The influence of heat treatments on order-disorder phenomena in a naturally occurring manganese-ferrite

Cherie R. Bluncson and B. J. Evans University of Michigan, Department of Chemistry, Ann Arbor, Michigan 48109-1055

Naturally occurring samples of iwakiite, $Mn_{1.1}Fe_{1.9}O_4$, have been annealed for 17 and 77 h at 1000 K in an investigation of order-disorder phenomena in Mn-Fe oxide spinels. Two distinct magnetic hyperfine patterns, which correspond to Fe- and Mn-rich phases, were observed in the ⁵⁷Fe Mössbauer spectra for the two annealed samples. While the spectrum of the Fe-rich phase corresponds to a magnetically and crystallographically highly ordered α -Fe₂O₃-like phase, the Mn-rich phase exhibits considerable disorder, as indicated by very broad linewidths and a low value for $H_{\rm eff}$. Annealing for longer periods of times leads to a considerable increase in the ordering of the Mn-rich spinel phase, as evidenced by the considerable narrowing of the lines, from Γ =2.40 mm s⁻¹ to Γ =1.17 mm s⁻¹, and an increase in $H_{\rm eff}$ from 318 to 417 kOe at 298 K. The Fe-rich,

 α -Fe₂O₃-like phase is not affected by further annealing, as expected. At 85 K, the ⁵⁷Fe Mössbauer spectrum of the Mn-rich, spinel phase continues to exhibit rather broad lines, indicating the presence

I. INTRODUCTION

The $Mn_xFe_{3-x}O_4$ system continues to present challenges concerning structure—property relationships, cation distributions, oxidation states, crystallographic order—disorder, and the electric and magnetic properties of oxide spinels exhibiting both variable cation distributions and mixed valence states.¹⁻⁴

of significant cation disorder.

Previous investigations of an iwakiite sample with the composition, Mn_{1.7}Fe_{1.3}O₄, ⁴ a naturally occurring mineral in the Mn₃O₄-Fe₃O₄ solid solution series, have demonstrated the existence of an unusually high degree of long-range order, as evidenced by a tetragonal distortion at a manganese content well below the limiting concentration observed in synthetic samples. In addition, results of 57Fe Mössbauer spectroscopic measurements on this sample at 92 K in an applied field of 30 kG have shown that Fe3+ occupies exclusively the octahedral sites. 5 The occurrence of Mn3+ clustering and the resulting Mn-rich and Mn-poor (or Fe-rich) regions, have been suggested, on the basis of the observation of two Fe³⁺ species on a single magnetic sublattice. It was proposed that the two Fe³⁺ species correspond to those occupying Mn-rich and Mn-poor (or Fe-rich) regions of the crystal.

Further confirmation and delineation of the characteristics of the clustering suggested above has been sought in the study reported here by investigating the thermal stability of

TABLE I. 57Fe hyperfine parameters at 298 K for iwakiite samples.

Sample	Subpattern	H _{eff} (kOe)	$\frac{\Delta E_Q}{(\text{mm s}^{-1})}$	δ ^a (mm s ⁻¹)	$\Gamma \atop \text{(mm s}^{-1}\text{)}$	Area (%)
Iwakiite Unannealed	•••	388	-0.007	0.392	0.920	100
Annealed 17 h	Fe rich	500	-0.109	0.390	0.458	38
Annealed 17 h	Mn rich	318	-0.020	0.347	2.400	62
Annealed 77 h	Fe rìch	513	-0.104	0.373	0.343	30
Annealed 77 h	Mn rich	417	-0.009	0.383	1.17	70

^aIsomer shifts are relative to Fe metal.

the naturally equilibrated iwakiite and monitoring the local and long-range crystal/chemical structures as a function of laboratory-based heat treatments. It is expected that the degree of order observed in iwakiite does not correspond to thermodynamically stable structures at temperatures at which ionic motion is observable on laboratory time scales, and that mild heat treatments will bring about significant changes in both the local and long-range crystal/chemical structures.

II. EXPERIMENT

Pure iwakiite samples from the Gozaisho Mine, Iwaki City, Japan, were prepared by pulverizing whole-rock, hand specimens and grinding them with an agate mortar and pestle to a fine powder that passed a 400 mesh standard sieve. Iwakiite was isolated from the associated minerals by suspending the powder in organic liquids and separating the iwakiite by means of a manually manipulated rare-earth permanent magnet. This process was repeated until a pure spinel phase was obtained, as determined by x-ray diffractometry.

The chemical composition of the separated sample was determined by means of electron microprobe analysis.

The iwakiite sample was separated into two fractions; one fraction was annealed in an evacuated quartz ampoule for 17 h at 1000 K, and the remaining fraction was annealed under the same conditions for 77 h. Both samples were cooled rapidly to 298 K following the heat treatments.

TABLE II. 57Fe hyperfine parameters at 85 K for iwakiite samples.

Sample	Subpattern	$H_{ m eff}$ (kOe)	ΔE_{Q} (mm s ⁻¹)	δ^a (mm s ⁻¹)	Γ (mm s ⁻¹)	Area (%)
Iwakiite	•••	503	0.011	0.500	0.623	100
Unannealed						
Annealed 17 h	Fe rich	525	-0.097	0.486	0.449	41
Annealed 17 h	Mn rich	496	-0.003	0.492	0.817	59
Annealed 77 h	Fe rich	529	-0.050	0,460	0.384	24
Annealed 77 h	Mn rich	500	-0.005	0.476	0.731	76

^aIsomer shifts are relative to Fe metal.

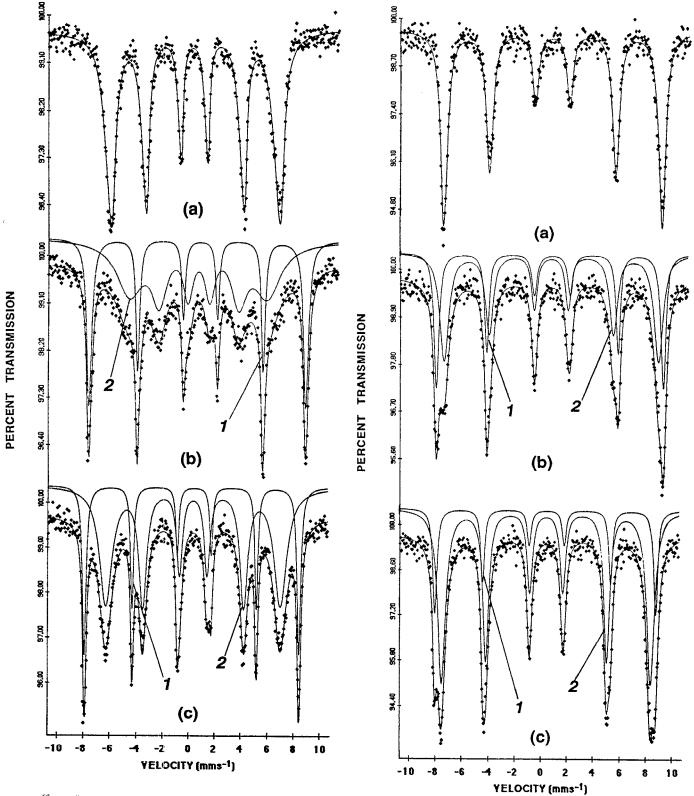


FIG. 1. 57 Fe Mössbauer spectra (298 K) of (a) unannealed iwakiite, (b) iwakiite annealed for 17 h at 1000 K, (c) iwakiite annealed for 77 h at 1000 K Fe-rich and Mn-rich regions are represented by 1 and 2, respectively.

FIG. 2. ⁵⁷Fe Mössbauer spectra (85 K) of (a) unannealed iwakiite, (b) iwakiite annealed for 17 h at 1000 K, (c) iwakiite annealed for 77 h at 1000 K Fe-rich and Mn-rich regions are represented by 1 and 2, respectively.

⁵⁷Fe Mössbauer spectra were obtained in a transmission geometry using a 512 channel, microcomputer-configured multichannel scaler and a constant acceleration electromechanical drive. Spectra were obtained with the absorber at

298 and 85 K in a vacuum cryostat. (See Tables I and II.) The source was always maintained at 298 K. The spectra were analyzed by means of a previously described least-mean-squares fitting procedure. ⁶ The filled dots in the spectra

are the experimental data and the solid lines are the results of the fitting procedure.

III. RESULTS AND DISCUSSION

The chemical composition of the iwakiite sample used in this investigation was found to be $Mn_{1.1}Fe_{1.9}O_4$. The manganese content is considerably lower than that of the previously studied sample, which had the composition $Mn_{1.7}Fe_{1.3}O_4$.^{4,5}

The 57 Fe Mössbauer spectrum at 298 K for the unannealed sample shown in Fig. 1 was fitted to one six line pattern corresponding to a magnetic hyperfine field, $H_{\rm eff}$, of 388 kOe. The samples annealed for 17 and 77 h were fitted to two six line patterns, where subpattern 1 with $H_{\rm eff}$ =500 kOe in Fig. 1(b), would be derived from the Fe-rich regions of the untreated iwakiite and subpattern 2, with $H_{\rm eff}$ =318 kOe, results from the Mn-rich regions of the untreated iwakiite. The parameters corresponding to these fits are given in Tables I and II.

The spectrum in Fig. 1(b), representing the sample annealed for 17 h can be understood as consisting of two components: subpattern 2, which is attributed to a poorly crystalline Mn-rich phase, as indicated by the large linewidths, Γ =2.40 mm s⁻¹, which has undergone phase separation from the unannealed iwakiite, and subpattern 1, which would correspond to a Fe-rich phase. From both the x-ray diffraction data and the 57Fe Mössbauer parameters, it is clear that the phase giving rise to subpattern 1 is hematite, α -Fe₂O₃. Upon further heating, as in the case of the 77 h anneal, there is considerable recrystallization and growth of the Mn-rich phase, giving rise to well-defined, long-range crystal/ chemical structures characteristic of the spinel structure. Clear evidence is provided for this by the remarkable decrease in the linewidths of subpattern 2 and the increase in the magnitude of $H_{\rm eff}$ from 318 to 417 kOe (cf. Table I). The width of the (404) line in the x-ray powder diffraction patterns undergoes a remarkable narrowing, in support of a decrease in the crystallographic disorder following the 77 h anneal at 1000 K.

Spectra of the iwakiite samples at 85 K exhibit similar trends to those obtained at 298 K, with the linewidths of subpattern 2, representing the Mn-rich phase, being broader for the sample annealed for 17 h than for the sample annealed for 77 h (cf. Table II).

A somewhat similar behavior has been observed for naturally occurring jacobsite, MnFe₂O₄, samples following

heat treatments between 675 and 1175 K.⁷ The findings of this investigation are quite similar to those reported for jacobsite, inasmuch as there is a facile separation of a highly ordered hematite-like phase. The phase-separated spinel phase in the annealed jacobsite exhibits considerably less crystallinity and magnetic order than observed in the case of annealed iwakiites. We believe this difference is due to the greater degree of clustering in the iwakiite samples than in the jacobsites; there was no evidence for clustering in the unannealed jacobsites.⁷

IV. CONCLUSION

Evidence for cation clustering and the occurrence of Mnand Mn-poor regions in naturally occurring Mn_xFe_{3-x}O₄ phases exhibiting high degrees of thermal equilibration has been provided by the ease with which such specimens undergo facile phase separations at temperatures well below those required for traditional solid state reactions. It is remarkable that the phase-separated, Fe-rich, α-Fe₂O₃-like phase is well ordered, even for short periods of thermal annealing; long-range order is well established for both the crystal and magnetic structures, as indicated by the close agreement between the hyperfine parameters of this pattern at 298 K and those for a pure hematite. By way of contrast, the Mn-rich spinel phase exhibits considerable cation disorder, even for very long annealing periods, and, as a consequence, has values of $H_{\rm eff}$ that are much lower than those seen in synthetic MnFe₂O₄. The persistence of this disorder is due, most likely, to insufficient thermal energy and or time for the Mn and Fe ions to achieve a narrow range of local cation configurations. The relative magnitude of the magnetic hyperfine fields are consistent with this interpretation.

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