Supporting Information for

"Seismicity During the Initial Stages of the Guy-Greenbrier, Arkansas, Earthquake Sequence"

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Section S1: Magnitude estimation

We report local magnitude M_L for all detected events in this study. To calibrate M_L , we first calculate the moment magnitude M_w for a selected group of 54 events with high-quality waveforms, located at different distances from station WHAR, within event clusters 1, 5, 6, and 7 (Figure S3a). To obtain M_w , we calculate seismic moment M_0 in the time domain for P, SH, and SV arrivals on displacement waveforms at WHAR (after removal of instrument response and rotation of the horizontal components to radial and transverse), following *Prejean and Ellsworth* [2001]:

$$M_0 = \frac{4\pi\rho v^3 r \left(\int u \, dt\right)}{2F} \tag{1}$$

where $\rho = 2.5 \text{ kg/m}^3$ is density, *v* is velocity (4.6 km/s for *P* and 2.7 km/s for *S* waves, estimated using travel times and hypocentral distances for these 54 events), and *r* is hypocentral distance. $(\int u \, dt)$ is the integrated area under the displacement pulse for the *P*, *SH*, or *SV* arrival. *F* is the radiation pattern correction term, which we approximate using the average radiation pattern term $\langle F \rangle = 0.52$, 0.41, and 0.48 for *P*, *SH*, and *SV*, respectively. Equation 1 has an extra factor of 2 in the denominator compared to *Prejean and Ellsworth* [2001] to approximately account for the amplitude at the free surface. To calculate *M*₀, we use only arrivals with high-quality displacement pulses: *P* and *SH* for Cluster 1 events (taking the average *M*₀), *SV* for Cluster 5 and 6 events, and *SH* for Cluster 7 events. We then calculate moment magnitude *M*_w from *M*₀ [*Bormann*, 2012]:

$$M_w = \frac{2}{3} \log_{10} M_0 - 6.07 \tag{2}$$

Next, we solve for the local magnitude distance correction using these 54 calibration events. We assume that the magnitude can be expressed as:

$$M_w = \log_{10} \left(A_{\text{peak}} R^k \right) + C$$

$$\Rightarrow M_w - \log_{10} A_{\text{peak}} = k \log_{10} R + C,$$
(3)

with $\log_{10} A_{\text{peak}}$ computed as:

=

$$\log_{10} A_{\text{peak}} = \frac{1}{2} \left(\log_{10} A_{\text{peak,East}} + \log_{10} A_{\text{peak,North}} \right), \tag{4}$$

where $A_{\text{peak,East}}$ and $A_{\text{peak,North}}$ are peak Wood-Anderson seismogram amplitudes, after applying a 1 Hz high-pass filter to remove low-frequency noise, from the east and north components, respectively [*Bormann*, 2012]. In Equation 3, *R* is epicentral distance, and the distance correction parameters to estimate are *k* (representing the effect of geometric spreading and attenuation) and *C* (constant for the base level, where we solve for a separate *C* for each station). Since we can obtain A_{peak} from Equation 4 at each station, with the objective of estimating 4 distance correction parameters

k, C_{WHAR} , C_{ARK2} , and C_{ARK1} , we express Equation 3 in matrix form (**d** = **Gm**), where the design matrix **G** has dimensions 54 events x 3 stations = 162 rows, by 4 columns, and *i* is the event index:

$M_{w}[i=1] - \log_{10} \left(A_{\text{peak,WHAR}}[i=1] \right)$	$\log_{10}\left(R_{\rm WHAR}[i=1]\right)$	1	0	0	
$M_{w}[i=2] - \log_{10} \left(A_{\text{peak},\text{WHAR}}[i=2] \right)$	$\log_{10}\left(R_{\rm WHAR}[i=2]\right)$	1	0	0	
÷	÷	÷	÷	÷	
$M_{w}[i = 54] - \log_{10} \left(A_{\text{peak},\text{WHAR}}[i = 54] \right)$	$\log_{10}\left(R_{\rm WHAR}[i=54]\right)$	1	0	0	
$M_{w}[i=1] - \log_{10} \left(A_{\text{peak},\text{ARK2}}[i=1] \right)$	$\log_{10} (R_{ARK2}[i = 1])$	0	1	0	k
$M_w[i=2] - \log_{10} \left(A_{\text{peak},\text{ARK2}}[i=2] \right)$	$\log_{10}\left(R_{\text{ARK2}}[i=2]\right)$	0	1	0	C _{WHAR}
÷	÷	÷	÷	÷	C _{ARK2}
$M_{w}[i = 54] - \log_{10} \left(A_{\text{peak},\text{ARK2}}[i = 54] \right)$	$\log_{10}\left(R_{\text{ARK2}}[i=54]\right)$	0	1	0	C_{ARK1}
$M_w[i = 1] - \log_{10} \left(A_{\text{peak},\text{ARK1}}[i = 1] \right)$	$\log_{10} (R_{ARK1}[i = 1])$	0	0	1	
$M_w[i=2] - \log_{10} \left(A_{\text{peak},\text{ARK1}}[i=2] \right)$	$\log_{10}\left(R_{\text{ARK1}}[i=2]\right)$	0	0	1	
÷	÷	÷	÷	÷	
$M_w[i = 54] - \log_{10} \left(A_{\text{peak,ARK1}}[i = 54] \right)$	$\log_{10} (R_{ARK1}[i = 54])$	0	0	1	
``````````````````````````````````````				-	(5)

Inverting for the best-fit distance correction parameters in a least-squares sense, we get k = 1.5273,  $C_{WHAR} = 1.7141$ ,  $C_{ARK2} = 0.4447$ , and  $C_{ARK1} = 0.5782$ . Plugging in these parameters into Equation 3, and assuming  $M_w = M_L$ , we calculate local magnitude at each station for every detected event, given peak amplitudes from 1 Hz high-pass filtered Wood-Anderson seismograms (Equation 4) and epicentral distances:

$$M_{\rm L,WHAR} = \log_{10} A_{\rm peak,WHAR} + k \log_{10} R_{\rm WHAR} + C_{\rm WHAR}$$

$$\Rightarrow M_{\rm L,WHAR} = \log_{10} A_{\rm peak,WHAR} + 1.5273 \log_{10} R_{\rm WHAR} + 1.7141$$
(6)

$$M_{\rm L,ARK2} = \log_{10} A_{\rm peak,ARK2} + k \log_{10} R_{\rm ARK2} + C_{\rm ARK2}$$
(7)

$$\Rightarrow M_{\rm L,ARK2} = \log_{10} A_{\rm peak,ARK2} + 1.5273 \log_{10} R_{\rm ARK2} + 0.4447$$

$$M_{L,ARK1} = \log_{10} A_{peak,ARK1} + k \log_{10} R_{ARK1} + C_{ARK1}$$

$$\Rightarrow M_{L,ARK1} = \log_{10} A_{peak,ARK1} + 1.5273 \log_{10} R_{ARK1} + 0.5782$$
(8)

For the 1,740 located events (Figure 6, blue), we report local magnitude  $M_L$  as the average of  $M_L$  estimates at the 3 stations:

$$M_{L} = \frac{1}{3} \left( M_{\rm L,WHAR} + M_{\rm L,ARK2} + M_{\rm L,ARK1} \right)$$
(9)

Figure S3b shows that the  $M_L$  values from Equation 9 are reasonably similar to  $M_w$  estimates for the 54 calibration events, although they sometimes deviate by as much as 0.7 magnitude units.

For the 6,508 assigned events (Figure 6, black) and 6,356 unassigned events (Figure 6, red), we report local magnitude  $M_L$  as equal to  $M_{L,WHAR}$  instead of using the average from Equation 9, because waveforms at stations ARK1 and ARK2 are unreliably noisy for these smaller events. For the assigned events, we calculate the epicentral distance  $R_{WHAR}$  using the average latitude and longitude from all located events in that cluster. For the unassigned events, we pick *P* and *S* phases on the Wood-Anderson seismograms at WHAR, compute the *S*-*P* time, and estimate  $R_{WHAR}$  as:

$$R_{\rm WHAR} = \frac{v_p v_s}{v_p - v_s} \left( t_s - t_p \right) \tag{10}$$

using the same velocities from Equation 2:  $v_p = 4.6$  km/s,  $v_s = 2.7$  km/s.

### Section S2: Phase picking procedure

We automatically pick *P*- and *S*-wave arrivals on the 3 stations (ARK1, ARK2, WHAR) with the Akaike Information Criteria (AIC) picker [*Maeda*, 1985], manually adjust them as needed, and assign integer weights for pick quality ranging from 0 (best) to 3 (worst). We pick *P* phases on the vertical component and *S* phases on the horizontal components. We estimate the origin time of each event using the *S*-*P* time on the station with the earliest arrivals, assuming a  $V_p/V_s$  ratio of  $\sqrt{3}$ .

### Section S3: VELEST inversion for quarry-constrained velocity model

The layer boundaries in the *Ogwari et al.* [2016] starting velocity model remain the same, but we allow  $V_p$  and  $V_s$  to change in each layer. We input the notch location (35.2928° N, 92.3973° W, 0 km depth), estimated origin times, and travel times for the 3 quarry blasts as shot data (with fixed coordinate locations), repeated 8 times, for a total of *nshot* = 24 shots, which balances the relative number of shots compared to the number of earthquakes used in the velocity model inversion. We do not fix the origin time of the shots (*nshfix* = 0), do not apply a shot correction (*nshcor* = 0), and do not invert for station corrections (*nsinv* = 0). To ensure a robust result, we run VELEST twice, each time using a different subset of 50 earthquakes with well-distributed locations, and confirm that we get a consistent velocity model for each case. We jointly invert for the velocity model and earthquake locations every *invertratio* = 3 iterations until we reach *ittmax* = 50 iterations, using damping parameters *othet* = *xythet* = *zthet* = 0.03 and *vthet* = 10.

#### Section S4: Differential travel time calculation for hypoDD

We first compute differential travel times from both catalog P and S picks and cross-correlation-based picks for the 1,229 events where we already have initial absolute locations from VELEST. The P

and *S* picks, originally determined for the VELEST locations, are input into the *ph2dt* program with parameters from Table S2 to get catalog differential times. The VELEST to *ph2dt* pick weight conversion is: 0 to 1 (best), 1 to 0.75, 2 to 0.5, 3 to 0.25 (worst). To get cross-correlation differential times, we cross-correlate short windows around the *P* and *S* arrivals for every event pair separated by less than 2 km within the 1,229 located earthquakes at all 3 stations, keeping only differential times with  $CC \ge 0.6$ . We use a 0.9 s window around the *P* arrival, 0.35 s before and 0.55 s after, on the vertical component. We use a longer 1.5 s window around the *S* arrival, 0.55 s before and 0.95 s after, on both horizontal components, and keep the differential time from the component that results in a higher CC. These time windows are tapered and filtered 2-20 Hz before cross-correlation in the time domain, where we achieve subsample precision of the differential time by fitting a parabola near the peak of the cross-correlation function and interpolating the time at the peak [e.g. *Deichmann et al.*, 1992; *Schaff et al.*, 2004].

We also compute cross-correlation differential times between each of the 1,229 initially located events and the 13,375 remaining unlocated events, with the goal of locating as many of the remaining events as possible. For the unlocated events, we automatically pick *P* phases on the vertical component and *S* phases on the horizontal components with the AIC picker. We cross-correlate short windows around the *P* and *S* arrivals at all 3 stations for each pair containing a located and an unlocated event, keeping only differential times shorter than 0.5 seconds with CC  $\geq$  0.6. For each unlocated event, we set its starting location as the initial VELEST location of the already-located event with the highest cross-correlation. We use the same window lengths, filter bands, and time-domain cross-correlation with subsample precision mentioned previously.

### Section S5: Assignment of unlocated events to existing clusters

We first generate a representative stack waveform at station WHAR for each event cluster 1-16 (Table S5) by stacking all located event waveforms at WHAR belonging to that cluster. Since all events within a cluster are close together relative to their distance to WHAR, their waveforms are reasonably similar. We designate the highest magnitude event in the cluster as the master event. For each located event in the cluster (including the master event), we take a time window of length 8 seconds, starting 1.5 seconds before the *P*-wave, on all 3 components at station WHAR; apply a 1-20 Hz bandpass filter; and normalize each waveform by dividing by its L2 norm. For each non-master event in the cluster, we cross-correlate its normalized waveform with the normalized master event waveform, and align the waveform to the master event using the delay time associated with the peak CC averaged over 3 components. We stack all aligned waveforms, including the master event, and

divide by the total number of events to get the representative stack waveform. Figure S6 shows these stack waveforms for the 11 clusters where the number of sub-clusters is 1 (Table S5).

Clusters 1-16 are defined by spatially compact regions, but sometimes waveforms at WHAR within a cluster exhibit significant variability, so combining all aligned waveforms results in a stack waveform that is not similar to most event waveforms in the cluster. In this case, we apply hierarchical clustering with average linkage to group events with similar waveforms together into sub-clusters (e.g. *Harris* [2006], *Leskovec et al.* [2014]). We first compute the total 3-component CC between each pair of located events in the cluster, to use as a similarity metric. We take an agglomerative approach to determine sub-clusters: we start with each event in its own sub-cluster, then we combine sub-clusters by grouping the most similar (highest 3-component CC) events together, then repeat the grouping step using the average CC between sub-clusters, until we reach a minimum CC cutoff. Events within a sub-cluster have more similar waveforms (above the minimum CC cutoff), compared to events from different sub-clusters. Clusters 3, 4, 5, 10, 13 have 5, 3, 3, 9, and 11 sub-clusters, respectively (Table S5). The number of sub-clusters depends on the minimum CC cutoff used in hierarchical clustering (Table S6), which we set empirically to get a relatively small number of sub-clusters. For the 5 clusters with multiple sub-clusters, the stack waveforms for sub-clusters within the same cluster exhibit obvious differences (Figures S7 to S11).

We cross-correlate each unlocated event waveform at WHAR on all 3 components with the representative stack waveform from each cluster (or each sub-cluster), and assign the unlocated event to the cluster where we get the highest CC between the unlocated event waveform and stack waveform, provided that the CC exceeds an empirically determined cluster-specific threshold listed as "Minimum CC threshold to assign" in Table S5. Figure S12 explains how to set this threshold to assign events to a cluster, for the example of Cluster 1 where the threshold is  $CC \ge 0.5$  (Table S5). We first calculate a CC distribution by cross-correlating the cluster stack waveform with each located event waveform belonging to the same cluster (Figure S12a) because we want the CC for these events in the same cluster to exceed the threshold. We then calculate a CC distribution between the cluster stack waveform and each located waveform belonging to a different cluster or sub-cluster. Since we have a total of 42 cluster or sub-cluster stack waveforms (Table S5; Figures S6 to S11), we compute 41 such distributions for every cluster. Figure S12b shows one out of these 41 CC distributions, between the Cluster 1 stack waveform and all located waveforms from Cluster 2; we want the CC for these events in a different cluster to remain below the threshold. We repeat this entire process for the stack waveform from the 41 remaining clusters or sub-clusters, obtaining a cluster-specific threshold that separates events belonging to the cluster from events belonging to other clusters. Once again, we

compute the average CC over all 3 components of station WHAR, using a time window of length 8 s, starting 1.5 s before the *P*-wave, filtered 1-20 Hz, and normalized by the L2 norm of the waveform.

Many previous studies have applied a CC threshold to group similar event waveforms into clusters, also called "multiplets" or "families". The choice of CC threshold depends on the time window length, frequency band, number of stations, noise content, and source characteristics of the event waveform. Most studies in induced seismicity [*Schultz et al.*, 2015a, 2016], volcano seismology [*Rowe et al.*, 2004; *Green and Neuberg*, 2006; *Petersen*, 2007], glacial seismicity [*Thelen*, 2013], and tectonic microseismicity [*Cattaneo et al.*, 1999; *Massa et al.*, 2006; *Bisrat et al.*, 2012] set a high CC threshold between 0.7 and 0.8 in order to include only the most similar events within each cluster. However, our goal was to include as many events as possible within a cluster, especially the noisier small events, so we set a lower CC threshold that ranged from 0.3 to 0.7 (Table S5). In addition to using the CC distributions of events within the same cluster and in different clusters as a guide to empirically set the threshold (Figure S12), we visually inspected all located and assigned waveforms in each of the 42 clusters or sub-clusters, as done for Cluster 1 in Figure 7, to check for waveform similarity within the cluster. *Harris and Dodge* [2011] used a low CC threshold of 0.316 to assign similar aftershocks to a cluster, because it was higher than most CC values in the distribution of unrelated event waveforms.





**Figure S1.** Map of Guy-Greenbrier area in central Arkansas (red box, inset at lower left) with ANSS catalog earthquake locations (75 events in Data Set S1), seismic stations, wastewater injection wells, and production wells with hydraulic fracturing stimulation during the time period 2010-06-01 to 2010-09-01. We plot ANSS catalog events after this time period (small gray dots) to delineate the location of the Guy-Greenbrier Fault. We also include the same red boxes B1-B5 from Figure 1. Fault traces are from *Horton* [2012].



**Figure S2.** Examples of (a) low-frequency, (b) high-frequency, narrowband noise signals above the FAST similarity event detection threshold of 0.33 (Table S1) that were automatically removed by post-processing. These signals were similar enough to at least another narrowband noise signal at different times in the continuous seismic data to exceed the detection threshold. Time windows are on the left and power spectra are on the right.



**Figure S3.** (a) Location of 54 selected high-quality events, belonging to one of the event clusters 1, 5, 6, 7 (blue rectangles) used to calibrate  $M_L$ . (b) Magnitude calibration results: comparison of  $M_L$  with  $M_w$  for the selected 54 events, colored by cluster number.



Figure S4. Comparison of our magnitude and depth estimates with those from the ANSS catalog. If an event lies on the dotted line, it matches the catalog value. Events are colored by cluster number (Table S5), with 0 indicating that the event does not belong to any of the 16 clusters. (a) Catalog magnitude  $M_d$ , computed from duration of the coda, versus our  $M_L$  local magnitude estimates, for 74 events. Although catalog  $M_d$  values have large uncertainties, they are reasonably consistent with our  $M_L$  estimates (Section S1). (b) Catalog depth versus our 3-station depth estimates for the 45 catalog events that we located (events before 2010-06-11 were not recorded on ARK1 or ARK2). Depths are positive downward, so the gray regions with negative depths are not physical. The depths generally agree, only because of the large depth uncertainties from the catalog.



**Figure S5.** *P*-wave polarization analysis at station ARK2 for three events from Cluster 1: (a) Event at 4235551.13 s, (b) Event at 4261764.56 s, (c) Event at 4998208.01 s, for the purpose of testing different velocity models. Event start times are defined by the number of seconds after UTC 2010-06-01 00:00:00. (Top) 0.4-second window on each component of ARK2, centered on the *P*-wave arrival. We use a 0.2-second time window after the *P* arrival (yellow rectangle) for polarization analysis of the horizontal components (Center), which shows the North (HHN, blue) versus East (HHE, red) particle motion (magenta). To determine the back-azimuth to the event location (black arrow away from ARK2), we estimate the best-fit orientation of the particle motion with singular value decomposition, then resolve the 180° ambiguity with the first motion polarities [*Havskov and Ottemoller*, 2010]. (Bottom) We compare the back-azimuth to the event (black arrow away from ARK2) with the event location determined using three different velocity models: *Chiu et al.* [1984] (red), *Ogwari et al.* [2016] (blue), and our new quarry-constrained model (cyan). For all three events, the locations obtained with the *Ogwari et al.* [2016] and quarry-derived models are reasonably consistent with the back-azimuth direction, but the location from the *Chiu et al.* [1984] model is in the wrong direction. The quarry-model locations shifted ~0.7 km southeast (cyan), as applied in Figure 9, agree with the back-azimuth locations.



**Figure S6.** Representative stack waveforms at each component of station WHAR for the 11 clusters from Table S5 with only one sub-cluster. Waveforms at WHAR for all located events in each cluster are normalized and aligned to maximum CC with the master event (with largest magnitude) before stacking. Unlocated events are cross-correlated with these stack waveforms in order to assign them to clusters.



**Figure S7.** Stack waveforms at each component of station WHAR for the 5 sub-clusters (3A-3E) from Cluster 3 (Table S5). Differences between the stack waveforms for each sub-cluster are evident. Events are assigned to Cluster 3 by cross-correlation with each sub-cluster stack waveform.



**Figure S8.** Stack waveforms at each component of station WHAR for the 3 sub-clusters (4A-4C) from Cluster 4 (Table S5). Differences between the stack waveforms for each sub-cluster are evident. Events are assigned to Cluster 4 by cross-correlation with each sub-cluster stack waveform.



**Figure S9.** Stack waveforms at each component of station WHAR for the 3 sub-clusters (5A-5C) from Cluster 5 (Table S5). Differences between the stack waveforms for each sub-cluster are evident. Events are assigned to Cluster 5 by cross-correlation with each sub-cluster stack waveform.



**Figure S10.** Stack waveforms at each component of station WHAR for the 9 sub-clusters (10A-10I) from Cluster 10 (Table S5). Differences between the stack waveforms for each sub-cluster are evident, though this is not surprising because this cluster includes events from a larger area. Events are assigned to Cluster 10 by cross-correlation with each sub-cluster stack waveform.



**Figure S11.** Stack waveforms at each component of station WHAR for the 11 sub-clusters (13A-13K) from Cluster 13 (Table S5). Differences between the stack waveforms for each sub-cluster are evident, especially because these waveforms are noisy. Events are assigned to Cluster 13 by cross-correlation with each sub-cluster stack waveform.



**Figure S12.** Description of how to set a CC threshold empirically to assign events to a specific cluster. In this example, the threshold is  $CC \ge 0.5$  (dotted red line) for Cluster 1 (Table S5). (a) CC distribution from cross-correlating the Cluster 1 stack waveform with all located event waveforms in Cluster 1, at 3 components on station WHAR. Almost all CC values exceed the threshold of 0.5 as desired. (b) CC distribution from cross-correlating the Cluster 1 stack waveform with all located event waveforms in Cluster 2, at 3 components on station WHAR. Since these events are in a different cluster from the stack waveform, most CC values are below the threshold of 0.5 as desired.

## Tables S1 to S8

**Table S1.** FAST input parameters [*Yoon et al.*, 2015] used to detect earthquakes in each component of continuous seismic data at station WHAR. The event detection threshold was applied to the total 3-component FAST similarity matrix.

FAST parameter	Value
Time series window length for spectrogram generation	300 samples (3 s)
Time series window lag for spectrogram generation	3 samples (0.03 s)
Spectral image window length	64 samples (1.92 s)
Spectral image window lag = fingerprint sampling period	10 samples (0.3 s)
Number of top $k$ amplitude standardized Haar coefficients	800 (out of 2048)
LSH: number of hash functions per hash table <i>r</i>	8
LSH: number of hash tables <i>b</i>	100
Initial pair threshold: number $v$ (fraction) of tables, pair in same bucket	4 (4/100 = 0.04)
Event detection threshold (total 3-component FAST similarity)	0.33
Similarity search: near-repeat exclusion parameter	5 samples (1.5 s)
Near-duplicate pair and event elimination time window	4 s
Template matching comparison time window	6 s

 Table S2.
 Input parameters for *ph2dt* program to compute catalog P and S differential travel times.

Parameter name	Value	Description
MINWGHT	0	Minimum pick weight allowed
MAXDIST	120	Maximum distance (km) between event pair and stations
MAXSEP	10	Maximum hypocentral separation (km)
MAXNGH	30	Maximum number of neighbors per event
MINLNK	3	Minimum number of links required to define a neighbor
MINOBS	1	Minimum number of links per pair saved
MAXOBS	6	Maximum number of links per pair saved

Table S3. Input	parameters for h	vpoDD program	to relocate earth	auakes.

Parameter name	Value	Description
IDAT	3	Cross-correlation and catalog
IPHA	3	P and S phases
DIST	400	Maximum distance (km) between cluster centroid and station
OBSCC	2	Minimum number of obs/pair for crosstime data
OBSCT	2	Minimum number of obs/pair for network data
MINDS	0	Minimum distance between individual event pairs and stations
MAXDS	150	Maximum distance between individual event pairs and stations
MAXGAP	-999	Maximum azimuthal gap between individual event pairs and stations (-999: not used)
ISTART	2	From network sources
ISOLV	2	LSQR
IAQ	2	Keep air-quakes, reset depths to those of previous (successful) iteration
NSET	5	Number of sets of iteration

**Table S4.** hypoDD data weighting parameters, LSQR mode, 24 iterations.

NITER	WTCCP	WTCCS	WRCC	WDCC	WTCTP	WTCTS	WRCT	WDCT	DAMP
4	0.01	0.01	-9	-9	1	1	-9	-9	180
4	0.01	0.01	-9	-9	1	1	5	4	180
4	1	1	-9	2	0.01	0.01	5	4	180
4	1	1	5	2	0.01	0.01	5	4	180
8	1	1	5	0.5	0.01	0.01	5	4	180

Table S5. Information for 16 earthquake clusters with at least 10 events: minimum and maximum latitude and longitude of box defining the cluster boundary; number of sub-clusters; minimum CC threshold to assign events to each cluster; number of located, assigned, and total events in cluster; production well permit numbers (Table S8) within 2 km of and associated with each cluster. For clusters with more than one sub-cluster (Clusters 3, 4, 5, 10, 13), the bottom table contains the minimum CC threshold to assign events to each sub-cluster, and the number of located, assigned, and total events in that sub-cluster.

	89, 43343, 43344, 43375, 43376	89, 43343, 43344, 43375, 43376	89, 43343, 43344, 43375, 43376				111			41, 43257, 43272, 43122, 43149	1 1	IIa	55, 43256	150	11]	
Nearby production wells: Permit number	42069, 42146, 42262, 423	42069, 42146, 42262, 4238	42069, 42146, 42262, 4238	42069, 43375, 43376	43114, 43304	43043	43153	43154, 43258	43154	43042, 43058, 43203, 432	43439	43433	43252, 43253, 43254, 432	1	43244	43219
Total number of events	3192	1078	714	288	57	520	255	384	12	460	134	14	1055	19	26	19
Number of assigned events	2525	778	486	266	12	440	210	365	6	428	79	5	848	15	24	18
Number of located events	667	300	228	22	45	80	45	19	33	32	55	6	207	4	2	1
Minimum CC threshold to assign	0.5	0.5	see table below	see table below	see table below	0.5	0.5	0.4	0.5	see table below	0.4	0.7	see table below	0.5	0.5	0.4
Number of sub-clusters	-	1	5	б	б	1		1	1	6		-1	11	1		1
Maximum Iongitude (deg)	-92.28	-92.28	-92.3	-92.31	-92.32	-92.26	-92.22	-92.203	-92.25	-92.15	-92.332	-92.345	-92.335	-92.3	-92.375	-92.235
Minimum longitude (deg)	-92.32	-92.32	-92.32	-92.325	-92.35	-92.29	-92.225	-92.24	-92.27	-92.25	-92.344	-92.355	-92.368	-92.32	-92.395	-92.26
Maximum latitude (deg)	35.36	35.34	35.3258	35.315	35.3	35.28	35.26	35.245	35.245	35.3	35.35	35.335	35.395	35.245	35.325	35.395
Minimum latitude (deg)	35.34	35.3258	35.315	35.3	35.27	35.26	35.245	35.225	35.235	35.26	35.335	35.325	35.365	35.233	35.3	35.38
Cluster number	-	2	б	4	5	9	7	8	6	10	11	12	13	14	15	16

Total number of events	45	88	260	320	1	181	99	41	L	4	46	75	LL	87	33	5	176	10	7	20	106	14	30	350	107	243	7	151	23	8	16
Number of assigned events	34	74	222	156	0	179	54	33	5	2	5	99	73	83	-1	2	169	6	9	19	94	12	28	346	4	171	5	131	12	0	5
Number of located events	Ξ	14	38	164	1	2	12	8	2	2	41	6	4	4	2	33	L	1	1	1	12	2	2	4	63	72	2	20	11	8	11
Minimum CC threshold to assign	0.6	0.5	0.5	0.6	0.5	0.6	0.6	0.7	0.5	0.4	0.6	0.5	0.5	0.5	0.8	0.5	0.5	0.5	0.4	0.5	0.4	0.3	0.4	0.4	0.4	0.4	0.3	0.5	0.4	0.35	0.3
Sub-cluster number	3A	3B	3C	3D	3E	4A	4B	4C	5A	5B	5C	10A	10B	10C	10D	10E	10F	10G	10H	101	13A	13B	13C	13D	13E	13F	13G	13H	131	13J	13K

Cluster number	Minimum CC cutoff	Number of sub-clusters
3	0.2	5
4	0.2	3
5	0.1	3
10	0.4	9
13	0.18	11

Table S6. Minimum CC cutoff for separating event clusters into sub-clusters; also see Table S5.

**Table S7.** Class 2 Underground Injection Control (UIC) wastewater injection wells (Figure 1, inverted triangles) active during the study period 2010-06-01 to 2010-09-01. Wells 1 and 5 (Figure 1, inverted triangles colored by depth), located nearest the Guy-Greenbrier Fault, started injecting during the study period. Volume and pressure are peak values observed during the injection period (not the study period). Data are taken from *Horton* [2012], *Ogwari et al.* [2016], and *AOGC* [2017a].

Well Number	Well Name	Permit	Volume (m ³ /month)	Pressure (MPa)	Start and Stop Dates	Injection Depth (m)
1	SRE	43266	62,622	11.8	2010-07-07 to 2011-03-03	1821 to 1969
2	Trammel	41079	54,058	15.8	2009-04-15 to 2011-06-20	1982 to 2009
3	Moore	39487	23,435	20.3	2009-06-15 to 2011-07-27	2365 to 3231
4	Underwood	42981	29,573	5.1	2010-01-15 to 2010-10-15	1713 to 1926
5	Edgmon	36380	19,580	19.6	2010-08-16 to 2011-03-03	2379 to 3344
6	Scroggins	42989	18,629	3.2	2010-04-05 to 2011-10-15	678 to 706

public Arkansas Oil and Gas Database [AOGC, 2017a,b]. The total depth, top depth, and bottom depth are measured depth values along the well trajectory, while the true vertical depth is Table S8. Data for 53 production wells with hydraulic fracture stimulation during the time period 2010-06-01 to 2010-09-01, located within the map area in Figure 1, retrieved from the

measured vertically. Bottom depth is deepest measured depth at the first stage of stimulation, while top depth is shallowest measured depth at the last stage of stimulation.

llons) (barrels)
10530 1519 10530 1519 9983 1349 10052 1349 9851 1349
10622 5940 10078 5880 10144 5935 9943 5736
5543 5381 5362 5359
758 664 664 676
-92.300343 -92.303790 -92.299977 -92.301942
35.335994 35.349317 35.349194 35.349272
-92.302812 -92.302812 -92.302342 -92.302577
35.336189 35.336171 35.336178
07-24-2010 07-29-2010 07-24-2010

### Data Set S1.

List of 75 catalog events from the ANSS catalog, with latitude between  $35.18^{\circ}$ N and  $35.42^{\circ}$ N and longitude between  $-92.52^{\circ}$ W and  $-92.08^{\circ}$ W, sorted by time. Figure S1 plots these catalog event locations, colored by depth. (To account for location uncertainty, we added  $0.02^{\circ} \approx 2$  km to the map boundaries in Figure S1 to define our catalog search box.) These events were recorded by the Cooperative New Madrid Seismic Network. One event (ID nm607354) occurred during a time gap in the continuous data at WHAR, so we did not detect it. The table below describes each column. Columns 16-19 match columns 1-4 in Data Set S2, and columns 5-8 in Data Set S3.

Column	Description
1	Catalog origin time in seconds since UTC 2010-06-01T00:00:00
2	Catalog origin time (UTC) in format YYYY-MM-DDTHH:MM:SS.SSSZ
3	Latitude (deg) from catalog
4	Longitude (deg) from catalog
5	Depth (km) from catalog
6	Magnitude from catalog: $M_d$ calculated from coda duration
7	Number of stations used for catalog location
8	Largest azimuthal gap between azimuthally adjacent stations (deg)
9	Horizontal distance from epicenter to nearest station (deg)
10	RMS travel time residual (s)
11	Catalog event ID
12	Catalog horizontal location error (km)
13	Catalog depth location error (km)
14	Catalog magnitude error
15	Number of stations used for magnitude calculation
16	Event id number = 100 * (Event detection time in seconds since UTC 2010-06-01T00:00:00)
17	Local magnitude $M_L$
18	Which algorithm detected this event? (0: both FAST and template matching, 1: only template matching, 2: only FAST)
19	Cluster number (from Table S5), including sub-cluster information. 0: does not belong to any cluster.

# Data Set S2.

List of 14,604 detected events from 2010-06-01 to 2010-09-01 at single station WHAR, sorted by time. The table below describes each column.

Column	Description
1	Event id number = 100 * (Event detection time in seconds since UTC 2010-06-01T00:00:00)
2	Local magnitude $M_L$
3	Which algorithm detected this event? (0: both FAST and template matching, 1: only template matching, 2: only FAST)
4	Cluster number (from Table S5), including sub-cluster information. 0: does not belong to any cluster. 17: quarry blast.

# Data Set S3.

List of 1,740 located events, sorted by time. The table below describes each column. Columns 5-8 match columns 1-4 in Data Set S2.

Column	Description
1	Latitude (deg)
2	Longitude (deg)
3	Depth (km)
4	Number of days since UTC 2010-06-01T00:00:00
5	Event id number = 100 * (Event detection time in seconds since UTC 2010-06-01T00:00:00)
6	Local magnitude $M_L$
7	Which algorithm detected this event? (0: both FAST and template matching, 1: only template matching, 2: only FAST)
8	Cluster number (from Table S5), including sub-cluster information. 0: does not belong to any cluster. 17: quarry blast.

### Movie S1.

Cumulative time evolution of seismicity in Cluster 1, and hydraulic fracturing stimulation at the 5 nearest production wells (labeled by permit number in Table S8), near north end of the Guy-Greenbrier Fault. We display the shifted event locations from Figure 9b. Earthquakes (circles sized by relative magnitude), as well as stimulated sections of production wells during each stage of hydraulic fracturing, are colored by time with Day 0 defined as 2010-07-16 00:00:00 UTC. Thick colored stages along the well path are currently being stimulated, while thin colored stages are past stimulations. The movie covers the 16-day time period shown in Figure 10e.