

# On the accuracy of adiabaticity parameter estimations using magnetospheric models

John D. Haiducek<sup>1</sup>, Natalia Y. Ganushkina<sup>1,2</sup>, Stepan Dubyagin<sup>2</sup>, Daniel T. Welling<sup>1</sup>

<sup>1</sup>Climate and Space Sciences, University of Michigan, Ann Arbor, MI, USA.

<sup>2</sup>Finnish Meteorological Institute, Helsinki, Finland

## Key Points:

- Neither MHD nor empirical models can give reliable estimates of the adiabaticity parameter  $K$ .
- A method is proposed to correct model estimates of  $K$  using concurrent observations in the magnetotail.
- Corrected estimates of  $K$  are close to the  $K = 8 - 10$  range expected for current sheet scattering.

This is the author manuscript accepted for publication and has undergone full peer review but has not been through the copyediting, typesetting, pagination and proofreading process, which may lead to differences between this version and the [Version of Record](#). Please cite this article as doi: [10.1029/2018JA025916](https://doi.org/10.1029/2018JA025916)

Corresponding author: John D. Haiducek, [jhaiduce@umich.edu](mailto:jhaiduce@umich.edu)

## Abstract

Recent studies have found that even during quiet times, observed proton isotropic boundaries (IB) are often projected to the region of high adiabaticity parameter ( $K \approx 30$ ), where  $K = \frac{R_c}{r_g}$  is the ratio of magnetic field line radius of curvature to the particle gyroradius. This contradicts the accepted hypothesis that current sheet scattering (CSS) is the dominant mechanism of IB formation because  $K \approx 8$  would be expected for this mechanism. We used magnetohydrodynamic (MHD) simulations and empirical models to compute  $K$  for 30 keV proton IB observations within three hours of local midnight. We found that neither class of model reliably estimates  $K$  unless supported by magnetic field observations in the current sheet. MHD simulations produced higher  $K$  values than expected for CSS ( $K = 15 - 30$ ), and empirical models gave lower values ( $K < 4$ ).

We obtained reliable estimates of  $K$  by controlling for the accuracy of the normal component and the gradient of the radial component in the neutral sheet, using observations from three THEMIS satellites. For the first time, we demonstrated that both these variables should be taken into account for the accurate estimation of the curvature radius. This greatly reduced the spread of  $K$  values, indicating that much of the previous spread was due to errors in the magnetic field, but also that these errors can be controlled. Most of the corrected values fall within the expected range for CSS, supporting the hypothesis that the IB's were formed by CSS. Accounting for all model results, we obtain an average corrected value of  $K = 6.0$ .

## 1 Introduction

Determining the structure of the Earth's magnetic field under various solar wind conditions is essential for understanding the relationship between ionospheric features and magnetospheric processes. This requires having magnetic field vectors throughout the relevant parts of the magnetosphere, so that field lines can be traced between the ionosphere and the magnetotail. At present, the available spacecraft observations are sparsely distributed, and magnetospheric models play a crucial role by providing estimations of the magnetic field throughout the geospace environment.

One way to gain insight into field line mappings is by studying the isotropic boundary (IB), a distinct feature that can be used to probe connections between the ionosphere and magnetosphere. The IB refers to a latitude in the auroral zone at which a substan-

44 tial change occurs in the flux of downwelling particles into the ionosphere. Equatorward  
45 of the IB, the flux in directions perpendicular to the local magnetic field well exceeds  
46 the downwelling flux parallel to the local magnetic field. Poleward of the IB, compara-  
47 ble fluxes are detected in the directions parallel and perpendicular to the field. This has  
48 been observed by many satellites, including Injun 1 and 3, ESRO IA and IB, NOAA, and  
49 DMSP [Søråas, 1972; Imhof *et al.*, 1977; Sergeev *et al.*, 1983; Newell *et al.*, 1998]. For  
50 protons, the IB is observed in all MLT sectors and at all activity levels [Sergeev *et al.*,  
51 1993].

52 The difference in loss-cone filling poleward and equatorward of the IB indicates that  
53 the particles observed at the IB originate from a transitional region within the magneto-  
54 sphere, in which the rate of pitch angle scattering changes significantly. On the night side,  
55 one mechanism for this is a transition from adiabatic to chaotic particle motion as parti-  
56 cles cross the current sheet, a process termed current sheet scattering (CSS) [West *et al.*,  
57 1978; Büchner and Zelenyi, 1987; Sergeev *et al.*, 1993]. This occurs when the radius of  
58 curvature of the local magnetic field line,  $R_c$ , becomes comparable to the effective parti-  
59 cle gyroradius  $r_g$  [Alfvén and Fälthammar, 1963; Tsyganenko, 1982; Büchner and Zelenyi,  
60 1987; Delcourt *et al.*, 1996], and the strength of this scattering process is parameterized by  
61 the ratio  $K = R_c/r_g$ . That the CSS mechanism results in isotropic precipitation from the  
62 plasma sheet is undisputed, because for  $K < 1$  the particle motion in the plasma sheet is  
63 chaotic [see e.g. Coroniti, 1980; Lyons and Speiser, 1982; Chen and Palmadesso, 1986;  
64 Büchner and Zelenyi, 1987; Sergeev and Gvozdevsky, 1995]. In general the IB may be  
65 formed by other processes, most importantly the interaction of particles with electromag-  
66 netic ion-cyclotron (EMIC) waves [Kennel and Petschek, 1966; Liang *et al.*, 2014; Sergeev  
67 *et al.*, 2015a,b]. The role of such interactions in particle precipitation has long been rec-  
68 ognized [see e.g. the review paper by Hultqvist, 1979]. However, EMIC waves cannot al-  
69 ways be responsible for IB formation, because the intensity of EMIC waves varies strongly  
70 with activity and MLT [e.g. Bräysy *et al.*, 1998; Halford *et al.*, 2010; Usanova *et al.*, 2012].  
71 A number of efforts have identified pitch-angle scattering due to CSS as the main source  
72 for particle precipitation from the magnetotail during quiet conditions [Sergeev and Tsy-  
73 ganenko, 1982; Sergeev *et al.*, 1993; Ganushkina *et al.*, 2005]. CSS does not require the  
74 presence of waves and can explain the fact that the IB is observed in all activity levels and  
75 all MLTs on the night side. However, scattering by waves can sometimes cause the IB to  
76 form at a different latitude than would occur for CSS, particularly during storms and sub-

77 storms [S raas *et al.*, 1980; Gvozdevsky *et al.*, 1997; Yahnin and Yahnina, 2007; Sergeev  
78 *et al.*, 2010; Dubyagin *et al.*, 2018], and there is evidence for IB formation by waves dur-  
79 ing quiet periods as well [Sergeev *et al.*, 2015b].

80 When the CSS mechanism is responsible for IB formation, the IB location is de-  
81 termined by the field geometry. This enables the IB latitude to be used to estimate the  
82 degree of magnetotail field stretching [Sergeev *et al.*, 1993; Sergeev and Gvozdevsky, 1995;  
83 Meurant *et al.*, 2007]. This motivates further study of the role of CSS in IB formation,  
84 in order to better determine the conditions under which CSS (as opposed to scattering by  
85 EMIC waves or some other process) is the controlling mechanism responsible for IB for-  
86 mation. One means to do so is by estimating the value of  $K$  associated with observed IB  
87 locations. Numerical tracing of particle trajectories from the current sheet [e.g. Delcourt  
88 *et al.*, 1996, 2000, 2006] has shown that CSS acts when  $K \lesssim 1 - 10$  (a range spanning an  
89 order of magnitude). Delcourt *et al.* [1996] noted significant dependence on the incident  
90 particle population. However, in analysis it is often useful to use a specific critical value  
91 (rather than a range) as the threshold for CSS, and for this purpose many researchers have  
92 used  $K_{crit} = 8$  as originally proposed by Sergeev *et al.* [1983].

93 To determine whether the IB is associated with CSS requires mapping IB obser-  
94 vations to the current sheet, and estimating  $K$  there. Both steps require a magnetic field  
95 model of some kind. A number of previous efforts, including Sergeev *et al.* [1993], Ganushk-  
96 ina *et al.* [2005], Sergeev *et al.* [2015b], and Dubyagin *et al.* [2018], have accomplished  
97 this using empirical models such as the Tsyganenko models. Such models have good  
98 traceability to observational data since they are constructed by fitting to available satel-  
99 lite measurements. However, such models tend to be limited to representing features that  
100 are resolved by the observational data used in their construction or represented in the form  
101 of the equations that are fit to that data. Global first-principles models such as magneto-  
102 hydrodynamic (MHD) and hybrid simulation codes offer an alternative. Such models have  
103 the potential to produce features that are governed by the physics incorporated in the mod-  
104 els, without necessarily requiring observational data that resolves those features directly.  
105 This makes first-principles models potentially useful in understanding the IB, which de-  
106 pends on the magnetic configuration in the magnetotail, a region that is highly dynamic  
107 and only sparsely covered by observational data.

108 To date, only *Gilson et al.* [2012] and *Ilie et al.* [2015] have used MHD models to  
109 explore the IB and its properties. Of these, *Ilie et al.* [2015] is of particular interest to us  
110 because they mapped the locations of in situ IB observations through the MHD fields to  
111 estimate  $K$  in much the same way as was previously done with empirical models. That  
112 effort focused on a one-day quiet interval on February 13, 2009, using IB observations  
113 obtained from the Medium Energy Proton and Electron Detector (MEPED) instruments  
114 [*Evans and Greer*, 2000] on the NOAA/POES and METOP spacecraft. A quiet interval  
115 was chosen in order to reduce the chance of particle scattering due to wave-particle inter-  
116 actions. Nonetheless, the estimates of  $K$  derived from MHD ranged from 27 to 44, unex-  
117 pectedly high values for  $K$  associated with quiet-time IB.

118 Many of the above studies produced  $K$  values covering a fairly wide range. More-  
119 over, since most use only a single model to map the IB to the magnetotail and to compute  
120  $K$ , it is generally not possible to tell what part of this wide variation is due to differences  
121 in the actual state of the magnetosphere (either in the field geometry or the action of other  
122 scattering mechanisms such as waves) and what part is due to differences between mod-  
123 els and observational methodologies. Using multiple models to study the same event in  
124 combination with in situ magnetic field observations will provide a means to distinguish  
125 variation in  $K$  due to model error from variation due to physical causes.

126 The goals of this paper are twofold: 1) Determine whether MHD is capable of es-  
127 timating  $K$  correctly for quiet-time IB observations 2) Provide better constraints on the  
128 range of  $K$  values associated with night-side IB formation during quiet time. The paper is  
129 organized as follows: Section 2 describes the event and the observations used to identify  
130 the IB locations from the observations. Section 3 gives the methodology used, including  
131 details of the three MHD simulations used in this paper, and the procedures used for field  
132 tracing and for computing  $K$ . Section 4 presents comparisons of the MHD simulations to  
133 magnetic field observations. Section 5 gives the results of the field line tracing from the  
134 IB locations using MHD, and gives the computed  $K$  values for all the IB crossings. Sec-  
135 tion 6 presents results of field line tracing and resulting  $K$  values obtained using the em-  
136 pirical models. Section 7 compares the two classes of models and shows results of the  $K$   
137 correction procedures. Section 8 discusses the implications of these results in the context  
138 of previous studies.

158 **Figure 1.** Solar wind observations and geomagnetic indices from 13 February, 2009. a) flow speed, b)  
 159 proton density, c) temperature, d)  $B_z$  (GSM), e) Sym-H, f) Kp, g) AL

## 139 2 Event and observations

140 We analyze the IB observations and their associated  $K$  values for the 24 hour inter-  
 141 val beginning at midnight UTC on 13 February, 2009. This interval was selected because  
 142 it is a quiet period in terms of solar wind driving and geomagnetic activity ( $AL > -150$  nT,  
 143  $Sym-H > -10$ ,  $Kp < 2$ ), and because of the availability of magnetic field observations in the  
 144 night-side magnetosphere from the five THEMIS spacecraft. THEMIS A, D, and E had  
 145 apogee at distances of 11.6-11.7  $R_E$ , near the region that is expected to map to the IB,  
 146 while THEMIS B and C had apogee at distances of 28.6 and 19.4  $R_E$ , respectively.

147 Figure 1 shows the solar wind driving conditions on this day, as well as the geo-  
 148 magnetic indices Sym-H, Kp, and AL. The solar wind velocity (Figure 1a) ranged from  
 149 298 to 335 km/s with no noticeable discontinuities, the proton density (Figure 1b) ranged  
 150 from 2.9 to 12  $\text{cm}^{-3}$ , and the  $B_z$  component of the IMF (Figure 1c) remained within  $\pm 3$   
 151 nT. The solar wind temperature (Figure 1d) ranged from  $1.1 \times 10^5$  to  $7.2 \times 10^5$  K. The  
 152 minimum Sym-H (Figure 1e) was -6 nT, the maximum Kp (Figure 1f) was 1.3, and the  
 153 minimum AL (Figure 1g) was -128 nT. Two negative diversions appear in the AL index  
 154 near the end of the day, but neither is strong enough to be considered a substorm. For in-  
 155 stance, neither the list from *Borovsky and Yakymenko* [2017] obtained using SML, nor the  
 156 list from the SuperMag website (<http://supermag.jhuapl.edu/substorms/>) produced using  
 157 the *Newell and Gjerloev* [2011] algorithm contain any substorms during this period.

160 Isotropic boundaries were identified using proton flux data from the MEPED instru-  
 161 ments on board several NOAA/POES and METOP spacecraft. The MEPED instrument,  
 162 which is described in *Evans and Greer* [2000], includes two telescopes, which measure  
 163 proton fluxes in four energy bands ranging from 30 to 6900 keV. The first telescope, called  
 164 the  $0^\circ$  telescope, is within  $10^\circ$  of the spacecraft's zenith direction (i.e. away from Earth).  
 165 At high latitudes this direction places the  $0^\circ$  telescope close to the direction of the local  
 166 magnetic field, so that it primarily detects precipitating particles. The second telescope,  
 167 termed the  $90^\circ$  telescope, is oriented nearly orthogonal to the  $0^\circ$  telescope. The  $90^\circ$  tele-  
 168 scope primarily detects locally trapped particles.

169 The IB locations used in this paper are the same as those from *Ilie et al.* [2015].  
170 These were determined from the MEPED P1 energy channel (30-80 keV). Although the  
171 nominal low energy limit of the P1 channel is 30 keV, it is actually somewhat higher  
172 and varies among the satellites due to detector degradation. In addition, the 90°- tele-  
173 scopes degrades more strongly than 0°-telescope. Table 1 shows the low energy limits  
174 for 0°- telescope as given by *Asikainen et al.* [2012]. To recalibrate 90°- telescope data to  
175 the 0°- telescope energy limit, we use procedure described in the Appendix of *Dubyagin*  
176 *et al.* [2018]. After this correction, we found IB crossings using the procedure in *Dubya-*  
177 *gin et al.* [2013] which identifies a poleward and equatorward limit for the IB location.  
178 Identifying a poleward and equatorward limit for the IB gives an uncertainty range of lat-  
179 itudes for each IB crossing, which was less than 0.3° for the selected events [*Ilie et al.*,  
180 2015]. The criteria for determining the equatorward limit of the IB were chosen to avoid  
181 identifying brief periods of isotropic or nearly isotropic fluxes near the low-latitude limit  
182 of the auroral oval, which may be the result of wave-particle interactions [*Gvozdevsky*  
183 *et al.*, 1997; *Yahnin and Yahnina*, 2007]. In total, 94 IB crossings were identified from  
184 five ionospheric satellites (NOAA 15-18 and METOP-02) using this procedure, of which  
185 27 were within three hours MLT of local midnight. As an additional measure to reduce  
186 the chances that the selected IB observations could be influenced by wave-particle interac-  
187 tions, only those IB observations that were of typical appearance were used. Typical ap-  
188 pearance means a sharp transition from an empty loss cone on the low-latitude side of the  
189 IB to a filled loss cone on the high-latitude side, with both the 0° and 90° fluxes reach-  
190 ing a maximum on the high-latitude side of the IB, followed by a monotonic decrease in  
191 fluxes going toward the polar cap. Rather than using all suitable IB observations, we in-  
192 clude only those for which the THEMIS A, D, or E spacecraft was within  $\pm 1$  hour MLT  
193 of the location of the IB observation, and within the radial distance range of  $r = 7 - 10$   
194  $R_E$  from the Earth. For the purpose of this paper we consider spacecraft satisfying these  
195 criteria as being conjugate with the IB location. We use observations from these conjugate  
196 spacecraft to test the accuracy of the model magnetic fields in the magnetotail, and to cor-  
197 rect for errors in those fields. After eliminating the IB observations that were of atypical  
198 appearance and those without suitable THEMIS observations, the final list consisted of 7  
199 IB observations, which are shown in Table 1.

201 An illustration of the locations of the THEMIS A, D, and E spacecraft is shown in  
202 Figure 2. Figure 2a shows the positions of THEMIS A (represented by blue squares) in

Satellite	Time	Detector energy (keV)	Mag. Lat.	MLT
METOP-02	2009-02-13/01:41:16	36	-68.06	22.77
METOP-02	2009-02-13/03:22:00	36	-67.97	0.04
NOAA-16	2009-02-13/02:23:16	45	-67.98	23.97
NOAA-16	2009-02-13/02:25:23	45	-68.11	22.82
NOAA-17	2009-02-13/02:41:34	45	-68.01	23.50
NOAA-17	2009-02-13/04:22:02	45	-68.34	0.91
NOAA-18	2009-02-13/16:54:17	30	67.83	2.53

200 **Table 1.** Times and locations of the IB observations, as well as detector cutoff energies for each spacecraft.

213 **Figure 2.** (a) THEMIS A location for the times of the IB observations. (b) and (c) Relative locations of  
 214 THEMIS A, D, and E at 4:22 UT.

203 the GSM  $z = 0$  plane at the time of selected IB observations. The spacecraft was located  
 204 near midnight for six of the seven IB observations. For the seventh, it was located closer  
 205 to dawn, though still more than  $5 R_E$  down-tail. Figures 2b and 2c show the relative loca-  
 206 tions of THEMIS A, D, and E at 4:22 UT. THEMIS D is represented by a purple upward-  
 207 pointing triangle, while THEMIS E is represented by a red downward-pointing triangle. It  
 208 is apparent that THEMIS A and E are separated significantly in the  $z$  direction (more than  
 209  $1 R_E$  apart), but are more closely spaced in the  $x$  and  $y$  directions. This enables us to es-  
 210 timate gradients in the  $z$  direction by comparing values at THEMIS A and E, which we  
 211 will use in Section 7 to estimate the influence of errors in  $R_c$  on the  $K$  values computed  
 212 by the models.

### 215 3 Methodology

216 Having obtained the list of IB observations in Table 1 which are conjugate with the  
 217 THEMIS A, D, or E spacecraft, we next proceed to computing  $K$  at a magnetotail location  
 218 corresponding with each IB observation. This will provide an indication of whether the  
 219 estimated field geometry is consistent with CSS for the observed IB crossings. Neither the  
 220 mapping nor the computation of  $K$  can be accomplished directly from the available obser-  
 221 vational data due to the small number of satellites operating in the magnetotail. Therefore



222 we require models to estimate the magnetic fields in order to do both the mapping and  
 223 the  $K$  calculation. We begin by tracing a field line from each IB observation using the  
 224 model-derived magnetic fields. Along this field line we find the location where  $|B|$  reaches  
 225 its minimum, and there compute  $K = R_c/r_g$  from the model-derived magnetic fields. The  
 226 tracing and the  $K$  calculation are accomplished using magnetic fields obtained from sev-  
 227 eral models, including three MHD simulations performed using the Space Weather Model-  
 228 ing Framework (SWMF) [Tóth *et al.*, 2005] and six empirical Tsyganenko models: *Tsyga-*  
 229 *nenko* [1995] (T96), *Tsyganenko* [2002] (T01), *Tsyganenko and Sitnov* [2005] (TS05), *Tsy-*  
 230 *ganenko and Andreeva* [2015] (TA15N and TA15B), and *Tsyganenko and Andreeva* [2016]  
 231 (TA16RBF). The SWMF simulations are described in detail later in this section, and the  
 232 Tsyganenko models are described in Section 6.

233 For each field line traced, we search for the point where  $|B|$  reaches its minimum.  
 234 At the location of  $|B|_{min}$ , the field line radius of curvature is computed as

$$R_c = \frac{1}{|(\mathbf{b} \cdot \nabla)\mathbf{b}|}, \quad (1)$$

235 where  $\mathbf{b}$  is the unit vector along the magnetic field direction, and  $\nabla\mathbf{b}$  is computed  
 236 using a two-point centered difference. We then compute the effective particle gyroradius  
 237  $r_g$ . When computing  $r_g$ , we take the low energy limit of the detector as the particle en-  
 238 ergy. As mentioned earlier, the detector energy limit varies among the different satellites  
 239 due to the degradation of the detectors over time [Asikainen *et al.*, 2012], and the values  
 240 used for each spacecraft are shown in Table 1.

241 Our SWMF simulations consist of the Block Adaptive Tree Solar-Wind, Roe-Type  
 242 Upwind Scheme (BATS-R-US) MHD model [Powell *et al.*, 1999; DeZeeuw *et al.*, 2000],  
 243 coupled with the Rice Convection Model (RCM) [Wolf *et al.*, 1982; Sazykin, 2000; Tof-  
 244 foletto *et al.*, 2003] and the Ridley Ionosphere Model (RIM) [Ridley and Liemohn, 2002;  
 245 Ridley *et al.*, 2004]. The inputs to the model were solar wind parameters obtained from  
 246 the 1-minute OMNI dataset provided by the NASA Goddard Spaceflight Center (GSFC),  
 247 and the F10.7 radio flux observed at Penticon, BC [Tapping, 2013].

248 We ran three SWMF simulations, with the same inputs but with differences in grid  
 249 resolution, numerical scheme, and coupling parameters. By testing different settings of  
 250 SWMF we are able to determine in a general sense how sensitive the results are to various

251 SWMF settings. The first SWMF simulation, henceforth referred to as “SWMFa,” used  
 252 settings based on those in *Ilie et al.* [2015]. The settings of SWMF differ from those used  
 253 in *Ilie et al.* [2015] in the following ways:

- 254 • The simulation in *Ilie et al.* [2015] paper used a dipole moment of  $31.1 \mu\text{T}$  oriented  
 255 at  $289.1^\circ$  geographic longitude and  $79.0^\circ$  latitude, which was the default in SWMF  
 256 at the time. The simulations for the present paper used the IGRF dipole param-  
 257 eters for February 13, 2009, which were  $287.86^\circ$  geographic longitude and  $79.96^\circ$   
 258 latitude, and a dipole moment of  $29.97 \mu\text{T}$ .
- 259 • Minimum values for pressure and density were set to improve numerical stability
- 260 • The numerical scheme was changed from an implicit-explicit scheme to a fully ex-  
 261 plicit one in order to improve stability. The switch to fully explicit in turn required  
 262 a reduction in the time step.

263 The grid of SWMFa is the same as was used in *Ilie et al.* [2015], and contains about  
 264 4 million cells. The minimum cell size is  $1/8 R_E$  near the Earth, and the maximum cell  
 265 size is  $2 R_E$  at the outflow boundaries.

266 The other two SWMF simulations used settings previously described in *Haiducek*  
 267 *et al.* [2017]. The first of these, termed “Hi-res” in *Haiducek et al.* [2017], is identified  
 268 as “SWMFb” in this paper. SWMFb used a grid with about 2 million cells. The grid  
 269 for SWMFb was the same as that used in SWMFa within the near-Earth region (out to  
 270 about  $60 R_E$  in each direction). Beyond  $60 R_E$ , SWMFa used a  $2 R_E$  resolution every-  
 271 where. SWMFb, on the other hand, used a  $1 R_E$  cell size in the current sheet region out  
 272 to  $120 R_E$ , and beyond  $120 R_E$  the cell size increases until reaching  $8 R_E$  near the out-  
 273 flow boundaries. As a result, SWMFb had a higher resolution in the current sheet re-  
 274 gion but a smaller total number of cells due to coarser resolution in the deep tail ( $120$   
 275  $R_E$  and beyond). Besides the grid refinement, SWMFb differed from SWMFa in terms  
 276 of the RCM settings. In the coupling between BATS-R-US and RCM, SWMFb used the  
 277 *Young et al.* [1982] composition model to determine the ratio of oxygen to hydrogen in the  
 278 coupling with RCM, where SWMFa used a fixed ratio. Finally, an ad hoc decay was ap-  
 279 plied to the RCM ring current in SWMFb, which is designed to improve agreement with  
 280 observations during storm recovery but is not expected to affect the quiet-time results sub-  
 281 stantially.

282 The final simulation, termed “SWMFc,” was also described in detail in *Haiducek*  
 283 *et al.* [2017], and in that paper was referred to as the “SWPC” simulation due to the set-  
 284 tings being largely the same as those used operationally at NOAA Space Weather Predic-  
 285 tion Center (SWPC). SWMFc used a coarser grid than either SWMFa or SWMFb, with a  
 286 minimum cell size of  $1/4 R_E$  near the Earth, a maximum cell size of  $8 R_E$  at the outflow  
 287 boundaries, and no additional refinement in the tail or current sheet. The RCM coupling  
 288 settings were the same as SWMFb except that a fixed oxygen to hydrogen ratio was used.

289 For all of the SWMF simulations, field lines were traced through the MHD domain  
 290 from the location of each IB crossing identified using the MEPED data. This was done  
 291 once every minute of simulation time. The inner boundary of the MHD domain lies at  
 292  $2.5 R_E$  from the center of the Earth (rather than at the surface). Since the altitudes of the  
 293 NOAA and METOP spacecraft were lower than this, the IB locations were mapped to  $2.5$   
 294  $R_E$  prior to tracing through the MHD domain. In order to minimize the influence of non-  
 295 dipole harmonics on the mapping, we implemented the mapping by first transforming the  
 296 IB locations into altitude adjusted corrected geomagnetic coordinates [AACGM, *Baker*  
 297 *and Wing*, 1989], with the reference height set to 0 km. After conversion to AACGM co-  
 298 ordinates, each IB location was mapped to  $2.5 R_E$  using a dipole field. Within the MHD  
 299 domain, the field lines were traced using a third order Runge-Kutta scheme with a second-  
 300 order error estimator and adaptive step size.

#### 301 **4 Validation of magnetic fields with magnetospheric satellite observations**

302 In order to verify the accuracy of the SWMF in estimating the magnetic field geom-  
 303 etry, we compared the magnetic fields estimated by SWMF along the orbits of the GOES  
 304 and THEMIS satellites with observations from the fluxgate magnetometers onboard the  
 305 spacecraft [*Singer et al.*, 1996; *Auster et al.*, 2008] during the time from 0000 to 1800 UT  
 306 on February 13, 2009. This time period includes all of the IB observation times listed in  
 307 Table 1. For the THEMIS spacecraft, we additionally restrict the analysis to points in time  
 308 for which the spacecraft was at least  $7 R_E$  from the Earth, since this was the minimum  
 309 distance used for including a THEMIS spacecraft in analysis of an IB event. As an ex-  
 310 ample, Figure 3 shows fields at THEMIS A. Plots for THEMIS B-E and GOES 11 and  
 311 12 are included in the supplemental data. In Figure 3, time series plots of the magnetic  
 312 fields estimated by each SWMF simulation are overlaid on top of a plot of the observed  
 313 magnetic field. The observational data is shown in light blue, SWMFa in medium blue,

314 SWMFb in orange, and SWMFC in green. The left hand column (Figures 3a, 3c, and 3e)  
 315 shows the  $B_x$ ,  $B_y$  and  $B_z$  (GSM) components of the total magnetic field. The right hand  
 316 column (Figures 3b, 3d, and 3f) shows the same components for the external magnetic  
 317 field. We obtain the external field by subtracting a dipole field from the total field, with  
 318 the parameters of the dipole being the IGRF parameters given in Section 3 that were used  
 319 within the SWMF simulations. The same dipole field was subtracted from both from the  
 320 model and from the observed total fields to obtain the respective external fields. Through-  
 321 out Figure 3, the times of the IB observations listed in Table 1 are denoted with vertical  
 322 dotted lines. Note that the IB observations at 02:23 and 02:25 UT were very close to-  
 323 gether in time, and while individual lines are drawn for those two events, they are difficult  
 324 to distinguish in the plot.

325 **Figure 3.** Magnetic field components in GSM coordinates at the THEMIS A satellite, observed and pre-  
 326 dicted, for February 13, 2009. Left column shows the total field, while the right column shows the external  
 327 field (intrinsic field of the Earth removed). Spacecraft locations in MLT and GSM coordinates are displayed  
 328 below the time scale.

329 The  $B_x$  component of the total field (Figure 3a) is consistently negative throughout  
 330 the time period shown. This indicates that the spacecraft was located south of the current  
 331 sheet. The largest discrepancy between the modeled and observed  $B_x$  is an overestimation  
 332 of the magnitude of  $B_x$  by SWMFb and SWMFC between 0800 and 1600, visible in both  
 333 Figures 3a and 3b, with the greatest overestimation being 12.0 nT by SWMFb at 10:44  
 334 UT. Since the difference is present in the external field, it must be due to differences in  
 335 magnetospheric currents. The underestimation of the magnitude of  $B_x$  indicates that the  
 336 the current sheet in the simulation was farther north than actual, or that the model current  
 337 sheet was thinner or contained stronger currents than the actual one, resulting in a stronger  
 338 gradient in  $B_x$  across the current sheet. The data in Figure 3a cannot distinguish between  
 339 these two explanations, but given the finite grid resolution of the model it is more likely  
 340 that the current sheet would be thicker than observed rather than thinner, in which case the  
 341 underestimation of the  $B_x$  magnitude indicates an error in the current sheet location.

342 The estimation of the  $B_y$  component (Figure 3d) is somewhat more accurate than  
 343 that of  $B_x$ . All three models miss two negative diversions of the observed  $B_y$  that occur  
 344 at 0800 and 1230 UT. The largest discrepancy is an underestimation (in magnitude) of 7.6

345 nT by SWMFc at 17:45 UT, at which time the external  $B_y$  in SWMFc is -2.1 nT while the  
346 observations and the estimations by SWMFa and SWMFb are all around -10 nT. At 17:45  
347 UT the spacecraft was moving toward the Earth, and the  $B_y$  component of the total field  
348 was around -50 nT (Figure 3c).

349 From the external  $B_z$  plot (Figure 3f) it is apparent that all three model configura-  
350 tions tend to overestimate  $B_z$  throughout the time period of the plot, though there are a  
351 few brief periods of underestimation by SWMFa and SWMFb. The largest discrepancy is  
352 an overestimation of 11.2 nT by SWMFc at 0131 UT. This occurs at a time when the  $B_z$   
353 component of the total field was around 70 nT (Figure 3e). An overestimation of  $B_z$  was  
354 previously reported for SWMF in *Ganushkina et al.* [2010].

355 Similar results were obtained for THEMIS D and E as for THEMIS A. For THEMIS  
356 B and C the behavior of the observed magnetic field was substantially different from THEMIS  
357 A, D, and E due to the THEMIS B and C spacecraft having apogee farther down-tail.  
358 The model delivered similarly close estimations for the GOES 11 and 12 magnetic fields:  
359 The greatest diversion from observations by any single component was 13.3 nT, and most  
360 SWMF estimations were within 5 nT of observations. A persistent overestimation of  $B_z$   
361 (like that seen at THEMIS A) was found at GOES 12 but not at GOES 11.

362 The SWMF simulations estimated the fields with reasonably good accuracy overall.  
363 Many differences are present in the behavior of transient features, but typically the differ-  
364 ences between the model and observations even during these transients are of a magnitude  
365 of only a few nT, and the differences only occasionally exceed 10 nT. The general behav-  
366 ior of the fields is captured well by the simulation, and we find the estimations to be of  
367 sufficient quality to warrant their use in studying the IB.

## 368 **5 Mapping locations of isotropic boundaries with MHD**

369 Having verified that the SWMF simulations give reasonably good estimations of the  
370 magnetic fields in the magnetotail, we proceed to tracing magnetic fields from the loca-  
371 tions of the IB observations. This will enable us to use the SWMF output to obtain infor-  
372 mation about the conditions leading to IB formation. Figure 4 shows the results of trac-  
373 ing field lines from two of the IB observations through the magnetic fields computed by  
374 the SWMFb MHD simulation. These two events were selected as representative examples  
375 from the total of 21 traces (7 IB observations and 3 model runs). The left column (panels

376 a-e) show the IB crossing at 0225 UT, while the right column (panels f-j) shows the IB  
 377 crossing at 0422 UT. Panels a and f show the location of the IB observation, the field line  
 378 traced from the IB location, and satellite locations in the GSM  $x$ - $y$  plane. Panels b and g  
 379 show the same in the  $x$ - $z$  plane. The Earth is denoted by a black circle, and surrounded  
 380 by a grey circle representing the inner boundary of the MHD domain. The location of the  
 381 observed IB, mapped to the inner boundary of the MHD domain, is denoted with a small  
 382 circle. The locations of the THEMIS A and E spacecraft are also shown. The field line  
 383 traced from the IB location is shown as a dashed line, and the minimum  $|B|$  point along  
 384 this field line is denoted by an “X.”

385 **Figure 4.** Example of tracing field lines for two IB crossings, using the MHD solution from the SWMFb  
 386 simulation. Left column: IB crossing of the NOAA-16 spacecraft at 02:25 UT; right column: IB crossing of  
 387 the NOAA-17 spacecraft at 04:22 UT. (a), (b), (f), and (g) show the IB location, field line traced from the IB  
 388 location, and satellite positions in the GSM  $x$ - $y$  and  $y$ - $z$  planes. The IB location (mapped to  $2.5 R_E$ ) is shown  
 389 as a small circle. A dashed line denotes the field line traced from this location, projected into the plane of the  
 390 figure. The minimum  $B$  point along this line is denoted with an “X.” Locations of THEMIS A and E space-  
 391 craft are shown using the same symbols as in Figure 2. A solid line denotes a radial line through the minimum  
 392 B point, projected into the minimum  $B$  surface and subsequently into the the plane of the figure. This line  
 393 crosses the field line at the X. The remaining plots show quantities computed along this line, as a function of  
 394 GSM  $x$ . (c) and (h) show  $|B|$ , (d) and (i) show magnetic field line radius of curvature  $R_c$ , and (e) and (j) show  
 395  $K$ .

396 By tracing the field lines within each SWMF simulation, we computed a surface  
 397 defined by  $|B| = |B|_{min}$  along each field line. The solid line extending outward from the  
 398 Earth in Figures 4a, 4b, 4f, and 4g denotes a radial line in this surface from the center of  
 399 the Earth through the point where the IB field line intersects the minimum  $|B|$  surface.  
 400 Figures 4c-4e and 4h-4j show simulation output along this minimum  $|B|$  line as a function  
 401 of  $x$  in GSM coordinates. The location where the IB field line intersects this surface is  
 402 denoted with an “X.”

403 Figures 4c and 4h show  $|B|$  along the minimum  $|B|$  line described above. Figures 4d  
 404 and 4i show  $R_c$ , the field line radius of curvature, along the minimum  $|B|$  line. In both  
 405 cases, the point that maps to the IB location (denoted with an “X”) occurs near the maxi-

406 mum of  $R_c$ . This local maximum indicates a transition from a dipolar to a stretched field  
 407 configuration. Within a dipolar field, the radius of curvature increases linearly with dis-  
 408 tance, so the presence of a local maximum indicates that the field has diverged signifi-  
 409 cantly from dipolar. At the same time,  $|B|$  is continuing to decrease rapidly, resulting in  
 410 an increase in the gyroradius  $r_g$  and making it increasingly likely that the conditions for  
 411 adiabatic motion will be violated. Similar plots of  $R_c$  as a function of  $x$  can be found in  
 412 *Sergeev and Tsyganenko* [1982] and *Yue et al.* [2014], in both cases produced using empir-  
 413 ical magnetic field models, and these exhibit generally similar behavior. The  $R_c$  plots in  
 414 both *Sergeev and Tsyganenko* [1982] and *Yue et al.* [2014] reaches a maximum at a some-  
 415 what closer distance to the Earth and drops off more rapidly compared to the plots in Fig-  
 416 ures 4c and 4h of the present paper. However, the behavior of  $R_c$  in the magnetotail is  
 417 quite volatile and depends strongly on local conditions in the plasma sheet and on the gen-  
 418 eral disturbance level of the magnetosphere.

419 The plots in Figures 4c and 4h show larger  $R_c$  than either *Sergeev and Tsyganenko*  
 420 [1982] or *Yue et al.* [2014]. The position at which the maximum  $R_c$  occurs is also slightly  
 421 farther from the Earth than in *Yue et al.* [2014]. This indicates that the MHD fields shown  
 422 in this figure are less stretched than those shown in the similar figures of *Sergeev and Tsy-*  
 423 *ganenko* [1982] and *Yue et al.* [2014].

424 A local maximum of  $R_c$  combined with a rapidly falling  $|B|$  implies a move toward  
 425 conditions favorable for pitch angle scattering. However, the point where pitch angle scat-  
 426 tering occurs is controlled more directly by  $K$ . Figures 4e and 4j show  $K$  as a function  
 427 of  $x$ .  $K$  is shown on a logarithmic scale; the motivation for the logarithmic scale will be  
 428 addressed later. For both of these,  $K$  decreases monotonically as a function of  $x$ . The  $K$   
 429 values at the points mapped from the IB locations (the points marked with an “X”) in Fig-  
 430 ure 4e and 4j are 18.9 and 26.0, respectively.

431 From Figures 3 and 4 we see that the SWMF simulations are tending to produce  
 432 magnetic fields that are more dipolar (less stretched) than actual, resulting in larger  $B_z$ ,  
 433 and likely causing an overestimate of  $K$ . We can quantify the discrepancy in the field  
 434 stretching by directly comparing with available in situ magnetic field measurements for  
 435 each IB observation. For this purpose, we use in situ  $B$  observations from the THEMIS  
 436 satellites that are conjugate with each IB observation under the criteria specified in Sec-  
 437 tion 2. We define  $\overline{\Delta B_z} = \overline{(B_{z,model} - B_{z,observations})}$  as the average of error in  $B_z$  at the

453 **Figure 5.**  $K$  as a function of the error  $\overline{\Delta B_z}$  averaged over the THEMIS spacecraft that were conjugate with  
 454 each IB. Black line denotes a fit to the SWMFa and SWMFc points.

438 spacecraft that are conjugate with each IB observation. Figure 5 shows  $K$  as a function  
 439 of  $\overline{\Delta B_z}$  for all simulations and all the IB observations shown in Table 1. The y-axis of  
 440 Figure 5 is shown on a logarithmic scale. It is apparent from the figure that the  $K$  values  
 441 obtained from SWMFa and SWMFc show a common dependence on  $\overline{\Delta B_z}$ . A least squares  
 442 fit to the combined  $K$  estimates from SWMFa and SWMFc is plotted on top of the points.  
 443  $K$  estimates from SWMFb have been omitted from the fit because they exhibited a sub-  
 444 stantially different dependence on  $K$ : Fitting to the SWMFb points produces a lower slope  
 445 and higher intercept, while the SWMFa and SWMFc data had similar slopes and inter-  
 446 cepts to each other. To assess the quality of the fits, we use coefficient of determination  
 447 ( $R^2$ ), defined as

$$R^2 = 1 - \frac{\sum_i (K_i - \bar{K})^2}{\sum_i (K_i - f(\overline{\Delta B_z}))^2}, \quad (2)$$

448 where  $f(\overline{\Delta B_z})$  is the regression curve.  $R^2$  represents the fraction of the variation in  
 449  $K$  that is explained by  $f(\overline{\Delta B_z})$ , and for a perfect fit we would get  $R^2 = 1$ . For SWMFb,  $R^2$   
 450 was only 0.02, while SWMFa and SWMFc had  $R^2$  values of 0.65 and 0.60, respectively.  
 451 The poor fit for SWMFb seems to be due in part to the influence of outlier points and in  
 452 part to the points being clustered in a fairly narrow range of  $\overline{\Delta B_z}$ .

455 The common dependence of  $K$  on  $\overline{\Delta B_z}$ , and the fact that it is similar for both SWMFa  
 456 and SWMFc, can be explained by first noting that  $K$  can be approximated by

$$K = \frac{R_c}{r_g} \approx \frac{qB_z^2}{\sqrt{2mE}dB_r/dz}, \quad (3)$$

457 where  $q$  denotes the particle charge,  $m$  the particle mass, and  $E$  the particle energy.  
 458  $B_r$  is the radial component of magnetic field in GSM coordinates, defined as

$$B_r = \frac{x B_x + y B_y}{\sqrt{x^2 + y^2}}. \quad (4)$$



459 Since we are working with protons,  $q$  and  $m$  are the elementary charge and the proton  
 460 mass.  $B_z$  and  $B_r$  denote the magnetic field  $z$  and radial components in GSM coordi-  
 461 nates. This expression indicates a quadratic relationship between  $B_z$  and  $K$ . We take the  
 462 logarithm of both sides to obtain  $\log K \propto \log B_z$ , and then linearize the right hand side of  
 463 the equation to obtain

$$\log K = A_1 + A_2 \overline{\Delta B_z}, \quad (5)$$

464 which is the expression for the fit curve shown in Figure 5.  $A_2$  represents the linear  
 465 dependence of  $K$  on  $B_z$ , which is the result of both the expression for  $K$  and the charac-  
 466 teristics of the model and the physical system it represents. By setting  $\overline{\Delta B_z} = 0$  we obtain  
 467  $K_0 = \exp(A_1) = 16$ , an estimate of the average  $K$  in the absence of  $B_z$  error. In order to  
 468 determine the uncertainty associated with  $A_1$ , we estimate the 95% confidence interval for  
 469 the intercept  $A_1$  as described in e.g. *Montgomery et al.* [2012]. The 95% confidence inter-  
 470 val for  $\exp(A_1)$  is [13, 19], indicating that the intercept value of  $K_0 = 16$  is significantly  
 471 lower than the value of  $K = 33$  reported in *Ilie et al.* [2015], but also significantly higher  
 472 than the commonly used threshold value  $K_{crit} = 8$  [*Sergeev et al.*, 1983].

## 473 6 Estimation of $K$ using empirical models

474 In this study we use the following empirical magnetospheric models: *Tsyganenko*  
 475 [1995] (T96), *Tsyganenko* [2002] (T01), *Tsyganenko and Sitnov* [2005] (TS05), *Tsyga-*  
 476 *nenko and Andreeva* [2015] (TA15N and TA15B), and *Tsyganenko and Andreeva* [2016]  
 477 (TA16RBF). For all these models, the magnetic field is built as a sum of analytical func-  
 478 tions of the model's driving parameters. These functions can represent separate current  
 479 systems (T96–TA15), or just be basis functions which have no counterpart among conven-  
 480 tional current systems in the magnetosphere. Below we briefly summarize the model's fea-  
 481 tures which can be critical for  $K$ -parameter estimation. Since we analyze a quiet event, the  
 482 ring current is very weak and the magnetic field at the region of expected IB formation as  
 483 well as ionosphere-magnetosphere mapping are mostly controlled by the tail current sys-  
 484 tem.

485 The distinctive feature of the T96 model is its optimization for an accurate iono-  
 486 sphere magnetosphere mapping: the angular difference between the modeled and observed  
 487 magnetic field vectors was minimized when T96 was fitted to observed magnetic field vec-

488 tors. For all other models, the squared difference of magnetic field components was min-  
489 imized. On the other hand, since T96 is the earliest model, its structure is rather simple.  
490 The T96 tail current thickness and position with respect to the Earth is fixed and only its  
491 intensity changes with activity which is parameterized by Akasofu epsilon parameter.

492 The T01 model represents a further development of the concept used for T96. The  
493 T01 tail model thickness does not vary with activity but the tail current inner edge moves  
494 towards the Earth when activity grows. In addition, the tail module includes two indepen-  
495 dent sub-modules which add a flexibility to the tail current radial profile.

496 The TS05 model was specially designed to describe the magnetosphere during storm-  
497 time disturbances. The model was fit to the measurements made during storm intervals  
498 and likely should not be used for modeling a quiet event. Its tail current thickness and  
499 position vary with activity which is parameterized by complex integral functions of the so-  
500 lar wind plasma and IMF parameters taking into account the prehistory of the conditions  
501 in the solar wind. TS05 has the most advanced parameterization of geomagnetic activity  
502 among all empirical models.

503 The TA15 model is a forecasting model which is parameterized entirely by exter-  
504 nal parameters (no ground-based indices are used). The tail current thickness and position  
505 vary with activity and the current intensity is assumed to fall off with distance as a power  
506 law. There are two versions of the model referred to as TA15N, parameterized by the uni-  
507 versal coupling function described in *Newell et al.* [2007], and TA15B, parameterized by  
508 an external driving function described in *Boynnton et al.* [2011].

509 It should be noted, that for T96–TA15 models, to a large extent, the current system  
510 geometries are determined in an ad hoc manner. For example: although the tail current  
511 thickness in TS05 and TA15 models can vary with activity, its variance with activity is  
512 parameterized for the current sheet as a whole, and its variance with X and Y is defined  
513 by a realistic but arbitrary function. This drawback was corrected in the TA16RBF model  
514 where the field is represented as a sum of elementary functions, each depending on ac-  
515 tivity parameters including the universal coupling function of *Newell et al.* [2007]. This  
516 approach minimizes the role of the authors in the current geometry definition. However,  
517 the amount of the existing spacecraft measurements does not allow the model to resolve  
518 fine spatial structures.

544 **Figure 6.** Results of  $K$  estimation for the four empirical models. Black line denotes a least-squares fit to all  
 545 points except those produced by TS05.

519 Using these models and geomagnetic dipole to represent the Earth's inner field, we  
 520 traced the field lines from IB locations (in AACGM coordinates). The standard field line  
 521 tracing subroutine from Geopack FORTRAN package was used. The package as well as  
 522 the model input driving parameters can be found at <http://geo.phys.spbu.ru/~tsyganenko/modeling.html>.  
 523  $K = \frac{R_c}{r_g}$  was estimated at the minimum magnetic field point (in the equatorial part of the  
 524 field line), using Equation 1 to estimate  $R_c$  and the cutoff energies from Table 1 to esti-  
 525 mate  $r_g$ . The output from the models was also computed at the locations of the THEMIS  
 526 probes and was compared with THEMIS observations.

527 Figure 6 shows the computed  $K$  values versus  $\overline{\Delta B_z}$  for all aforementioned models  
 528 in the same format as in Figure 5 for MHD. As before,  $\overline{\Delta B_z}$  represents the model error  
 529 on  $B_z$ , averaged over the THEMIS satellites that were conjugate with each IB observation.  
 530 The conjugate THEMIS satellites were selected as described in Section 2. It is apparent  
 531 that all models except for TS05 exhibit a common  $K$  vs.  $\overline{\Delta B_z}$  dependence. It can be seen  
 532 that T96 and T01 show the best agreement with observations, based on their comparatively  
 533 low values of  $\overline{\Delta B_z}$ . TA16RBF shows the worst agreement, with observations systemat-  
 534 ically underestimating the  $B_z$  values. The TS05 model produced unrealistically low  $K$ -  
 535 values for 6 of 7 IBs. It can be speculated that the deviation of TS05 from the common  
 536 dependence is a result of its narrow specification. Since it was fitted to storm-time obser-  
 537 vations, quiet periods perhaps comprised a negligible part of the data. For this reason, the  
 538 TS05 model output is an extrapolation in the region of quiet external conditions. Equa-  
 539 tion 5 has been fit to the points of the plot, excluding TS05 due to the substantial number  
 540 of outliers present, and the resulting fit curve is plotted in black. The y-axis intercept of  
 541 this fit occurs at  $K = K_0 = 2.6$ , with a 95% confidence interval of [2.1, 3.2]. This is sig-  
 542 nificantly lower than the  $K_0 = 16$  obtained with SWMF.  $K_0 = 2.6$  falls within the range  
 543 expected for CSS, but is appreciably below  $K_{crit} = 8$ .

## 546 7 Comparison of empirical and MHD results

547 Comparing between Figures 5 and 6, large differences are apparent between the  
 548 SWMF simulations and the empirical models. In a preliminary effort to account for the

549 influence of magnetic field errors  $\overline{\Delta B_z}$ , we fitted Equation 5 to each dataset and noted  
 550 the y-axis intercepts, which provide an estimate for the average  $K$  in the absence of these  
 551 errors. Even these values differed by nearly an order of magnitude between the SWMF  
 552 simulations and the empirical models. This indicates a significant difference in the depen-  
 553 dency between  $K$  and  $B_z$  between the different models. One possible reason for this is a  
 554 systematic difference in  $R_c$  within the current sheet, which depends on the model current  
 555 sheet thickness. Given the sparsity of the available observational data, we have no means  
 556 to determine the true value of  $R_c$ . However, the THEMIS configuration allows a rough  
 557 estimation of  $G = \frac{dB_r}{dz}$ , which appears in the denominator of Equation 3, and is a major  
 558 contributor to the value of  $R_c$ . As was mentioned in Section 2, it is evident from Figures  
 559 2c and 2d that THEMIS A and E are closely spaced in the  $x$  and  $y$  directions but are sig-  
 560 nificantly displaced in the  $z$  direction. This enables us to estimate  $G$  as

$$G \approx \frac{B_{rTHE} - B_{rTHA}}{z_{THE} - z_{THA}}, \quad (6)$$

561 where the subscripts  $THA$  and  $THE$  refer to the spacecraft THEMIS A and E, and  
 562 the values  $z_{THA}$  and  $z_{THE}$  refer to the GSM  $z$  coordinates of the spacecraft.

563 To test the validity of this estimate, we define

$$\Delta r_{xyE-A} = \sqrt{(x_{THE} - x_{THA})^2 + (y_{THE} - y_{THA})^2} \quad (7)$$

564 and

$$\Delta z_{E-A} = z_{THE} - z_{THA}, \quad (8)$$

565 where again the positions are given in GSM coordinates and the subscripts  $THA$   
 566 and  $THE$  refer to the spacecraft THEMIS A and E. By taking the satellite positions at  
 567 the seven selected IB observations shown in Table 1, we find that the ratio  $\frac{\Delta z_{E-A}}{\Delta r_{xyE-A}}$  ranges  
 568 from 1.44 and 2.3, indicating that the displacement in  $z$  is consistently greater than the  
 569 horizontal displacement  $\Delta r_{xyE-A}$ . The ratio 1.44 not ideal since it indicates the horizontal  
 570 distance between the spacecraft is significant compared with their displacement in  $z$ . This  
 571 may contribute to the uncertainty in estimates of  $G$ , since any horizontal ( $x$  or  $y$ ) gradient  
 572 in  $B_r$  will influence the estimate of  $G$ . Another source of uncertainty is the distance of the

575 **Figure 7.** Observed  $B_r$  and  $B_z$  (GSM) for THEMIS A and E. IB crossing times are denoted with verti-  
 576 cal dashed lines. Both spacecraft had an elliptical orbit with apogee around  $11 R_E$  from the Earth, but the  
 577 THEMIS A spacecraft was located approximately  $1 R_E$  southward at apogee relative to THEMIS E.

573 spacecraft from the center of the current sheet, since we will apply this  $G$  estimation to  
 574 correct  $K$  estimations computed in the current sheet.

578 To obtain reasonable estimates for  $G$ , we need  $B_r$  to differ appreciably between the  
 579 THEMIS A and E spacecraft. Figure 7 shows the observed  $B_r$  and  $B_z$  magnetic field com-  
 580 ponents at THEMIS A and E. Both spacecraft observed positive  $B_r$ , indicating that they  
 581 are south of the current sheet. However, the  $B_r$  component is comparatively weak ( $<10$  nT  
 582 around apogee) at THEMIS E, while it is 15-20 nT greater at THEMIS A. This indicates  
 583 that THEMIS E was located fairly close to the current sheet (for most of the day), while  
 584 THEMIS A was farther below it. For the last IB crossing (16:54 UT), the  $B_r$  component  
 585 at THEMIS E is 30.0 nT, significantly stronger than for the earlier IB crossings, indicat-  
 586 ing a more significant displacement from the current sheet. However, even for the 16:54  
 587 UT IB crossing the field at THEMIS E is still significantly weaker (by 11.6 nT) than that  
 588 observed by THEMIS A at that time. It is worth noting that  $B_z$  also differs significantly  
 589 between the two spacecraft. This means that the configuration is not 1-D. In a 1-D mag-  
 590 netic field (one in which  $\frac{d\mathbf{B}}{dx} = \frac{d\mathbf{B}}{dy} = 0$ ),  $B_z$  cannot vary with  $z$  without violating  $\nabla \cdot \mathbf{B} = 0$ .  
 591 While the current sheet can sometimes be approximately 1-D, the fact that  $B_z$  varies with  
 592  $z$  in the THEMIS observations implies that gradients in  $x$  and  $y$  are present, which will  
 593 contribute to errors in estimating  $G$ .

594 The SWMF and empirical model results are shown together in Figure 8. Like Fig-  
 595 ures 5 and 6, Figure 8 shows  $K$  plotted as a function of  $\overline{\Delta B_z}$ . However, the points are now  
 596 colored according to  $\Delta G = G_{model} - G_{obs}$ . For both the observations and the models, we  
 597 estimate  $G$  using the difference between the THEMIS A and THEMIS E locations accord-  
 598 ing to Equation 6. As shown in Figure 2, the two spacecraft were located relatively close  
 599 in GSM  $x$ , but were displaced in  $z$ , and this resulted in significantly different  $B_r$  compo-  
 600 nents as shown in Figure 7.

604 Fits to the SWMF points and the empirical model points (the same fits shown in  
 605 Figures 5 and 6) are drawn in black in Figure 8. It is apparent that the simulations and

601 **Figure 8.** Estimated  $K$  as a function of  $\overline{\Delta B_z}$  for the SWMF simulations (except SWMFb) and the empirical  
 602 models (except TS05). Colors denote  $\Delta G$ , the estimation error in the derivative  $G = \frac{dB_z}{dz}$ . Black lines denote  
 603 least-squares fits to the SWMF simulations and empirical models.

606 empirical models are producing distinctly different results, with the SWMF consistently  
 607 producing higher  $K$  values. At the same time, the empirical models tend to overestimate  
 608  $G$ , while the SWMF simulations tend to underestimate  $G$ . From Equation 3 we expect that  
 609 overestimations of  $G$  will lead to smaller  $K$  values, and vice versa. Thus, the systematic  
 610 underestimation of  $G$  by SWMF contributes to its comparatively large  $K$  estimates, and  
 611 the systematic overestimation of  $G$  by the empirical models contributes to their smaller  $K$   
 612 estimates.

613 We now attempt to correct the  $K$  estimates based on the available observations.  
 614 We first apply a correction for  $\overline{\Delta B_z}$  using the same procedure used in *Dubyagin et al.*  
 615 [2018]. Equation 5 provides an average relationship between  $K$  and  $\overline{\Delta B_z}$  for each of the  
 616 two classes of models (MHD and empirical). We use this relationship to correct particular  
 617  $K$  values by replacing the fitted y-axis intercept  $A_1$  with a value  $\log K^*$ , where  $K^*$  is the  
 618 value for a particular  $K$  that would be computed in the absence of  $B_z$  error, assuming the  
 619 linear relationship of Equation 5 holds. This produces the relationship

$$\log K = \log K^* + A_2 \overline{\Delta B_z}, \quad (9)$$

620 from which we compute  $K^*$  as

$$K^* = K \exp(-A_2 \overline{\Delta B_z}). \quad (10)$$

621 Furthermore, motivated by the trends in  $G$  noted in Figure 8, we construct a new  
 622 linearization of Equation 3, incorporating both  $\overline{\Delta B_z}$  and  $\Delta G$ :

$$\log K = C_1 + C_2 \overline{\Delta B_z} + C_3 \Delta G. \quad (11)$$

623  $C_1$ ,  $C_2$ , and  $C_3$  are obtained using least squares minimization using all of the points  
 624 shown in Figure 8 (that is, the output from SWMFa, SWMFc, and all the empirical mod-

625 els except TS05). Solving this equation for  $K$  in the case of  $\overline{\Delta B_z} = \Delta G = 0$  gives  $K_0 =$   
 626  $\exp(C_1) = 6.0$  with a 95% confidence interval of [4.7, 7.8]. This can be regarded as an  
 627 average corrected  $K$  for the case of zero magnetic field error. It falls within the expected  
 628 range for CSS, but somewhat below the  $K_{crit} = 8$  threshold.

629 For the general case (non-zero field error), we obtain a new correction  $K^{**}$ , which  
 630 we derive from Equation 11 by substituting  $K^{**}$  for  $C_1$  giving

$$\log K = \log K^{**} + C_2 \overline{\Delta B_z} + C_3 \Delta G, \quad (12)$$

631 which we can solve for  $K^{**}$  to obtain

$$K^{**} = K \exp(-C_2 \overline{\Delta B_z} - C_3 \Delta G). \quad (13)$$

632 Figure 9 shows the overall effect of applying these two correction schemes. The fig-  
 633 ure shows probability density functions (PDFs) of  $K$ ,  $K^*$ , and  $K^{**}$ , obtained through ker-  
 634 nel density estimation [Parzen, 1962]. Kernel density estimation approximates a PDF as  
 635 a sum of Gaussian kernel functions centered at each data point. The resulting plot can  
 636 be interpreted as a normalized histogram. The mean, median, and interquartile range  
 637 (IQR) of the distributions shown in Figure 9 can be found in Table 2, along with the y-  
 638 axis intercepts (with 95% confidence intervals) of the fits used to compute  $K^*$  and  $K^{**}$ .  
 639 Figure 9a shows PDFs of  $K$  for SWMFa and SWMFc, and for all empirical models ex-  
 640 cept TS05. The two distributions overlap negligibly, with the SWMF  $K$  values uniformly  
 641 greater than those of the empirical models. Figure 9b shows PDFs of  $K^*$ , with the fit to  
 642 Equation 5 having been performed separately for the empirical models and for SWMF.  
 643 This correction leads to both distributions becoming more narrow. The many small  $K$  val-  
 644 ues of the empirical models are shifted higher, while many of the larger SWMF values are  
 645 shifted down. There is still no overlap between the two distributions, however. Figure 9c  
 646 shows PDFs of  $K^{**}$ . The two classes of models did not individually have a clear linear re-  
 647 lationship between  $K$  and  $G$ , so to compute  $K^{**}$  a single fit was performed for all models  
 648 together, once again omitting TS05 and SWMFb. Some overlap can be seen in the dis-  
 649 tributions of  $K^{**}$  in Figure 9c, although the corresponding interquartile ranges shown in  
 650 Table 2 for  $K^{**}$  do not overlap.

651 **Figure 9.** Kernel density plots of (a)  $K$ , (b)  $K^*$ , and (c)  $K^{**}$  for SWMFa, SWMFb, and for all empirical  
 652 models except TS05.

Model type	25th percentile	Median	75th percentile	y-axis intercept	Intercept 95% CI
<i>Uncorrected K</i>					
SWMF	18	21	27	–	–
Empirical	0.93	2.3	4.6	–	–
Both	1.2	4.3	17	–	–
<i>K* correction</i>					
SWMF	14	15	19	16	[13, 19]
Empirical	2.0	2.8	3.7	2.6	[2.1, 3.2]
Both	2.5	3.5	13	–	–
<i>K** correction</i>					
SWMF	7.6	9.9	11	6.0	[4.7, 7.8]
Empirical	4.4	5.4	7.6	6.0	[4.7, 7.8]
Both	4.7	6.2	9.5	6.0	[4.7, 7.8]

653 **Table 2.** Summary of the results of the correction schemes. The 25th percentile, median, and 75th per-  
 654 centile are shown for  $K$ ,  $K^*$ , and  $K^{**}$ . For  $K^*$  and  $K^{**}$  the y-axis intercept (i.e., the correction value for the  
 655 special case of  $\overline{\Delta B_z} = \Delta G = 0$ ) and its 95% confidence interval are shown. Note that in the case of  $K^{**}$  the  
 656 y-axis intercept is the same for both SWMF and empirical models since a fit is obtained by fitting to both  
 657 types of models simultaneously and the same fit is used to correct both types of models.



## 8 Discussion and conclusions

The goal of this study is to test the viability of using MHD for estimating  $K$  values associated with IB observations, and to provide better constraints on the range of  $K$  associated with the IB during quiet time. To accomplish this, we used multiple models, both physics-based and empirical, to compute  $K$  values corresponding to IB observations. We ran three SWMF simulations using different simulation parameters, as well as six different empirical models. We used the same time interval and the same IB observations that were used by *Ilie et al.* [2015]. A quiet interval was chosen in order to reduce the chance of particle scattering due to wave-particle interactions, which occur primarily during active periods [*Bräysy et al.*, 1998; *Halford et al.*, 2010; *Usanova et al.*, 2012], and to reduce possible inaccuracies in the model mapping. In addition, *Ilie et al.* [2015] noted an absence of observational evidence for EMIC waves during the interval chosen. We first traced fields from the locations where the IB was observed through the magnetic fields of each model. We computed  $K$  at the point of minimum  $|B|$  along each of these field lines. However, this computed  $K$  is still subject to errors in the model field relative to reality. To address this, we corrected each  $K$  value based on the error in  $B$  relative to observations by nearby THEMIS spacecraft. By computing  $K$  with multiple models, and correcting for the errors in those models, we are able to better constrain the possible range of  $K$  associated with CSS driven IB formation, and to test the hypothesis that these quiet-time IB observations are associated with CSS.

Comparison between the SWMF simulations and in situ observations indicates a tendency toward an under-stretched night side magnetic field geometry in the simulations. This same tendency was reported in *Ilie et al.* [2015] for the same event with SWMF, and other previous papers have also reported under-stretched fields with MHD [e.g. *Welling and Ridley*, 2010; *Ganushkina et al.*, 2010], though this can be reduced through the use of non-ideal MHD [e.g. *Meng et al.*, 2013]. The empirical models exhibited a range of behaviors in terms of magnetotail stretching: The T01 fields tend to be under-stretched and the TS05, TA16RBF, and TA15N fields are consistently over-stretched, while the T96 and TA15B fields are sometimes over-stretched and sometimes under-stretched. This reflects a variety of observational datasets used to construct the models, as well as different strategies used for fitting. There has been some evolution in terms of field stretching in empirical models. The T96 model has been previously reported as producing over-stretched fields, but the over-stretching was most pronounced during disturbed periods

[Tsyganenko, 2001; Ganushkina et al., 2002; Huang et al., 2008]. Huang et al. [2008] reported TS05 as producing less field stretching than T96, but this was for a storm interval and at closer distances than our THEMIS observations. The fact that our results do not show this tendency in T96 may be due in part to the use of a quiet period for this study. The TA16RBF model was previously reported as having similar tail stretching as the older T89 model in the 10-12  $R_E$  region [Tsyganenko and Andreeva, 2016], and T89 in turn has been reported as producing over-stretched fields in the magnetotail [Tsyganenko, 1989; Peredo et al., 1993].

At the same time, the SWMF consistently produces negative values for  $\Delta G$ , implying a thicker and/or weaker than actual current sheet. The empirical models, on the other hand, uniformly overestimate  $G$ , which implies a current sheet that is thinner and/or stronger than actual. The combined effect of the discrepancies in  $B_z$  and  $G$  is that the SWMF consistently produces large  $K$  values, while the empirical models generally produce smaller  $K$ .

The large range of  $K$  values we obtain, and their association with magnetic field errors, highlights the importance of quantifying the influence of magnetic field errors on the  $K$  estimates, as we did in computing  $K^*$  and  $K^{**}$ . However, it should be noted that  $K^*$  and  $K^{**}$  should be regarded as rough estimates of the true value of  $K$ . The corrections are subject to a number of sources of uncertainty. These include errors due to linearization of the nonlinear  $K$  formula, the position of the satellites relative to the  $|B|_{min}$  point, and the fact that we perform linear fits of multiple models and multiple IB observations, effectively averaging them together. Estimations of the dependence of  $K$  on  $G$  are especially sensitive to the satellite positions, because  $G$  has a larger gradient in  $z$  (compared to  $B_z$ ), and because the errors due to the positions of the spacecraft relative to the  $|B|_{min}$  point are compounded by the errors contributed by the positions of the spacecraft relative to each other (on which the estimation of  $G$  depends). A more complex  $K$  correction that accounts for the position of the satellites relative to the  $|B|_{min}$  point might be possible, but this could be error prone as it would require introducing new assumptions about the variation of  $B_r$  and  $B_z$  with position. It is possible that some effects of the relative distances and directions to spacecraft are somehow corrected for as a result of the fitting process used to obtain the correction factors, but this is by no means certain. A final source of uncertainty is the fact that we include only the  $G$  and  $B_z$  contributions in the correction of

723  $K$ . Including additional field components would likely result in over-fitting given the small  
724 number of points used in our analysis, but might be appropriate for a larger scale study.

725 In addition, the criteria for selecting conjugate spacecraft may impact the estima-  
726 tions of  $G$  and  $B_z$ . One could conceive of alternate criteria, using either wider or narrower  
727 ranges in MLT or distance from the Earth, or using locations relative to the traced field  
728 lines as part of the selection criteria. The simple mean of  $\Delta B_z$  over all conjugate space-  
729 craft could also be replaced with another measure of central tendency, perhaps a weighted  
730 average applying higher weights to spacecraft that were closer to the field line mapping to  
731 the NOAA or METOP spacecraft. However, it is not obvious that the spacecraft closest to  
732 the mapped location will provide a better error estimate in practice, and *Dubyagin et al.*  
733 [2018] attempted several of these alternative methods and found no significant improve-  
734 ment in the correlation between  $K$  and  $\Delta B_z$ .

735 Despite the aforementioned sources of uncertainty, the narrow range of values ob-  
736 tained for  $K^{**}$  (compared with the range of the uncorrected  $K$ ), and the overlap between  
737 the two classes of models, indicate that the procedure is largely successful in accounting  
738 for the influence of magnetic field errors on  $K$  estimates in the models. This means that  
739 both  $B_z$  and  $G = \frac{dB_z}{dz}$  contribute significantly to errors in estimating  $K$ , and that the two  
740 corrections in combination account for a large fraction of the variation in model  $K$  estima-  
741 tions.

742 The relatively narrow ranges we obtained for  $K^{**}$  also suggests that the IB events we  
743 analyzed were formed by a common mechanism which depends primarily on field geome-  
744 try. A majority of the  $K^{**}$  values (53% for SWMF, and 91% for the empirical models) are  
745 below  $K = 10$ , which implies that CSS is the likely mechanism forming the IB in these  
746 cases. The fact that we obtain this using two classes of models, with very different under-  
747 lying assumptions, lends additional confidence to this result. By comparing the IQR we  
748 obtained for  $K$  with the results of *Ilie et al.* [2015], we find that our results and systemati-  
749 cally smaller, and slightly more tightly distributed, than those obtained by *Ilie et al.* [2015]  
750 for the same event. The IQR of  $K$  was 30-42 in the data reported by *Ilie et al.* [2015]; the  
751 fact that this range does not overlap with the IQR of our data indicates that the difference  
752 between their results and ours is larger than the degree of variation within each dataset.  
753 That we obtain smaller values of  $K$  over a narrower range compared with *Ilie et al.* [2015]  
754 indicates that more accurate dipole parameters and other changes to the simulation set-

755 tings had a substantial effect. We obtain a much narrower IQR for  $K^{**}$ , indicating that a  
756 large part of the variation in SWMF-derived  $K$  can be attributed to correctible errors in  
757 the model's magnetic field estimation, rather than to differences in the IB formation pro-  
758 cess over time. The range of  $K^{**}$  is also narrower than the range of  $K$  values obtained by  
759 *Sergeev et al.* [2015b]. In the present work we aimed to only study IB's that were likely  
760 to have been formed by CSS. To that end we chose a quiet-time interval, checked for the  
761 absence of EMIC wave observations, and only used events that had typical appearance.  
762 *Sergeev et al.* [2015b] studied a longer time period which was mostly, but not exclusively,  
763 quiet time. Since they do not describe any special efforts to restrict their analysis to "typ-  
764 ical" IB observations, it is likely they used all available events that were conjugate with  
765 THEMIS spacecraft. Since *Sergeev et al.* [2015b] used adaptive models that attempted to  
766 minimize the magnetic field error relative to the THEMIS observations, the fact that we  
767 obtain a narrower range for  $K^{**}$  is likely due to the use of a longer time period, more IB  
768 observations, and a wider range of conditions in *Sergeev et al.* [2015b] compared to the  
769 present work.

770 The  $K = 6.0$  value we obtained is marginally lower than the commonly used  $K_{crit} =$   
771  $8$  threshold. Since the  $K_{crit} = 8$  is outside the 95% confidence interval for  $K$ , our result  
772 is significantly lower in a statistical sense. However, this result depends on the use of both  
773 the SWMF and the empirical models together in the fit. The y-axis intercept values in the  
774 fourth and fifth rows of Table 2 (2.6 for the empirical models and 16 for SWMF) can be  
775 regarded as probable lower and upper bounds for the range of variability that could result  
776 if we changed the mixture of models used in the study. That is, we would expect the inter-  
777 cept for the combined  $K$  estimate (using all models together) to range from roughly 2.6 to  
778 roughly 16 if we were to modify the mix of models included in the  $K^{**}$  correction.

779 Our work shows that even after restricting the analysis to a few quiet-time IB ob-  
780 servations with ideal spatial distribution of fluxes across the IB, and after correcting the  
781 results for errors in the magnetic field models, a substantial uncertainty remains in the  
782 range of  $K$ . The remaining uncertainty is consistent with the dependency on the incident  
783 particle population reported by *Delcourt et al.* [1996], so there is no need to invoke an ad-  
784 ditional scattering mechanism to explain it. Therefore, obtaining a more precise threshold  
785 condition than the range we obtained will likely require not only an improved magnetic  
786 field model but also an accounting for the phase space distribution of particles prior to  
787 scattering.

788 The conclusions of the paper can be summarized as follows:

- 789 1. Prior to correction, the MHD simulations often (though not always) produce  $K$  val-  
790 ues above the expected range for CSS, while the empirical models produce  $K$  val-  
791 ues at the low end of the expected range for CSS.
- 792 2. The corrected  $K^*$  and  $K^{**}$  values have much narrower spread than the uncorrected  
793  $K$  values. This implies that much of the spread in  $K$  was due to errors in the esti-  
794 mated magnetic field, but also shows that these errors can be corrected.
- 795 3. The distributions of  $K^{**}$  overlap substantially for the MHD and for the empirical  
796 models.
- 797 4. A majority of the corrected  $K^{**}$  values (53% for SWMF, and 91% for the empirical  
798 models) fall within the expected range for CSS, which supports the hypothesis that  
799 the IB was formed by CSS in those cases.
- 800 5. We estimate  $K_0$ , the average value of  $K$  in the absence of magnetic field error,  
801 to be 6.0. This is within the expected range for CSS, and somewhat lower than  
802  $K_{crit} = 8$ . However, the uncertainty range is estimated as [2.6, 16], which extends  
803 above  $K_{crit} = 8$ .

#### 804 **Acknowledgments**

805 The research of S. Dubyagin and N. Ganushkina leading to these results was partly  
806 funded by the European Union's Horizon 2020 research and innovation programme un-  
807 der grand agreement 637302 PROGRESS. The work of N. Ganushkina and J. Haiducek  
808 in Michigan was also partly supported by the National Aeronautics and Space Administra-  
809 tion under grant agreements NNX14AF34G, NNX17AB87G and 80NSS17K0015 issued  
810 through the ROSES-2013 and ROSES-2016 programmes. Portions of the results presented  
811 here were achieved under the framework of the Finnish Centre of Excellence in Research  
812 of Sustainable Space (FORESAIL) funded by the Academy of Finland, decision number  
813 312351.

814 The NOAA/POES particle data were downloaded from National Geophysical Data  
815 Center website: <https://www.ngdc.noaa.gov/stp/satellite/poes/index.html>.

816 THEMIS magnetic field data were loaded and calibrated using official software at  
817 <http://themis.ssl.berkeley.edu>.

818 GOES magnetic field data were obtained from CDAWeb (<https://cdaweb.sci.gsfc.nasa.gov/>).

819 The precomputed input parameters for the empirical models were downloaded from  
820 <http://geo.phys.spbu.ru/~tsyganenko/modeling.html>.

821 Full output from the three model runs presented in this paper can be found at the  
822 following URLs:

823 [http://vmr.engin.umich.edu/Model/\\_swmf\\_mag/plot?run=IB\\_quiet\\_feb2009\\_SWMFa](http://vmr.engin.umich.edu/Model/_swmf_mag/plot?run=IB_quiet_feb2009_SWMFa)

824 [http://vmr.engin.umich.edu/Model/\\_swmf\\_mag/plot?run=IB\\_quiet\\_feb2009\\_SWMFB](http://vmr.engin.umich.edu/Model/_swmf_mag/plot?run=IB_quiet_feb2009_SWMFB)

825 [http://vmr.engin.umich.edu/Model/\\_swmf\\_mag/plot?run=IB\\_quiet\\_feb2009\\_SWMFC](http://vmr.engin.umich.edu/Model/_swmf_mag/plot?run=IB_quiet_feb2009_SWMFC)

## 826 References

827 Alfvén, H., and C. G. Fälthammar (1963), *Cosmical electrodynamics: Fundamental Principles (Volume 1)*, International series of monographs on physics, Clarendon Press.

828  
829 Asikainen, T., K. Mursula, and V. Maliniemi (2012), Correction of detector noise and re-  
830 calibration of NOAA/MEPED energetic proton fluxes, *Journal of Geophysical Research: Space Physics*, *117*(A9), A09,204, doi:10.1029/2012JA017593.

831  
832 Auster, H. U., K. H. Glassmeier, W. Magnes, O. Aydogar, W. Baumjohann, D. Constanti-  
833 nescu, D. Fischer, K. H. Fornacon, E. Georgescu, P. Harvey, O. Hillenmaier, R. Kroth,  
834 M. Ludlam, Y. Narita, R. Nakamura, K. Okrafka, F. Plaschke, I. Richter, H. Schwarzl,  
835 B. Stoll, A. Valavanoglou, and M. Wiedemann (2008), The THEMIS Fluxgate Magne-  
836 tometer, *Space Science Reviews*, *141*(1-4), 235–264, doi:10.1007/s11214-008-9365-9.

837 Baker, K. B., and S. Wing (1989), A new magnetic coordinate system for conjugate stud-  
838 ies at high latitudes, *Journal of Geophysical Research: Space Physics*, *94*(A7), 9139–  
839 9143, doi:10.1029/JA094IA07P09139.

840 Borovsky, J. E., and K. Yakymenko (2017), Substorm occurrence rates, substorm recur-  
841 rence times, and solar wind structure, *Journal of Geophysical Research: Space Physics*,  
842 *122*(3), 2973–2998, doi:10.1002/2016JA023625.

843 Boynton, R. J., M. A. Balikhin, S. A. Billings, H. L. Wei, and N. Ganushkina (2011), Us-  
844 ing the NARMAX OLS-ERR algorithm to obtain the most influential coupling func-  
845 tions that affect the evolution of the magnetosphere, *Journal of Geophysical Research: Space Physics*, *116*(A5), A05,218, doi:10.1029/2010JA015505.

846  
847 Bräysy, T., K. Mursula, and G. Marklund (1998), Ion cyclotron waves during a great mag-  
848 netic storm observed by Freja double-probe electric field instrument, *Journal of Geo-*

- 849 *physical Research: Space Physics*, 103(A3), 4145–4155, doi:10.1029/97JA02820.
- 850 Büchner, J., and L. M. Zelenyi (1987), Chaotization of the electron motion as the cause of  
851 an internal magnetotail instability and substorm onset, *Journal of Geophysical Research*,  
852 92(A12), 13,456–13,466, doi:10.1029/JA092iA12p13456.
- 853 Chen, J., and P. J. Palmadesso (1986), Chaos and Nonlinear Dynamics of Single-Particle  
854 Orbits in a Magnetotaillike Magnetic Field, *Journal of Geophysical Research*, 91(1),  
855 1499–1508, doi:10.1029/JA091iA02p01499.
- 856 Coroniti, F. V. (1980), On the tearing mode in quasi-neutral sheets, *Journal of Geophysical*  
857 *Research*, 85(A12), 6719, doi:10.1029/JA085iA12p06719.
- 858 Delcourt, D. C., J.-A. Sauvaud, R. F. Martin, and T. E. Moore (1996), On the nonadia-  
859 batic precipitation of ions from the near-Earth plasma sheet, *Journal of Geophysical*  
860 *Research: Space Physics*, 101(A8), 17,409–17,418, doi:10.1029/96JA01006.
- 861 Delcourt, D. C., T. E. Moore, B. L. Giles, and M.-C. Fok (2000), Quantitative modeling  
862 of modulated ion injections observed by Polar-Thermal Ion Dynamics Experiment in the  
863 cusp region, *Journal of Geophysical Research: Space Physics*, 105(A11), 25,191–25,203,  
864 doi:10.1029/2000JA000034.
- 865 Delcourt, D. C., H. V. Malova, and L. M. Zelenyi (2006), Quasi-adiabaticity in  
866 bifurcated current sheets, *Geophysical Research Letters*, 33(6), L06,106, doi:  
867 10.1029/2005GL025463.
- 868 DeZeeuw, D. L., T. I. Gombosi, C. P. T. Groth, K. G. Powell, and Q. F. Stout (2000), An  
869 Adaptive MHD Method for Global Space Weather Simulations, *IEEE Trans. Plasma*  
870 *Sci.*, 28(6), 1956–1965, doi:10.1109/27.902224.
- 871 Dubyagin, S., N. Ganushkina, S. Apatenkov, M. Kubyshkina, H. Singer, and M. Liemohn  
872 (2013), Geometry of duskside equatorial current during magnetic storm main phase  
873 as deduced from magnetospheric and low-altitude observations, *Annales Geophysicae*,  
874 31(3), 395–408, doi:10.5194/angeo-31-395-2013.
- 875 Dubyagin, S., N. Y. Ganushkina, and V. Sergeev (2018), Formation of 30 KeV Proton  
876 Isotropic Boundaries During Geomagnetic Storms, *Journal of Geophysical Research:*  
877 *Space Physics*, doi:10.1002/2017JA024587.
- 878 Evans, D. S., and M. S. Greer (2000), Polar Orbiting Environmental Satellite Space En-  
879 vironment Monitor -2 Instrument Descriptions and Archive Data Documentation, *Tech.*  
880 *rep.*, Boulder, Colorado.

- 881 Ganushkina, N. Y., T. I. Pulkkinen, M. V. Kubyshkina, H. J. Singer, and C. T. Russell  
882 (2002), Modeling the ring current magnetic field during storms, *Journal of Geophysi-*  
883 *cal Research*, 107(A7), 1092, doi:10.1029/2001JA900101.
- 884 Ganushkina, N. Y., T. I. Pulkkinen, M. V. Kubyshkina, V. A. Sergeev, E. A. Lvova, T. A.  
885 Yahnina, A. G. Yahnin, T. Fritz, N. Yu Ganushkina, T. I. Pulkkinen, M. V. Kubyshkina,  
886 V. A. Sergeev, E. A. Lvova, T. A. Yahnina, A. G. Yahnin, and T. Fritz (2005), Proton  
887 isotropy boundaries as measured on mid- and low-altitude satellites, *Annales Geophysi-*  
888 *cae*, 23(5), 1839–1847, doi:10.5194/angeo-23-1839-2005.
- 889 Ganushkina, N. Y., M. W. Liemohn, M. V. Kubyshkina, R. Ilie, and H. J. Singer (2010),  
890 Distortions of the magnetic field by storm-time current systems in Earth’s magneto-  
891 sphere, *Annales Geophysicae*, 28(1), 123–140, doi:10.5194/angeo-28-123-2010.
- 892 Gilson, M. L., J. Raeder, E. Donovan, Y. S. Ge, and L. Kepko (2012), Global simulation  
893 of proton precipitation due to field line curvature during substorms, *Journal of Geophys-*  
894 *ical Research: Space Physics*, 117(A5), A05,216, doi:10.1029/2012JA017562.
- 895 Gvozdevsky, B. B., V. A. Sergeev, and K. Mursula (1997), Long lasting energetic proton  
896 precipitation in the inner magnetosphere after substorms, *Journal of Geophysical Re-*  
897 *search: Space Physics*, 102(A11), 24,333–24,338, doi:10.1029/97JA02062.
- 898 Haiducek, J. D., D. T. Welling, N. Y. Ganushkina, S. K. Morley, and D. S. Ozturk  
899 (2017), SWMF Global Magnetosphere Simulations of January 2005: Geomagnetic  
900 Indices and Cross-Polar Cap Potential, *Space Weather*, 15(12), 1567–1587, doi:  
901 10.1002/2017SW001695.
- 902 Halford, A. J., B. J. Fraser, and S. K. Morley (2010), EMIC wave activity during geomag-  
903 netic storm and nonstorm periods: CRRES results, *Journal of Geophysical Research:*  
904 *Space Physics*, 115(A12), A12,248, doi:10.1029/2010JA015716.
- 905 Huang, C.-L., H. E. Spence, H. J. Singer, and N. A. Tsyganenko (2008), A quantitative  
906 assessment of empirical magnetic field models at geosynchronous orbit during mag-  
907 netic storms, *Journal of Geophysical Research: Space Physics*, 113(A4), n/a–n/a, doi:  
908 10.1029/2007JA012623.
- 909 Hultqvist, B. (1979), The hot ion component of the magnetospheric plasma and some re-  
910 lations to the electron component-observations and physical implications, *Space Science*  
911 *Reviews*, 23(4), 581–675, doi:10.1007/BF00212357.
- 912 Ilie, R., N. Ganushkina, G. Toth, S. Dubyagin, and M. W. Liemohn (2015), Testing the  
913 magnetotail configuration based on observations of low-altitude isotropic boundaries



- 914 during quiet times, *Journal of Geophysical Research: Space Physics*, *120*(12), 10,557–  
915 10,573, doi:10.1002/2015JA021858.
- 916 Imhof, W. L., J. B. Reagan, and E. E. Gaines (1977), Fine-scale spatial structure in the  
917 pitch angle distributions of energetic particles near the midnight trapping boundary,  
918 *Journal of Geophysical Research*, *82*(32), 5215–5221, doi:10.1029/JA082i032p05215.
- 919 Kennel, C. F., and H. E. Petschek (1966), Limit on stably trapped particle fluxes, *Journal*  
920 *of Geophysical Research*, *71*(1), 1–28, doi:10.1029/JZ071i001p00001.
- 921 Liang, J., E. Donovan, B. Ni, C. Yue, F. Jiang, and V. Angelopoulos (2014), On an  
922 energy-latitude dispersion pattern of ion precipitation potentially associated with mag-  
923 netospheric EMIC waves, *Journal of Geophysical Research: Space Physics*, *119*(10),  
924 8137–8160, doi:10.1002/2014JA020226.
- 925 Lyons, L. R., and T. W. Speiser (1982), Evidence for current sheet acceleration  
926 in the geomagnetic tail, *Journal of Geophysical Research*, *87*(A4), 2276, doi:  
927 10.1029/JA087iA04p02276.
- 928 Meng, X., G. Tóth, A. Gloer, M.-C. Fok, and T. I. Gombosi (2013), Pressure anisotropy  
929 in global magnetospheric simulations: Coupling with ring current models, *Journal of*  
930 *Geophysical Research: Space Physics*, *118*(9), 5639–5658, doi:10.1002/jgra.50539.
- 931 Meurant, M., J.-C. Gérard, C. Blockx, E. Spanswick, E. F. Donovan, B. Hubert,  
932 V. Coumans, and M. Connors (2007), EL - a possible indicator to monitor the mag-  
933 netic field stretching at global scale during substorm expansive phase: Statistical  
934 study, *Journal of Geophysical Research: Space Physics*, *112*(A5), A05,222, doi:  
935 10.1029/2006JA012126.
- 936 Montgomery, D. C., E. A. Peck, and G. G. Vining (2012), *Introduction to linear regression*  
937 *analysis*, 645 pp., Wiley.
- 938 Newell, P. T., and J. W. Gjerloev (2011), Evaluation of SuperMAG auroral electrojet in-  
939 dices as indicators of substorms and auroral power, *Journal of Geophysical Research:*  
940 *Space Physics*, *116*(A12), A12,211, doi:10.1029/2011JA016779.
- 941 Newell, P. T., V. A. Sergeev, G. R. Bikkuzina, and S. Wing (1998), Characterizing the  
942 state of the magnetosphere: Testing the ion precipitation maxima latitude (b<sub>2i</sub>) and the  
943 ion isotropy boundary, *Journal of Geophysical Research: Space Physics*, *103*(A3), 4739–  
944 4745, doi:10.1029/97JA03622.
- 945 Newell, P. T., T. Sotirelis, K. Liou, C.-I. Meng, and F. J. Rich (2007), A nearly univer-  
946 sal solar wind-magnetosphere coupling function inferred from 10 magnetospheric

- 947 state variables, *Journal of Geophysical Research: Space Physics*, 112(A1), doi:  
948 10.1029/2006JA012015.
- 949 Parzen, E. (1962), On estimation of a probability density function and mode, *Ann. Math.*  
950 *Statist.*, 33(3), 1065–1076, doi:10.1214/aoms/1177704472.
- 951 Peredo, M., D. P. Stern, and N. A. Tsyganenko (1993), Are existing magnetospheric  
952 models excessively stretched?, *Journal of Geophysical Research*, 98(A9), 15,343, doi:  
953 10.1029/93JA01150.
- 954 Powell, K. G., P. L. Roe, T. J. Linde, T. I. Gombosi, and D. L. De Zeeuw (1999), A  
955 Solution-Adaptive Upwind Scheme for Ideal Magnetohydrodynamics, *Journal of Com-*  
956 *putational Physics*, 154(2), 284–309, doi:10.1006/jcph.1999.6299.
- 957 Ridley, A. J., and M. W. Liemohn (2002), A model-derived storm time asymmetric  
958 ring current driven electric field description, *Journal of Geophysical Research: Space*  
959 *Physics*, 107(A8), SMP 2–1–SMP 2–12, doi:10.1029/2001JA000051.
- 960 Ridley, A. J., T. I. Gombosi, and D. L. De Zeeuw (2004), Ionospheric control of the mag-  
961 netosphere: conductance, *Annales Geophysicae*, 22(2), 567–584, doi:10.5194/angeo-22-  
962 567-2004.
- 963 Sazykin, S. Y. (2000), Theoretical Studies of Penetration of Magnetospheric Electric  
964 Fields to the Ionosphere, Ph.D. thesis, Utah State University, Logan, Utah.
- 965 Sergeev, V., and N. Tsyganenko (1982), Energetic particle losses and trapping boundaries  
966 as deduced from calculations with a realistic magnetic field model, *Planetary and Space*  
967 *Science*, 30(10), 999–1006, doi:10.1016/0032-0633(82)90149-0.
- 968 Sergeev, V., E. Sazhina, N. Tsyganenko, J. Lundblad, and F. Søråas (1983), Pitch-angle  
969 scattering of energetic protons in the magnetotail current sheet as the dominant source  
970 of their isotropic precipitation into the nightside ionosphere, *Planetary and Space Sci-*  
971 *ence*, 31(10), 1147–1155, doi:10.1016/0032-0633(83)90103-4.
- 972 Sergeev, V. A., and B. B. Gvozdevsky (1995), MT-index – a possible new index to char-  
973 acterize the magnetic configuration of magnetotail, *Annales Geophysicae*, 13(10), 1093–  
974 1103, doi:10.1007/s00585-995-1093-9.
- 975 Sergeev, V. A., M. Malkov, and K. Mursula (1993), Testing the isotropic boundary algo-  
976 rithm method to evaluate the magnetic field configuration in the tail, *Journal of Geo-*  
977 *physical Research: Space Physics*, 98(A5), 7609–7620, doi:10.1029/92JA02587.
- 978 Sergeev, V. A., T. A. Kornilova, I. A. Kornilov, V. Angelopoulos, M. V. Kubyshkina,  
979 M. Fillingim, R. Nakamura, J. P. McFadden, and D. Larson (2010), Auroral signatures

- 980 of the plasma injection and dipolarization in the inner magnetosphere, *Journal of Geo-*  
981 *physical Research: Space Physics*, 115(A2), A02,202, doi:10.1029/2009JA014522.
- 982 Sergeev, V. A., S. A. Chernyaeva, S. V. Apatenkov, N. Y. Ganushkina, and S. V. Dubyagin  
983 (2015a), Energy-latitude dispersion patterns near the isotropy boundaries of energetic  
984 protons, *Annales Geophysicae*, 33(8), 1059–1070, doi:10.5194/angeo-33-1059-2015.
- 985 Sergeev, V. A., I. A. Chernyaev, V. Angelopoulos, and N. Y. Ganushkina (2015b), Mag-  
986 netospheric conditions near the equatorial footpoints of proton isotropy boundaries, *An-*  
987 *nales Geophysicae*, 33(12), 1485–1493, doi:10.5194/angeo-33-1485-2015.
- 988 Singer, H., L. Matheson, R. Grubb, A. Newman, and D. Bouwer (1996), Monitoring space  
989 weather with the GOES magnetometers, in *Proceedings of SPIE*, vol. 2812, edited by  
990 E. R. Washwell, pp. 299–308, International Society for Optics and Photonics, doi:  
991 10.1117/12.254077.
- 992 Søråas, F. (1972), ESRO IA/B Observations at High Latitudes of Trapped and Precipi-  
993 tating Protons with Energies above 100 keV, in *Earth's Magnetospheric Processes*, pp.  
994 120–132, Springer, Dordrecht, doi:10.1007/978-94-010-2896-7\_12.
- 995 Søråas, F., J. Lundblad, N. Maltseva, V. Troitskaya, and V. Selivanov (1980), A compari-  
996 son between simultaneous I.P.D.P. groundbased observations and observations of ener-  
997 getic protons obtained by satellites, *Planetary and Space Science*, 28(4), 387–405, doi:  
998 10.1016/0032-0633(80)90043-4.
- 999 Tapping, K. F. (2013), The 10.7 cm solar radio flux ( F 10.7 ), *Space Weather*, 11(7),  
1000 394–406, doi:10.1002/swe.20064.
- 1001 Toffoletto, F., S. Sazykin, R. Spiro, and R. Wolf (2003), Inner magnetospheric mod-  
1002 eling with the Rice Convection Model, *Space Sci. Rev.*, 107(1-2), 175–196, doi:  
1003 10.1023/A:1025532008047.
- 1004 Tóth, G., I. V. Sokolov, T. I. Gombosi, D. R. Chesney, C. R. Clauer, D. L. De Zeeuw,  
1005 K. C. Hansen, K. J. Kane, W. B. Manchester, R. C. Oehmke, K. G. Powell, A. J. Ri-  
1006 dley, I. I. Roussev, Q. F. Stout, O. Volberg, R. A. Wolf, S. Sazykin, A. Chan, B. Yu,  
1007 and J. Kóta (2005), Space Weather Modeling Framework: A new tool for the space  
1008 science community, *Journal of Geophysical Research: Space Physics*, 110(A12), doi:  
1009 10.1029/2005JA011126.
- 1010 Tsyganenko, N. A. (1982), Pitch-angle scattering of energetic particles in the current sheet  
1011 of the magnetospheric tail and stationary distribution functions, *Planetary and Space*  
1012 *Science*, 30(5), 433–437, doi:10.1016/0032-0633(82)90052-6.

- 1013 Tsyganenko, N. A. (1989), A magnetospheric magnetic field model with a warped tail cur-  
1014 rent sheet, *Planetary and Space Science*, 37(1), 5–20, doi:10.1016/0032-0633(89)90066-  
1015 4.
- 1016 Tsyganenko, N. A. (1995), Modeling the Earth’s Magnetospheric Magnetic Field Con-  
1017 fined Within a Realistic Magnetopause, *J. Geophys. Res.*, 100(A4), 5599–5612, doi:  
1018 10.1029/94JA03193.
- 1019 Tsyganenko, N. A. (2001), Empirical Magnetic Field Models for the Space Weather Pro-  
1020 gram, pp. 273–280, American Geophysical Union (AGU), doi:10.1029/GM125p0273.
- 1021 Tsyganenko, N. A. (2002), A model of the near magnetosphere with a dawn-dusk asym-  
1022 metry 1. Mathematical structure, *Journal of Geophysical Research: Space Physics*,  
1023 107(A8), SMP 12–1–SMP 12–15, doi:10.1029/2001JA000219.
- 1024 Tsyganenko, N. A., and V. A. Andreeva (2015), A forecasting model of the magnetosphere  
1025 driven by an optimal solar wind coupling function, *Journal of Geophysical Research:*  
1026 *Space Physics*, 120(10), 8401–8425, doi:10.1002/2015JA021641.
- 1027 Tsyganenko, N. A., and V. A. Andreeva (2016), An empirical RBF model of the magneto-  
1028 sphere parameterized by interplanetary and ground-based drivers, *Journal of Geophysi-*  
1029 *cal Research: Space Physics*, 121(11), 10,786–10,802, doi:10.1002/2016JA023217.
- 1030 Tsyganenko, N. A., and M. I. Sitnov (2005), Modeling the dynamics of the inner magne-  
1031 tosphere during strong geomagnetic storms, *Journal of Geophysical Research*, 110(A3),  
1032 A03,208, doi:10.1029/2004JA010798.
- 1033 Usanova, M. E., I. R. Mann, J. Bortnik, L. Shao, and V. Angelopoulos (2012), THEMIS  
1034 observations of electromagnetic ion cyclotron wave occurrence: Dependence on AE,  
1035 SYMH, and solar wind dynamic pressure, *Journal of Geophysical Research: Space*  
1036 *Physics*, 117(A10), A10,218, doi:10.1029/2012JA018049.
- 1037 Welling, D. T., and A. J. Ridley (2010), Validation of SWMF magnetic field and plasma,  
1038 *Space Weather*, 8(3), S03,002, doi:10.1029/2009SW000494.
- 1039 West, H. I., R. M. Buck, and M. G. Kivelson (1978), On the configuration of the magne-  
1040 totail near midnight during quiet and weakly disturbed periods: Magnetic field model-  
1041 ing, *Journal of Geophysical Research*, 83(A8), 3819, doi:10.1029/JA083iA08p03819.
- 1042 Wolf, R. A., M. Harel, R. W. Spiro, G.-H. Voigt, P. H. Reiff, and C. K. Chen (1982),  
1043 Computer simulation of inner magnetospheric dynamics for the magnetic storm  
1044 of July 29, 1977, *Journal of Geophysical Research*, 87(A8), 5949–5962, doi:  
1045 10.1029/JA087iA08p05949.

- 1046 Yahnin, A. G., and T. A. Yahnina (2007), Energetic proton precipitation related to ion-  
1047 cyclotron waves, *Journal of Atmospheric and Solar-Terrestrial Physics*, 69(14), 1690–  
1048 1706, doi:10.1016/j.jastp.2007.02.010.
- 1049 Young, D. T., H. Balsiger, and J. Geiss (1982), Correlations of magnetospheric ion com-  
1050 position with geomagnetic and solar activity, *Journal of Geophysical Research*, 87(A11),  
1051 9077–9096, doi:10.1029/JA087iA11p09077.
- 1052 Yue, C., C.-P. Wang, L. Lyons, J. Liang, E. F. Donovan, S. G. Zaharia, and M. Hender-  
1053 son (2014), Current sheet scattering and ion isotropic boundary under 3-D empirical  
1054 force-balanced magnetic field, *Journal of Geophysical Research: Space Physics*, 119(10),  
1055 8202–8211, doi:10.1002/2014JA020172.

Figure 1.

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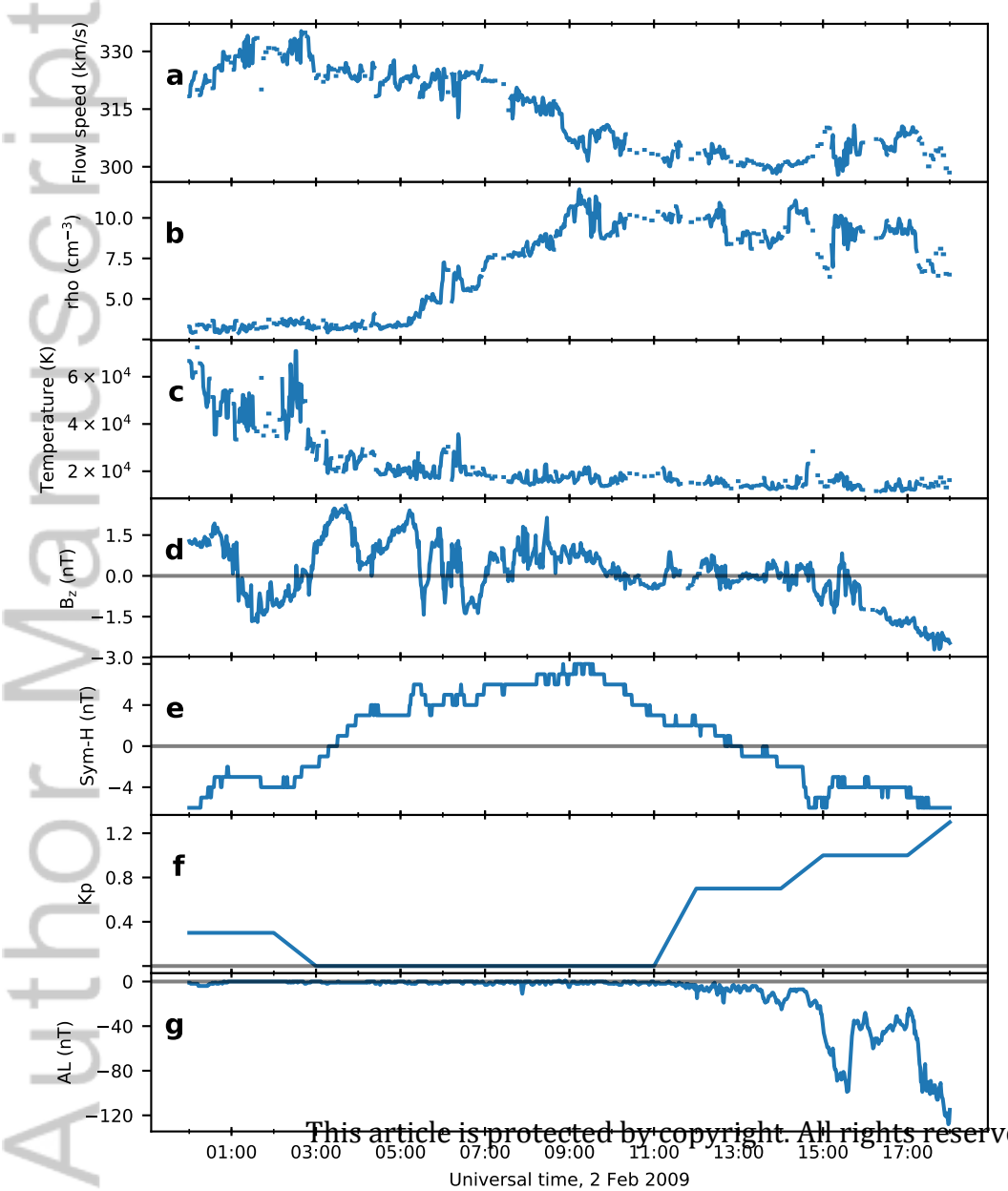


Figure 2.

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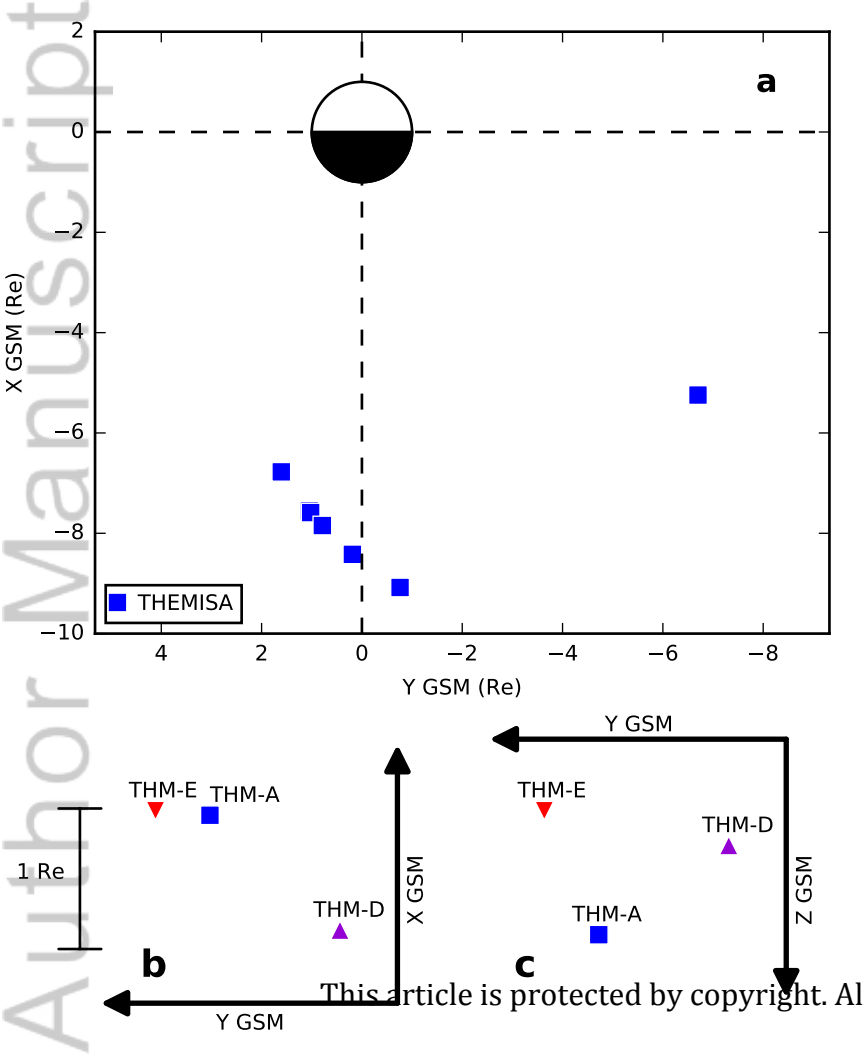


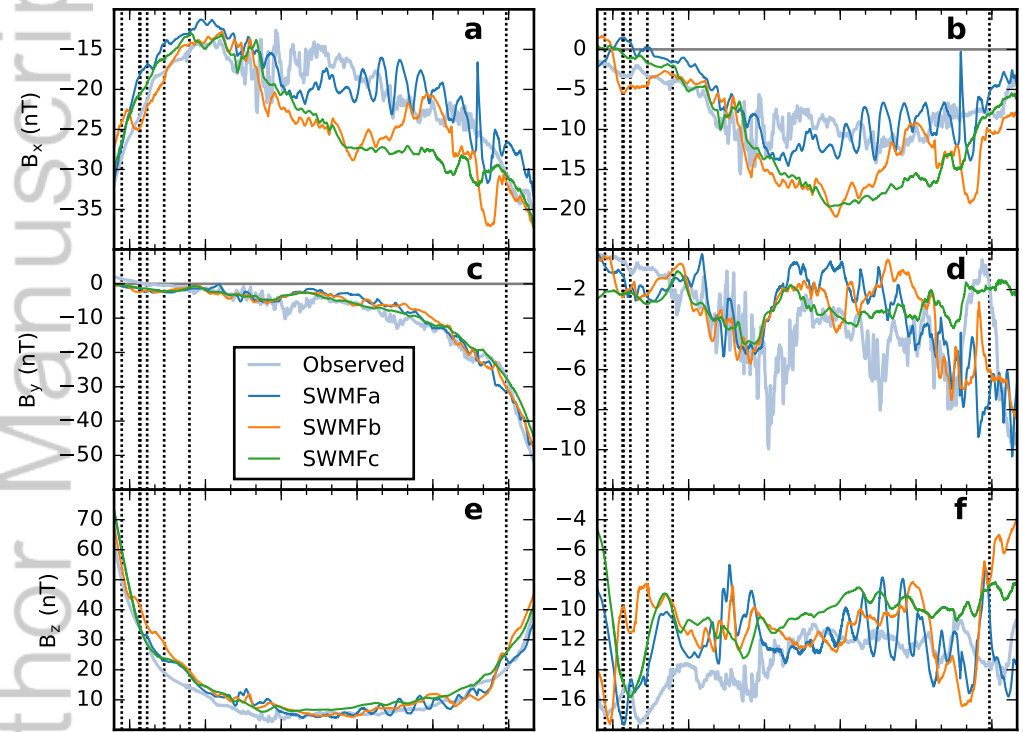
Figure 3.

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# THEMISA

## Total Field

## External Field



	UT	02:00	05:00	08:00	11:00	14:00	17:00		UT	02:00	05:00	08:00	11:00	14:00	17:00
	MLT	23.3	0.553	1.48	2.14	2.74	3.49		MLT	23.3	0.553	1.48	2.14	2.74	3.49
X GSM		-7.14	-9.4	-10.1	-9.51	-7.89	-5.14		X GSM	-7.14	-9.4	-10.1	-9.51	-7.89	-5.14
Y GSM		1.37	-1.37	-4.09	-5.98	-6.87	-6.68		Y GSM	1.37	-1.37	-4.09	-5.98	-6.87	-6.68
Z GSM		-2.68	-3.91	-3.81	-3.17	-2.77	-2.49		Z GSM	-2.68	-3.91	-3.81	-3.17	-2.77	-2.49

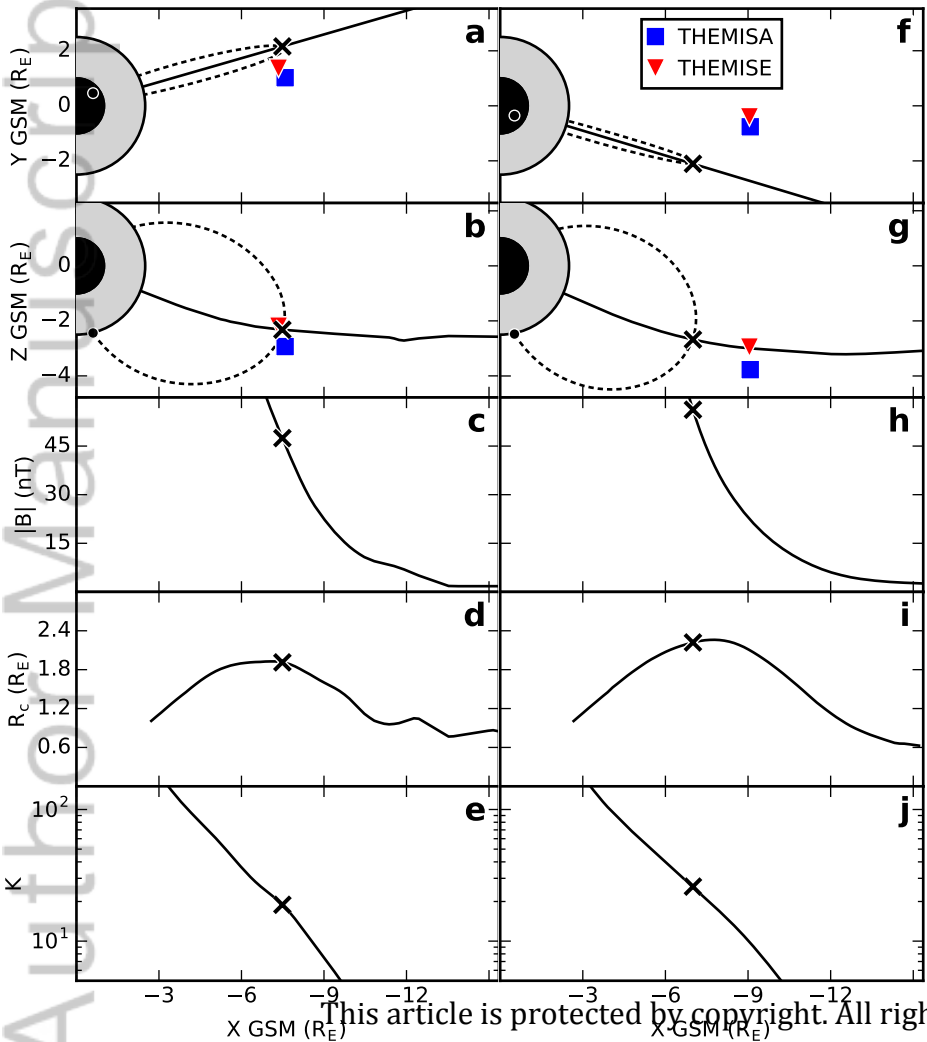
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Figure 4.

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Figure 5.

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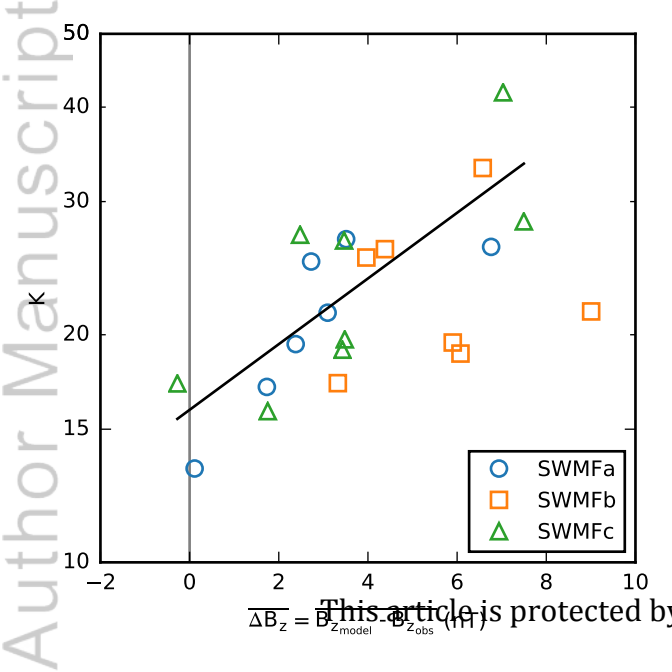


Figure 6.

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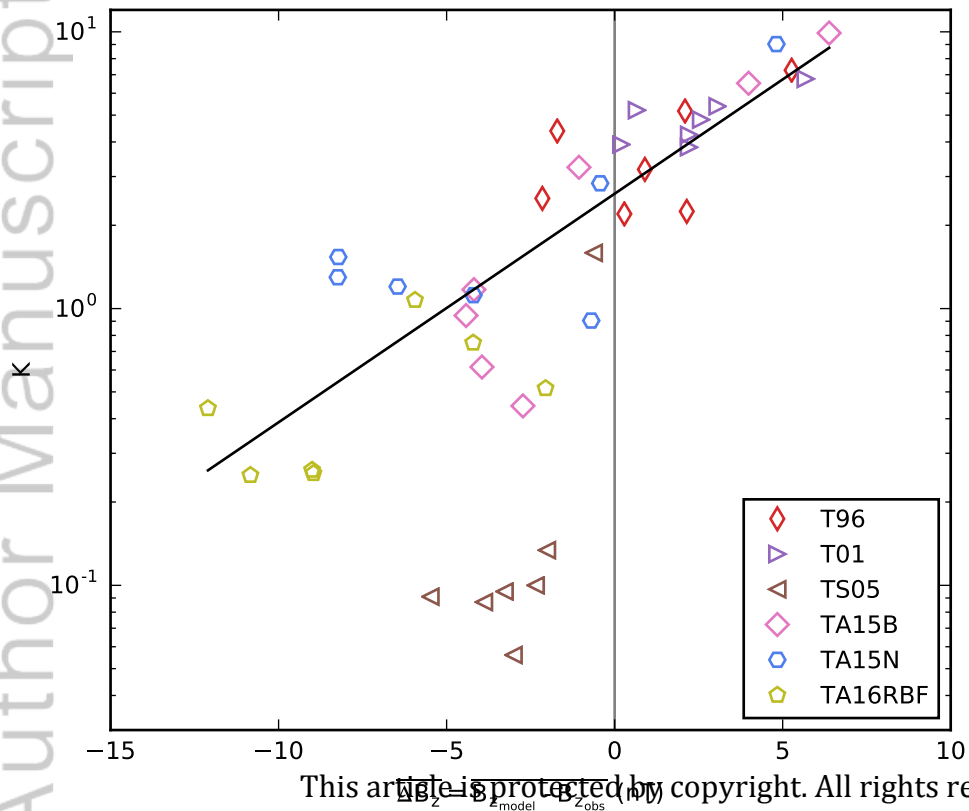


Figure 7.

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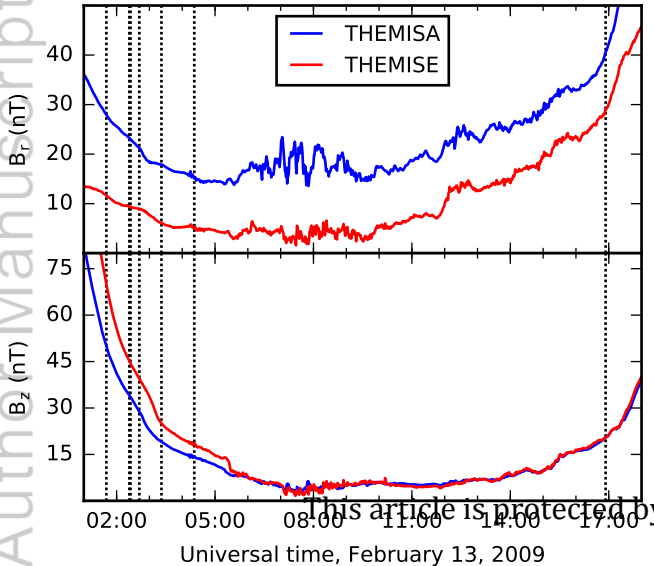


Figure 8.

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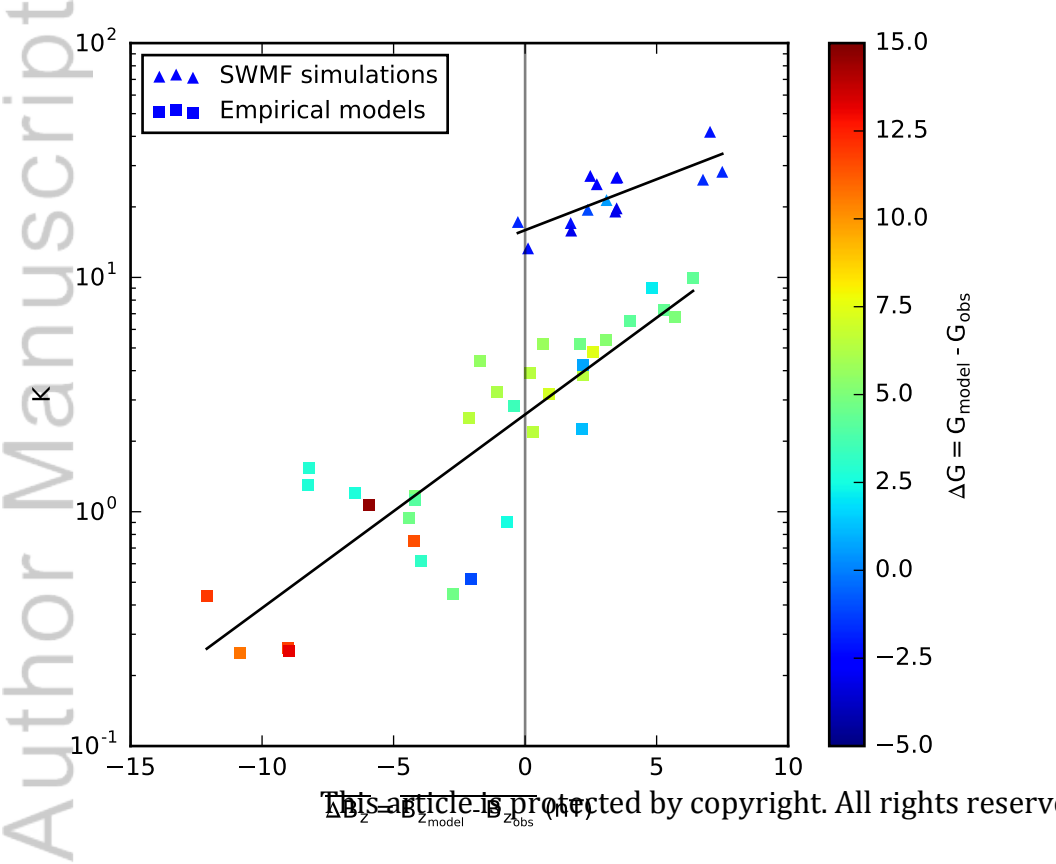
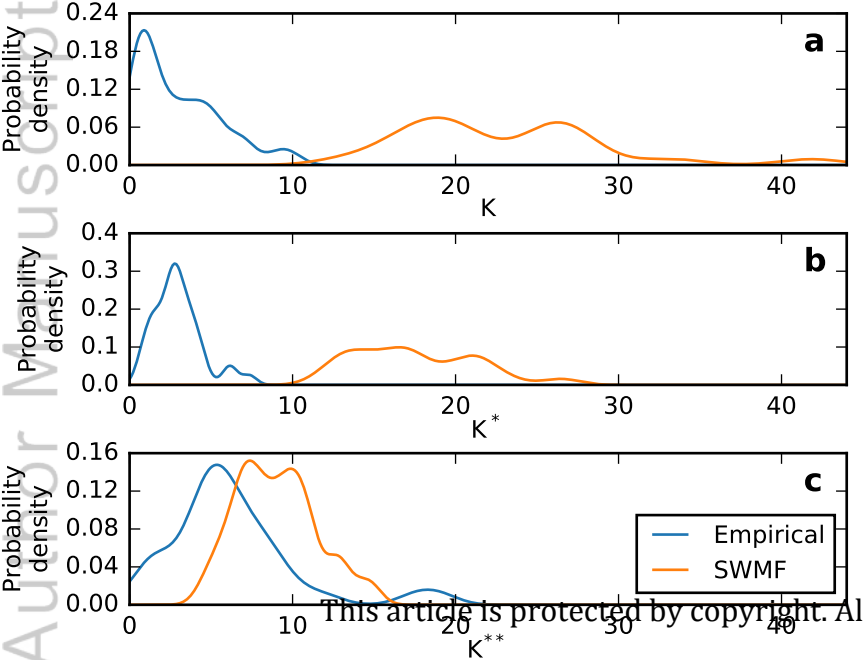
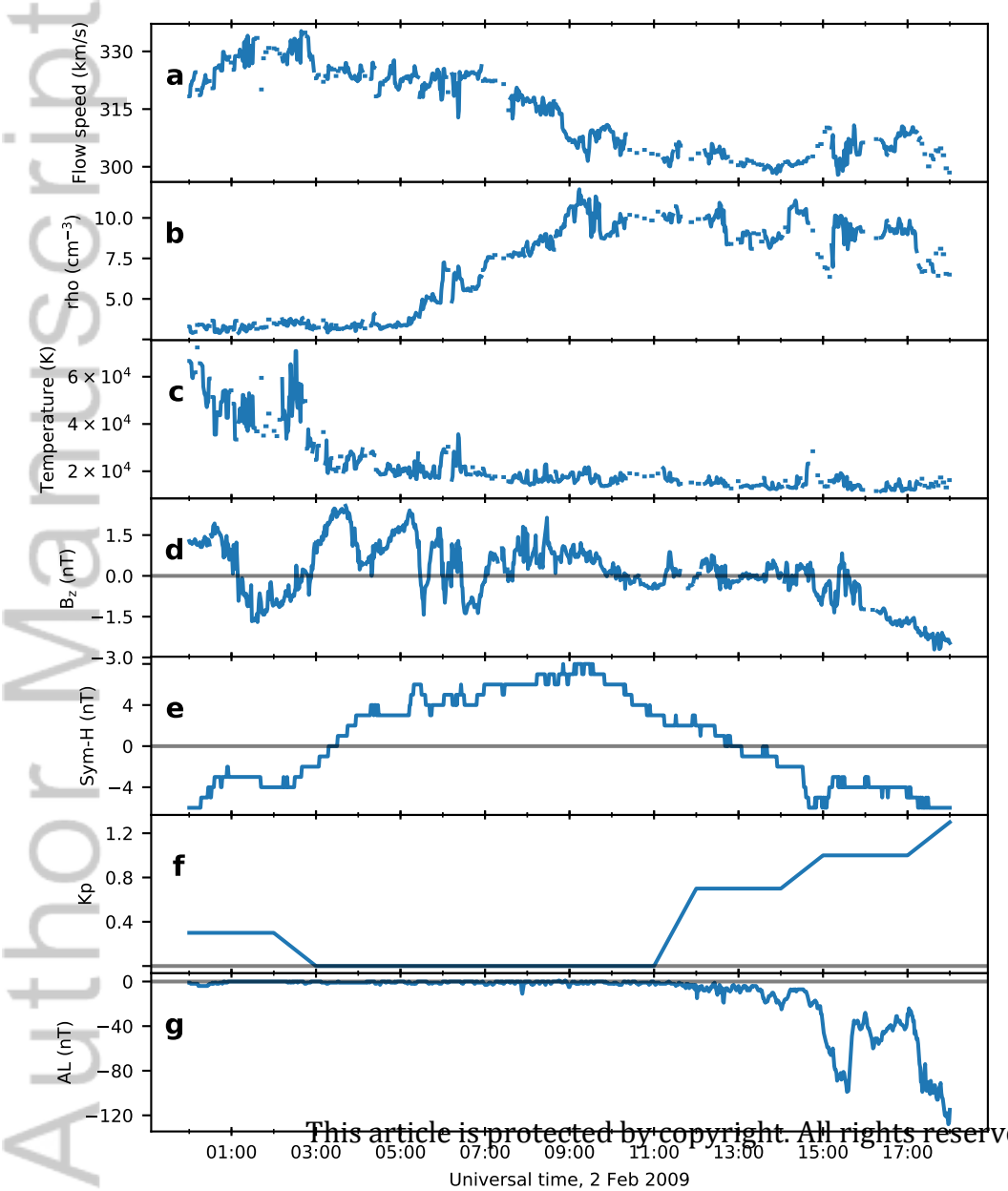


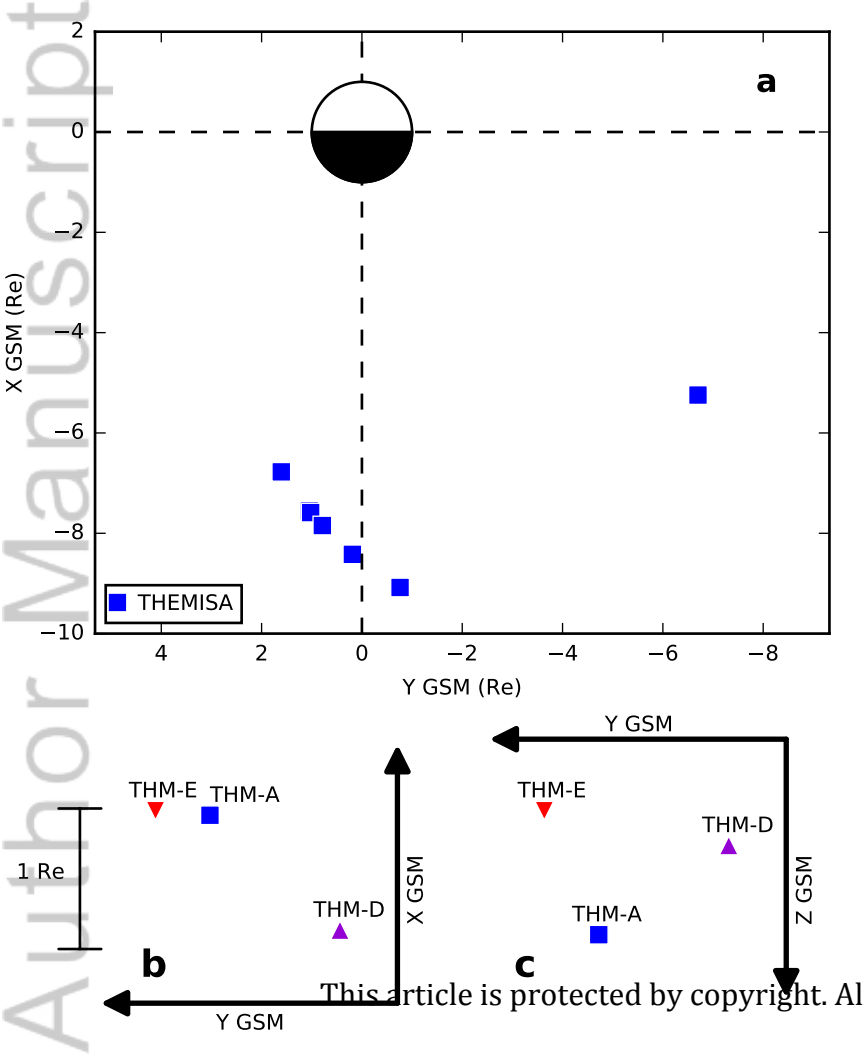
Figure 9.

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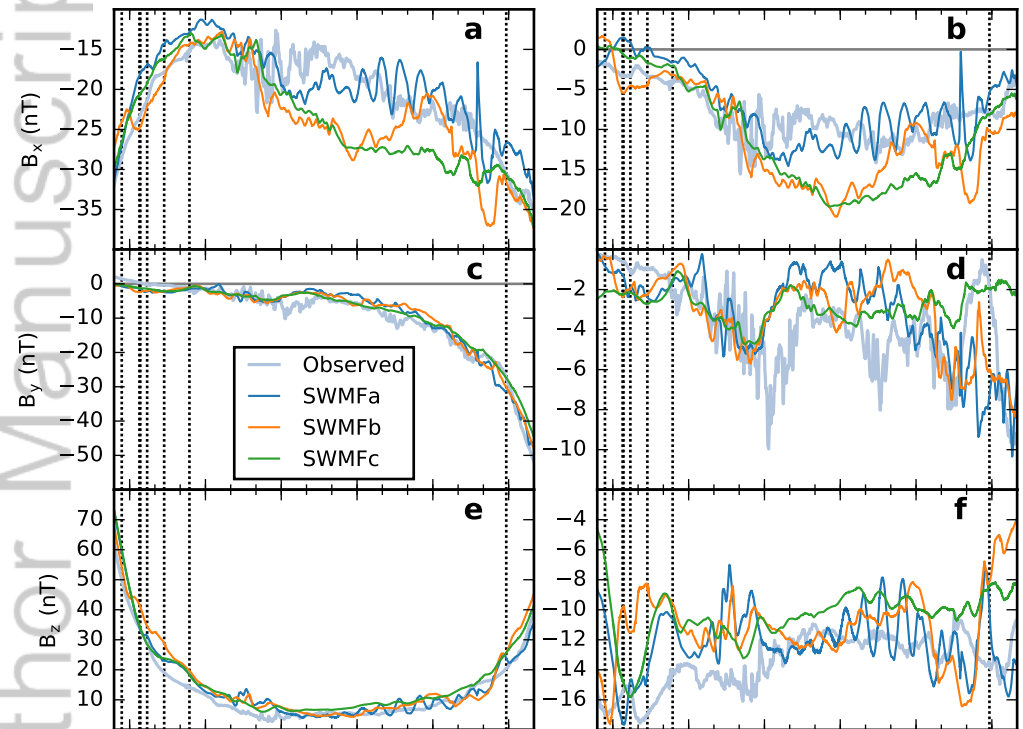




# THEMISA

## Total Field

## External Field

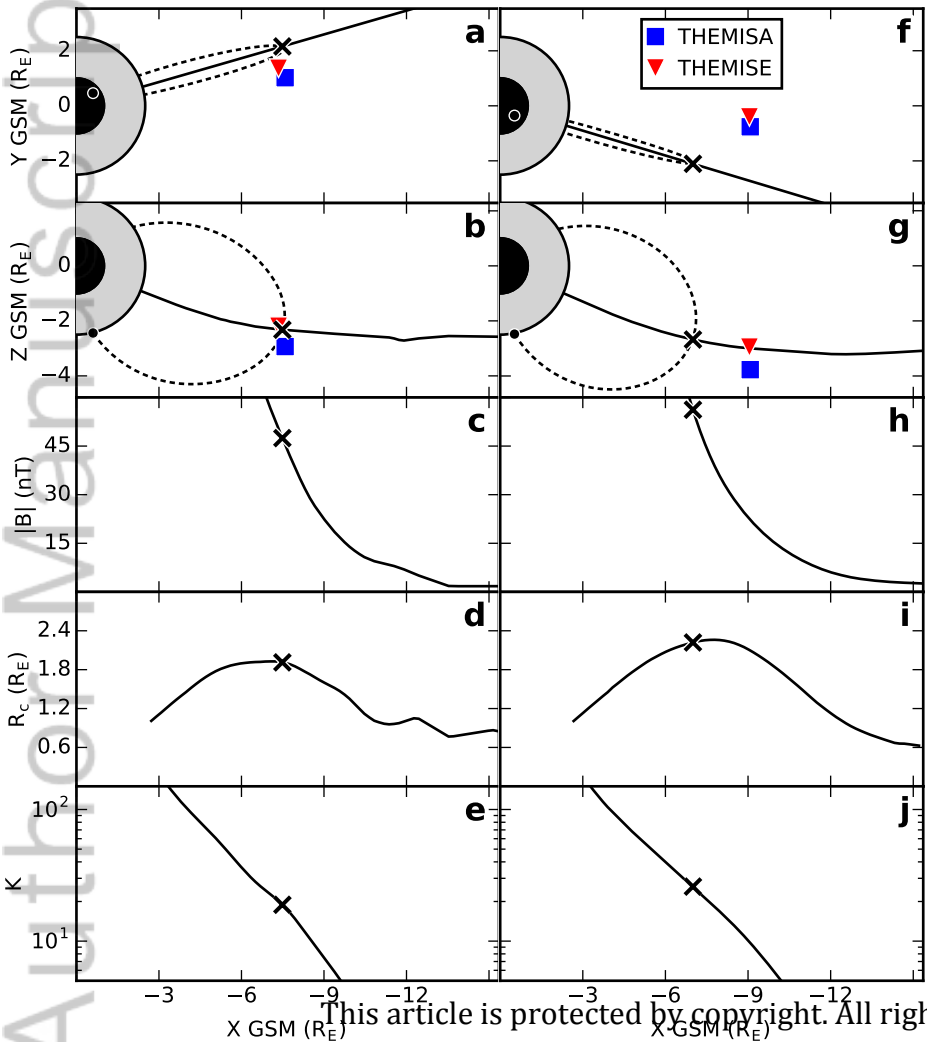


—	Observed
—	SWMFa
—	SWMFb
—	SWMFc

	UT	02:00	05:00	08:00	11:00	14:00	17:00		UT	02:00	05:00	08:00	11:00	14:00	17:00
MLT		23.3	0.553	1.48	2.14	2.74	3.49		23.3	0.553	1.48	2.14	2.74	3.49	
X GSM		-7.14	-9.4	-10.1	-9.51	-7.89	-5.14		-7.14	-9.4	-10.1	-9.51	-7.89	-5.14	
Y GSM		1.37	-1.37	-4.09	-1.98	-6.87	-6.68		1.37	-1.37	-4.09	-1.98	-6.87	-6.68	
Z GSM		-2.68	-3.91	-3.81	-3.17	-2.77	-2.49		-2.68	-3.91	-3.81	-3.17	-2.77	-2.49	

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