statement that no transitional forms are known between these taxa. Within Compsocrinina, Permo-Carboniferous Hexacrinitacea are reasonably distinct from Carpocrinacea and Periechocrinacea along axis 2. Within Permo-Carboniferous Hexacrinitacea (Fig. 6), Acrocrinidae, Camptocrininae, and Dichocrininae are not mutually distinct, reflecting the unity of structure of these taxa (Wachsmuth and Springer, 1897; Springer, 1926; Moore and Strimple, 1969; Broadhead, 1981, 1985). Characters were not chosen specifically to reflect structural complexity. Because the intercalary plates that make acrocrinids appear complex account for only two out of 78 characters, structurally simple dichocrinids (Broadhead, 1981, 1985) and structurally complex acrocrinids (Moore and Strimple, 1969) do not differ greatly in form as quantified here. Within Flexibilia (Fig. 7), Permo-Carboniferous Lecanocrinacea are separate from other taxa along the second axis. Otherwise, the flexible subtaxa overlap extensively (Springer, 1920: p. As for pre-Carboniferous flexibles, Permo-Carboniferous Sagenocrinida are more 76). dispersed than Taxocrinida. This agrees with Springer's (1920: p. 116) assessment of the Taxocrinida as rather homogeneous. In addition, Lecanocrinacea are quite variable morphologically. Within Cladida (Fig. 8), Cyathocrinina and Poteriocrinina are reasonably distinct, as are Cyathocrinitacea and Codiacrinacea within Cyathocrinina. The overlap between Dendrocrinina and Poteriocrinina is perhaps to be expected, since these two suborders share many defining characters (Kammer and Ausich 1992). Within post-Devonian Poteriocrinina (Fig. 9), the various superfamilies overlap extensively, although Scytalocrinacea seem coherent along axes 1 and 3. Disparida in the Carboniferous and Permian show less diversity and morphological dispersion than they had before the Carboniferous (Fig. 10).

Changes in Morphospace Occupation.-That morphological disparity did not increase between the Ordovician and Permian may suggest that major constraints on crinoid form were reached early (Ausich, 1988; Foote, 1994a,b). However, it is possible for morphological distributions to evolve without a corresponding change in disparity. A single comparison between pre-Carboniferous and post-Devonian distributions gives poor resolution, but it does show some of the major features of crinoid evolution. Much of the morphological space occupied before the Carboniferous by Monobathrida (primarily Glyptocrinina) (Fig. 5), Dendrocrinina (Fig. 8), Disparida (primarily Cincinnaticrinacea and Homocrinacea) (Fig. 10), and Diplobathrida (Fig. 11), had been evacuated by the Carboniferous. The main areas of post-Devonian expansion seen in Figures 4-10 correspond with the diversification of the poteriocrine cladids (coinciding with environmental expansion; Lane 1971), and the armless codiacrinaceans (Fig. 8). In addition, on higher principal-coordinate axes there are new extremes colonized by the monobathrid superfamily Hexacrinitacea (especially Acrocrinidae, Camptocrininae, and Dichocrininae) and a few Carboniferous Flexibilia (Fig. 12). Thus, much of the occupation of new morphological space coincides with the diversification of particular higher taxa (Lane, 1967; Moore and Strimple, 1969; Broadhead, 1981, 1985).

Apparent changes in morphological distribution illustrate a limitation of empirical morphospaces (Raup, 1966, 1967; Gould, 1991; McGhee, 1991). A large group of very similar species has the potential to "pull" the axes because of its sheer numbers. To investigate this possibility, a culling procedure (described in detail in Foote, 1995) was used to thin down the samples so that each genus would be separated from all other contemporaneous genera by a morphological distance at least as great as the median distance among congeneric species. Recomputed principal coordinates based on the culled data show that many pre-Carboniferous Dendrocrinina, Glyptocrinina, Disparida, Rhodocrinitacea, and some Flexibilia occupy regions



FIG. 4—Distribution of higher taxa along first three principal-coordinate axes. A, Ordovician-Devoniar; B, Carboniferous and Permian. Key: C, Cladida; D, Diplobathrida; F, Flexibilia; H, Hybocrinida; M, Monobathrida; X, Disparida. Compare with Figures 5-11. Areas in morphological space essentially evacuated before the Carboniferous and Permian include those corresponding to dendrocrine cladids, disparids (especially Homocrinacea and Cincinnaticrinacea), diplobathrids (especially Rhodocrinitacea), and certain monobathrids (Glyptocrinina). Areas in morphospace newly colonized in the Carboniferous and Permian include those corresponding to poteriocrine cladids (with low values of principal coordinate 1), and codiacrinacean cladids and some flexibles (with high values of principal coordinate 3). Sum of first three eigenvalues is equal to 38% of trace of distance matrix.



Monobathrida

FIG. 5—Species of Monobathrida along first three principal-coordinate axes. A, Ordovician-Devonian;
B, Carboniferous and Permian. Key: upper case letters, Compsocrinina (C, Carpocrinacea; H, Hexacrinitacea; P, Periechocrinacea; X, Xenocrinacea); numerals, Glyptocrinina (1, Eucalyptocrinitacea; 2, Glyptocrinacea; 3, Melocrinitacea; 4, Patelliocrinacea; 5, Platycrinitacea); a, *Stipatocrinus*;
b, *Colpodecrinus*. Note that Compsocrinina and Glyptocrinina occupy somewhat different morphological regions. Comparing pre-Carboniferous with Permo-Carboniferous, field of Compsocrinina is expanded by diversification of Hexacrinitacea, while field of Glyptocrinina is reduced by loss of all superfamilies but Platycrinitacea.



Carboniferous-Permian Hexacrinitacea

FIG. 6—Species of Carboniferous and Permian Hexacrinitacea along first three principal-coordinate axes. Key: A, Acrocrinidae; C, Camptocrininae (Dichocrinidae); D, Dichocrininae (Dichocrinidae); P, *Pterotocrinus* (Talarocrininae, Dichocrinidae). Families and subfamilies do not seem to occupy distinct morphological fields.

in morphological space that were not occupied in the Carboniferous and Permian (Figs. 13, 14). Newly colonized morphospace still reflects mainly the Codiacrinacea and Hexacrinitacea as before, but Poteriocrinina to a lesser extent. The expansion of morphospace corresponding with the poteriocrine radiation apparently reflects, at least in part, a pull on the principal-coordinate axes because of the sheer number of poteriocrine species, rather than their extreme morphology.

Character Contribution to Principal-Coordinate Axes.—Because the principal-coordinate axes do not represent all the variation in the original morphological data, it is worth asking to what extent the characters themselves are represented by the axes. As an analog to loadings in Rmode principal-components analysis, the strength of association between the discrete characters and the principal coordinates was calculated. For binary and ordered multistate characters, the gamma coefficient (Siegel and Castellan, 1988: p. 291) was used, and for unordered multistate characters the Cramér coefficient was used (Siegel and Castellan, 1988: p. 225), with the principal coordinates arbitrarily divided into four equal intervals. Although the apparent statistical significance of these associations is exaggerated by the nonindependence of species (Raup and Gould, 1974; Felsenstein, 1985), the coefficients still allow the relative strengths of association to be assessed (Table 3). About half the characters contribute substantially to one or more of the first three principal-coordinate axes, suggesting that the principal-coordinate space represents a fair proportion of the information in the larger set of characters. However, for the rest of the characters, information is spread out over many principal coordinates rather than being concentrated. It was previously shown (Foote, 1994b) that the temporal pattern of variability in the first twenty principal coordinates reflects that in the discrete characters rather well. The same is found for the revised data (result not presented here).

On the first principal-coordinate axis, higher coordinates are associated with the following features: meric stem, homeomorphic stem, fewer radials, closed radials (but radials more likely to be open by basals if open), monocyclic cup (but basals more likely to be unequal if cup di-



Flexibilia

FIG. 7—Species of Flexibilia along first three principal-coordinate axes. A, Ordovician-Devonian; B, Carboniferous and Permian. Key: numerals, Sagenocrinida (1, Icthyocrinacea; 2, Lecanocrinacea; 3, Sagenocrinitacea); T, Taxocrinida. Sagenocrinida and Taxocrinida overlap extensively. Sagenocrinids are more dispersed than taxocrinids, and Lecanocrinacea is especially variable in form.

cyclic), radial prongs, unequal radials, compound radials, fewer plates in basalmost circlet, basalmost plates unequal, fewer anals in cup, higher cup, convex base, cup diminutive relative to stem, cup cylindrical or globose, cup bilateral or pentaradial without strong bilateral overprint, lack of arms, more arms when present, multiple arms on a single radial, shorter arms, unbranched arms (but greater degree of branching and more likely to be heterotomous when branched), presence of fixed brachials, lack of pinnules, symmetric uniserial brachials,



Cladida

FIG. 8—Species of Cladida along first three principal-coordinate axes. A, Ordovician-Devonian; B, Carboniferous and Permian. Key: D, Dendrocrinina; P, Poteriocrinina; numerals, Cyathocrinina (1, Codiacrinacea; 2, Cyathocrinitacea; 3, Gasterocomacea; 4, Thalamocrinidae). Comparing A and B, note contraction in field occupied by Dendrocrinina and expansion of fields occupied by Poteriocrinina and Codiacrinacea.



FIG. 9-Species of Carboniferous and Permian Poteriocrinina along first three principal-coordinate axes. A, data plotted at same scale as Figures 4-8 and 10-11; B, data plotted at finer scale. Key: G, Agassizocrinacea; A, Apographiocrinacea; C, Cromyocrinacea; X, Calceolispongiacea; D, Decadocrinacea; E, Erisocrinacea; H, Hydreionocrincaea; L, Lophocrinacea; I, Pirasocrinacea; P, Poteriocrinitacea; R, Rhenocrinacea; S, Scytalocrinacea; T, Texacrinacea; Z, Zeacrinitacea; ?, superfamily uncertain. Note that Scytalocrinacea tend to cluster together, but on the whole, superfamilies overlap extensively.



Disparida

FIG. 10—Species of Disparida along first three principal-coordinate axes. A, Ordovician-Devonian; B, Carboniferous and Permian. Key: numerals, Ordovician-Devonian Allagecrinacea (1, Acolocrinidae; 2, Agostocrinus; 3, Anamesocrinidae); A, Allagecrinidae (Allagecrinacea); T, Catillocrinidae (Allagecrinacea); B, Belemnocrinidae (Belemnocrinacea); S, Synbathocrinidae (Belemnocrinacea); C, Calceocrinidae (Calceocrinacea); N, Cincinnaticrinacea; H, Homocrinidae (Homocrinacea); R, Ramseyocrinus; other symbols, Myelodactylacea (*, Eustenocrinidae; O, Iocrinidae; +, Myelodactylidae). Disparids seem more dispersed morphologically before the Carboniferous.



Diplobathrida

FIG. 11—Species of Diplobathrida along first three principal-coordinate axes. A, Ordovician-Devonian;
B, Carboniferous and Permian. Key: D, Dimerocrinitacea; N, Nyctocrinacea; R, Rhodocrinitacea;
Z, Zygodiplobathrina. Comparing A and B, note great contraction in morphological field of Diplobathrida, corresponding with loss of most of the subtaxa.



Stratigraphic Interval

152

M. FOOTE

anus through dorsal cup, anal tube absent (but more likely to occupy posterior position and lack ridges when present), and tegmen consisting of a few plates. In addition to the obvious complements to the foregoing, lower coordinates on the first axis correspond to pentagonal and pentastellate stem, bowl-shaped cup, cup pentaradial with strong bilateral overprint, and multiplated tegmen. Axes 2 and 3 show similar associations with some of the characters (Table 3). In addition, species with elongate stem ossicles tend to have low coordinates on the third axis. Much of the pattern of association between characters and principal coordinates makes sense in terms of the positions of higher taxa along the axes (cf. Figs. 4-11, Appendix 1).

Effect of Unknown Characters.—In calculating a pairwise dissimilarity between species, characters that are logically applicable but whose states are unknown in either or both of the species make no contribution to the distance. The validity of this convention rests upon the assumption that, as far as morphological disparity is concerned, the missing characters are drawn from the same underlying distribution as the known characters. A disproportionately large part of the missing data reflects characters of the stem (Table 4), whose temporal pattern of disparity was shown not to match that of overall morphology (Donovan, 1986, 1989a,b; Foote, 1994b). Because nearly complete stems become far less common after the Devonian (Table 4), it is conceivable that a bias could be imparted to the pattern of disparity over the entire Paleozoic. (Tegmina are also slightly less complete on average after the Devonian [see Moore and Laudon, 1943: p. 35; Strimple, 1973], but this difference is not as great as for stems [Table 4].) The history of disparity barely changes if characters of the stem and tegmen are omitted (leaving an average of only 1-2 missing data do not appear greatly to bias the patterns of disparity documented here (Foote 1995).

Variation in Character Weighting.—As discussed above, characters were weighted equally in the calculation of morphological distances. One might reasonably ask how sensitive the temporal pattern of disparity is to the convention of equal character weighting. To explore this question, a set of random character weights was assigned and the disparity curve recomputed. This procedure was repeated 100 times to gain an idea of the range of patterns that would result from arbitrary character weighting. In order to yield a mean weight of unity, random weights were sampled from the uniform distribution on (0,2). The curves differ in height (Fig. 16), but show roughly the same pattern amongst themselves and compared to the pattern derived from equal weighting (Fig. 1). The major differences are that, with random weights, disparity is not always greatest in the mid Ordovician, and estimated disparity in Permian-3 sometimes

FIG. 12-Ordination of Paleozoic crinoids along first twenty principal-coordinate axes. Each point represents a species. Stratigraphic intervals 1-18 correspond with Lower Ordovician through Permian-3 (see Table 2 and Fig. 1). For clarity, only every third interval is labeled. Vertical scale is not the same in all panels. Some exceedance of Carboniferous and Permian species beyond Ordovician-Devonian extremes, and of Ordovician-Devonian species beyond Carboniferous and Permian extremes can be noted, as follows. Extremely low values on axis 1 are mainly Carboniferous and Permian Poteriocrinina. Extremely high values on axis 2 are mainly Ordovician-Devonian Extremely low values on axis 3 are mainly Ordovician-Devonian Disparida. Dendrocrinina. Extremely high values on axis 3 and extremely low values on axis 4 are mainly Carboniferous and Permian Codiacrinacea. Extremely high values on axis 5 are mainly Carboniferous Flexibilia. Extremely low values on axis 6 are mainly Ordovician-Devonian Diplobathrida (primarily Extremely high values on axis 6 are mainly Carboniferous and Permian Rhodocrinitacea). Hexacrinitacea. Extremely low values on axis 7 are mainly Ordovician-Devonian Disparida and Diplobathrida (all Rhodocrinitacea). Extremely high values on axis 8 are mainly Calceocrinacea. Extreme values on remaining axes are much less homogeneous taxonomically.

TABLE 3—Patterns of association between principal coordinates and morphological characters. See table 1 and Foote (1994a) for description of characters. Binary, ordered multistate, and unordered multistate characters are denoted by B, O, and U, respectively. Symbols show relative strength of association (\blacksquare for p < 0.05, \blacksquare for p < 0.01, and \blacksquare for p < 0.001, assuming independence of species.) For binary and ordered characters, strength of association is evaluated with gamma coefficient; sign gives direction of correlation. For unordered characters (for which association lacks direction), strength of association is evaluated with Cramér coefficient. Very few of the coefficients of association for higher axes are significant; therefore, these are not presented.

		Princi	pal coordina	te axis
	Character	Axis 1	Axis 2	Axis 3
Pelma				
1(U)	Form of pelma			
2(B)	Xenomorphy			
3(B)	Heteromorphy			
4(B)	Coiled column			
5(B)	Meric columnals	+		
6(U)	Shape of columnals			
7(U)	Shape of lumen			
8(B)	Relative height of columnals			+
9(B)	Columnal articulation			
10(B)	Cirri			
11(B)	Regular cirral arrangement			
12(B)	Form of cirral arrangement			
13(B)	Specialized distal structure			
14(U)	Form of distal structure			
Dorsal cu	D			
15(B)	Regular plating			
16(O)	Number of radials	-	-1	
17(B)	Open radial circlet		-	
18(Ú)	Nature of opening			
19(B)	Number of circlets	-	-1	
20(B)	Radial prongs	+	+	
21(B)	Relative development of radials	+	+	
22(B)	Compound radials	+	+	
23(O)	Number of compound radials			
24(O)	Relative size of radial and basal-most circlets			
25(O)	Number of basal-most plates			
26(B)	Relative development of basal-most plates	+	+	+
27(B)	Open intermediate circlet			
28(U)	Nature of opening			
29(0)	Number of plates in intermediate circlet			
30(B)	Relative development of plates in intermediate circlet	+		
31(O)	Relative size of intermediate and radial circlets			
32(0)	Number of anal plates in cup			
33(U)	Cup shape (sag.)			
34(0)	Relative cup height			
35(U)	Cup shape (trans.)			
36(U)	Cup symmetry (trans.)			
37(B)	Concave base	-		
38(B)	Relative diameter of cup and stem			
39(B)	Median ray ridges			
40(B)	Stellate ridges			
40A(B)	Intercalary plates			
40B(O)	Number of circlets of intercalaries			

154

CRINOID MORPHOLOGY

TABLE 3 (continued)

		Princi	pal coordinat	te axis
	Character	Axis 1	Axis 2	Axis 3
Arms				_
41(B)	Arms present			-
42(0)	Number of distinct arms	+		
43(0)	Number of arms on single radial	+		
44(0)	Relative development of arms			
45(0)	Separation of arms	. 🛥		
46(B)	Arm fusion between rays	+	+	
47(B)	Arm branching			
48(U)	Degree of branching	+		
49(B)	Isotomy or neterotomy	+		
50(U)	Form of neterolomy			
51(B)	Detailed process			
52(B)	Patenoid process			
53(B)	Bracman asymmetry Belative breachiel height			
54(U)	Arm fusion within rays			
55(B)	Arm Iusion within rays			
50(U)	Arm autilude Besumbert arma			
57(B)	Eined brochiele			
58(B)	Fixed Dracmais	T		
39(U)	Number of fixed oracinals			
60(B)	Interorachials			
61(B)	Porm of interoracinais		_	_=
62(B)	Pinnules			
63(U)	Number of pinnules per oracinal			
64(B)	Neumbert ambulacra			
65(U)	Number of recumbent ambulacra			
66(B)	Extent of recumbent amountera		_=	
67(U)	Ratio of arm length to cup height			
07B(B)	Torred arms			
Anal and	tegminal features			
68(B)	Anal opening through cup	+	+	
69(B)	Anal tube or sac	-		
70(O)	Position of tube or sac			-
71(B)	Extent of tube or sac			
72(B)	Ridges on tube or sac	-		
73(B)	Regular plating of tube or sac			
74(B)	Multiplated tegmen			-1
75(B)	Extent of tegmen			



Data culled to increase separation of species

FIG. 13—Distribution of higher taxa along first three recomputed principal-coordinate axes, based on data culled to yield morphological separations among contemporaneous genera at least as great as the average separation among congeneric species. A, Ordovician-Devonian; B, Carboniferous and Permian. Key to higher taxa as in Figure 4. Higher taxa that evacuated areas in morphospace before the Carboniferous or colonized areas anew after the Devonian are largely the same as for the unculled data. The main difference is that poteriocrine cladids no longer define the extreme values of the first principal coordinate and no longer appear to represent as substantial an expansion of occupied morphological space for crinoids as a whole. This suggests that the extent to which poteriocrines expanded into new morphological space is exaggerated by the pull this clade exerts, by virtue of its sheer numbers rather than its extreme morphology, on the principal-coordinate axes (see also Figs. 4, 12, and 14).

Stratigraphic interval	Stem	Cup	Arms	Tegmen	Total
Lower Ordovician	4.2	1.8	1.0	5.0	12.0
Ordovician-2	3.5	0.8	13	29	8 5
Ordovician-3	3.4	1.2	0.1	2.9	7.6
Llandoverian	3.3	0.8	1.0	3.6	87
Wenlockian	3.9	0.5	0.1	37	82
Upper Silurian	4.9	0.9	0.1	3.1	9.0
Lower Devonian	5.5	1.0	0.5	4.1	11.1
Middle Devonian	5.4	0.5	0.4	2.8	9 1
Upper Devonian	5.2	1.4	0.4	5.0	12.0
Tournaisian	5.4	0.7	0.7	2.7	9.6
Visean	5.9	0.9	0.9	4.0	11 7
Serpukhovian	7.8	0.8	0.4	4.7	13.7
Bashkirian	9.0	1.1	1.2	6.7	18.0
Moscovian	7.9	0.6	1.2	5.2	14.9
Stephanian	6.4	0.8	1.0	5.0	13.2
Permian-1	6.1	1.2	1.4	4.9	13.5
Permian-2	8.6	0.6	2.0	3.8	14.9
Permian-3	10.0	0.5	2.0	2.8	15.2

TABLE 4-Mean number of unknown characters per species.

exceeds the level of the Ordovician. However, since the Permian-3 sample contains only five species, this last result should be treated with caution.

The foregoing approach allows a limited variation in character weight between 0 and 2. An alternative is to allow greater variation in weight by bootstrapping the characters, i.e. resampling the characters with replacement. In this way, a character can, in principle, be weighted between 0 and M, where M is the total number of characters, while still having an average weight of unity over a large number of character samples. This procedure results in more variation among the disparity curves (Fig. 17), but these curves still have the same general shape.

Although the decline in disparity between the Visean and Stephanian may be a sampling artifact (see above), it is reasonable to conclude that the major feature of crinoid history documented here, namely a lack of increase in disparity through most of the Paleozoic, despite extensive taxonomic diversification and some 200 million years of post-Ordovician evolution, is not an artifact of character weighting.

None of the approaches to character weighting used here attempts to assign weight to character complexes. It is probable that many characters evolve nonindependently as complexes. According to all three weighting schemes explored above, such complexes would tend to be weighted in proportion to the number of unit characters into which they are decomposed for the sake of analysis. One might want to devise a weighting scheme in which each character complex was given unit weight. However, at least partly because the concept of a character complex is so variable, we do not yet have an objective and fully general way of recognizing such complexes (see Olson and Miller, 1958, Campbell and Barwick, 1990, and Zelditch et al., 1992 for some discussions of this issue). Although much remains to be said on this subject, the fact that cup characters and arm characters alone each show sequences of disparity broadly similar to that based on all characters (Foote, 1994b) suggests that the large-scale history of disparity in crinoids does not simply reflect undue weight assigned to a small number of characters.



Data culled to increase separation of species

M. FOOTE

158

Stratigraphic Interval



FIG. 15—Morphological disparity in Paleozoic crinoids, omitting characters of the stem and tegmen, the morphological regions that account for the majority of missing data (Table 4). Pattern is very similar to that based on all characters, suggesting that overall pattern is not an artifact of temporal changes in missing data.

DISCUSSION AND SUMMARY

The foregoing analyses of discrete characters complement previous work in suggesting that most of the spectrum of crinoid form evolved fairly early in the history of the clade (Ausich, 1988; Foote, 1994a,b, 1995). This result is robust, not being sensitive to character weighting or missing data. The same result is shown in a cruder way by the fact that only three of 78 characters that vary among crinoids as a whole did not vary among species in the pre-Carboniferous sample. Other aspects of the distribution of morphology are considered elsewhere (Foote 1995).

Although morphological disparity did not increase substantially after the Ordovician, the distribution of forms was not static. In addition to well noted evolutionary trends within higher taxa (e.g., Springer, 1920; Moore and Laudon, 1943; Moore, 1952; Lane and Webster, 1966;

FIG. 14—Ordination of Paleozoic crinoids along first twenty recomputed principal-coordinate axes, based on data culled to yield separations among contemporaneous genera at least as great as the average separation among congeneric species. See Figure 12 for explanation of graphs. Ordovician-Devonian taxa exceeding morphological extremes of the Carboniferous and Permian are still Dendrocrinina (low values on axis 2), Disparida (low values on axis 3 and low values on axis 4), and Diplobathrida (mostly Rhodocrinitacea; high values on axis 6). In addition, many Ordovician-Devonian Glyptocrinina (high values on axis 2) and Flexibilia (low values on axis 11) exceed Carboniferous and Permian extremes. Carboniferous and Permian taxa exceeding Ordovician-Devonian extremes are still Codiacrinacea (high values on axes 3 and 5) and Hexacrinitacea (low values on axis 6). However, very few of the 135 remaining species of Carboniferous and Permian Poteriocrinina lie beyond the range of Ordovician-Devonian data (three species with low values on axis 1, two with high values on axis 2, three with low values on axis 6, and a scattering of species along the higher axes).



FIG. 16—Time series of morphological disparity with randomly assigned character weights. For each time series, each character is assigned a weight uniform on (0,2). This yields an average weight of unity over a large number of time series. Panels A-G each show only five different time series, so that each trajectory can be followed clearly. Panel H shows 100 time series. Although curves vary

Webster and Lane, 1967; Moore and Strimple, 1969; Broadhead, 1985, 1988a; Simms, 1990; Kammer and Ausich, 1992), changes in diversity within higher taxa corresponded with the evacuation and colonization of whole regions in morphological space. An attempt to explore the effects of differential sampling of crinoid forms suggests that the apparent expansion of poteriocrine cladids into new morphological space is exaggerated by the density with which poteriocrines are sampled.

On the whole, the observed changes in the morphological distribution of crinoids seem small compared with the expectation of evolutionary models of unbounded diffusion (Stanley, 1973; Raup and Gould, 1974; Fisher, 1986; Gould, 1988; McKinney, 1990; Carlson, 1992; Foote, 1993; McShea, 1993a; Valentine et al., 1994), which would predict a rather steady increase in morphological dispersion. Not all morphological extremes reached before the Carboniferous represent absolute limits (Figs. 4, 12). Nevertheless, the importance of functional and structural constraints (Ausich, 1988; Broadhead, 1988b; Kendrick, 1992, 1993) in the large-scale evolution of crinoids is suggested by the small extent to which morphological extremes were exceeded after the Devonian (Figs. 12, 14; see Foote 1995). Although the Lower Carboniferous is often regarded as the acme of crinoids (Wachsmuth and Springer, 1897: p. 163; Springer, 1920: p. 97; Lane and Sevastopulo, 1990), this view reflects their abundance and diversity much more than their variety of form.

ACKNOWLEDGMENTS

For advice, comments and discussion on this and related work I thank J. Alroy, W. I. Ausich, T. K. Baumiller, B. E. Bodenbender, J. C. Brower, L. S. Collins, T. M. Collins, D. C. Fisher, P. D. Gingerich, T. E. Guensburg, P. Holterhoff, D. Jablonski, D. C. Kendrick, D. W. McShea, D. L. Meyer, F. M. Richter, J. J. Sepkoski, Jr., R. Terry, and M. L. Zelditch. C. Stachnik helped greatly in obtaining literature. G. D. Webster kindly provided unpublished portions of his bibliography. I thank T. K. Baumiller for reading the manuscript. F. M. Richter suggested the approach of assigning random character weights, and J. Alroy suggested the bootstrap resampling of characters. This work was supported by the National Science Foundation (DEB-9207577) and by the Donors of the Petroleum Research Fund, administered by the American Chemical Society.

in height, most show the same shape, suggesting that the pattern documented in Figure 1 is not an artifact of an arbitrary weighting scheme. Note that the occurrence of maximal disparity in Ordovician-2 is not completely robust to variation in character weighting. In this and the following figure, high values in Lower Ordovician and Permian-3 may be unreliable because of very small sample size.



FIG. 17—Time series of morphological disparity based on bootstrap resampling of characters (with replacement). This yields greater variation in character weights while still giving a mean weight of

LITERATURE CITED

ALLMON, W. D. 1992. Genera in paleontology: definition and significance. Historical Biology, 6: 149-158.

ANSTEY, R. L., and J. F. PACHUT. 1992. Cladogenesis and speciation in early bryozoans. Geological Society of America Abstracts with Programs, 24: A139.

AUSICH, W. I. 1980. A model for niche differentiation in Lower Mississippian crinoid communities. Journal of Paleontology, 54: 273-288.

——. 1983. Functional morphology and feeding dynamics of the Early Mississippian crinoid *Barycrinus asteriscus*. Journal of Paleontology, 57: 31-41.

. 1986. Paleoecology and history of the Calceocrinidae (Palaeozoic Crinoidea). Palaeontolgy, 29: 85-99.

——. 1988. Evolutionary convergence and parallelism in crinoid calyx design. Journal of Paleontology, 62: 906-916.

BAUMILLER, T. K. 1990. Physical modeling of the batocrinid anal tube: functional analysis and multiple hypothesis testing. Lethaia, 23: 399-408.

------. 1993. Survivorship analysis of Paleozoic Crinoidea: effect of filter morphology on evolutionary rates. Paleobiology, 19: 304-321.

——— and R. E. PLOTNICK. 1989. Rotational stability in stalked crinoids and the function of wing plates in *Pterotocrinus depressus*. Lethaia, 22: 317-326.

BREIMER, A. 1969. A contribution to the paleoecology of Paleozoic stalked crinoids. Proceedings, Koninklijke Nederlandse Akademie van Wetenschappen B, 72: 139-150.

—— and N. G. LANE. 1978. Ecology and paleoecology. *In* R. C. Moore and C. Teichert (eds.), Treatise on Invertebrate Paleontology, Part T, Echinodermata 2. Geological Society of America, Boulder, Colorado, and University of Kansas, Lawrence, Kansas, p. T316-T347.

and G. D. WEBSTER. 1975. A further contribution to the paleoecology of fossil stalked crinoids. Proceedings, Koninklijke Nederlandse Akademie van Wetenschappen B, 78: 149-167.

BRETT, C. E. 1981. Terminology and functional morphology of attachment structures in pelmatozoan echinoderms. Lethaia, 14: 343-370.

BRIGGS, D. E. G., and R. A. FORTEY. 1989. The early radiation and relationships of the major arthropod groups. Science, 246: 241-243.

, ____, and M. A. WILLS. 1992. Morphological disparity in the Cambrian. Science, 256: 1670-1673.

BROADHEAD, T. W. 1981. Carboniferous camerate crinoid subfamily Dichocrininae. Palaeontographica A, 176: 81-157.

——. 1985. Evolution of Carboniferous Hexacrinitacea (Crinoidea, Camerata). Proceedings, Ninth International Congress on Carboniferous Stratigraphy and Geology, 5: 205-215.

. 1988a. The evolution of feeding structures in Palaeozoic crinoids. *In* C. R. C. Paul and A. B. Smith (eds.), Echinoderm Phylogeny and Evolutionary Biology. Clarendon, Oxford, p. 257-268.

. 1988b. Heterochrony-- A pervasive influence in the evolution of Paleozoic Crinoidea. In R. D. Burke, R. V. Mladenov, P. Lambert, and R. L. Parsley (eds.), Echinoderm Biology. Balkema, Rotterdam, p. 115-128.

unity over many time series. Despite greater variation among the curves, it is still difficult to make a case for a steady increase in disparity over the Paleozoic. However, certain smaller scale features, such as the occurrence of maximal disparity in Ordovician-2, are sensitive to variation in character weighting.

BROWER, J. C. 1966. Functional morphology of Calceocrinidae with description of some new species. Journal of Paleontology, 40: 613-634.

—. 1973. Crinoids from the Girardeau Limestone (Ordovician). Palaeontographica Americana, 7: 259-499.

——. 1982. Phylogeny of primitive calceocrinids. In J. Sprinkle (ed.), Echinoderm Faunas from the Bromide Formation (Middle Ordovician) of Oklahoma. University of Kansas Paleontological Contributions, Monograph 1, p. 90-110.

-----. 1988a. Ontogeny and phylogeny in primitive calceocrinid crinoids. Journal of Paleontology, 62: 917-934.

. 1988b. Allometry of the food-gathering system in calceocrinid crinoids. In R. D. Burke, P. V. Mladenov, P. Lambert, and R. L. Parsley (eds.), Echinoderm Biology. Balkema, Rotterdam, p. 73-80.

——. 1992a. Cupulocrinid crinoids from the Middle Ordovician (Galena Group, Dunleith Formation) of northern Iowa and southern Minnesota. Journal of Paleontology, 66: 99-128.

-----. 1992b. Hybocrinid and disparid crinoids from the Middle Ordovician (Galena Group, Dunleith Formation) of northern Iowa and southern Minnesota. Journal of Paleontology, 66: 973-993.

------. 1994. Camerate crinoids from the Middle Ordovician (Galena Group, Dunleith Formation) of northern Iowa and southern Minnesota. Journal of Paleontology, 68: 570-599.

—— and J. VEINUS. 1974. Middle Ordovician crinoids from southwestern Virginia and eastern Tennessee. Bulletins of American Paleontology, 66(283): 1-125.

——— and ———. 1978. Middle Ordovician crinoids from the Twin Cities area of Minnesota. Bulletins of American Paleontology, 74(304): 369-506.

CAMPBELL, K. S. W., and R. E. BARWICK. 1990. Paleozoic dipnoan phylogeny: functional complexes and evolution without parsimony. Paleobiology, 16: 143-169.

CAMPBELL, K. S. W., and C. R. MARSHALL. 1987. Rates of evolution among Palaeozoic echinoderms. In K. S. W. Campbell and M. F. Day (eds.), Rates of Evolution. Allen and Unwin, London, p. 61-100.

CARLSON, S. J. 1992. Evolutionary trends in the articulate brachiopod hinge mechanism. Paleobiology, 18: 344-366.

CHERRY, L. M., S. M. CASE, J. G. KUNKEL, J. S. WYLES, and A. C. WILSON. 1982. Body shape metrics and organismal evolution. Evolution, 36: 914-933.

DERSTLER, K. L. 1981. Morphological diversity of early Cambrian echinoderms. In M. E. Taylor (ed.), Short papers for the second International Symposium on the Cambrian System. United States Geological Survey Open File Report 81-743, p. 71-75.

-----. 1982. Estimating the rate of morphological change in fossil groups. Proceedings, Third North American Paleontological Convention, 1: 131-136.

DONOVAN, S. K. 1986. Pelmatozoan columnals from the Ordovician of the British Isles, part 1. Palaeontographical Society Monograph, 138(568): 1-68.

. 1988. Functional morphology of synarthrial articulations in the crinoid stem. Lethaia, 21: 169-175.

——. 1989a. Pelmatozoan columnals from the Ordovician of the British Isles, part 2. Palaeontographical Society Monograph, 142(580): 69-120.

——. 1990. Functional morphology of synostosial articulations in the crinoid column. Lethaia, 23: 291-296.

EFRON, B. 1982. The jackknife, the bootstrap, and other resampling plans. Society for Industrial and Applied Mathematics, Philadelphia.

FELSENSTEIN, J. 1985. Phylogenies and the comparative method. American Naturalist, 125: 1-15.

FISHER, D. C. 1986. Progress in organismal design. In D. M. Raup and D. Jablonski (eds.), Patterns and Processes in the History of Life. Springer, Berlin, p. 99-117.

FOOTE, M. 1991. Morphological and taxonomic diversity in a clade's history: the blastoid record and stochastic simulations. Contributions from the Museum of Paleontology, University of Michigan, 28: 101-140.

——. 1992. Paleozoic record of morphological diversity in blastozoan echinoderms. Proceedings of the National Academy of Sciences USA, 89: 7325-7329.

^{. 1989}b. The significance of the British Ordovician crinoid fauna. Modern Geology, 13: 243-255.

——. 1993. Discordance and concordance between morphological and taxonomic diversity. Paleobiology, 19: 185-204.

——. 1994a. Morphology of Ordovician-Devonian crinoids. Contributions from the Museum of Paleontology, University of Michigan, 29: 1-39.

——. 1994b. Morphological disparity in Ordovician-Devonian crinoids and the early saturation of morphological space. Paleobiology, 20: 320-344.

——. 1995. Morphological diversification of Paleozoic crinoids. Paleobiology, 21: 273-299.

and S. J. GOULD. 1992. Cambrian and Recent morphological disparity. Science, 258: 1816.
GOULD, S. J. 1988. Trends as changes in variance: a new slant on progress and directionality in evolution. Journal of Paleontology, 62: 319-329.

—. 1989. Wonderful Life: The Burgess Shale and the Nature of History. Norton, New York.

——. 1991. The disparity of the Burgess Shale arthropod fauna and the limits of cladistic analysis: why we must strive to quantify morphospace. Paleobiology, 17: 411-423.

——. 1993. How to analyze Burgess Shale disparity--a reply to Ridley. Paleobiology, 19: 522-523.

GOWER, J. C. 1966. Some distance properties of latent root and vector methods used in multivariate analysis. Biometrika, 53: 325-338.

GUENSBURG, T. E. 1984. Echinodermata of the Middle Ordovician Lebanon Limestone, central Tennessee. Bulletins of American Paleontology, 87(319): 1-100.

— and J. SPRINKLE. 1990. Early Ordovician crinoid-dominated echinoderm fauna from the Fillmore Formation of western Utah. Geological Society of America Abstracts with Programs, 22: A220.

and ———. 1992. Rise of echinoderms in the Paleozoic evolutionary fauna: Significance of paleoenvironmental controls. Geology, 20: 407-410.

HARLAND, W. B., R. L. ARMSTRONG, A. V. COX, L. E. CRAIG, A. G. SMITH, and D. G. SMITH. 1990. A geologic time scale 1989. Cambridge University Press, New York.

HICKMAN, C. S. 1993a. Biological diversity: elements of a paleontological agenda. Palaios, 8: 309-310.

——. 1993b. Theoretical design space: a new program for the analysis of structural diversity. Neues Jahrbuch für Geologie und Paläontologie, Abhandlungen, 190: 169-182.

JACOBS, D. K. 1990. Selector genes and the Cambrian radiation of the Bilateria. Proceedings of the National Academy of Sciences USA, 87:4406-4410.

KAMMER, T. W. 1985. Aerosol filtration theory applied to Mississippian deltaic crinoids. Journal of Paleontology, 59: 551-556.

— and W. I. AUSICH. 1987. Aerosol suspension feeding and current velocities: distributional controls for late Osagean crinoids. Paleobiology, 13: 379-395.

— and — . 1992. Advanced cladid crinoids from the middle Mississippian of the east-central United States: primitive-grade calyces. Journal of Paleontology, 66: 461-480.

— and — . 1993. Advanced cladid crinoids from the middle Mississippian of the east-central United States: intermediate-grade calyces. Journal of Paleontology, 67: 614-639.

——— and ———. 1994. Advanced cladid crinoids from the middle Mississippian of the east-central United States: advanced-grade calyces. Journal of Paleontology, 68: 339-351.

KELLY, S. M. 1982. Origin of the crinoid orders Disparida and Cladida: Possible inadunate cup plate homologies. Proceedings, Third North American Paleontological Convention, 1: 285-290.

KENDRICK, D. C. 1992. Crinoid arm branching topology, pinnulation, and the convergence of crinoid arm designs. Geological Society of America Abstracts with Programs, 24: A225.

——. 1993. Computer modelling of crinoid calyx morphologies and comparisons with real forms. Geological Society of America Abstracts with Programs, 25: A103.

KNAPP, W. D. 1969. Declinida, a new order of late Paleozoic inadunate crinoids. Journal of Paleontology, 43: 340-391.

KOENIG, J. W. 1965. Ontogeny of two Devonian crinoids. Journal of Paleontology, 39: 398-413.

KOLATA, D. R. 1975. Middle Ordovician echinoderms from northern Illinois and southern Wisconsin. Paleontological Society Memoir, 7: 1-74.

LANE, N. G. 1963a. Meristic variation in the dorsal cup of monobathrid camerate crinoids. Journal of Paleontology, 37: 917-930.

——. 1963b. The Berkeley crinoid collection from Crawfordsville, Indiana. Journal of Paleontology, 37: 1001-1008.

——. 1967. Revision of suborder Cyathocrinina (class Crinoidea). University of Kansas Paleontological Contribtuions, Paper, 24: 1-13.

——. 1971. Crinoids and reefs. Proceedings, North American Paleontological Convention, 2: 1430-1443.

— and G. D. SEVASTOPULO. 1981. Functional morphology of a microcrinoid: Kallimorphocrinus punctatus n. sp. Journal of Paleontology, 55: 13-28.

—— and ——. 1982a. Microcrinoids from the Middle Pennsylvanian of Indiana. Journal of Paleontology, 56: 103-115.

—— and ——. 1982b. Growth and systematic revision of Kallimorphocrinus astrus, a Pennsylvanian microcrinoid. Journal of Paleontology, 56: 244-259.

—— and ——. 1990. Biogeography of Lower Carboniferous crinoids. Geological Society (London) Memoir, 12: 333-338.

——— and G. D. WEBSTER. 1966. New Permian crinoid fauna from southern Nevada. University of California Publications in Geological Sciences, 63: 1-87.

LEE, M. S. Y. 1992. Cambrian and Recent morphological disparity. Science, 258: 1816-1817.

MACURDA, D. B., Jr. 1974. A quantitative phyletic study of the camerate crinoid families Actinocrinitidae and Periechocrinitidae and its taxonomic implications. Journal of Paleontology, 48: 820-832.

McGHEE, G. R. 1991. Theoretical morphology: the concept and its applications. *In* N. L. Gilinsky and P. W. Signor (eds.), Analytical Paleobiology (Short Courses in Paleontology 4). The Paleontological Society, Knoxville, Tennessee, p. 87-102.

MCKINNEY, M. L. 1990. Classifying and analysing evolutionary trends. In K. J. McNamara (ed.), Evolutionary Trends. University of Arizona Press, Tucson, Arizona, p. 28-58.

McSHEA, D. W. 1993a. Evolutionary change in the morphological complexity of the mammalian vertebral column. Evolution, 47: 730-740.

------. 1993b. Arguments, tests, and the Burgess Shale--a commentary on the debate. Paleobiology, 19: 399-402.

MEYER, D. L. 1973. Feeding behavior and ecology of shallow-water unstalked crinoids (Echinodermata) in the Caribbean Sea. Marine Biology, 22: 105-129.

-----. 1979. Length and spacing of the tube feet in crinoids (Echinodermata) and their role in suspension feeding. Marine Biology, 51: 361-369.

——. 1983. Food and feeding mechanisms: Crinozoa. In M. Jangoux and J. M. Lawrence (eds.), Echinoderm nutrition. Balkema, Rotterdam, p. 25-42.

MOORE, R. C. 1952. Evolutionary rates among crinoids. Journal of Paleontology, 26: 338-352.

——. 1962. Ray structures of some inadunate crinoids. University of Kansas Paleontological Contributions, Echinodermata, Article, 5: 1-47.

—— and L. R. LAUDON. 1943. Evolution and classification of Paleozoic crinoids. Geological Society of America Special Paper, 46: 1-153.

and F. B. PLUMMER. 1938. Upper Carboniferous crinoids from the Morrow subseries of Arkansas, Oklahoma, and Texas. Denison University Scientific Laboratories Journal, 32: 209-313.

—— and ——. 1940. Crinoids from the Upper Carboniferous and Permian strata in Texas. University of Texas Publication, 3945: 1-459.

----- and H. L. STRIMPLE. 1969. Explosive evolutionary differentiation of unique group of Mississippian-Pennsylvanian camerate crinoids (Acrocrinidae). University of Kansas Paleontological Contributions, Paper, 39: 1-44.

and ———. 1973. Lower Pennsylvanian (Morrowan) crinoids from Arkansas, Oklahoma, and Texas. University of Kansas Paleontological Contributions, Article, 60: 1-84.

and C. TEICHERT (eds.). 1978. Treatise on Invertebrate Paleontology, Part T, Echinodermata
 Geological Society of America, Boulder, Colorado and University of Kansas, Lawrence, Kansas.

OLSON, E. C., and R. L. MILLER. 1958. Morphological Integration. University of Chicago Press, Chicago.

PAUL, C. R. C. 1977. Evolution of primitive echinoderms. In A. Hallam (ed.), Patterns of Evolution. Elsevier, Amsterdam, p. 123-157.

——. 1979. Early echinoderm radiation. In M. R. House (ed.), The Origin of Major Invertebrate Groups. Academic Press, London, p. 415-434.

RAUP, D. M. 1966. Geometric analysis of shell coiling: general problems. Journal of Paleontology, 40: 1178-1190.

——. 1967. Geometric analysis of shell coiling: coiling in ammonoids. Journal of Paleontology, 41: 43-65.

------ and S. J. GOULD. 1974. Stochastic simulation and evolution of morphology-- towards a nomothetic paleontology. Systematic Zoology, 23: 305-322.

RIDDLE, S. W., J. I. WULFF, and W. I. AUSICH. 1988. Biomechanics and stereomic microstructure of the *Gilbertsocrinus tuberosus* column. *In* R. D. Burke, R. V. Mladenov, P. Lambert, and R. L. Parsley (eds.), Echinoderm Biology. Balkema, Rotterdam, p. 641-648.

RIDLEY, M. 1993. Analysis of the Burgess Shale. Paleobiology, 19: 519-521

- ROY, K. 1994. Effects of the Mesozoic Marine Revolution on the taxonomic, morphologic, and biogeographic evolution of a group: aporrhaid gastropods during the Mesozoic. Paleobiology 20:274-296.
- RUNNEGAR, B. 1987. Rates and modes of evolution in the Mollusca. In K. S. W. Campbell and M. F. Day (eds.), Rates of Evolution. Allen and Unwin, London, p. 39-60.
- SAUNDERS, W. B., and A. R. H. SWAN. 1984. Morphology and morphologic diversity of mid-Carboniferous (Namurian) ammonoids in time and space. Paleobiology, 10: 195-228.
- SEPKOSKI, J. J., Jr., and D. M. RAUP. 1986. Periodicity in marine extinction events. In D. K. Elliott (ed.), Dynamics of Extinction. Wiley, New York, p. 3-36.
- SIEGEL, S., and N. J. CASTELLAN, Jr. 1988. Nonparametric Statistics for the Behavioral Sciences, Second Edition. McGraw-Hill, New York.
- SIMMS, M. J. 1990. Crinoids. In K. J. McNamara (ed.), Evolutionary Trends. University of Arizona Press, Tucson, Arizona, p. 188-204.

-----. 1993. Reinterpretation of the cal plate homology and phylogeny in the Class Crinoidea. Lethaia, 26: 303-312.

- SMITH, A. B. 1990. Evolutionary diversification of echinoderms during the early Palaeozoic. In P. D. Taylor and G. P. Larwood (eds.), Major Evolutionary Radiations. Clarendon Press, Oxford, p. 265-286.
- -----. 1994. Systematics and the Fossil Record: Documenting Evolutionary Patterns. Blackwell, Oxford.
- SNEATH, P. H. A., and R. R. SOKAL. 1973. Numerical Taxonomy. W. H. Freeman, San Francisco. SPRINGER, F. 1920. The Crinoidea Flexibilia. Smithsonian Institution Publication, 2501: 1-486.
- ------. 1924. A remarkable fossil echinoderm fauna in the East Indies. American Journal of Science, series 5, 8: 325-335.
- ——. 1926. Unusual forms of fossil crinoids. Proceedings, United States National Museum, 67(5): 1-137.
- SPRINKLE, J. 1980. An overview of the fossil record. In T. W. Broadhead and J. A. Waters (eds.), Echinoderms: notes for a short course. University of Tennessee, Knoxville, Tennessee, p. 15-26.
- (ed.). 1982. Echinoderm Faunas from the Bromide Formation (Middle Ordovician) of Oklahoma. University of Kansas Paleontological Contributions Monograph, 1: 1-369.
 - —. 1983. Patterns and problems in echinoderm evolution. Echinoderm Studies, 1: 1-18.

——. 1990. New echinoderm fauna from the Ninemile Shale (Lower Ordovician) of central and southern Nevada. Geological Society of America Abstracts with Programs, 22: A219.

——. 1992. Radiation of Echinodermata. In J. H. Lipps and P. W. Signor (eds.), Origin and Early Evolution of the Metazoa. Plenum, New York, p. 375-398.

----- and T. E. GUENSBURG. 1991. Origin of echinoderms in the Paleozoic evolutionary fauna: New data from the Early Ordovician of Utah and Nevada. Geological Society of America Abstracts with Programs, 23: A278.

——— and G. P. WAHLMAN. 1994. New echinoderms from the Early Ordovician of west Texas. Journal of Paleontology, 68: 324-338.

STANLEY, S. M. 1973. An explanation for Cope's rule. Evolution, 27: 1-26.

STRIMPLE, H. L. 1973. Fossil crinoid studies, Part 7. Tegminal structures of some inadunate crinoids. University of Kansas Paleontological Contributions, Paper, 66: 27-32.

UBAGHS, G. 1978a. Skeletal morphology of fossil crinoids. In R. C. Moore and C. Teichert (eds.), Treatise on Invertebrate Paleontology, Part T, Echinodermata 2. Geological Society of America, Boulder, Colorado, and University of Kansas, Lawrence, Kansas, p. T58-T216.

——. 1978b. Evolution of camerate crinoids. In R. C. Moore and C. Teichert (eds.), Treatise on Invertebrate Paleontology, Part T, Echinodermata 2. Geological Society of America, Boulder, Colorado, and University of Kansas, Lawrence, Kansas, p. T281-T292.

VALENTINE, J. W. 1986. Fossil record of the origin of Baupläne and its implications. In D. M. Raup and D. Jablonski (eds.), Patterns and Processes in the History of Life. Springer, Berlin, p. 209-222.
——. 1991. Major factors in the rapidity and extent of the metazoan radiation during the Proterozoic-Phanerozoic transition. In A. Simonetta and S. Conway Morris (eds.), The Early Evolution of Metazoa and the Significance of Problematic Taxa. Cambridge University Press, Cambridge, U. K., p. 11-13.

—. 1992. The macroevolution of phyla. *In* J. H. Lipps and P. W. Signor (eds.), Origin and Early Evolution of the Metazoa. Plenum, New York, p. 525-553.

——, A. G. COLLINS and C. P. MEYER. 1994. Morphological complexity increase in metazoans. Paleobiology, 20: 131-142.

VAN VALEN, L. 1974. Multivariate structural statistics in natural history. Journal of Theoretical Biology, 45:235-247.

WACHSMUTH, C., and F. SPRINGER. 1897. The North American Crinoidea Camerata. Harvard College Museum of Comparative Zoology Memoir, 20: 1-897, 21: plates 1-83.

WAGNER, P. J. 1993. Temporal patterns of morphologic disparity among early Paleozoic "archaeogastropods." Geological Society of America Abstracts with Programs, 25: A51.

WANNER, J. 1916. Die permischen Echinodermen von Timor, Teil 1. Paläontologie von Timor, part 6, number 11: 1-329.

——. 1924. Die permisheen Krinoiden von Timor. Jaarboek van het Mijnwezen in Nederlandsch Oost-Indië, Verhandelingen, 50(3): 1-348.

WARD, P. D. 1980. Comparative shell shape distributions in Jurassic-Cretaceous ammonites and Jurassic-Tertiary nautilids. Paleobiology, 6: 32-43.

WARN, J. M. 1975. Monocyclism vs. Dicyclism: A primary schism in crinoid phylogeny? Bulletins of American Paleontology, 67(287): 423-441.

WEBSTER, G. D. 1969. Bibliography and index of Paleozoic crinoids, 1942-1968. Geological Society of America Memoir, 137: 1-341.

——. 1977. Bibliography and index of Paleozoic crinoids, 1969-1973. Geological Society of America Microform Publication, 8: 1-235.

-----. 1981. New crinoids from the Naco Formation (Middle Pennsylvanian) of Arizona and a revision of the family Cromyocrinidae. Journal of Paleontology, 55: 1176-1199.

——. 1986. Bibliography and index of Paleozoic crinoids, 1974-1980. Geological Society of America Microform Publication, 16: 1-405.

——. 1987. Permian crinoids from the type-section of the Callytharra Formation, Callytharra Springs, Western Australia. Alcheringa, 11: 95-135.

——. 1988. Bibliography and index of Paleozoic crinoids and coronate echinoderms, 1981-1985. Geological Society of America Microform Publication, 18: 1-235.

---. 1990. New Permian crinoids from Australia. Palaeontology, 33: 49-74.

——. 1993. Bibliography and index of Paleozoic crinoids, 1986-1990. Geological Society of America Microform Publication, 25: 1-204.

— and P. A. JELL. 1992. Permian echinoderms from Western Australia. Memoirs of the Queensland Museum, 32: 311-373.

——— and N. G. LANE. 1967. Additional Permian crinoids from southern Nevada. University of Kansas Paleontological Contributions, Paper 27: 1-32.

WILLS, M. A., D. E. G. BRIGGS, and R. A. FORTEY. 1994. Disparity as an evolutionary index: a comparison between Cambrian and Recent arthropods. Paleobiology, 20: 93-130.

ZELDITCH, M. L., F. L. BOOKSTEIN, and B. L. LUNDRIGAN. 1992. Ontogeny of integrated skull growth in the cotton rat *Sigmodon fulviventer*. Evolution, 46: 1164-1180.

APPENDIX 1

Character Data for Carboniferous and Permian Crinoid Species Used in this Study

Character data for crinoid species used in this study are tabulated on the following pages. Also given are pre-Carboniferous species not included in Foote (1994a), and emendations for species included in Foote (1994a). Unknown character states are indicated by ?; inapplicable states are indicated by N. See Table 1 and Foote (1994a) for explanations of characters and character states. Sources for data are listed in Appendix 2; where no source is given, see Foote (1994a). Two-letter code preceding each species indicates higher taxonomic affinity: CL, Cladida; DB, Diplobathrida; DI, Disparida; FL, Flexibilia; HY, Hybocrinida; and MO, Monobathrida.

Genus	Species	Sources	Char.no. 1111111122222222223333333344 4 12345678901234567890123456789012345678900 0 A B	4 4 44444455555555 1 2 345678901234567	5 5 666 8 9 012	6 666666777777 3 4567789012345 A
Lower Ordovician CL Elpasocrinus	radiatus	105	2710100070NN77151110111250125013117401010 N	1 5 0720170N0001010	ONN N O	N 011720777
<u>Ordovician-2</u> DI Peltacrinus DI Tunguskocrinus HY Cornucrinus	sculptatus ivanovae mirus	-	2710100070NN77150N0000N250NNNNN0207501000 N 27777777777150N000132550NNNN1217401000 N 2000007070NNN77150N00111241NNNNN0312001000 N	1 5 0010120N0001010 1 5 0000110N00000010 0 N NNNNNNNNNNN	ONN N NNN N O	N 0NN3001070077 N 0NN3001770077 N 130NN1NNN00
<u>Ordovician-3</u> DB Cotylacrima DB Simplococrinus MO Eopatelliocrinus CL Aithriocrinus DI Cincinnaticrinus DI Isotomocrinus DI Niocrinus	sandra persculptus s scyphogracil s trahani t varibrachial tenuis laxus	16 is us	2010007010NN1?15131000N25?0N5011300411000 N 2?10067000NN??15131000N150135011?10401110 N 2?10006010NN??150N0000N231NNNNN011?401010 N 2?10700070NN0N15??1070N25???5?0011??00000 N 201010?0?0NN12150N00012250NNNN1111?400000 N 2?101660?0NN??7150N00012250NNNN1111?400000 N 2?101660?0NN??7150N000122550NNNN1111?400000 N	1 20 00200NNN1NN0010 1 20 00100NNN1NN0010 1 10 00100NNN0011010 1 5 000121100002010 1 5 00101200002010 1 5 000011100002010	11111111111111111111111111111111111111	1 DNN?00DNNN10 1 DNN2002727212 1 DNN2002727272 1 DNN2002727272 1 DNN200100000 2 DNN2001010072 2 N DNN3001110072 2 N DNN3001110072
<u>Llandoverian</u> DB Silfonocrinus CL Euspirocrinus CL Kanabinocrinus	siluricus heliktos thyaros		2?1000?0?0NN??15131000N2500N5011110411100 N 2?100000?0NN??15111000N2500N501311?401000 N 2?000000?0NN??15111010N1500N501211?500000 N	1 20 00100NNN0011010 1 5 0010120N0001010 1 5 0210120N0001010	1 2 111 0 N NNO 0 N NNO	? ONN300?????? N ONN2001?100?? N ONN2001?11???
<u>Wenlockian</u> MO Promelocrinus CL Enallocrinus DI Pisocrinus	anglicus scriptus pillula		2?100070?0NN??150N0000N241NNNNN011?501100 N 2?0?0000?????15111000N1500N5021100400000 N 20000000?0NN??150N01111251NNNNN0310501000 N	1 20 021012121NN0010 1 5 0000120N0000010 1 5 00000NNN0003010	1 10 111 0 N NNO 0 N NNO	1 ONN200?????? N ONN300?????10 N ONN3001010000
Upper Silurian DB Anthemocrinus FL Pepitaxocrinus CL Syndetocrinus DI Cremacrinus DI Pisocrinus	minor svobodai dartae tubuliferus quinquelobus	98 98 98 98 98	2?????????????????????????????????????	1 10 0010100N1NN0010 1 5 001012110000011 1 30 12001201000001 1 4 020012100001010 1 5 00000NN0003011	01 4 111 00 N NU 01 10 0N0 00 N NN0 00 N NN0	1 DNN200777777 N DNN300777777 N DNN20077771 N DNN2007777710 N DNN3001010077 N DNN100777700
Lower Devonian MO Ctenocrinus FL Clidochirus FL Pycnosaccus CL Bridgerocrinus CL Pagecrinus	arduennensis schucherti tenuibrachia sp. gracilis	itus	2?1000?010NN??150N0000N24?NNNNN0111401110 N 2?1000?0?0NN??15111011123?0N5011110401000 N 2???????????15111000N2510N5012217401000 N 2??????????15111000N2500N5013117401000 N 2?10?0?0?0NN??15111000N1500N502112?400000 N	1 15 021012100001010 1 35 000017000000010 1 5 002012000000011 1 5 07101000002011 1 5 000012100002011	01 5 111 01 12 010 00 N NN0 00 N NN1 00 N NN0	Z ONNZOO10??110 N ONN100?????? N ONN100??????? 1 ONN3001?110?? 1 ONN3001?110??

170

CRINOID MORPHOLOGY

CL Parisangulocrinu DI Calycanthocrinus	.zeaeformis decadactylus		2010?0?0?080NN1115111000N2500N501311?400000 2000?0?0?0NN??190N00111231NNNNN011?401000	zz		5 0020120N00010100 9 00100NNN00030100	ONN NNN N	N 0NN3001011077 N 0NN3007777777	
Middle Devonian DB Opsiocrinus DB Rhipidocrinus MO Thamnocrinus FL Taxocrinus CL Proctothylacocri CL Zostocrinus	mariana crenatus springeri lobatus .longus ornatus	31,63	2717060010NN771515131000N2500N5011100411000 2707007017777715131000N21N135001210501000 271070007010NN7715110000N230NNNNN1111401100 27107000707020NN7715121000N2310N510071401000 2710000070NN7715111000N2500N5013101400000 271706707777715111000N1500N5003010401010	zzzzz	-0	0 00100NNN1NN00101 0 021012101NN00101 0 001012111NN00101 5 0010120N01010101 5 0010120N00000100 5 020017??000?0100	4 111 4 111 2 1110 N NN0 N NN0	1 0NN300????17 1 0NN2000NNN10 1 0NN20011100710 1 0NN30010?001? N 0NN30010110?? ? 0NN70017207?	
<u>Upper Devonian</u> MO Lenneocrinus CL Hallocrinus	ventanillensis ornatissimus		2710000071777715110000N230NNNNN1717401110 2710077010NN7715111000N2500N5012117401000	z z		0 0010120N1NN10?01 5 0010121100100100	4 111 N NN1	2 ONN2007777777 1 ONN3001711077	
<u>Tournaisian</u> DB Cribanocrinus DB Gilbertsocrinus DB Rhodocrinites	wortheni typus kirbyi	58, 142 142 142	27700070700NN7715131000N2500N5011310501000 271000707070NN7715131000N2500N5011010511000 20100070700NN0N1515151000N2500N5011200501110	2223	-0-0	0 0220110N1NN00101 0 0720120N1NN00301 0 0010110N1NN0030101	4 4 111 4 4 111 7 111	1 ONN200777710 7 ONN7000NNNN10 1 ONN2000NNNN10 2 ONN2000NNNN10	
MO Abatocrinus MO Actinocrinites MO Agaricocrinus	macbridei multiradiatus inflatus	141, 142 142 142	2010005020NN9215110000N230NNNN1100401000 2210002020NN9215110000N230NNNN1112401010			0 00100NN1NN00101	4 111	7 ONN3001177110	
MO Aorocrinus MO Aryballocrinus	immaturus whitei	142 142	2010007070NN015110000N230NNNN1107401000 2710007010NN715110000N230NNNNN1217401000	z z :		0 00100NN1NN00101 0 0020120N1NN00101	4 111 4 111		
MO Auliskocrinus MO Batocrinus	crassitestus subaequalis	13, 142 77, 142	2710707070707070777715110000N220NNNNN1220401000 271700707777715110000N130NNNNN1100401000	z z :	~~~	00100001010100 2 0010000101010100 2	N N N N N N N N N N N N N N N N N N N	1 0002001110110	
MO Cactocrinus MO Cusacrinus	<pre>proboscidiali. nodobrachiatu.</pre>	142	2210002020NN7715110000N230NNNN11107401010 2210002020NNNN15110000N230NNNN1110401010			00100NN1NNN00101 00100NN1NNN00101	111 111 111	1 0NN3001120110	
MO Cytidocrinus MO Dialutocrinus	sculptus aculeatus	55,142 169	2710007070NN7715110000NZ30NNNN1117401010 2010007070NN7715110000NZ30NNNN1700401010	zz	- ~	0 02201210101010101 0 0210101111000101	4 111 4 111	1 0NN2001110110	
MO Dichocrinus MO Displodocrinus	cinctus divergens	13,82 79,142,156	2710007070101511000002200000000000000000	z z 0 0		5 0020100N00000100 5 002012111NN00201	NNN 1 111 1	7 0NN3001220111	
MO Dorycrinus	unicornis remibrachiatu	142	2210002020NN2215110000N230NNNN1100401000 2210002020NN2215110000N230NNNN1600401000	2 Z	~~~	4 00100NNN1NN00101 0 00100NNN1NN00201	4 111 6 111	1 0NN3000NNNN10 1 0NN30011???10	
MO Eucladocrinus	millebrachiat.	142	2200010020NN??150N0000N21NNNNN0210501000			5 001012101NN00200	N NN1	1 0NN3000NNN10	
MO Eutrochocrinus MO Macrocrinus	christyi vernenilianus	142	271000202020070771511000001320000000222200060200002220000602000022000000			8 00100NNN1NN00101	4 111	1 0NN3001110110	
MO Megistocrinus	nobilis	142	27177077777777715110000N230NNNN1270401000	z :		C 0010120N1NN00101	9 111	1 0NN7000NNN10	
MO Paradichocrinus MO Deveatocrinus	liratus ventricosus	13 12 142	U10102007070000151100000023000001050000000000000000	z z		0 00100NN1NN00101	7 111	1 0NN2000NNN10	
MO Platycrinites	burlingtonens.	79, 142	2710007070NN77150N0000N231NNNNN0211501000	z	-	5 002012171NN00100	L NN N	1 0NN2001000710	
MO Plemnocrinus	beebei	61	2?0001?0?0NN??150N0000N21NNNNN0201501000	z •		5 002012101NN00200	SNN N	2 ONNZOCONNNN10	
MO Protacrocrinus	primitivus	71,91	2?????????????????????????????????????	- 2		0 002012101NN00201	N NN N	1 0NN2001122210	
MU Steganocrinus MO Strimolecrinus	pentagonus inornatus	13, 142	2010002020NN1115110000N220NNNN1211401000	: z		5 0020100N1NN00100	N NN1	1 0NN2001000110	
MO Strotocrinus	perumbrosus	7,142	27100070700N7715110000N230NNNNN1600401110	z =	112	0 07100NN1NN00101	12 111 5 111	1 0NN7000NNN10	
MO Sunwaptacrinus	brazeauensis	23	ΛΙ ΛΙ ΛΗΡΟΖΙ ΝΝΝΝΝΟΣΡΝΟΛΟΟΙ LCL222222020202022	z	-			ALLO LINDANNO I	

171

172

M. FOOTE

1 0NN2001110110	N 00010100100010	N 0NN20022222222	N ONN110777777	N 0NN100???????	N 0NN2007277717	N 0NN200107777	N 0NN1001020022	N 0NN1001000010	N 0NN200777777	1 ONN300???????	N 0NN200777777	1 0NN3001??10??	2 ONN3007777777	1 ONN5001710077	7 UNN5UU777777	2200101002002000 2	220012100CNN0 N		innining N	1 0NNS007777777	2 0NN3002222222	2 00010000 2	1 ONN3001?110??	N 0NN3001??00??	1 ONN3001?100??	1 ONN3001??00??	1 0NN300????????	? ONN3001?110??	1 ONN3001?10???	1 0NN3001?110??	1 ONN3001?110??	1 ONN3001?110??	2 ONN3002222222	N 0NN2001??00??	N 0NN3001?100??	N 0NN2001020072	N 0NN3001010000	N 0NN3001?100??		1 0NN1000NNN10	1 0NN3001110110	1 0110111002000 1	SINNNNNNNNNNNNN I	
10 111		2 110	N NNO	6 ONO	ONN N	2 110	12 110	3 110	3 110	N NN1	N NNO	NN N	NNN N	NN N					N NNC	LNN N	NN N	NNN N	NNN N	N NNO	NN N	N NN N	NN N	N NN N	NN N	NN N	NN N	N NN 1	inn n	NNN N	ONN N	N NNO	ONN N	ONN N		4 111	11 111	6 111		
B0 00000NNN1NN00101		5 0010120N0000101	5 0000120N00000100	20 00000NNN00000101	5 0020120N00000100	5 0020121000000101	40 0000100N00000101	10 0010120N00000101	10 0010121100000101	5 0000110N00000100	4 0710121000010100	5 0010120N00100100	5 0200120N00000100	5 0010100N00020100	5 0100121100000100	5 0000121100110100	5 0010120N00020100	2 00101777700000100	5 0020120N00020100	5 0010100N00120100	5 0200110N00010100	5 0200121100000100	5 00100NNN00120100	5 0010111000110100	5 0010100N00110100	5 0010100N00120100	5 0000100N00000100	5 00100NNN00120100	5 0220120N00110100	5 0000100N00120100	5 0010120N00010100	5 0010110N00110100	5 0200120N00010100	5 0210121000110100	53 20000NNN00020100	3 0200121000020100	5 00000NNN00020100	5 0210121000110100		10 0010120N1NN00301	26 00100NNN1NN00101	30 0210110N1NN00101	12 00100NNINNN00200	20 10000001 NUNNUNNUN
л. Х	- •	- . z z		: _	z	r r	л Г	r z	- v	Л	۲ ۲	, Z	r z	r z	z	z	- · z :	z	z	- z	, z	с Х	r z	r Z	z	, N	z Z	z L	z	z	r L	х Г	r L	л Г	N 1	N 1	z	N L	:	N L	z	z	z	20 1
271000707000N777151100000N230NNNNN1110401010	2200000070NN7715110000N150NNNNN16104U10U0	2000007010NN77151210U0N250UN51011117400000	210100/010N/??15121000N2310N51122102401000	210007010002150010000231005000201000	2000002020NN22150N1000N21N0N5000200501000	21000020200NV2215121000N2310N5100102401000	2000002010NN??15121010N2310N5100200401000	2200002010NN2215121000N2300N5100200401000	2100002010NN22150N1000N2310N5000101501000	221000202000N2215111000N2500N5013112401000	22222620222222215111010N2500N5013112401000	200000070NN??15111000N2500N5013112401000	2?1000?0?0NN??15111000N2500N501311?401000	2????0?0??????15111000N2??20N501311?401000	27777077777777715111000N2500N5013107401000	271000707010007151110000250005013107401000	2210202020NN2215111000N2500N5013112401000	2210002020101215111000N2500N5013112400000	220020222200002500N2215111000N2500N5011310401000	20100020200002215111000N2222005013202401010	2710007077777715111000N2770N5013207411000	2????0?????????15111000N2500N5013201411000	20000020700NN7715111000N2500N5023117401000	2210006021122215111000N2500N5012112401000	2010002021112215110000N2220N5111202401000	201000202000N2215111000N2500N5013112401000	22000020200N2215111000N2500N5111102401000	220000702000022151110000150005013112401000	220006202000N022111000N22220N5013202401000	2210062020NN2215221000N2500N5012112401000	2200000010NN2215111000N2500N5013112400000	2712202022222215111000N2520N5013112401010	222222062222222215111000N222202401000	2710007070NN7715110000N050NNNN1027401000	2200002020NN22150N0010N231NNNN0100100000	220000202080221400001122310000012101000	2210000020150000000231000000000000000000	2210062021122215110000N150NNNN1212401000		2010002010NN??15131000N2500N5011010511000	22100020200NN2215110000N230NNNN1110401000	271000707070NN?715110000N230NNNNN1117401010	2710007070NN7715110000N277NNNN1002411000	2?1000?0?0NN??15110000N220NNNNN15Z0401001
	42										Q	61			04,160	04,160			69			3.104		72 82	141	22 84	5						51	140.160		103	201	160		142	142	38,142	142	142
142	1.62	101	5	55	5101	101	101	101	101	2	0 12	88	169	2	43.1	33,1	r	2	26.1	R	2	32.4	19	7.6		7	2	140	2 2 2	. 99	1,00	56		202	6	2 2	5 8	120	ĥ	138	138	6,1	82,	91,
umbrosus	pyriformis	fletcheri	nobilis	konincki bualiastonone	uachemithi	diversus	nobilis	intermedius	thiemei	madisonensis	robustus	iesuni	conicus	fairvensis	bursaeformis	ventricosus	warreni	jeffersonensi.	planus	brazeauensis	logani	elegans	inuensis	sculntilis	londicirrifer	maccahai	ni + i due	ni ciuus dactvi aidae	uactytotaes incignic			superhus	dranil osus		turbinatus	doctvi ue	uacty.us upphemiithi	flonifor flonifor		tiherosiis	arandis	gibsoni	splendens	amphora
10 Teleiocrinus	10 Uperocrinus	-L Eutaxocrinus	L Forbesiocrinus	FL Mespilocrinus	TL MELICIILIIYUUI IIIUS Il Ninternorinis	T Dovehocripus	si Parichthvocrinus	El Tavorrinus	si Wachsmithicrinus	al Arbelecribule	de Aprieteer muss of Atalastocriphis	ol Riothrocrinus	ri Bollandocrinus	cl Bridgerocrinus	CL Cercidocrinus	CL Coeliocrinus	CL Cradeocrinus	cL culmicrinus	ri Cvathocrinites	cl Decadocrinus	ri Dipotocrinis	ri Fratocrinis	ct tilmocrinus	cL di unoci indo	ct volcocrinus	ct nutcoti mus	ct Nootooninus	ct Nactocrinus ci Oatimooniano	ct upniurocrinus	ct Phonologniaus	ct Flacetoci illus Ci Botoriocrinites	CL FOLELIOCI III CES	CL 303 (1 01 0C1 11 143	ut belemnorrinus	DI Detemiou indo	DI CALITIOCITTUS	UI Hatystoch Hius Di Gimbathaanimin	UI SYNDALNOCTINUS	DI WUITEOCTINUS	<u>Visean</u> DB_Gilbertsocrinus	MO Abatocrinus	MO Actinocrinites	MO Agaricocrinus	MO Amphoracrocrinus

CRINOID MORPHOLOGY

0NN3007277777 ONN2001110110 0NN2007777777 0NN2000NNN77 ONN2001110210 0NN7000NNNN10 0NN7000NNN10 ONN2001110110 0NN3001000110 ONN2007777777 ONN2000NNN77 ONN2007777777 ONN1001112122 ONN1007777777 ONNO107777777 ONN2001000010 ONN1001000012 ONN3001111077 ONN3001210022 ONN200777777 DNN3001?100?? 0NN2007272727 DNN3001222022 ONN30011100?? 0NN300??????? ONN?001????1? ONN2001??00?? 00120022 ONN3001711077 0NN1007777777 ONN2001070077 0NN100???????? 000003222222 ONN1001020022 0NN300???????? 0NN7207272727 ONN200100??10 0NN3001?110?? 0NN3001?1???? 0NN3001?10??? 0NN2002222222 0NN3007777777 0NN300777777 ONN200777777 0NN200272727 0NN3001110073 NN NN **NN1** NN1 110 NN0 NNO ς NN NN 2NN NNO 111 NN SNN? 111 NN NN **NN** NN NN NN. NN1 **NN1** NN Ň N NN NN Ň 002012121NN00100 0020110N1NN10100 0010100N00110100 0010100N00110100 00100NN1NN00101 0010100N0001010C 0010100N0000010C 0000100N1NN0010101 00100NN1NN00101 02100NN1NN00201 00100NN1NN00101 0710100N1NN0010C 00100NN1NN00101 0200100N00000100 0000120N00010100 0210120N00100100 0200120N00100100 0210100N00120100 0210100N00000100 0200100N00000100 00700NNN1NN00100 0000120N00010100 0010121000010100 02002222222222200 0?20????00010100 0000120N00100100 0010120N00110100 0200120N00000100 0020120N00010100 0010121100100100 0010100N00110100 0010110000110100 0200121100000100 02001???1NN00100 0210110N00110100 00100NN1NN00101 101011101NN00101 0000120N00000101 0000120N0000010' 0000171700000101 0000121100000101 001012110100010' 0020121000000101 0010120N00000101 0010120N00000010C 0710100N00120100 **6**16 20 <u>2005</u> 20 ខ្ល 20 20 9 29 20 2 ഗഗ z 2?1??000?????15110000N2??NNNNN1200401000 210100?0?0NN??15110000N220NNNN121?401000 201000?0?0NN??15110000N220NNNN121?401000 2?0000?0?0NN??15110000N130NNNN1112401000 2?10?0?0?0?0NN??15110000NZ20NNNN1111401000 22070070700N7715121010N2770N5101210501000 2?0000?010NN??15121000N2??0N5100200501000 2000002010NN0N15121000N2310N5100200401000 2?10006010NN??15110000N130NNNN110?401000 2?100000?0NN??15110000N230NNNN110?401100 2?1000?0?0NN??15110000N130NNNN111?401000 2?1000?0?0NN??15110000N220NNNN1112401000 2?1000?010NN??150N0000N231NNNNN0201501000 2?1000?0?0NN??15110000N220NNNN122?401010 2?00?0?0?080N??150N1000N2310N5000200501000 2??0?0?0?0NN??15121000N2310N5101210401000 2000007010NN??15121000N2310N5102210401000 2?????0????????150N1000N21N0N5010200501000 2?0100?0?0NN??15121000N2310N5100200401000 2000007010NN?715121000N2310N5100202401000 2?0000?0?0NN??15121000N23?0N511010?401000 2?1006?010NN??15111000N2??0N501320?401000 2????0??1?????15111000N250135103001411000 2?1??0????????15111000N2??0N5102200411000 ONNNNNNNNNN15111000N11N0N502331040N000 2???00?0??????15111000N2500N501111?401000 2?1?0060?????15111000N2500N5013110401000 2?1000?010NN??15111000N2500N5003101401000 20100020200002500N22500N5013112401000 2210060010NN??15111000N2??0N501320?411010 2?1000?0?0NN??15111000N2500N5012202401000 2????0??1?????15111000N21N135000102501000 2?1000?0?0NN??15121000N1500N511211?400000 2000007070NN7715110000N220NNNN1217401001 2210202020N2215111000N1500N5021112401000 2?0000?0?0NN??15111000N2500N501311?401000 2????0??1?????15111000N25?0N5023200411000 2?1000?0?0NN??15111000N2500N501121?401000 2?1000?0?0NN??15111000N2??0N5013102401000 2?1??0?0?111??15111000N2500N5001200411000 22102020200N2215111000N2220N5013200411000 2???00?0??????15111000N2500N5012201411000 2??????????????15111000N2??0N5013200411000 2????0?????????15111000N2500N501311?401000 2?1000?0?0NN??15111000N2500N5003201411000 2210007070NN7715111000N2770N5013200411000 43,81,88,104 91,104 13,104,142 39, 127 74, 138, 140 36, 138, 166 59,93,140 myelodactylus 104,142 101,169 38, 142 13, 138 93,169 37, 138 38, 140 24,122 27 52,124 8,140 13, 14 20,45 36,60 37,82 4,138 36,51 7,85 mississippien.142 169 69 redesdalensis 169 69 69 69 169 multibrachiat.101 20 5 auatuordecimb.42 multibrachiat. trautscholdi pol ydactylus maxvillensis originarius cantonensis Springeracrocrin.intermedius spiniferus biserialis asteriscus commaticus subovatus springeri compactus concinnus tielensis scoticus diabolus mundulus strenuus gracilis pringlei anomalus bullatus ^ragilis rotundus georgii striatus coxanus grandis ulrichi ulrichi nodosus dunlopi epidus obatus sorosus curtus ungula unicus mundus bellus Artichthyocrinus Parichthyocrinus Paradichocrinus Strimplecrinus Forbesiocrinus Agassizocrinus Carlopsocrinus Cyathocrinites Harmostocrinus **Platycrinites** Caldenocrinus Hyrtanecrinus Mespilocrinus Anartiocrinus Cryphiocrinus Aenigmocrinus Anemetocrinus Corythocrinus Cosmetocrinus Eupachycrinus Dizygocrinus Enascocrinus Camptocrinus **Onychocrinus** Abrotocrinus Ampelocrinus Aphelecrinus Forthocrinus Dinotocrinus Cymbiocrinus Uperocrinus Amphicrinus Gaulocrinus Macrocrinus Adinocrinus Dichocrinus Cestocrinus Cydrocrinus Eratocrinus Fifeocrinus Dorycrinus Aorocrinus Batocrinus Taxocrinus Aulocrinus Barycrinus ş ę £ ş ş ş ş ş £ £ ð ş £ Ľ Ч 교 ц ដ ರ ರ 5 Ч ರರ ರರರ ರರ 555555 ц ដ ರರ

173

174

1 0NN3001711777 1 0NN3001711777 1 0NN3001711777 1 0NN300171777777 1 0NN300171077 1 0NN300171077 1 0NN3001710077 1 0NN3001710077 1 0NN3001710077 1 0NN3001710077 1 0NN3001710077 1 0NN3001710077 1 0NN30017107777777 2 0NN3001710177777777777777777777777777777	
NNNNNNNNNNNNNNNNNNNNNNNNNNNNNNNNNNNNNN	
<pre>5 021010000110100 5 00101212100000100 5 021010000010010 6 0210100000100100 8 0010120100100100 5 0000120000100100 5 0000120000100100 5 000012000001000 5 0000120000010100 5 000010000010100 5 000010000010100 5 000010000010100 5 000010000010100 5 00001000001000 5 00001000001000 5 000012000000100 5 000012000000100 5 000012000001000 5 000012000000100 5 000012000000100 5 00001000001000 5 000012000000100 5 000012000000100 5 000012000000100 5 0000100000100 5 000012000000100 5 000012000000100 5 000012000000100 5 000012000000100 5 000012000000100 5 000012000000100 5 000012000000100 5 000012000000100 5 0000120000000000000000000000000 5 00001200000000000000000000000000000000</pre>	
Z?1000?0?0NN??15111000N2500N501321?401000 Z?1000?0?0NN??15111000N2500N501320?401010 Z?177070?7?77715111000N250N501320?401010 Z?10707010N17715111000N2770N5013207411000 Z?10707010NN7715111000N2770N5013207411000 Z?10707010NN7715111000N2770N5013207411000 Z?10707010NN7715111000N2770N5013207411000 Z?10707010NN7715111000N2770N5013207411000 Z?10707010NN7715111000N270N5013207411000 Z?100000700NN7715111000N2500N5013207411000 Z?1000007070NN7715111000N2500N5013207411000 Z?10007070NN7715111000N2500N5013207411000 Z?7777077777715111000N2500N500N513207411000 Z?7777077777715111000N2500N570013207411000 Z?710007070NN7715111000N2500N513207411000 Z?710007070NN7715111000N2500N513207411000 Z?710007070NN7715111000N2500N513207411000 Z?710007070NN7715111000N2500N513207411000 Z?710007070NN7715111000N2500N513207411000 Z?710007070NN7715111000N2500N513207411000 Z?710007070NN7715111000N2500N513207411000 Z?710007070NN7715111000N2500N513207411000 Z?710007070NN7715111000N2500N513207411000 Z?710007070NN7715111000N2500N57013217401000 Z?710007070NN7715111000N2500N5701320411000 Z?710007070NN7715111000N2500N5701320411000 Z?700007070NN7715111000N2500N5701320411000 Z?700007070NN7715111000N2500N5701320411000 Z?000007070NN7715111000N2500N5701320411000 Z?000007070NN7715111000N2500N5701320411000 Z?000007070NN7715111000N2500N5701320411000 Z?000007070NN7715111000N2500N5701320411000 Z?000007070NN7715111000N2500N5701320411000 Z?000007070NN7715111000N2500N5701320400000 Z?000007070NN7715111000N2500N570110000200000002 Z700007070NN7715111000N2500N5500N501100000000000000000	
138 169 36, 138 36, 138 36, 138 46 35, 138 35, 99, 138 37, 79, 138 37, 79, 138 37, 79, 138 37, 79, 138 37, 79, 138 169 169 169 17, 20 36 17, 20 17, 20 17, 20 17, 20 102, 138 102, 138 102, 138 87, 103	
coreyi amplus briareus hoveyi hirmani seminulum depressus grandis praemorsus minutus aequalis validus konincki konincki konincki tyriensis agnatus bulgeri agnatus bulgeri ildesdalensi magnliaeform austinii .carpenteri bradleyi nodosus	
CL Histocrinus CL Hydreionocrinus CL Hydreionocrinus CL Hypselocrinus CL Intermediacrinus CL Lageniocrinus CL Lanecrinus CL Lebetocrinus CL Lebetocrinus CL Lophocrinus CL Parazeacrinites CL Parazeacrinites CL Parazeacrinus CL Secrinus CL Scotiacrinus CL Staphylocrinus CL Staphylocrinus CL Staphylocrinus CL Staphylocrinus CL Staphylocrinus CL Vecorinus CL Vecorinus CL Staphylocrinus CL Staphylocrinus CL Vecorinus CL Vecorinus CL Vecorinus CL Vecorinus CL Vecorinus CL Vecorinus CL Staphylocrinus CL Vecorinus CL Vecori	Serpukhovian

142	2710007010NN??15110000N220NNNN1220401001	20 1	20	1010100N1NN00201	5 ON1		0NN2000NNN1
	2111002020NN2215110000N220NNNN1212401000	۲ ۲	ŝ	0010100N00100100	N NN N	~	0NN20077777
	2220202222222215110000N220NNNN1102401000	۲ ۲	20	20100NN1NN00101	3 ON1	-	0NN2000NNN1
	22222202010NN2215111000N2210N5011200401000	r L	Ś	0000100N0000100	ONN N	z	.¿¿???????????????????????????????????
	2000002010NN2215111000N2310N5101202401000	r 7	Ś	0010121000000101	1 110	z	0100200102202
	22000020200NN2215131000N2310N5000200401000	N L	20	0010110N0001010101	5 110	z	0NN100107001
• 0•	2222000012222215111000N2520N5003200411000	N N	Ś	0000121100000100	NN N		0NN30012122
	2010002021112215111000N2500N5001202401000	۲ N	Ś	0010110N00110100	N NN N		.¿¿¿??????????????????????????????????
. 4	22222222222222222215111000N2500N5013110401000	л Г	Ś	0200100N00000100	N NN1		.¿¿???????????????????????????????????
	271776777777715111000N2500N5013102401000	N L	Ś	0010120N00110100	N NN1	-	22222022222
.110	221220202222215111000N2520N5012300411000	۲ ۲	Ś	0000100N1NN00100	NN N	~ ·	0NN20077777
	2210002020NN2215111000N1500N5013112401000	۲ ۲	Ś	0210170N00110100	NN N	-	00N300121002
	2010006071222715111000N2220N5001200411000	N L	Ś	0010100N00110100	INN N.	-	0NN30077777
	2?1?00?0??????15111000N2??0N5013200411000	л Г	Ś	0200110N1NN00100	NN N	-	0NN30077777
23	2?1?07001?????15111000N25?0N5013201411010	N N	Ś	0210120N00110100	NN N	<u>~</u> .	0NN30012122
ł	2710007070NN7715111000N277135003207401000	N.	Ś	000012121NN00100	N NN 1	¢.	0NN300111127

2	ukhovian		11 10 10
<	crocrinus	snumardi	24, 71, 14
C	amptocrinus	cirrifer	20,104
۵.	terotocrinus	acutus	24
C	ibolocrinus	africanus	134
0	nychocrinus	pulaskiensis	101
-	axocrinus	whitfieldi	24,101
4	lcimocrinus	ornatus	20,109
<	mpelocrinus	kaskaskiensis	52,124
<	nartiocrinus	lyoni	45,104
<	phelecrinus	randolphensis	20,56
ß	ronaughicrinus	figuratus	19,64,11
C	ulmicrinus	vagulus	24,104
C	ymbiocrinus	grandis	24,57
ш	upachycrinus	boydi i	42,80
Ŧ	lel iosocrinus	aftonensis	110,123
Ŧ	lydre i onocr i nus	goniodactylus	169

CRINOID MORPHOLOGY

? 0NN3001117777 2 0NN3007777777 1 0NN3007777777 1 0NN3001770077 ? 0NN3001710077 ? 0NN30017100777 ? 0NN3001711777	7 0NN3001710077 1 0NN3001717777 1 0NN3001717777 7 0NN3001717777 7 0NN3001110177 7 0NN3001710177 1 0NN3007777777 1 0NN300777777777777777777777777777777777	1 0NN300777777 2 0NN300777777 2 0NN3007777777 2 0NN2007777777 1 0NN3007777777 1 0NN3007777777 1 0NN3007777777 1 0NN3007777777 1 0NN3007777777 2 0NN3007777777 1 0NN3007777777 2 0NN300777777777777777777777777777777777	N 0NN000777777 N 0NN200107077777 7 0NN3001117777 7 0NN3001117777 1 0NN3001011077 7 0NN300111077 1 0NN300111077777 7 0NN3001000077 7 0NN3001777777 1 0NN30077777777 1 0NN300777777777777777777777777777777777	1 0NN200777777 2 0NN3001117777 1 0NN3001717077 7 0NN3007777777 1 0NN3007777777 7 0NN30077777777777777777777777777777777
		N N N N N N N N N N N N N N N N N N N	N N N N N N N N N N N N N N N N N N N	N NN1 N NN1 N NN1 N NN1 N NN1 N NN1
<pre>5 000120N0000100 5 0010120N000020100 5 0010120N00000100 5 00101000001100 5 0010100N000011010 5 0001000001000 5 00001000001000 5 000010000001000</pre>	5 0010121100100100 5 00101000000100 5 00001210100000100 5 02001211100000100 5 020011000000100 5 020011000000100 5 000011000000100	<pre>5 0000100N1NN00100 5 0000100N10N00100 5 0000100N10N00100100 5 0000100N10N00000100 5 0000100N00000100 5 0000100N00000100 5 0000100N00000100</pre>	<pre>5 0700110N00000100 0 0010121100000101 5 0200100N1NN00100 5 001012110000010100 5 0700100N1NN00100 5 0700100N1NN00100 5 0700100N100000100 5 0000100100100100 5 0000100N0100100 5 0000100N0100100 5 0000100N0100100</pre>	5 000100N1NN00100 5 000120N0000100 5 0010110N00110100 5 000100N00000100 5 0710120N00000100 5 0000100N1NN00100
			~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	
2????000??????5111000N2500N5003101401000 N 2?????????????515111000N2500N5001200411000 N 2??????????????515111000N2500N501370?401000 N 2????????????????515111000N2500N5013210411000 N 2?10005010NN??15111000N2500N5013110401000 N 2?100070??????515111000N2500N5013210?401000 N 2?100070???????15111000N2500N5013210?401000 N	20100/07/07/08/27151110000827208/01507401000 N 27777077777775151110000826008501300401000 N 2777707777777151110000826008501300401000 N 2717766707777771511100008260085013013200401000 N 271776070777771511100008250085013117401000 N 2717000707018/7151110008250085013107400000 N 271700707078877777151110008250085013107400000 N	2????0????????515111000N2500N5012300411000 N 2????0??????515111000N2570N5012300411000 N 2????0?01?????515111000N2570N5011200411000 N 2????0??????515111000N2500N5013110401000 N 2????0?0??????515111000N2500N5013110401000 N 2????0?0???????515111000N2500N5011210401000 N 0NNNNNNNNN151111000N2500N5011210401000 N 2??????????????????111000N2500N5011210401000 N 2????????????????????????????????????	2?????????????????????????????????????	Z717007017777715111000N2500N5012200401000 N 27770707717777715111000N2500N5013200401000 N 2010006010NN?715111000N2500N5013202411000 N 2717000017?7?711000N2500N5013200401000 N 2710007070NN?715111000N2500N5013110401000 N 2777777777777777777777777777777777777
110,116 19 70,24 220,24 24,34 224,34	24,55 20 20,111 34,44,104,155 104,169 34,88,104	89, 106, 128 43, 88, 106, 128 89, 92, 130 92, 118, 151 32, 150 28 92 28	90 101 135 115 135 116 170 170 135 135 135 135	11, 116 3, 170 170 170 90
aptus ormandi stereosoma gracilis longidactylus maniformis campanulus	spinosus delicatus aftonensis casspes spinosus chesterensis macrodactylus wortheni	oklahomaensis girtyi croneisi gillumi defendus spicata tarri glenisteri crassibrachia.	punctatus incurvus konecnyorum toddi bulbosus bassus regulatus simplex ornatus grossus acutus kickapoensis	holdenvillens. spiniferus trautscholdi geminatus multiplex brachiatus
L Paianocrinus L Paracymbiocrinus L Pelecocrinus L Pentaramicrinus L Phanocrinus L Phanocrinus	L Knopocrinus L Scammatocrinus L Scytalocrinus L Telikosocrinus L Tholocrinus L Ulrichicrinus L Woodocrinus L Zeacrinites	ashkirian L Aglaocrinus L Alcimocrinus L Diphuicrinus L Mathericrinus L Myrnowcrinus L Oxynocrinus L Paragassizocrinus L Proallosocrinus L Scytalocrinus	oscovian L Cibolocrinus L Synerocrinus L Anchicrinus L Anchicrinus L Anobasicrinus L Araeocrinus L Cricocrinus L Cromyocrinus L Dicromyocrinus L Birmocrinus L Graffhamicrinus L Geofhamicrinus L Lecobasicrinus	L Mttacromyocrinus L Mttaperimestocri L Miatshkovocrinus L Mooreocrinus L Moscovicrinus L Neoprotencrinus
0000000000		៙៲៰៰៰៰៰៰៰៰៰៰		00000000

~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~	0 ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~	
2222	7272 7272 1102 7272 7272 7272 7002 7272 727	2102 2722 2222 2222 2222 2222 2222 2222
01222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 02222 022222 02222 022222 02222 02222 02222 02222 022222 02222 022222 022222 022222 022222 022222 022222 022222 022222 022222 022222 022222 022222 022222 0222222	002222000121120002222000002222000000000	012100221
	NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300 NN300	N30 N30 N30 N30 N30 N30 N30 N30 N30 N30
		588888888
~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~		~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~
Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	* * * * * * * * *
010 0000 0000 0020 0020 0020 0000 0000		10000000000000000000000000000000000000
		22000000000000000000000000000000000000
		012(2)2(2)
22000000000000000000000000000000000000		
~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	νυνανανανανανανανανανανανα	ດເບັບເບັບເບັບ
222222222222222		
4011 4011 4011 4011 4011 4011 4011 4011	2001 2001 2001 2001 2001 2001 2001 2001	
2200011022000110220001102200001102200001102200001100001100000110000011000001100000110000		22010 22010 22010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20010 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 20000 2000000
013 003 003 003 003 003 003 003 003 003	NU000000000000000000000000000000000000	
00000000000000000000000000000000000000		
22222222222222222222222222222222222222		NZ5 NZ5 NZ5 NZ5 NZ5 NZ5 NZ5 NZ5 NZ5 NZ5
22222222222222222222222222222222222222	00000000000000000000000000000000000000	222 222 222 222 222 222 222 222 222 22
2222 2120000000000000000000000000000000	22222222222222222222222222222222222222	222 222 222 222 222 222 222 222 222 22
	<b>,                                    </b>	<u>,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,</u>
135	. M	
	132	<u>8</u> 6.
1170 1170 1135 1135 90, 1135 1135		221 221
90, 111, 111, 111, 111, 111, 111, 111, 1	288886666666666666666666666666666666666	
<b>(A) (A)</b>	s s cr	sis is
sor cus tus lus sis nis sis us	icias sec sector ens sector ens sector ens sector ens sector ens sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector sector s s s s s s s s s s s	i rus us us
at in a second s	and the second sec	tus ire ife us ili gon
and the second s	nuc nuc nuc nuc nuc nuc nuc nuc	oli cgu urt irg exa
	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	ŬĔŎŴŨ>Ĕ
snu snusnus snus	suntinution survey surv	sn sn sn s
s crii inu: inu: inu: inu: inu:	nus nus nus nus s s s s s s s s s s s s	us rin rin rin inu
inu: voci inu: voci inu: voci inu: voci inu: voci inu: voci inu: voci inu: voci inu: voci inu: voci voci voci voci voci voci voci voci	Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contraction Contra	rin ric cri toc toc toc
zea ocr ist ist f l l l l l l c in c in c c in c c i c c c c c c c c	<u> </u>	xoc uso ble ble nop
Scharter Scharter Scharter Scharter	CD CC CC CC CC CC CC CC CC CC	Prol Ste
		ರರರರರರ

CL Texacrinus CL Ulocrinus DI Kallimorphocrinu	irradiatus convexus lasallensis	114, 132 133 133	Z?1000?010NN??15111000NZ5?0N5013200411000 N 1 5 0010 Z?0000701110??15111000NZ500N501Z31?401000 N 1 5 0010 Z???0000?????150N0000NZ31NNNNN0302400000 N 1 8 12000	110N00100100 100N1NN00100 0NNN00030100	NN NN1 NNN NN1 NNN NN0	1 ONN3007777777
Permian-1 MO Neocamptocrinus FL Cibolocrinus FL Nevadacrinus FL Trampidocrinus CL Arroyocrinus CL Eribatocrinus CL Elibatocrinus CL Elibatocrinus CL Erisocrinus CL Microcaracrinus CL Microcaracrinus CL Moscovicrinus CL Moscovicrinus CL Neozeacrinus CL Neozeacrinus CL Stuartwellercrinus CL Stuartwellercrinus CL Stuartwellercrinus CL Stuartwellercrinus CL Stuartwellercrinus CL Tundracrinus CL Tundracrinus CL Stuartwellercrinus CL Stuartwellercrinus CL Stuartwellercrinus CL Stuartwellercrinus CL Stuartwellercrinus CL Stuartwellercrinus CL Stuartwellercrinus	millerensis patriciae geniculatus phiala popenoei cuneatus ekpansus elongatus longwelli moorei colubrosus inornatus bipinnatus tourteloti wanneri piutae corbatoi permicus	164 21,22 68,155 68,155 68,155 68,155 68,155 68,155 68,155 68,155 68,155 119,155 119,155	2?20001702110?715110000N220NNNN1210401000 N 1 10 0710 2?720070177771511000N2310N5001200401000 N 1 5 0710 2100007010NN?715111000N2310N5001201401000 N 1 5 0710 2100007010NN?715111000N2770N5013217401000 N 1 5 0200 277200777777715111000N2570N5013217401000 N 1 5 0200 2772007010NN?715111000N2570N50131702411000 N 1 5 0200 2710000700NN?715111000N2570N50131702411000 N 1 5 0200 2710000700NN?715111000N2500N5013117401000 N 1 5 0010 2710000700NN?715111000N2500N5013117401000 N 1 5 0000 2710000700NN?715111000N2500N5013117401000 N 1 5 0710 2710000700NN?715111000N2500N5013117401000 N 1 5 0700 2710000700NN?715111000N2500N5013117401000 N 1 5 0700 2710000700NN?715111000N2500N5013117401000 N 1 5 0700 2710007070NN?715111000N2500N5013117401000 N 1 5 0700 2710007070NN?715111000N2500N5013117401000 N 1 5 0700 27700707077777715111000N2500N5013117401000 N 1 5 0700 277007070NN?715111000N2500N5013117401000 N 1 2 2200	777707010101 100N00000100 120N0101010101 170N010101010101 170N0100110200 170N0110200 170N0110100100 17212000010100 12210000010100 12210000010100 12210000010100 12210000010100 12210000010100 12210000010100 12210000010100 12210000010100 12210000010100 12210000010100 12210000010100 12210000010100 12210000010100 12210000010100 12210000010100 12210000010100 1221000000100 1221000000100 1221000000100 1221000000100 1221000000100 1221000000100 1221000000100 1221000000100 1221000000100 1221000000100 1221000000100 1221000000100 1221000000100 1221000000100 1221000000100 12210000001000 12210000001000 12210000000100 122100000000		<pre>&gt; ONN370ONNNN00 &gt; ONN3007777777777777777777777777777777777</pre>
Permian-2 MO Dichocrinus? MO Neocamptocrinus MO Stomiocrinus FL Ancistrocrinus FL Calycocrinus FL Loxocrinus FL Petrocrinus FL Petrocrinus FL Prophyllocrinus FL Prophyllocrinus FL Syntomocrinus FL Syntomocrinus CL Abrachiocrinus CL Abrachiocrinus CL Cacriaiocrinus CL Coenocystis CL Coenocystis CL Cosmetocrinus CL Cosmetocrinus CL Dichostreblocrin CL Embryocrinus	australis wardenensis feruginus depressus curvatus globulus beyrichi torynocrinoid. dentatus sundaicus timoritus clavulus clavulus abundans concinnus angulosus middalyaensis timorensis hanieli	165 165 154 154 154 147 143 144 144 145 153 153 155 153 154 154 154 154 154 154	Z??000?0?0?070777777515110000NZZ0NNNN1717401000 N 1 5 0?201 Z???700?7?7777515110000NZZ0NNNN11710401000 N 1 5 07101 Z????057777777775150N10100NZ20NNNN1117401000 N 1 5 07101 Z????0671777775150N1010NZ2N0N51100100000 N 1 5 02001 Z???7077067177777150N1010NZ1N0N511010100000 N 1 5 02001 Z???70777777775150N1010NZ2N0N50113104010000 N 1 5 02001 Z???7077777777775150N10100NZ1N0N5011111501000 N 1 5 02001 Z?7770777777777775150N1000NZ1N0N5011111501000 N 1 5 00001 Z?7770777777777775150N1000NZ1N0N5011111501000 N 1 5 00001 Z?7777077777777775150N1000NZ1N0N5010310511000 N 1 5 00001 Z?7777077777777777150N1000NZ310N5001300511000 N 1 5 00001 Z?777007777777777715000NZ510N5001300511000 N 1 5 00001 Z?77700777777777710NN1NNNN310N50N0214501000 N 0 N NNNNN Z?71005071777777710NN1NNNNN1N0N50N0214501000 N 0 N NNNNN Z?7100077777777710NN1NNNN310N50N0214501000 N 0 N NNNNN Z?710007777777710NN1NNNNN1NNNS10N50N0214501000 N 0 N NNNNN Z?710007777777710NN1NNNNN310N50N0214501000 N 0 N NNNNN	7???00110100 17??000101010 17??0000101010 170000010100 110000000100 110000000100 11000000		<ul> <li>0NN70077777777</li> <li>0NN70077777777</li> <li>0NN700077777777</li> <li>0NN700077777777</li> <li>0NN70077777777</li> <li>0NN70077777777</li> <li>0NN100077777777</li> <li>0NN100077777777</li> <li>0NN100077777777</li> <li>0NN1000077777777</li> <li>0NN1000077777777</li> <li>0NN10000077777777</li> <li>0NN10000077777777</li> <li>0NN100000177777777</li> <li>0NN1000000777777777777</li> <li>0NN100000007777777777777777777777777777</li></ul>

## CRINOID MORPHOLOGY

OOO N O N NNNNNNNNNNNNNN N O N NNN	010 N 1 5 0?20120N00010100 N NNO N 0NN?001???	000 N 1 5 0000120N00020100 N NN? ? 0NN30011700	OOO N O N NNNNNNNNNNNNNN N O N NNN	000 N 1 5 00100NNN00110100 N NN1 1 0NN30010000	000 N O N NNNNNNNNNNNNNN N ONNN N ONNNNOONNNN	000 N D N NNNNNNNNNNNNNN N DNNN N DNNNOONNNN	000 N 1 5 0200121100000100 N NN1 4 0NN30011102	010 N 1 5 0710121000110100 N NNO N 0NN30010070	000 N 1 5 0212122222202100 N NN2 2 0NN22010222	000 N 1 5 020022200000100 N NN2 2 0NN20022722	000 N 1 5 0000120N00000100 N NN? ? 0NN300?????	000 N 1 5 0000120N00030100 N NN? ? 0NN300?????	000 N 1 5 00001???00020100 N NN? ? 0NN30010101	000 N 1 5 0210121200010100 N NN1 1 0NN30022722	000 N 1 5 0010120N00110100 N NN1 1 0NN3002222	000 N 1 5 0000122200000100 N NN2 2 0NN30012122		000 N 1 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	1000 N N N N N N N N N N N N N N N N N N	CCCCCUUCHNU C CHN N COLOCODOCCOCC /C F H CCC	1000 N 1 24 2700777700020100 N NN: 2 00070011111	TOUD N N N NUCLOANNANDADANA N NN N NN N NN NANDADA N NN			ООО N 1 TO U?TUT???ОООООТОТ Z TT? ? ОМИ??ООИИИ ОООО N 1 5 5 ООООТОЛИОЛОООТОТ N NNO N ONNOOO??????	NUNNNNNNNNNNNNNNNNNNNNNNNNNNNNNNNNNNNN	1000 N 1 5 0710170N00110100 N NN1 1 0NN3007777	1000 N O N NNNNNNNNNNNNNN N NNN N ONNNNNNNN
ONNNNNNNNNNNN 10NN 1NNNNN 1NON 50N00105NN	22222202022272215111000N2310N5011212401	2222220222222222215111000N2500N5011200411	27720071777777710NN1NNNNN310N50N0210501	27100020200NN2215111000N2500N5021312411	222222222222222222220NN1NNNNN310N50N0101501	22222222222222222222222222222222222222		22111000001110111110000000000000000000	2/1000/01/1/2/2/2/2/11/10/00/2/2/2/2/2/2	222220062122222215111000N2500N5011301401	2222200012222215011000N2500N5010201511	22222022222222222222222222222222222222	22222222222222222222222222222222222222	222220201222222215111000N2520N5013200611	2111100011111111000112111100011210112200111	20120120120002000001110122222222222222					2201002020NN??715UNUU1UNZ??NNNNNUUUU1U1	10c1010NNNN%22N000001222220000002122			2000012021107715110000N220NNNNN120740U	COTICS COCNOL CONDULITED 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5	22222062012222221511000N2500N5011200411	22222222222222222221513100000010005020111501
л к С	163	143_149		137		10	2,0	140	105	871 771 271	147, 144, 140	144,147	1/8	140		154	145, 149	145	143	.152,154,162	144,146	144,152	162		164	66 70	50 154	2 66
harbari	voisevi	numilus	pumitus abrachiatum	bostocki	mavimie	ilian Ililus ahi ahamanai a	snicnanensis	peramptus	ornatus	austratis	. exornatus	illa La Lanus immetrisrie		ber III cus	granutosus	australis	timoricus	sulcatus	mirabilis	.callytharraen	.granulatus	campanulatus	cephal onus		bundanooens i s	tunisiensis	rıchardsoni cherrehinensi	variahilie
· conilidiocrimus	L EOPILIAIOUTINAS	d Granhiocrinus:	i lemistrenterron	u nemistreptation 1 Jimharrinus		L Lagentocrinus	L Neolageniocrinus	L Neozeacrinus	L Nowracrinus	L Occiducrinus	L Paragraphiocrinu.	L Parastacnyocrinu.	L Permiocrinus	L Prolobocrinus	CL Skarocrinus	CL Spheniscocrinus	CL Tapinocrinus	CL Tenagocrinus	CL Timorechinus	<b>Motiocatillocrin</b> .	<b>DI Paracatillocrinu</b>	<b>DI Synbathocrinus</b>	01 Taucatillocrinus	bermian-3	40 Neocamptocrinus	FL Trinalicrinus	CL Coenocystis	ct tympioti inus 21 Embryocripie

.

### APPENDIX 2: Sources for Morphological Data on Crinoid Species

- 1. Arendt, Yu. A. 1963. Krona morskoy lilii iz srednego Ordovika r. Podkamennoy Tunguski. Paleontologicheskiy Zhurnal, 1963(4): 131-136.
- Arendt, Yu. A. 1970. Morskie lilii gipokrinidy. Akademia Nauk SSSR, Trudy Paleontologicheskogo Instituta, 128: 1-220.
- Arendt, Yu. A. 1981. Trekhrukie morskie lilii. Akademia Nauk SSSR, Trudy Paleontologicheskogo Instituta, 189: 1-195.
- 4. Ausich, W. I. 1983. Functional morphology and feeding dynamics of the Early Mississippian crinoid *Barycrinus asteriscus*. Journal of Paleontology, 57: 31-41.
- Ausich, W. I., and T. W. Kammer. 1990. Systematics and phylogeny of the late Osagean and Meramecian crinoids *Platycrintes* and *Eucladocrinus* from the Mississippian stratotype region. Journal of Paleontology, 64: 759-778.
- 6. Ausich, W. I., and T. W. Kammer. 1991a. Late Osagean and Meramecian Actinocrinites (Echinodermata: Crinoidea) from the Mississippian stratotype region. Journal of Paleontology, 65: 485-499.
- Ausich, W. I., and T. W. Kammer. 1991b. Systematic revisions to Aorocrinus, Dorycrinus, Macrocrinus, Paradichocrinus, Strotocrinus, and Uperocrinus: Mississippian camerate crinoids (Echinodermata) from the stratotype region. Journal of Paleontology, 65: 936-944.
- 8. Ausich, W. I., and T. W. Kammer. 1992. *Dizygocrinus*: Mississippian camerate crinoid (Echinodermata) from the midcontinental United States. Journal of Paleontology, 66: 637-658.
- 9. Ausich, W. I., and N. G. Lane. 1982. Crinoids from the Edwardsville Formation (Lower Mississippian) of southern Indiana. Journal of Paleontology, 56: 1343-1361.
- 10. Bowsher, A. L. 1955. New genera of Mississippian camerate crinoids. University of Kansas Paleontological Contributions, Echinodermata, Article, 1: 1-23.
- Branson, C. C. 1962. Additional illustrations of some Oklahoma crinoids. Oklahoma Geology Notes, 22: 162-163.
- 12. Brett, C. E. 1981. Terminology and functional morphology of attachment structures in pelmatozoan echinoderms. Lethaia, 14: 343-370.
- 13. Broadhead, T. W. 1981. Carboniferous camerate crinoid subfamily Dichocrininae. Palaeontographica, Abt. A, 176: 81-157.
- 14. Broadhead, T. W., and H. L. Strimple. 1980. *Hyrtanecrinus*, a new Carboniferous camerate crinoid genus from eastern North America. Journal of Paleontology, 54: 35-44.
- 15. Brower, J. C. 1965. The genus, Steganocrinus. Journal of Paleontology, 39: 773-793.
- 16. Brower, J. C. 1994. Camerate crinoids from the Middle Ordovician (Galena Group, Dunleith Formation) of northern Iowa and southern Minnesota. Journal of Paleontology, 68: 570-599.
- 17. Burdick, D. W., and H. L. Strimple. 1969. Revision of some Chesteran Inadunate crinoids. University of Kansas Paleontological Contributions, Paper, 40: 1-14.
- Burdick, D. W., and H. L. Strimple. 1971. Faunal studies of the type Chesteran, Upper Mississippian of southwestern Illinois, Part 2. Crinoids from the Beech Creek Limestone, lower Golconda Group, St. Clair County, Illinois. University of Kansas Paleontological Contributions, Paper, 51: 15-47.
- 19. Burdick, D. W., and H. L. Strimple. 1973. New Late Mississippian crinoids from northern Arkansas. Journal of Paleontology, 47: 231-242.
- Burdick, D. W., and H. L. Strimple. 1983. Genevievian and Chesterian crinoids of Alabama. Geological Survey of Alabama, Bulletin, 121: 1-277.
- 21. Burke, J. J. 1975. A new Permian Cibolocrinus from Bolivia. Kirtlandia, 20: 1-7.
- 22. Burke, J. J., and R. K. Pabian. 1978. Two crowns of *Cibolocrinus patriciae* (Crinoidea, Flexibilia) from the lower Permian of Bolivia. Journal of Paleontology, 52: 1065-1069.
- Carpenter, P. H., and R. Etheridge, Jr. 1881. Contributions to the study of the British Paleozoic crinoids.-No. 1. On Allagecrinus, the representative of the Carboniferous limestone series. Annals and Magazine of Natural History, series 5, 7: 281-298.
- 24. Chesnut, D. R., Jr., and F. R. Ettensohn. 1988. Hombergian (Chesterian) Echinoderm paleontology and paleoecology, south-central Kentucky. Bulletins of American Paleontology, 95(330): 1-102.
- Donovan, S. K., and C. J. Veltkamp. 1993. Crinoids from the Upper Ashgill (Upper Ordovician) of Wales. Journal of Paleontology, 67: 604-613.
- Donovan, S. K., N. T. J. Hollingworth, and C. J. Veltkamp. 1986. The British Permian crinoid 'Cyathocrinites' ramosus (Schlotheim). Palaeontology, 29: 809-825.

- 27. Ettensohn, F. R. 1975. The autecology of Agassizocrinus lobatus. Journal of Paleontology, 49: 1044-1061.
- 28. Ettensohn, F. R. 1980. Paragassizocrinus: systematics, phylogeny and ecology. Journal of Paleontology, 54: 978-1007.
- 29. Frederickson, E. A., and D. E. Waddell. 1960. An unusual crinoid from the Pennsylvanian of Oklahoma. Oklahoma Geology Notes, 20: 172-174.
- 30. Girty, G. H. 1908. The Guadalupian fauna. United States Geologicl Survey Professional Paper, 58: 1-512.
- 31. Goldring, W. 1954. Devonian crinoids: new and old, II. New York State Museum Circular, 37: 1-51.
- 32. Hall. J. 1858. Palaeontology of Iowa. Report of the Geological Survey of Iowa, 1(2): 473-724.
- 33. Hall. J. 1861. Descriptions of new species of Crinoidea from the Carboniferous rocks of the Mississippi valley. Boston Society of Natural History Journal, 7: 261-328.
- 34. Horowitz, A. S. 1965. Crinoids from the Glen Dean Limestone of southern Indiana and Kentucky. Indiana Geological Survey Bulletin, 34: 1-52.
- 35. Jaekel, O. 1895. Beiträge zur Kenntnis der paläozoischen Crinoiden Deutschlands. Palaeontologishe Abhandlungen, new series, 3: 1-116.
- Kammer, T. W., and W. I. Ausich. 1992. Advanced cladid crinoids from the middle Mississippian of the east-central United States: primitive-grade calyces. Journal of Paleontology, 66: 461-480.
- 37. Kammer, T. W., and W. I. Ausich. 1993. Advanced cladid crinoids from the middle Mississippian of the east-central United States: intermediate-grade calyces. Journal of Paleontology, 67: 614-639.
- Kammer, T. W., and W. I. Ausich. 1994. Advanced cladid crinoids from the middle Mississippian of the east-central United States: advanced-grade calyces. Journal of Paleontology, 68: 339-351.
- 39. Kirk, E. 1929a. Cryphiocrinus, a new genus of free-swimming crinoids. American Journal of Science, series 5, 17: 153-171.
- 40. Kirk, E. 1929b. The fossil crinoid genus Vasocrinus Lyon. Proceedings of the United States National Museum, 74(15): 1-4.
- 41. Kirk, E. 1937a. *Clistocrinus*, a new Carboniferous crinoid genus. Journal of the Washington Academy of Sciences, 27: 105-111.
- 42. Kirk, E. 1937b. *Eupachycrinus* and related Carboniferous crinoid genera. Journal of Paleontology, 11: 598-607.
- 43. Kirk, E. 1938. Five new genera of Carboniferous Crinoidea Inadunata. Journal of the Washington Academy of Sciences, 28: 158-172.
- 44. Kirk, E. 1939. Two new genera of Carboniferous inadunate crinoids. Journal of the Washington Academy of Sciences, 29: 469-473.
- 45. Kirk, E. 1940a. Anartiocrinus, a new crinoid genus from the Mississippian. American Journal of Science, 238: 47-55.
- Kirk, E. 1940b. A redescription of Lageniocrinus de Koninck. American Journal of Science, 238: 129-139.
- 47. Kirk, E. 1940c. *Cestocrinus*, a new fossil inadunate crinoid genus. Proceedings of the United States National Museum, 88(3080): 220-224.
- 48. Kirk, E. 1940d. Lebetocrinus, a new crinoid genus from the Upper Borden of Indiana. Journal of Paleontology, 14: 74-77.
- 49. Kirk, E. 1940e. Seven new genera of Carboniferous Crinoidea Inadunata. Journal of the Washington Academy of Sciences, 30: 321-344.
- 50. Kirk, E. 1941a. Dinotocrinus, a new fossil inadunate crinoid genus. Proceedings of the United States National Museum, 89(3103): 513-517.
- 51. Kirk, E. 1941b. Four new genera of Mississippian Crinoidea Inadunata. Journal of Paleontology, 15: 82-88.
- 52. Kirk, E. 1942a. Ampelocrinus, a new crinoid genus from the Upper Mississippian. American Journal of Science, 240: 22-28.
- 53. Kirk, E. 1942b. *Rhopocrinus*, a new fossil inadunate crinoid genus. Proceedings of the United States National Museum, 92(3144): 151-155.
- 54. Kirk, E. 1942c. Sarocrinus, a new crinoid genus from the Lower Mississippian. Journal of Paleontology, 16: 382-386.
- 55. Kirk, E. 1943. A revision of the genus Steganocrinus. Journal of the Washington Academy of Sciences, 33: 259-265.
- 56. Kirk, E. 1944a. Aphelecrinus, a new inadunate crinoid genus from the upper Mississippian. American Journal of Science, 242: 190-203.

- 57. Kirk, E. 1944b. Cymbiocrinus, a new inadunate crinoid genus from the upper Mississippian. American Journal of Science, 242: 233-245.
- 58. Kirk, E. 1944c. Cribanocrinus, a new rhodocrinoid genus. Journal of the Washington Academy of Sciences, 34: 13-16.
- 59. Kirk, E. 1945. *Gaulocrinus*, a new inadunate crinoid genus from the Mississippian. Journal of the Washington Academy of Sciences, 35: 180-182.
- 60. Kirk, E. 1946a. Corythocrinus, a new inadunate crinoid genus from the Lower Mississippian. Journal of Paleontology, 20: 269-274.
- 61. Kirk, E. 1946b. *Plemnocrinus*, a new crinoid genus from the Lower Mississippian. Journal of Paleontology, 20: 435-441.
- 62. Kirk, E. 1947. Three new genera of inadunate crinoids from the Lower Mississippian. American Journal of Science, 245: 287-303.
- 63. Kirk, E. 1948. Two new inadunate crinoid genera from the Middle Devonian. American Journal of Science, 246: 701-710.
- 64. Knapp, W. D. 1969. Declinida, a new order of late Paleozoic inadunate crinoids. Journal of Paleontology, 43: 340-391.
- 65. Lane, N. G. 1975. The anal sac of *Aesiocrinus*, a Pennsylvanian inadunate crinoid. Journal of Paleontology, 49: 638-645.
- 66. Lane, N. G. 1979. Upper Permian crinoids from Djebel Tebaga, Tunisia. Journal of Paleontology, 53: 121-132.
- 67. Lane, N. G., and R. M. Howell. 1986. Unusual crinoids from the Ramp Creek Formation (Mississippian), Indian Creek, Montgomery County, Indiana. Journal of Paleontology, 60: 898-903.
- 68. Lane, N. G., and G. D. Webster. 1966. New Permian crinoid fauna from southern Nevada. University of California Publications in Geological Sciences, 63: 1-87.
- 69. Laudon, L. R. 1933. The stratigraphy and paleontology of the Gilmore City Formation of Iowa. University of Iowa Studies in Natural History, 15(2): 1-74.
- 70. Laudon, L. R. 1941. New crinoid fauna from the Pitkin Limestone of northeastern Oklahoma. Journal of Paleontology, 15: 384-391.
- 71. Laudon, L. R., and B. H. Beane. 1937. The crinoid fauna of the Hampton Formation at LeGrand, Iowa. University of Iowa Studies in Natural History, 17: 226-273.
- 72. Laudon, L. R., and J. L. Severson. 1953. New crinoid fauna, Mississippian, Lodgepole formation, Montana. Journal of Paleontology, 27: 505-536.
- 73. Laudon, L. R., J. M. Parks, and A. C. Spreng. 1952. Mississippian crinoid fauna from the Banff formation, Sunwapta Pass, Alberta. Journal of Paleontology, 26: 544-575.
- Lyon, S. S., and S. A. Casseday. 1859. Description of Nine new species of Crinoidea from the Subcarboniferous rocks of Indiana and Kentucky. American Journal of Science, series 2, 28: 233-246.
- Lyon, S. S., and S. A. Casseday. 1860. Description of Nine new species of Crinoidea from the Subcarboniferous rocks of Indiana and Kentucky. American Journal of Science, series 2, 29: 68-79.
- Marez Oyens, F. A. H. W. de. 1940. Neue permische Krinoiden von Timor, mit Bemerkungen über deren Vorkommen im Basleogebiet. Geological Expedition to the Lesser Sunda Islands under Leadership of H. A. Brouwer, 1: 285-348.
- 77. McChesney, J. H. 1868. Descriptions of fossils from the Palaeozoic rocks of the western states, with illustrations. Transactions of the Chicago Academy of Sciences, 1: 1-57.
- 78. Meek, F. B., and A. H. Worthen. 1866. Descriptions of invertebrates from the Carboniferous system. Illinois Geological Survey, 2(2): 143-411.
- 79. Meek, F. B., and A. H. Worthen. 1873. Palaeontology. Descriptions of invertebrates from Carboniferous system. Illinois Geological Survey, 5(2): 323-619.
- 80. Miller, S. A. 1879. Remarks upon the Kaskaskia group, and descriptions of new species of fossils from Pulaski County, Kentucky. Cincinnati Society of Natural History Journal, 2: 31-42.
- Miller, S. A. 1892. A description of some Lower Carboniferous crinoids from Missouri. Geological Survey of Missouri, Bulletin, 4: 1-40.
- 82. Miller, S. A., and W. F. E. Gurley. 1890. Description of some new genera and species of Echinodermata from the Coal Measures and Subcarboniferous rocks of Indiana, Missouri, and Iowa. Indiana Department of Geology and Natural History, Sixteenth Annual Report, p. 327-373.
- Miller, S. A., and W. F. E. Gurley. 1893. Descriptions of some new species of invertebrates from the Palaeozoic rocks of Illinois and adjacent states. Illinois State Museum of Natural History Bulletin, 3: 1-81.
- Miller, S. A., and W. F. E. Gurley. 1894. New genera and species of Echinodermata. Illinois State Museum of Natural History Bulletin, 5: 1-53.

- 85. Miller, S. A., and W. F. E. Gurley. 1896. New sepcies of crinoids from Illinois and other states. Illinois State Museum of Natural History Bulletin, 9: 1-66.
- 86. Moore, R. C. 1939. New crinoids from the Upper Pennsylvanian and Lower Permian rocks of Oklahoma, Kansas, and Nebraska. Denison University Scientific Laboratories Journal, 34: 171-294.
- 87. Moore, R. C. 1962a. Revision of Calceocrinidae. University of Kansas Paleontological Contributions, Echinodermata, Article, 4: 1-40.
- 88. Moore, R. C. 1962b. Ray structures of some inadunate crinoids. University of Kansas Paleontological Contributions, Echinodermata, Article, 5: 1-47.
- 89. Moore, R. C., and F. B. Plummer. 1938 [for 1937]. Upper Carboniferous crinoids from the Morrow subseries of Arkansas, Oklahoma and Texas. Denison University Scientific Laboratories Journal, 32: 209-313.
- 90. Moore, R. C., and F. B. Plummer. 1940. Crinoids from the Upper Carboniferous and Permian strata in Texas. University of Texas Publication, 3945: 1-459.
- 91. Moore, R. C., and H. L. Strimple. 1969. Explosive evolutionary differentiation of unique group of Mississippian-Pennsylvanian camerate crinoids (Acrocrinidae). University of Kansas Paleontological Contributions, Paper, 39: 1-44.
- 92. Moore, R. C., and H. L. Strimple. 1973. Lower Pennsylvanian (Morrowan) crinoids from Arkansas, Oklahoma, and Texas. University of Kansas Paleontological Contributions, Article, 60: 1-84.
- 93. Moore, R. C., H. W. Rasmussen, N. G. Lane, G. Ubaghs, H. L. Strimple, R. E. Peck, J. Sprinkle, R. O. Fay, and H. Sieverts-Doreck. 1978. Systematic Descriptions, Pp. T403-T1027 in R. C. Moore and C. Teichert, eds. Treatise on invertebrate paleontology, Part T, Echinodermata 2. Geological Society of America, Boulder, Colorado, and University of Kansas, Lawrence, Kansas.
- 94. Pabian, R. K., and H. L. Strimple. 1974a. Crinoid studies. Bulletins of American Paleontology, 64(281): 249-337.
- 95. Pabian, R. K., and H. L. Strimple. 1974b. Fossil crinoid studies. University of Kansas Paleontological Contributions, Paper, 73: 1-54.
- 96. Pabian, R. K., and H. L. Strimple. 1979. Notes on biometrics, paleoecology and biostratigraphy of Cibolocrinuis conicus Strimple from Oklahoma, Kansas and Nebraska. Journal of Paleontology, 53: 421-437.
- 97. Pabian, R. K., and H. L. Strimple. 1980. Some crinoids from the Argentine Limestone (Late Pennsylvanian--Missourian) of southeastern Nebraska and southwestern Iowa. Transactions of the Nebraska Academy of Sciences, 8: 155-186.
- 98. Prokop, R. J., and V. Petr. 1989. New Silurian and Devonian crinoid taxa from Bohemia. Časopis pro Mineralogii a Geologii, 34: 45-58.
- 99. Schmidt, W. E. 1930. Die Fauna des deutschen Unterkarbons, 1. Die Echinodermen. Abhandlungen der Preussischen Geologischen Landesanstalt, new series, 122: 1-92.
- 100. Springer, F. 1911. Some new American fossil crinoids. Harvard College Museum of Comparative Zoology Memoir 25: 117-161.
- 101. Springer, F. 1920. The Crinoidea Flexibilia. Smithsonian Institution Publication 2501: 1-486, plates A-C, 1-76.
- 102. Springer, F. 1923. On the fossil crinoid family Catillocrinidae. Smithsonian Miscellaneous Collections, 76(3): 1-41.
- 103. Springer, F. 1926a. American Silurian crinoids. Smithsonian Institution Publication 2871: 1-239.
  104. Springer, F. 1926b. Unusual forms of fossil crinoids. Proceedings, United States National Museum 67(5): 1-137.
- 105. Sprinkle, J. and G. P. Wahlman. 1994. New echinoderms from the Early Ordovician of west Texas. Journal of Paleontology, 68: 324-338.
- 106. Strimple, H. L. 1940. Some new crinoid species from the Morrow subseries. Bulletins of American Paleontology, 25(91): 1-11.
- 107. Strimple, H. L. 1947. Three new crinoid species from the Virgil series of southeastern Kansas. Bulletins of American Paleontology 31(124): 1-12.
- 108. Strimple, H. L. 1949a. Crinoid studies, parts III-VII. Bulletins of American Paleontology, 32(133): 1-42.
- 109. Strimple, H. L. 1949b. Studies of Carboniferous crinoids. Palaeontographica Americana, 3(23): 1-40.
- 110. Strimple, H. L. 1951a. New Carboniferous crinoids. Journal of Paleontology, 25: 669-676. 111. Strimple, H. L. 1951b. New crinoids from the Pitkin of Oklahoma. Journal of the Washington Academy of Sciences, 41: 260-263.
- 112. Strimple, H. L. 1951c. Notes on Phanocrinus cylindricus and description of new species of Chester crinoids. Journal of the Washington Academy of Sciences, 41: 291-294.

### CRINOID MORPHOLOGY

- 113. Strimple, H. L. 1952a. The arms of *Polusocrinus*. Journal of the Washington Academy of Sciences, 42: 12-14.
- 114. Strimple, H. L. 1952b. Notes on *Texacrinus*. Journal of the Washington Academy of Sciences, 42: 216-220.
- 115. Strimple, H. L. 1952c. The arms of Perimestocrinus. Journal of Paleontology, 26: 784-788.
- 116. Strimple, H. L. 1961. Late Desmoinesian crinoid faunule from Oklahoma. Oklahoma Geological Survey Bulletin, 93: 1-189.
- 117. Strimple, H. L. 1962. Crinoids from the Oologah Formation (Pennsylvanian), Tulsa County, Oklahoma. Oklahoma Geological Survey Circular, 60: 1-75.
- 118. Strimple, H. L. 1966a. New species of cromyocrinids from Oklahoma and Arkansas. Oklahoma Geology Notes, 26: 3-12.
- 119. Strimple, H. L. 1966b. Some notes concerning the Allagecrinidae. Oklahoma Geology Notes, 26: 99-111.
- Strimple, H. L. 1971a. Crinoids from the Vinland Shale (Virgilian) of Kansas. Journal of Paleontology, 45: 998-1000.
- 121. Strimple, H. L. 1971b. Fossil crinoid studies, Part 3. The occurrence of *Hydriocrinus* in Oklahoma and Russia. University of Kansas Paleontological Contributions, Paper, 56: 16-19.
- 122. Strimple, H. L. 1973a. Fossil crinoid studies, Part 3. Aenigmocrinus, a new Chesteran inadunate crinoid genus. University of Kansas Paleontological Contributions, Paper, 66: 15-18.
- 123. Strimple, H. L. 1973b. Fossil crinoid studies, Part 4. The inadunate crinoid genus *Heliosocrinus*. University of Kansas Paleontological Contributions, Paper, 66: 18-21.
- 124. Strimple, H. L. 1973c. Fossil crinoid studies, Part 5. Notes on Mississippian Ampelocrinidae. University of Kansas Paleontological Contributions, Paper, 66: 21-24.
- 125. Strimple, H. L. 1975a. Middle Pennsylvanian (Atokan) crinoids from Oklahoma and Missouri. University of Kansas Paleontological Contributions, Paper, 76: 1-30.
- 126. Strimple, H. L. 1975b. New Chesterian (Upper Mississippian) crinoids from Illinois. University of Kansas Paleontological Contributions, Paper, 79: 1-9.
- 127. Strimple, H. L. 1977. The inadunate crinoid genus Cryphiocrinus Kirk. Geological Magazine, 114: 209-214.
- 128. Strimple, H. L. 1982. An exceptional inadunate crinoid crown from the Morrowan (Lower Pennsylvanian) of Oklahoma. Oklahoma Geology Notes, 42: 200-203.
- 129. Strimple, H. L., and D. R. Boardman, II. 1971. Fossil crinoid studies, Part 6. Notes on Stenopecrinus and Perimestocrinus. University of Kansas Paleontological Contributions, Paper, 56: 27-30.
- 130. Strimple, H. L., and W. D. Knapp. 1966. Lower Pennsylvanian fauna from eastern Kentucky; Part 2, crinoids. Journal of Paleontology, 40: 309-314.
- 131. Strimple, H. L., and M. R. McGinnis. 1969. Fossil crinoid studies, Part 5. New crinoid from the Gilmore City Formation, Lower Mississippian of Iowa. University of Kansas Paleontological Contributions, Paper, 42: 21-22.
- 132. Strimple, H. L., and R. C. Moore. 1971a. Crinoids of the Francis Shale (Missourian) of Oklahoma. University of Kansas Paleontological Contributions, Paper, 55: 1-20.
- 133. Strimple, H. L., and R. C. Moore. 1971b. Crinoids of the LaSalle Limestone (Pennsylvanian) of Illinois. University of Kansas Paleontological Contributions, Article, 55: 1-48.
- Strimple, H. L. and C. Pareyn. 1982. *Cibolocrinus* from the Namurian of North Africa with notes on the genus. Journal of Paleontology, 56: 226-232.
- 135. Strimple, H. L., and W. T. Watkins. 1969. Carboniferous crinoids of Texas with stratigraphic implications. Palaeontographica Americana, 6(40): 141-275.
- 136. Teichert, C. 1949. Permian crinoid *Calceolispongia*. Geological Society of America Memoir, 34: 1-132.
- 137. Teichert, C. 1954. A new Permian crinoid from Western Australia. Journal of Paleontology, 28: 70-75.
- 138. Van Sant, J. F., and N. G. Lane. 1964. Crawfordsville (Indiana) crinoid studies. University of Kansas Paleontological Contributions, Echinodermata, Article, 7: 1-136.
- 139. Wachsmuth, C., and F. Springer. 1877. Revision of the genus *Belennocrinus* and description of two new species. American Journal of Science, series 3, 13: 253-259.
- Wachsmuth, C., and F. Springer. 1880-1886. Revision of the Palaeocrinoidea, parts 1-3. Academy of Natural Sciences of Philadelphia, Proceedings, 1879: 226-378 [part 1]; 1881: 175-411 [part 2]; 1885: 223-364 [part 3, section 1]; 1886: 64-226 [part 3, section 2].
- 141. Wachsmuth, C., and F. Springer. 1890. New species of crinoids and blastoids from the Kinderhook Group of the Lower Carboniferous rocks at LeGrand, Iowa, and a new genus from the Niagaran Group of Western Tennessee. Illinois Geological Survey, 8(2): 155-208.

- 142. Wachsmuth, C., and F. Springer. 1897. The North American Crinoidea Camerata. Harvard College Museum of Comparative Zoology Memoir 20: 1-897, 21: plates 1-83.
- 143. Wanner, J. 1916. Die permischen Echinodermen von Timor, Teil 1. Paläontologie von Timor, part 6, number 11: 1-329.
- 144. Wanner, J. 1924 [for 1921]. Die permischen Krinoiden von Timor. Jaarboek van het Mijnwezen in Nederlandsch Oost-Indië, Verhandelingen, 50(3): 1-348.
- 145. Wanner, J. 1929. Neue Beiträge zur Kenntnis der permischen Echinodermen von Timor. II. Hypocrinites. Dienst van den Mijnbouw in Nederlandsch-Indië, Wetenschappelijke Mededeelingen, 11: 36-117.
- 146. Wanner, J. 1930a. Neue Beiträge zur Kenntnis der permischen Echinodermen von Timor. III. Hypocrinidae, Paracatillocrinus und Allagecrinus dux. Dienst van den Mijnbouw in Nederlandsch-Indië, Wetenschappelijke Mededeelingen, 13: 1-31.
- 147. Wanner, J. 1930b. Neue Beiträge zur Kenntnis der permischen Echinodermen von Timor. IV. Flexibilia. Dienst van den Mijnbouw in Nederlandsch-Indië, Wetenschappelijke Mededeelingen, 14: 1-61.
- 148. Wanner, J. 1937. Neue Beiträge zur Kenntnis der permischen Echinodermen von Timor. VIII-XIII. Palaeontographica Supplement-Band IV, 4(2): 57-212.
- 149. Wanner, J. 1949. Neue Beiträge zur Kenntnis der permischen Echinodermen von Timor, XVI. Poteriocrinidae, 4. Teil. Palaeontographica Supplement-Band IV, 5(1): 1-56.
- 150. Washburn, A. T. 1968. Early Pennsylvanian crinoids from the south-central Wasatch Mountains of central Utah. Brigham Young University Geology Studies, 15(1): 115-132.
- 151. Webster, G. D. 1981. New crinoids from the Naco formation (Middle Pennsylvanian) of Arizona and a revision of the family Cromyocrinidae. Journal of Paleontology, 55: 1176-1199.
- 152. Webster, G. D. 1987. Permian crinoids from the type-section of the Callytharra Formation, Callytharra Springs, Western Australia. Alcheringa, 11: 95-135.
- 153. Webster, G. D. 1990. New Permian crinoids from Australia. Palaeontology, 33: 49-74.
- 154. Webster, G. D., and P. A. Jell. 1992. Permian echinoderms from Western Australia. Memoirs of the Queensland Museum, 32: 311-373.
- 155. Webster, G. D., and N. G. Lane. 1967. Additional Permian crinoids from southern Nevada. University of Kansas Paleontological Contributions, Paper, 27: 1-32.
- 156. Webster, G. D., and N. G. Lane. 1987. Crinoids from the Anchor Limestone (Lower Mississippian) of the Monte Cristo Group, southern Nevada. University of Kansas Paleontological Contributions, Paper, 119: 1-55.
- 157. Weller, S. 1909. Description of a Permian crinoid fauna from Texas. Journal of Geology, 17: 623-635.
- 158. Wetherby, A. G. 1880. Descriptions of new crinoids from the Cincinnati Group of the Lower Silurian and Subcarboniferous of Kentucky. Cincinnati Society of Natural History Journal, 2: 245-253.
- 159. Wetherby, A. G. 1881. Descriptions of crinoids from the upper Subcarboniferous of Pulaski County, Ky. Cincinnati Society of Natural History Journal, 3: 324-330.
- 160. White, C. A. 1865 [for 1862]. Description of new species of fossils from the Devonian and Carboniferous rocks of the Mississippi Valley. Proceedings, Boston Society of Natural History, 9: 8-33.
- 161. Whitfield, R. P. 1881. Description of a new species of crinoid from the Burlington lime stone, at Burlington, Iowa. Bulletin, American Museum of Natural History, 1: 7-9.
- 162. Willink, R. J. 1978. Catillocrinids from the Permian of eastern Australia. Alcheringa, 2: 83-102.
- 163. Willink, R. J. 1979. Some conservative and some highly-evolved Permian crinoids from eastern Australia. Alcheringa, 3: 117-134.
- 164. Willink, R. J. 1980a. A new coiled-stemmed camerate from the Permian of eastern Australia. Journal of Paleontology, 54: 15-34.
- 165. Willink, R. J. 1980b. Two new camerate crinoid species from the Permian of eastern Australia. Alcheringa, 4: 227-232.
- 166. Worthen. A. H. 1883. Description of fossil invertebrates. Illinois Geological Survey, 7: 269-322.
- 167. Worthen, A. H., and F. B. Meek. 1875. Descriptions of invertebrates. Illinois Geological Survey 6(2): 489-532.
- 168. Wright, J. 1937. Scottish Carboniferous crinoids. Geological Magazine, 74: 385-411.
- Wright, J. 1950-1960. A monograph of the British Carboniferous Crinoidea. Volumes 1 (parts 1-5), 2 (parts 1-5). Palaeontographical Society, London.
- 170. Yakovlev, N. N., and A. P. Ivanov. 1956. Morskie lilii i blastoidei kamennougol'nykh i permskikh otlozheniy SSSR. Trudy Vsesoyuznogo Nauchno-issledovatel'skogo Geologicheskogo Instituta, new series, 11: 1-142.