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Instrumentation and Data Processing

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Instrumentation

Two recent developments in seismometry allow the study of seismic phenomena over a broad period range. *Block and Moore* [1970] developed a quartz accelerometer with the required stability and sensitivity to record seismic waves as well as earth tides. The instrument, especially useful at periods greater than 80 sec, recorded free oscillations of the earth associated with an earthquake of magnitude 6.5 [*Block et al.*, 1970]. *Pomeroy et al.* [1969] described three-component wide-band long-period seismographs that operate at maximum magnifications up to 500,000 in the period range of 30 to 60 sec. Relatively noise-free output at these magnifications was achieved by using rigid environmental control (i.e., an air-tight chamber in a deep mine). A displacement transducer output on the same instrumentation permits concurrent recording of the solid-earth tides.

A symmetrical triaxial seismometer, in which the orthogonal directions of response form equal angles (approximately 55°) with the vertical, minimizes some problems of natural period variation and allows matched response characteristics [*Melton and Kirkpatrick*, 1970]. *Farrell* [1969] described a gyroscopic seismometer consisting of two counter-rotating pendulous gyroscopes that permits the separation of tilts from horizontal displacements. The Ocean Bottom Seismometer (OBS) off the coast of northern California produced a variety of geophysical data [*Latham et al.*, 1967; *Latham and Nowroozi*, 1968; *Nowroozi et al.*, 1968; *Anderson and Latham*, 1969]. *Bradner* [1970] described measurements with a free-floating midwater seismometer. A system for counting seismic events was designed by *Decker* [1968]. *Mereu and Kovach* [1970] described a portable seismic system for crustal studies. Response characteristics and seismograph design criteria were discussed by *Rodgers* [1968], *Russell et al.* [1968], *Bollinger* [1969], *Mitchell and Landisman* [1969], *Reasenber* [1969], and *Rodgers* [1969]. *Rihn* [1969] presented data on the design of electromagnetic damping circuits.

Research in instrumentation for the measurement of

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earth strains produced several new instruments. A laser interferometer 1000 meters long was described by *Vali et al.* [1966, 1968], *Bostrom and Vali* [1968a, b], and *Vali and Bostrom* [1968a, b]. A laser strain meter with an 800 meter baseline was described by *Berger and Lovberg* [1969]. A general review of laser developments for measurement of long distances was presented by *Bender* [1967]. Errors in electronic distance measurements were discussed by *Thompson and Janes* [1967]. A diamagnetic suspension tiltmeter [*Simon et al.*, 1968; *McConnell and Simon*, 1969; *Simon and Strong*, 1968] is now being adapted for borehole use. The use of confined fluids in the earth as strain measuring devices was discussed by *Bodvarsson* [1970] and *Bredehoeft* [1967]. *Shopland and Kirklin* [1969, 1970] discussed the discrimination of seismic waves through the use of surface extensometers. *Hade et al.* [1968] presented a strain calibration system that utilizes a laser interferometer.

A history of the development of seismometry up to 1900 was published by *Dewey and Byerly* [1969].

Data Processing

Significant progress was made in improving existing techniques or in introducing new ones for processing of seismological data to elucidate evidence for tectonic processes of the earth, to infer earth structure, to explain features of the seismograms, and to determine hypocentral data for earthquakes more precisely.

Ben-Menahem et al. [1968] described a computer program for fast reduction of multistation surface-wave and body-wave data to determine focal mechanisms of earthquakes. *Ben-Menahem et al.* [1970] published universal tables for fast determination of focal mechanisms from isolated surface-wave signals. *Kanamori* [1970] used syntheses of long-period surface waves to infer focal mechanisms. *Tsai and Aki* [1970] analyzed surface waves at a small number of stations to determine seismic moment and dislocation at the focus, provided the focal mechanism of the earthquake was known independently. *Bollinger* [1968] used amplitude spectra of *P* waves to infer time duration of the earthquake.

Comprehensive seismicity maps based on U.S. Coast and Geodetic Survey epicenter data were constructed by *Barazangi and Dorman* [1969, 1970] and by *Tarr* [1970]. *Levy et al.* [1970] described a technique for displaying earthquake occurrence in space and time in a motion-picture format.

Landisman et al. [1969], *Dziewonski et al.* [1969], *Bloch and Hales* [1968], and *Savage* [1969] discussed new techniques for obtaining surface-wave dispersion curves from seismograms. Universal dispersion tables were computed by *Anderson and Harkrider* [1968] for

Love waves and by *Anderson and Kovach* [1969] for free oscillations of the earth. *Randall* [1967], *Watson* [1970], and *Schwab and Knopoff* [1970] proposed schemes for rapid and precise computation of theoretical dispersion curves. *Backus and Gilbert* [1967] and *Dziewonski* [1970] investigated the uniqueness of models obtained from inversion of seismological data. *Chander et al.* [1968] determined leaking-mode dispersion curves from synthesis of the dispersed waves associated with *S*, *SS*, and *SSS* phases, on the basis of the shear-coupled *PL*-wave hypothesis.

Tsai and Aki [1970] proposed a scheme for determining focal depths of seismic events from analyses of surface waves at a small number of stations, provided that the focal mechanism of the event was known independently. *Freedman* [1967] presented a modified scheme for computing mean body-wave magnitude for an earthquake from the magnitudes reported by individual stations.

Arrays

Numerous geophysical studies were based on the seismological data obtained from the Large Aperture Seismic Array (Lasa) in eastern Montana. By using various data-processing techniques [e.g., techniques of *Capon et al.*, 1967], signal-to-noise ratio was improved by $N^{1/2}$ (where N is the number of sensors used) for short-period seismic signals [*Capon et al.*, 1968] and by 12–20 db for long-period body and surface waves [*Capon et al.*, 1969]. *Capon et al.* [1969] showed that surface waves from earthquakes with body-wave magnitude (m_b) greater than 4.9 at epicentral distances between 50–60° were detected at Lasa. *Capon* [1970] argued that, in some cases, the surface waves recorded at Lasa had followed paths other than the great-circles to the epicenters.

The structure of the seismic background noise at Lasa was investigated by several authors. *Lacoss et al.* [1969] and *Toksöz and Lacoss* [1968] showed that, at periods smaller than 5 sec, the microseisms consisted of body waves and higher-mode surface waves. *Haubrich and McCamy* [1969] showed that the peaks centered at 7 and 14 sec in the energy spectra of microseisms consisted predominantly of fundamental-mode Rayleigh waves, with contributions in some cases from Love waves, higher-mode Rayleigh waves, and long-period body waves. In both studies the body waves were associated with distant storms over open oceans. *Capon* [1969] demonstrated that noise in the 20–40 sec period range consisted of Rayleigh waves and nonpropagating noise generated by local atmospheric conditions.

Mack [1969] showed that crustal variations within Lasa were responsible for variations in the *P*-wave signals at different sensors and that delayed summation of individual sensor outputs could yield a better estimate of the *P*-wave pulse at the base of the crust. *Glover and Alexander* [1969] presented further evidence of lateral variations in the crust under Lasa. Amplitude corrections were computed by *Syed* [1969], and dipping interface corrections for $dT/d\Delta$ were computed by *Zengeni*

[1970]. Applications of arrays were discussed by *Evernden* [1969] and *Baker* [1970]. Array design was discussed by *Haubrich* [1968]. *Johnson* [1968] outlined one problem in array utilization. *Niazi* [1968] suggested the use of source arrays to study regional crustal structure.

Johnson [1967], using inversion of $dT/d\Delta$ measurements [discussed by *Husebye*, 1969] on array data, presented a *P*-wave velocity structure for the upper 750 kilometers of the mantle, with a low-velocity zone between 70 and 150 kilometers and a high-velocity gradient near the lower boundary of the zone. *Johnson* [1969] presented data that indicated the presence of several other regions in the mantle with high-velocity gradients.

Strain Measurements

In earlier seismological studies of underground explosions, the tacit assumption was made that any release of pre-existing tectonic strain occurred in a small region (a few square kilometers) surrounding a nuclear explosion. Recently, the possibility was investigated that large explosions can affect tectonic strain and trigger earthquakes at distances up to a few hundred kilometers. *Smith et al.* [1969b] estimated from extensometer records that at Isabella, 260 km from the site of the Boxcar explosion ($m_b = 6.3$), the static strain was less than 2×10^{-10} and concluded that significant strain changes associated with large explosions are confined to distances less than 100 km. *Smith et al.* [1969a] reported a strain step of 1.2×10^{-7} at a distance of 29 km from Benham ($m_b = 6.3$) and concluded that permanent and quasi-static strains could affect occurrence of local earthquakes to a distance of 15 km from the explosion. From a series of strain measurements in conjunction with Benham, *Romig et al.* [1969] reported step-like residual strains of 1.8×10^{-7} and 0.35×10^{-7} at distances of 28 and 71 km, respectively, from the explosion; the strains decayed to about zero within half an hour. The strains measured by *Romig et al.* are comparable to the strains observed from earthquakes of similar body-wave magnitude at comparable distances, but the earthquake strains decay over several days, if at all. *Dickey* [1969] reported on strain measurements from geodetic observations. *McKeown and Dickey* [1969] and *Houser* [1969] discussed explosion-triggered motions on faults adjacent to the explosions.

Wideman and Major [1967] concluded from extensive measurements of residual strain steps associated with earthquakes in the 3–8.5 magnitude range, that strain steps of about 10^{-9} may be expected at distances of 20 times the fault length. The strain steps propagate at 3.3 ± 0.3 km/sec under continents and at 3.6 km/sec under oceans; the amplitude decreases with distance according to the $R^{-3/2}$ law.

Based on repeated precision leveling, *Tryggvason* [1968, 1970] presented evidence of vertical crustal movements in Iceland, whereas *Decker* [1968] initiated a study of horizontal movements in Iceland. *Whitten* [1969] reported vertical and horizontal movements of

up to 15 meters south of Montague Island during the great Alaskan earthquake of 1964. Meister et al. [1968] reduced USCGS triangulation data for the Dixie Valley–Fairview Peak earthquakes of 1954 and inferred that the epicentral area suffered NW–SE extension, with parallel contraction in the surrounding region. Scholz and Fitch [1969, 1970] concluded, from analysis of geodimeter data for long lines across faults on which creep is observed in California, that the observed creep accounts for a quarter of the total strain measured and that slippage or creep on secondary faults is required if there is no net accumulation of strain in the fault zones. Rayleigh and Burford [1969] reported that areal dilation is the predominant form of strain at San Andreas Lake, California. Hoffmann [1968] reported on geodimeter measurements in California. Hoffmann [1969] prescribed criteria for adequate geodimeter measurements and attempted earthquake prediction on the basis of observations of anomalous movements on faults. Major and Romig [1969] reported a cumulative shear in the Denver area of 2.5×10^{-5} in a 9-month period, and Marshall and Stephens [1968] reported tilts prior to some of the Denver events. Smith and Van de Lindt [1969] presented a new model for calculating strain energy release and suggested net increase of stored strain in parts of California.

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