

1 **Comparison of ICON/MIGHTI and TIMED/TIDI**
2 **Neutral Wind Measurements in the Lower**
3 **Thermosphere**

4 **Manbharat S. Dhadly¹, Christoph R. Englert¹, Douglas P. Drob¹, John T.**
5 **Emmert¹, Rick Niciejewski², and Kate A. Zawdie¹**

6 ¹Space Science Division, U.S. Naval Research Laboratory, Washington DC, USA

7 ²Climate and Space Sciences and Engineering, University of Michigan, Ann Arbor, Michigan, USA

8 **Key Points:**

- 9 • Cross-compare ICON/MIGHTI and TIMED/TIDI MLT region neutral winds.
10 • Overall, MIGHTI and TIDI neutral wind measurements are in agreement.
11 • TIDI coldside measurements in forward flight show systematic bias.

Corresponding author: Manbharat Dhadly, manbharat.dhadly@nrl.navy.mil

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Abstract

This study cross-compares ICON/MIGHTI and TIMED/TIDI MLT region neutral winds from middle Northern Hemisphere to low Southern Hemisphere latitudes. We utilized MIGHTI level-2.2 (v4) and TIDI level-3 (v11) neutral winds from January 2020 to November 2020 and found their conjunctions using a space-time window of $LST \pm 15\text{min}$, $\text{latitude} \pm 4$ degrees, and $\text{longitude} \pm 4$ degrees around each TIDI wind measurement. Due to the nature of their orbital geometry, frequent conjunctions occurred between MIGHTI and TIDI. These conjunctions are spread in longitudes and they occur at approximately fixed LSTs and latitudes, which allows to compare their observed diurnal variability. MIGHTI and TIDI wind observations agree well (except on the TIDI coldside during forward flight) and show similar large amplitude longitudinal variations that can reach more than 100 m/s. MIGHTI and TIDI zonal and meridional winds show moderate correlations of 0.60 and 0.55, respectively. The slopes of regression fits for zonal and meridional winds are 0.92 and 0.91, respectively. The root mean square differences in zonal and meridional winds are 56 m/s and 66 m/s respectively. We found that TIDI coldside measurements in forward flight show a systematic bias and this behavior is repetitive as the instrument pointing direction is changed by the periodic TIMED yaw maneuver. The nature of this systematic bias suggests that the TIDI zero-wind references (at least for the coldside telescopes) need revision. This investigation can provide guidance towards improving the TIDI data analysis. In addition, the results of this study act as a validation of MIGHTI MLT winds.

1 Introduction

The Earth's upper atmosphere is a dynamic environment and neutral winds are a critical component of it. Neutral winds control a large part of the dynamics within the coupled upper atmosphere and ionosphere system and thus play an important role in determining the state of the ionosphere-thermosphere (I-T) system at all latitudes. Due to the geometry of the Earth's geomagnetic field, neutral winds at lower latitudes can generate electric fields via the dynamo effect and push the ionospheric plasma upward and downward along the magnetic field lines (e.g., Rishbeth, 1972; Kelly, 1989; Immel et al., 2021). At high latitudes, they feed back into the ionospheric convection and can transport energy and momentum from high to low latitudes (e.g., Richmond et al., 2003; Killeen & Roble, 1986; Killeen, 1987; Dhadly & Conde, 2017; Wang et al., 2021). In the

44 mesosphere and lower thermosphere (MLT, 60-110 km), where most tides and waves dis-
45 sipate, neutral winds control the energy and momentum entering the upper atmosphere
46 via gravity waves, tides and planetary waves from the underlying regions (including the
47 troposphere, stratosphere, and lower mesosphere) (e.g., Lindzen, 1981; Smith, 1996, 2004;
48 Forbes et al., 2006; Forbes, 2007; Yigit et al., 2016). The changes in tidal dissipation and
49 gravity wave breaking/filtering by neutral winds can lead to changes in the whole atmo-
50 sphere circulation (e.g., Smith, 2004; England, 2012). Therefore, accurate global mon-
51 itoring of upper thermospheric, as well as MLT region neutral winds is crucial for ad-
52 vancing global I-T space weather research and forecasting capabilities. Despite the known
53 importance of neutral winds and their role in I-T electrodynamics, the global coverage
54 and monitoring of neutral winds has remained one of the most undersampled state pa-
55 rameters in the earth's upper atmosphere (e.g., Meriwether, 2006; Emmert et al., 2008;
56 Drob et al., 2015; Dhadly et al., 2017; Dhadly, Emmert, Drob, Conde, et al., 2018; Heelis
57 & Maute, 2020). In addition, relatively few wind monitoring missions have been conducted
58 in past decades. Given our ever-increasing reliance on space-based technology, and rec-
59 ognizing the control of neutral winds on upper atmospheric and ionospheric dynamics
60 and electrodynamics, the accurate knowledge of the neutral wind system is crucial for
61 humanity's emerging space needs.

62 Various space-based missions using different wind measuring techniques have been
63 launched in the past. Although some of them co-existed for short periods of time, they
64 never presented an opportunity to cross-compare/calibrate with each other. Dynamic
65 Explorer 2 (DE2), launched in 1981, was the first to monitor upper thermospheric neu-
66 tral winds from space using optical Doppler spectroscopy using a Fabry-Perot interfer-
67 ometer (FPI) (e.g., Hays et al., 1981; Spencer et al., 1981; Killeen & Roble, 1988). WINDII
68 (Wind Imaging Interferometer - based on Michelson Interferometer) for lower and up-
69 per thermospheric winds and HRDI (High-Resolution Doppler Imager - based on FPI)
70 for MLT and below, were launched on UARS (Upper Atmosphere Research Satellite) in
71 1991 and stayed operational until 2005 (e.g., G. G. Shepherd et al., 1993; G. Shepherd
72 et al., 2012). TIMED/TIDI (Thermosphere, Ionosphere, Mesosphere Energetics & Dy-
73 namics/TIMED Doppler Interferometer) primarily focused on monitoring MLT winds
74 was launched in 2001, but did not reach optimum performance at least until 2005 (e.g.,
75 Killeen et al., 1999; Niciejewski et al., 2006). GOCE and CHAMP measured in-situ cross-
76 track upper thermospheric winds using onboard accelerometers with life spans from 2000

77 to 2010 and 2009 to 2013, respectively (e.g., Förster et al., 2008; Doornbos et al., 2010).
78 GOCE and CHAMP wind measurements overlapped between 2009 and 2010, but the
79 differences in their orbital geometries and measurements of only one wind component
80 (cross-track) made any cross-comparison between them extremely challenging. Thus, even
81 though some overlap between these past missions existed, none of them presented good
82 opportunities for in-space cross-comparison/calibration. Their wind observations (mostly
83 upper thermospheric) have been cross-validated in the past, mostly using ground-based
84 wind measurements - climatologically as well as using conjunctions (e.g., Killeen et al.,
85 1984; Gault et al., 1996; Duboin, 1997; Dhadly et al., 2017; Dhadly, Emmert, Drob, Conde,
86 et al., 2018; Dhadly et al., 2019; A. Aruliah et al., 2019). However, large-scale neutral
87 wind cross-validation studies are extremely rare due to the lack of continuous and simul-
88 taneous vector wind measurements.

89 Out of these space-based missions, TIMED/TIDI is still monitoring MLT winds
90 after ~ 20 years in orbit. TIDI was launched in 2001 into a $\sim 74^\circ$ inclination orbit at 625
91 km altitude to investigate the dynamics and energetics of the Earth's MLT region (Killeen
92 et al., 1999; Yee et al., 1999; Niciejewski et al., 2006). TIDI has produced the most ex-
93 tensive global MLT neutral wind archive ever collected by the aeronomy community. Re-
94 cently, a new space-based sensor - MIGHTI (Michelson Interferometer for Global High-
95 resolution Thermospheric Imaging) on board ICON (Ionospheric Connection Explorer)
96 - was launched in 2019 into a 27° inclination orbit at ~ 575 km altitude to observe lower
97 and upper thermospheric neutral wind simultaneously (Englert et al., 2015, 2017). Be-
98 cause TIDI is in high and MIGHTI is in a low inclination orbit, frequent conjunctions
99 spread over wide geographic regions are occurring between them (see details in Section
100 3). In addition, both of them provide fully resolved MLT region horizontal wind vector
101 observations. The frequent conjunctions and fully resolved horizontal winds provide us
102 an unprecedented opportunity to perform a large-scale cross-calibration study of MLT
103 neutral winds from two independent space-based optical sensors. Such frequent conjunc-
104 tions are hard to achieve between a space-based and ground-based stations, unless a spe-
105 cial orbit is selected that would limit its geospatial coverage. Because of the large geospa-
106 tial and temporal spread of conjunctions between MIGHTI and TIDI, this study serves
107 as a large-scale cross-validation of ICON/MIGHTI and TIMED/TIDI. TIDI local solar
108 time (LST) coverage does not vary much each day and conjunctions are spread in lon-
109 gitude (see Section 3). Thus, this study, for the first time, allows us to compare longi-

110 tudinal variability at a fixed latitude and LST from two independent platforms. Migrat-
111 ing tides are responsible for creating strong day-night variations in the upper atmosphere,
112 whereas nonmigrating tides create longitudinal variations at a fixed local time. Thus,
113 the longitudinal organization of the MLT at fixed LST is preferentially controlled by non-
114 migrating tides (e.g., Oberheide et al., 2006). Forcing due to zonally asymmetric latent
115 heat release by deep tropical convective cells is considered the major source of non-migrating
116 tides (e.g., Hagan et al., 2007, 2009). The longitudinal signatures of nonmigrating tides
117 can only be detected with distributed networks of ground-based observations or space-
118 craft (e.g., England, 2012). MIGHTI and TIDI are well suited for these kinds of obser-
119 vations. Because of the MIGHTI and TIDI conjunction geometry, this study provides
120 the first validation of the longitudinal fluctuations that TIDI has been seeing for the past
121 two decades.

122 MIGHTI upper thermospheric (~ 240 km) and MLT winds (between 94-104km) have
123 been validated recently by Makela et al. (2021) and Harding et al. (2021), respectively.
124 Makela et al. (2021) used two ground-based FPIs and Harding et al. (2021) used four
125 ground-based meteor radars in the validation process. However, the latitudinal and lon-
126 gitudinal geographic coverage of winds has remained scarce due to an uneven distribu-
127 tion of suitable ground-based instruments. Here, we investigate the large-scale cross-validation
128 of ICON/MIGHTI and TIMED/TIDI MLT horizontal neutral winds from low southern
129 hemisphere ($\sim 15^\circ\text{S}$) to middle Northern Hemisphere latitudes ($\sim 42^\circ\text{N}$).

130 This study is organized as follows. Section 2 discusses the details of instruments
131 and their neutral wind measurements. Section 3 describes the data used, the details of
132 the criteria used for finding conjunctions between MIGHTI and TIDI, and the method-
133 ology for inter-comparison. Section 4 presents the results of the inter-comparisons. Fi-
134 nally, the major findings are summarized in Section 5.

135 2 Instrumentation

136 2.1 ICON/MIGHTI

137 MIGHTI uses the Doppler Asymmetric Spatial Heterodyne (DASH) technique (Englert
138 et al., 2007) for remote sensing of the lower and upper thermospheric neutral winds be-
139 tween the altitudes of 90 km and 300 km. It is similar to the Michelson Interferometers
140 used on WINDII, but MIGHTI has no moving interferometer parts and interferogram

141 samples are recorded simultaneously for different emission lines (Englert et al., 2015, 2017).
142 MIGHTI records interferometric limb images of the naturally occurring thermospheric
143 green line (557.7 nm) and red line (630.0 nm) atomic oxygen emissions. Using these im-
144 ages, altitude profiles of LOS winds are retrieved from the observed Doppler shift. Dur-
145 ing daytime, MIGHTI measures wind profiles continuously between 90 and 300 km, us-
146 ing green line data between 90 km and 190 km and red line data between 180 km and
147 300 km. However, at nighttime, green line emissions stay confined to the lower thermo-
148 sphere and red line emissions are confined to the upper thermosphere. Thus, at night-
149 time, MIGHTI wind profiles are observed using the green line between 90 and 109 km
150 and the red line between 210 km and 300 km, leaving a gap between 109 and 210 km
151 (Englert et al., 2017; Harding et al., 2021).

152 MIGHTI employs two separate DASH interferometers (MIGHTI A and MIGHTI
153 B) aligned orthogonal to each other (refer to Figure 1). Currently they are both north-
154 viewing and pointed 45 and 135 in azimuth from the spacecraft velocity direction. In this
155 northward limb viewing geometry, they cover latitudes from $\sim 12\text{S}$ to 42N . MIGHTI day-
156 time and nighttime single exposure times are 30 sec and 60 sec, respectively, which cor-
157 responds to 250 km to 500 km along the satellite track (Englert et al., 2017; Wu et al.,
158 2020). MIGHTI B observes virtually the same atmospheric tangent point regions as MIGHTI
159 A, parallel to the satellite track and within about 8 minutes. The orthogonality between
160 MIGHTI A and B sensors allows to them to construct a full horizontal wind vector from
161 the LOS winds retrieved from each sensor, assuming neutral winds are constant over that
162 time period. The details of MIGHTI A and B sensors are discussed by Englert et al. (2017),
163 and their wind retrieval process is discussed by Harding et al. (2017). Although the sam-
164 pling rate of each MIGHTI sensor is high, two LOS winds at a single common location
165 are measured ~ 8 mins apart. Thus, the estimated horizontal vector winds are analo-
166 gous to high cadence data but susceptible to any temporal wind variations within the
167 8 minute window. The vertical resolution of MIGHTI in the MLT is ~ 3 km. MIGHTI
168 covers all LSTs for a given latitude in ~ 27 days (Immel et al., 2018).

169 For this comparison, we used the MIGHTI level 2.2 data product (v4) from Jan-
170 uary to November 2020. The accurate retrieval of winds from the observed Doppler spec-
171 trum depends on the zero-wind calibration (i.e. the determination of the interferogram
172 signature expected for zero wind speed), because the error in the zero-wind leads to a
173 bias in the retrieved winds. For MIGHTI version 4, level 2.2. wind data, the zero-wind

174 is estimated by comparing a 60-day average of MIGHTI data with a 60-day average of
175 the empirical wind climatology Horizontal Wind Model 2014 (HWM14 - Drob et al. (2015)).
176 This approach constrains the average MIGHTI winds to the average HWM14 winds, but
177 preserves any measured wind variations. A MIGHTI zero-wind determination that is in-
178 dependent of HWM14 is planned for implementation in version 5.

179 2.2 TIMED/TIDI

180 The TIMED TIDI instrument is a limb-scanning FPI. It has four vertically scan-
181 ning orthogonally oriented telescopes in a cross-shaped pattern, 45 degrees to the satel-
182 lite velocity vector and observes the limb simultaneously at four orthogonal directions
183 (refer to Figure 1). This configuration is similar to the MIGHTI, but MIGHTI has only
184 two viewing directions. TIDI telescopes make two orthogonal observations of the same
185 tangent location from two satellite orbital positions, each ~ 8 minutes apart as the satel-
186 lite moves forward. The two observations are paired and decomposed into the meridional
187 and zonal wind components. Because of four viewing directions, winds are measured at
188 two different LSTs from each spacecraft location. Four orthogonal telescopes, two on ei-
189 ther side of the TIMED orbital path, view the limb simultaneously in two narrow swaths
190 separated by ~ 3300 km (corresponding to ~ 30 degrees longitude) at low latitudes (Niciejewski
191 et al., 2006).

192 The TIDI telescopes facing the sun-ward side are referred to as “warmside” and the
193 ones facing away from the sun are referred to as “coldside” telescopes. The TIMED satel-
194 lite performs one 180° yaw maneuver roughly every 60 days to ensure that the warm and
195 cold sides of the satellite are always pointing toward and away from the sun, respectively.
196 That is, the warmside telescopes always remain on the sunward side and coldside tele-
197 scopes remain on the shadow side. TIDI measures MLT region winds between 70 and
198 120 km at daytime and 80 and 105 km at nighttime using O_2 (0,0) atmospheric band
199 lines. The precession rate of the TIDI satellite is about 12 minutes per day, so that the
200 full LST coverage is achieved every 60 days. The TIDI orbit is such that it returns to
201 almost the same LST and latitude coordinates it viewed exactly 1 year earlier (Niciejewski
202 et al., 2006).

203 Right after the launch in 2001, the throughput of TIDI rapidly decreased due to
204 the formation of frost on the detector windows. However, two roll maneuvers were per-

205 formed in 2003 to melt the frost; they helped with the sublimation of a large fraction of
206 frost and throughput continued to increase before stabilizing around 2005. The TIDI and
207 throughput restoring details are discussed in Killeen et al. (2006) and Niciejewski et al.
208 (2006)

209 It is important to note that there are two versions of TIDI MLT winds processed
210 independently by two institutions - University of Michigan (UMich) and National Cen-
211 ter for Atmospheric Research (NCAR). This study utilizes UMich TIDI level 3 (v11) of
212 vector winds. The zero-reference wind for UMich TIDI consists of a 4x2 array correspond-
213 ing to each of the four individual telescope views, one set for forward flight and the sec-
214 ond for backward flight. The backward and forward flight represent the pointing direc-
215 tions. This was necessary to accommodate the change of view directions following each
216 yaw maneuver cycle. The zero-wind baseline for UMich TIDI data were calculated us-
217 ing three parameters: the spacecraft velocity, the Earth rotation velocity, and a refer-
218 ence velocity. The reference velocity has a slow time dependence, and is different (or set
219 to zero) for each of the telescopes. TIDI zero-wind baselines were calculated after the
220 launch of TIDI when it was recovering from the frost deposition on the optics; the winds
221 are estimated assuming that the earlier baselines for each telescope are still valid. The
222 TIDI zero-wind baseline calculation details are discussed in Killeen et al. (2006). The
223 NCAR TIDI dataset and analysis details will be discussed in a follow-on study.

224 3 Data and Conjunction Selection Criteria

225 This study includes MIGHTI data from January to November 2020. For finding
226 coincident wind measurements (or conjunctions) between MIGHTI and TIDI, we used
227 a space-time window of $LST \pm 15\text{min}$, $\text{latitude} \pm 4$ degrees, and $\text{longitude} \pm 4$ degrees around
228 each TIDI wind measurement. Any MIGHTI data within this TIDI space-time window
229 was considered a conjunction. Figure 2 illustrates the conjunction locations as well as
230 the all wind measurement locations as a function of latitude and longitude from MIGHTI
231 and TIDI on 01 January 2020. Also shown are the conjunctions as a function of longi-
232 tude, latitude, LST, and day of year 2020. The space-time conjunctions between TIDI
233 and MIGHTI are spread in longitudes around fixed latitudes and LSTs each day (because
234 TIDI local solar time coverage varies only by 12 minutes per day). Due to high spatiotem-
235 poral resolution of MIGHTI and TIDI data, sometimes multiple TIDI and MIGHTI mea-
236 surements fall within each space time conjunction window (as shown in the 01 January

237 example- top panel). Thus, multiple MIGHTI/TIDI data points exist at a single con-
 238 junction. To simplify the comparisons, we have binned and averaged the TIDI and MIGHTI
 239 winds at each conjunction into 0.5 UT wide, 10 degree wide latitude, and 5 km wide al-
 240 titude bins for each day.

241 Both MIGHTI and TIDI measurements suffer from twilight contamination due to
 242 the limb sounding geometry. MIGHTI data quality is controlled by the “wind quality fac-
 243 tor”, which can be used to select appropriate data. For TIDI, we exclude TIDI data when
 244 the solar beta angle (the angle between the orbital plane and the Earth-Sun vector) is
 245 greater than 55 (that is, when TIDI observations are near the terminator). Refer to Fig-
 246 ure S1 of Dhadly, Emmert, Drob, McCormack, and Niciejewski (2018) for the variation
 247 of solar beta angle as a function of day of year. This TIDI data elimination leaves a gap
 248 of roughly 3 weeks centered at the middle of February, April, August, and October each
 249 year.

250 For finding conjunctions, we only used samples for which the MIGHTI “wind qual-
 251 ity factor” is greater or equal to 0.5 (i.e., moderate to highest quality). We have also tested
 252 our statistical results (discussed in Section 4) using only data with wind quality factor=1,
 253 but did not see significant improvement in the overall results. We have removed the data
 254 points corresponding to the South Atlantic Anomaly (SAA) region from any statistical
 255 analyses in Section 4; however, we have kept them in individual day visualizations. Based
 256 on the above conjunction selection criteria (excluding SAA), a total of ~ 8200 conjunc-
 257 tions (binned and averaged) occurred between 90 km and 115 km from January to Novem-
 258 ber 2020 covering various seasons, local time, latitudes, and longitudes. Individually, the
 259 number of data points at conjunctions before binning and averaging for MIGHTI and
 260 TIDI are ~ 62000 and ~ 36000 respectively.

261 In Figure 2, the green vertical lines mark the days when TIDI made yaw changes
 262 to keep the coldside telescopes facing away from the sun. There are six TIDI yaw cycles
 263 between January and November 2020, each covering ~ 60 days. TIDI yaw cycle 1 cov-
 264 ered day 1-56, yaw cycle 2 covered 56-199, yaw cycle 3 covered 119-175, yaw cycle 4 cov-
 265 ered 175-239, yaw cycle 5 covered 239-301 and so on. TIDI flew in backward configura-
 266 tion during the yaw cycles 1, 3, and 5 and in forward configuration during yaw cycles
 267 2, 4, 6. Some of the gaps in the conjunction plots are because of no overlapping MIGHTI

268 and TIDI measurements. Others are due to the gaps in TIDI wind measurements when
269 its solar beta angle is greater than 55 (as discussed above).

270 For reference, we have also included Horizontal Wind Model (HWM14 - Drob et
271 al. (2015)) winds in this study. After finding the conjunctions between MIGHTI and TIDI,
272 we calculated HWM14 winds at MIGHTI and TIDI conjunction locations and processed
273 HWM14 winds in the same manner as discussed above for MIGHTI and TIDI.

274 4 Results and discussion

275 Figures 3, 4, and S1 show MIGHTI and TIDI horizontal wind components (zonal
276 and meridional) at conjunctions on 02 and 03 January as a function of universal time
277 (UT) and altitude. The variation in LST, geographic latitude, and geographic longitude
278 with UT is also shown for reference. The last panel (panel f) in each of these figures dis-
279 plays information on the TIDI telescope directions (coldside or warmside). On these days,
280 the MIGHTI and TIDI conjunctions occurred at daytime as well as nighttime (as shown
281 in the LST subpanel). The daytime conjunctions occurred both on the warmside and
282 the coldside of TIDI, while the nighttime conjunctions occurred only on the coldside (as
283 shown in the last subpanel of each figure). Daytime conjunctions (warm and coldside -
284 Figure 3 and S1) occurred around equatorial and low southern latitudes. Nighttime con-
285 junctions (coldside) occurred only at middle latitudes (Figure 4). The TIDI nighttime
286 data is shown only for altitudes below 100 km because at nighttime its altitude cover-
287 age is limited as discussed in the instrumentation section. HWM14 winds are included
288 for reference. As illustrated in these figures, MIGHTI and TIDI conjunctions are almost
289 locked in local solar time and latitude. Thus, there is no cross-contamination by LST
290 or latitude in the depicted winds; it allows a time-dependent comparison between MIGHTI
291 and TIDI. In other words, these figures show the longitudinal variation of winds at a fixed
292 LST. The longitudes between the two grey horizontal dashed lines (panel f) corresponds
293 to the SAA. SAA data were included only to avoid any unnecessary gaps/breaks in the
294 graphical representation. In the SAA region, the contaminated wind measurements from
295 MIGHTI and TIDI sometimes are different as the two instruments react differently in
296 this region to the enhanced flux of energetic particle precipitation.

297 As shown in these figures, MIGHTI and TIDI wind observations agree well with
298 each other overall, and show similar large amplitude longitudinal variations in MLT winds

299 that match in amplitude and phase. These longitudinal fluctuations can reach more than
300 100 m/s. In both data sets, with increasing altitude, the overall shape of longitudinal
301 variation is retained, but it shows phase progression; this is a signature of upward prop-
302 agating tidal/wave activity in winds. Because these longitudinal fluctuations are at fixed
303 LST, they are most likely the signatures of non-migrating tides or stationary planetary
304 waves in MLT winds (e.g., Oberheide et al., 2006; England, 2012). It is interesting to
305 note that HWM14, which is based on empirical climatology of decades of previous wind
306 measurements, shows almost non-existent longitudinal variation in winds with a mean
307 around 0 m/s. This is because HWM14 does not fit nonmigrating tides. The absence of
308 these large amplitude longitudinal fluctuations in HWM14 is thus consistent with a non-
309 migrating tidal source of the longitudinal variations seen in MIGHTI and TIDI.

310 The root mean square difference (RMSD) between MIGHTI and TIDI on dayside/nightside
311 for zonal and meridional winds are 30/34 m/s and 38/29 m/s respectively. It must be
312 noted that MIGHTI and TIDI measurements are not expected to precisely agree because
313 MIGHTI and TIDI conjunctions are generally not at exactly at the same time and lo-
314 cation, as described above and the instruments are integrating along different viewing
315 geometries and some of the scatter in their measurements can be attributed to the dif-
316 ferent gradients in the parameters along those viewing geometries. However, we deem
317 the agreement between TIDI and MIGHTI estimated winds to be more than satisfac-
318 tory. Figures S2 and S3 illustrate a similar example but from day 37 to 38.

319 The figures discussed above are cases with good agreement. However, we found that
320 MIGHTI and TIDI winds do not always agree. For example, Figures 5 and 6 illustrate
321 comparisons from days 72 to 73. On these days, MIGHTI and TIDI conjunctions occurred
322 mostly on the dayside at lower (~ 5 -10 degree) and lower-middle (20-25 degree) latitudes
323 (as shown in latitude variation panels). Figure 5 shows TIDI warmside conjunctions and
324 Figure 6 shows TIDI coldside conjunctions. On these days, the variance of the TIDI warm-
325 side winds is more than double that of MIGHTI. Nevertheless, TIDI wind variations ap-
326 pear to be in phase without any systematic bias. However, on the TIDI coldside, there
327 is a systematic difference between MIGHTI and TIDI winds. TIDI mean zonal and merid-
328 ional winds are 71 m/s more westward and 96 m/s more southward than MIGHTI. The
329 bias between TIDI MIGHTI appears consistent at all the altitudes shown. The interest-
330 ing feature of the TIDI and MIGHTI wind time series is that their wind fluctuations are
331 in phase (vary together) with each other. Figures S4 and S5 show a similar example, but

332 on a different day. The systematic difference between MIGHTI and TIDI wind, the wind
333 bias, appears to be additive, given that the direction of the bias is the same whether the
334 wind component is positive or negative. The additive nature of the bias and similarly
335 varying MIGHTI and TIDI wind fluctuations at all altitudes suggests that the bias orig-
336 inates in the zero-wind calibration of one of the instruments (MIGHTI or TIDI). MIGHTI
337 winds follows HWM14 baseline which is not surprising because MIGHTI zero-wind base-
338 line is based on HWM14 winds.

339 The nature of the bias between TIDI and MIGHTI suggests that TIDI zero-wind
340 baselines may no longer be valid. The TIDI zero-wind baselines were calculated after the
341 launch and have not been revised after. They contain a slow time-dependent component
342 which was necessary at the start of the mission. At the beginning of the TIDI mission,
343 the overall throughput of the experiment was compromised for the first year decreasing
344 rapidly during 2002. It was hypothesized that the reduction in detector performance was
345 due to a buildup of frost on the outside detector window, and not on the detector itself.
346 In early 2003, two “heating” experiments designed to warm the detector window by point-
347 ing the detector side of TIMED towards Earth in order to sublimate away the frost were
348 performed. These led to an immediate improvement in detector capability. Calibration
349 measurements suggest that changes in throughput eased between 2015 and 2020 and that
350 TIDI is in a stable configuration. Consequently, those zero-reference winds which have
351 a time-dependency may require adjustment.

352 We found that the TIDI systematic bias (as illustrated in Figures 6, S5, and S7)
353 occurs only on the coldside and occurs only every other yaw cycle. TIDI yaw cycles are
354 shown in Figure 2; the green lines mark the days (written in purple) when TIDI conducted
355 yaw maneuvers. The bias in TIDI winds occurred on the coldside during yaw cycles 2,
356 4 and 6. During these yaw cycles, TIDI was flying in “forward” configuration. The data
357 used in this study covers only six TIDI yaw cycles of 2020, but they are enough to high-
358 light the repetitive character of the TIDI bias. Because the issue is most likely connected
359 to the zero-wind baseline estimation, we expect this bias trend to be persistent. Simi-
360 lar examples from TIDI yaw cycles 4 (day 189) and 6 (day 304-305) are demonstrated
361 in Figures S4 to S7.

362 Fortunately, conjunctions between TIDI and MIGHTI also occurred when TIDI con-
363 ducted a yaw maneuver from yaw cycle 2 to yaw cycle 3. This allowed us to investigate

364 the TIDI bias problem further. Before the TIDI yaw maneuver on day 119, the winds
365 in the regions where MIGHTI and TIDI conjunctions occurred were observed by TIDI
366 coldside telescopes and TIDI was flying in forward configuration. However, after the yaw
367 maneuver on day 119, TIDI started flying in backward configuration and warmside tele-
368 scopes started to observe the regions that were earlier observed by TIDI coldside tele-
369 scopes. Figure 7 illustrates this case by showing MLT winds from day 118 to 120. Be-
370 fore day 119, as shown in Figure 7, 15 LST at around 40 N latitude, was observed by
371 TIDI coldside. After the yaw maneuver, TIDI warmside telescopes started to measure
372 the same region. A sharp decrease in TIDI zonal wind at all altitudes is evident on day
373 119 when TIDI performed the yaw maneuver and the coldside became the warmside. This
374 is a clear demonstration of the bias in TIDI winds. Interestingly, when compared to winds
375 from days 72-73 (Figure 6), 102-103 (Figure S8), 116-117 (Figure S9), it appears that
376 the bias in meridional winds subsided from day 72 to 117, but persisted in the zonal winds.

377 Figure 8 shows four examples (from two different days) of continuous TIDI wind
378 measurements (rather than conjunctions) as a function of day number (in the form of
379 $\text{day} + \text{UT}/24$) showing TIDI wind measurements before and after a yaw maneuver. The
380 figure specifically shows the yaw cycle transitions from 2 to 3 and from 4 to 5. During
381 these yaw cycles, TIDI continuously measured coldside and warmside winds for more than
382 ± 10 days around the yaw maneuvers without any gap associated with the solar beta an-
383 gle. TIDI measures two LSTs simultaneously (one on the coldside and other on the warm-
384 side), so there are two panels in this figure for each day group (days 96-133, days 220-
385 262). The fixed latitudes shown in the figure are selected so that they match with the
386 MIGHTI and TIDI conjunction comparisons shown earlier. In both the examples, LST
387 is changing slowly due to slow precession of the TIDI orbit. The only other change is the
388 day number (seasonal change). However, in both the examples, a discontinuity in the
389 TIDI zonal wind is present when the TIDI coldside telescopes point towards the new cold-
390 side (previously it was warmside) after a yaw maneuver. No such feature exists when the
391 TIDI warmside telescopes point towards the new warmside (previously it was coldside)
392 after a yaw maneuver. As discussed earlier, the bias in meridional winds subsided from
393 day 72 to 117; however, an overall shift in meridional wind baseline is visible in the case
394 when the yaw cycle changes from 4 to 5. As illustrated in latitude, longitude, and LST
395 plots below each wind data example, TIDI warmside telescopes measured almost the same
396 space-time region that was measured earlier by the coldside telescopes and vice versa.

397 Thus, in an ideal situation, the wind transition from coldside to warmside (and vice versa)
398 would be continuous.

399 It is important to note that the coldside winds are not always biased; they are bi-
400 ased only during yaw cycles 2, 4, and 6. TIDI was flying in “backward” configuration dur-
401 ing the yaw cycles 1, 3, and 5 and in “forward” configuration during yaw cycles 2, 4, 6.
402 Therefore, we find that TIDI coldside winds are biased when TIDI flies in a forward con-
403 figuration and we expect this bias trend to continue in future.

404 Further, we compared MIGHTI and TIDI conjunction winds statistically using the
405 scatter plot graphical representation shown in Figure 9. As both datasets of wind dis-
406 tributions are linearly related, theoretically they should form a straight line. The top two
407 panels include all the MIGHTI and TIDI conjunctions. The fitted linear curve using or-
408 thogonal distance regression (ODR) with wind errors involved in the fitting process is
409 also shown. The ODR involving wind errors is the appropriate approach for comparing
410 two stochastic data sets; note that the TIDI wind errors (typically 12 m/s) are typically
411 much larger than MIGHTI (typically 4 m/s). The slope, y-intercept of the fitted linear
412 curve, and Pearson correlation coefficients for wind are also shown in each panel. The
413 slope of the linear fit provides a quantitative measure of how different/similar the wind
414 magnitudes are in the two datasets, whereas the correlation coefficient informs about the
415 common variance in the two datasets. Although MIGHTI and TIDI wind variations seem
416 highly correlated visually, their zonal and meridional winds show moderate correlations
417 of 0.60 and 0.55, respectively. The RMSD in zonal and meridional winds are 56 m/s and
418 66 m/s respectively. The slopes of fit are almost the same for zonal (0.92) and merid-
419 ional winds (0.91). The slopes of less than 1 and moderate correlations are likely due to
420 the presence of the bias in TIDI winds which skews the wind distribution but could also
421 include additional systematic errors not accounted for in the TIDI and/or MIGHTI data
422 sets. Furthermore Figure 9 shows scatter plots for conjunctions on TIDI coldside and
423 warmside (middle panels), and for TIDI forward and backward configuration (bottom
424 panels) separately. The data points that deviate the most from slope 1 are from the cold-
425 side (shown in blue in second row) and when TIDI was flying in forward configuration
426 (shown in orange in the third row). The correlation for the backward case is stronger
427 for zonal wind and weaker for meridional wind. However, we found earlier that TIDI winds
428 are biased on the coldside only when TIDI moves in forward configuration; therefore, we
429 also investigated MIGHTI and TIDI conjunction winds statistically when TIDI flew in

430 forward configuration as shown in Figure S11; the coldside data shown in blue clearly
431 deviates from the slope 1. The warmside slopes are much closer to 1 and correlations are
432 stronger than for the coldside. The zonal and meridional RMSD for forward warm/coldside
433 are 50/81 m/s and 60/81 m/s respectively. Similarly, the zonal and meridional RMSD
434 for backward warm/coldside are 48/41 m/s and 67/57 m/s respectively.

435 Figures 9 and S11 show significant scatter around the lines of slope 1, and there
436 can be several reasons for this behavior. The major reasons for the scatter likely include
437 the fact that the MIGHTI and TIDI observations are not taken at the exact same time
438 and location and that there is a difference in MIGHTI and TIDI observation geometries.
439 TIDI and MIGHTI face the same challenges as ground-based and space-based compar-
440 ison studies such as (Killeen et al., 1984; Gault et al., 1996; Harding et al., 2021; Makela
441 et al., 2021). For ground-based and space-based conjunctions, their viewing geometries
442 are different and both instruments look through different media affecting the signal recorded
443 by the sensor differently. The same is true for TIDI and MIGHTI. Because of the dif-
444 ferences in TIDI and MIGHTI orbits, they are looking at approximately the same tan-
445 gent point location but through different media along their line of sights; the difference
446 in viewing geometries likely accounts for some of the scatter in the comparison plots. Some
447 of the scatter can also be attributed to the statistical noise in each instrument's mea-
448 surements. Figure 10 shows the histograms of differences in MIGHTI and TIDI winds.
449 The rows and columns are in the same format as shown in Figure 9, but the depicted
450 data is the difference between MIGHTI and TIDI wind components. Overall the zonal
451 wind differences (ZWD - shown on top) peak around 0 m/s, but the ZMD distribution
452 is skewed toward the left. As manifested in the coldside and warmside ZWD (shown in
453 first column, middle panel), and forward and backward ZWD (shown in first column, bot-
454 tom panel), the skewing in ZWD is caused by the TIDI coldside measurements when fly-
455 ing in forward configuration. However, the meridional wind differences (MWD) peak around
456 -33 m/s (top panel). The distribution appears normal with major differences in the tail
457 on the right. The manifestation in coldside MWD (shown in second column, middle panel)
458 and forward MWD (shown in second column, bottom panel) demonstrates that the dif-
459 ferences in the tail are caused by the TIDI coldside measurements when flying in the for-
460 ward configuration. The same discussion as for Figures 9 and S11 also applies to Fig-
461 ure 10.

462 As discussed earlier (in the discussion associated with Figures 6, S5-S9), the cor-
463 related wind fluctuations and systematic differences (on TIDI coldside and forward mov-
464 ing) between MIGHTI and TIDI at all altitudes suggest that the TIDI bias origin lies
465 in the zero-wind calibration. The zero-wind baseline calculation for optical instruments
466 has remained a topic of discussion since their inception (e.g., A. L. Aruliah & Rees, 1995).
467 The TIDI zero-baseline for each telescope were designed to address instrument perfor-
468 mance issues post-launch and were never revised later. Rapid sublimation of frost from
469 the outside of the TIDI detector window required that some of the 4×2 zero-wind es-
470 timates have a time-dependency. This study demonstrates that the old zero-wind base-
471 lines are no longer valid (at least for the coldside telescopes) and a proper revision is re-
472 quired. The current TIDI/MIGHTI study provides guidance towards improving the TIDI
473 analysis software package.

474 An alternate way to remove the TIDI bias is to adjust the winds using HWM14
475 as a baseline. The current version of MIGHTI winds utilize HWM14 winds as baseline.
476 Figure 11 demonstrates an example of the result of removing the TIDI bias with HWM14
477 winds as a baseline on days 72-73 (check Figure 6). In this process, we first calculated
478 a bias baseline separately for zonal and meridional winds by removing mean daily HWM14
479 winds from mean daily TIDI winds (bias baseline = daily TIDI wind average - daily HWM14
480 wind average). Then the calculated bias baseline was removed from the TIDI winds (TIDI
481 Corrected = TIDI - bias baseline). The corrected TIDI winds are shown in Figure 11.
482 This quick method shows remarkable improvement in MIGHTI and TIDI wind agree-
483 ment. Thus, fixing TIDI winds based on HWM14 can serve as a fairly straightforward
484 and quick fix until the TIDI zero-wind calibration of the TIDI observations is revised.
485 However, there may be some limitations of using this method because of the known dis-
486 crepancies between observations and HWM14 (e.g., Meriwether et al., 2016; Malki et al.,
487 2018).

488 5 Conclusions

489 This investigation is the first to cross-compare ICON/MIGHTI and TIMED/TIDI
490 MLT region neutral wind measurements from the low Southern Hemisphere (12°S) to
491 middle Northern Hemisphere latitudes (42°N). We used MIGHTI level 2.2 (v4) data and
492 University of Michigan TIDI level 3 (v11) from January 2020 to November 2020 for this
493 investigation. For finding conjunctions between MIGHTI and TIDI, we used a space-time

494 window of $LST \pm 15 \text{ min}$, latitude ± 4 degrees, and longitude ± 4 degree around each TIDI
495 wind measurement. A total of ~ 8200 conjunctions (binned and averaged) occurred be-
496 tween MIGHTI and TIDI during this period covering various seasons, local time, lati-
497 tudes, and longitudes. TIDI is in a high inclination (74°) orbit and MIGHTI is in a low
498 inclination (27°) orbit. Because of their orbital geometry, frequent conjunctions of their
499 measurement locations, that are spread over wide geographic regions, are occurring be-
500 tween them. The frequent conjunctions and their fully resolved horizontal winds provide
501 us an unprecedented opportunity to perform the first ever, large-scale cross-calibration
502 study of MLT neutral winds from two independent space-based optical sensors. In ad-
503 dition, the nature of their orbits is such that conjunctions are spread in longitudes and
504 occurring at around fixed LST and latitudes. Therefore, their conjunctions provide an
505 opportunity to compare longitudinal variability from two independent space-based op-
506 tical platforms. In addition, the results of this study act as a validation of MIGHTI lower
507 thermospheric winds.

508 Individual day comparisons show that MIGHTI and TIDI wind observation agree
509 well (except on the TIDI coldside during forward flight) and show similar large ampli-
510 tude longitudinal variations in MLT winds that can reach amplitudes of more than 100
511 m/s. Comparisons show that with increasing altitude, these longitudinal variations in
512 MLT winds at fixed LST retain their shape but show phase progression. These behav-
513 iors suggest that these longitudinal variations are likely the signatures of upward prop-
514 agating non-migrating tides in MLT winds. Such longitudinal signatures of nonmigrat-
515 ing tides can only be detected with space-based instruments or distributed networks of
516 ground-based instruments.

517 We found that TIDI coldside measurements in forward flight show a systematic bias
518 and this behavior is repetitive as the instrument pointing direction is changed by a yaw
519 maneuver. The similarity between MIGHTI and TIDI wind fluctuations and a consis-
520 tent offset between them at all the altitudes suggests that the offset origin lies in the zero-
521 wind calibration of TIDI the data. Fortunately, some MIGHTI and TIDI conjunctions
522 occurred on the days when TIMED made yaw maneuvers. An abrupt shift in winds mea-
523 sured before and after the yaw maneuvers confirms that the TIDI zero-wind baseline is
524 the root of the problem. The TIDI zero-wind baselines for each of the four individual
525 telescopes (one set for forward flight and the second for backward flight) were calculated
526 at the start of the mission. Rapid sublimation of frost on the outside of the TIDI detec-

527 tor window required zero-wind estimates to have a time-dependency. Calibration mea-
528 surements suggest that changes in throughput eased between 2015 and 2020 and that
529 TIDI is now in a stable configuration. The nature of the bias between TIDI and MIGHTI
530 suggests that the TIDI zero references, which have a time-dependence, may require ad-
531 justment. The current TIDI/MIGHTI study can provide guidance towards improving
532 the TIDI data analysis.

533 MIGHTI and TIDI zonal and meridional winds show moderate correlations of 0.60
534 and 0.55, respectively. The slopes of regression fits using ODR for zonal and meridional
535 winds are 0.92 and 0.91, respectively. The RMSD in zonal and meridional winds are 56
536 m/s and 66 m/s respectively. The zonal and meridional RMSD for forward warm/coldside
537 are 50/81 m/s and 60/81 m/s respectively. Similarly, the zonal and meridional RMSD
538 for backward warm/coldside are 48/41 m/s and 67/57 m/s respectively. It must be noted
539 that MIGHTI and TIDI measurements are not expected to always be identical because
540 the TIDI and MIGHTI observations compared here are not taken at the exact same time
541 and location and because the instruments are integrating along the different viewing ge-
542 ometries. Thus, some of the scatter in the observations can be attributed to the tem-
543 poral and spatial variability, in addition to the effect of different gradients in the param-
544 eters along the differing viewing geometries.

545 The offset in TIDI winds is associated with its zero-wind calibration, thus we ex-
546 pect this bias trend to continue in the future until the TIDI zero-wind calibration is re-
547 visited. An alternate way to fix the TIDI bias is to constrain TIDI mean winds to match
548 HWM14 as illustrated in detail with an example in the results section of this paper. So
549 far TIDI data is the most extensive archive of global MLT neutral winds collected by the
550 aeronomy community. The recently launched ICON/MIGHTI has dramatically increased
551 our equatorial to middle latitude coverage of neutral winds. MIGHTI and TIDI collec-
552 tively can provide much needed global monitoring of MLT neutral winds with better LST
553 coverage than single space-based instruments. This is required for a more comprehen-
554 sive understanding of macroscale and mesoscale dynamics and further enhances our knowl-
555 edge of the contribution of terrestrial weather to ionospheric and thermospheric variabil-
556 ity. Thus, a future TIDI wind data revision would be highly beneficial for future space
557 weather research.

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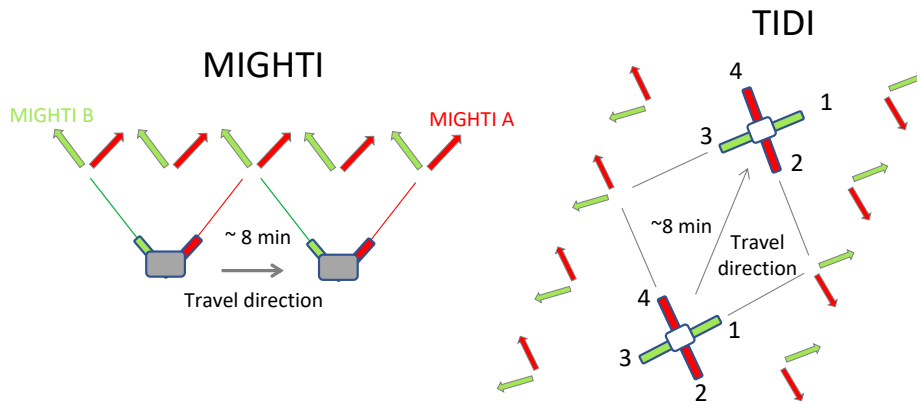


Figure 1. A graphic illustration of MIGHTI and TIDI viewing geometry.

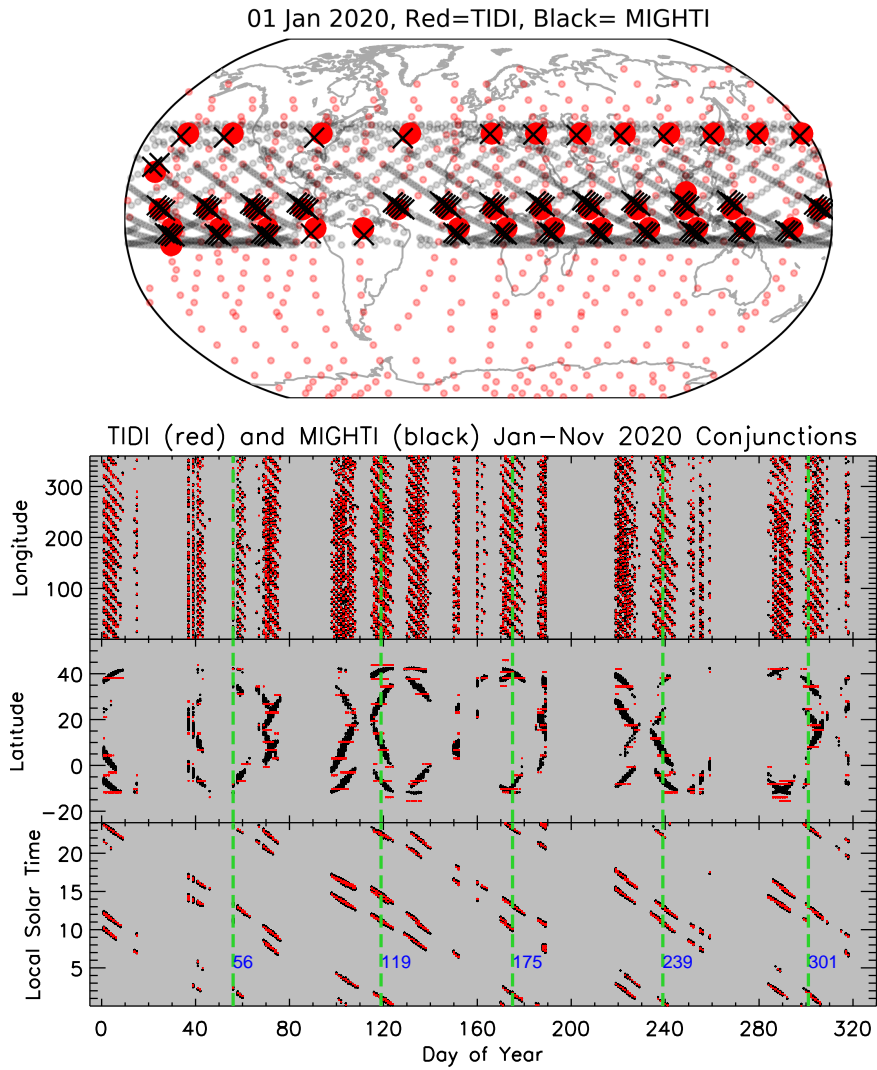


Figure 2. Top panel shows mapped MIGHTI and TIDI measurement conjunction locations on 01 January 2020. The conjunction locations are marked with large symbols (black crosses for MIGHTI and red dots for TIDI), while the small symbols show all the MIGHTI and TIDI measurements locations 01 January 2020. The lower three panels show MIGHTI and TIDI measurement conjunctions between January and November 2020, as a function of longitude, latitude, local solar time, and day of year. The green vertical lines mark the days when TIDI make yaw maneuvers and blue numbers next to the green lines indicate their day number.

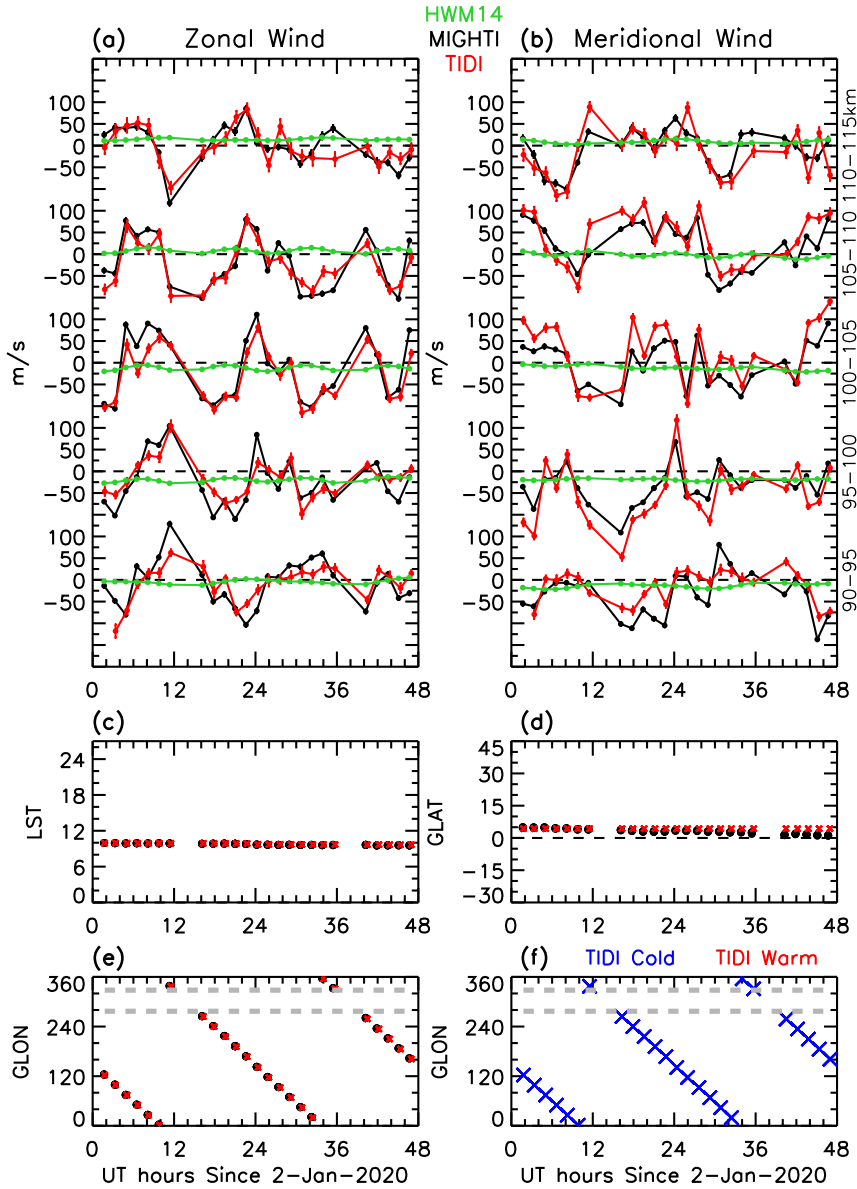


Figure 3. Comparisons between MIGHTI (black) and TIDI (red) wind components (zonal in panel (a) and meridional in panel (b)) for days 2 and 3 of 2020. The comparisons are shown as a function of UT at various 5 km wide altitude bins (shown on the right in the panel (b)). HWM14 winds (green) are also shown for reference. The variation in local solar time (LST), geographic latitude (GLAT), and geographic longitude (GLON) with universal time (UT) are shown in panels (c), (d), and (e), respectively. The panel (f) indicate which TIDI telescopes (warmside or coldside) measured the shown winds in panels (a) and (b). These measurements are from day-side equatorial latitudes and from TIDI coldside telescopes. The longitudes between two grey horizontal dashed lines (in panels (e) and (f)) corresponds to the SAA.

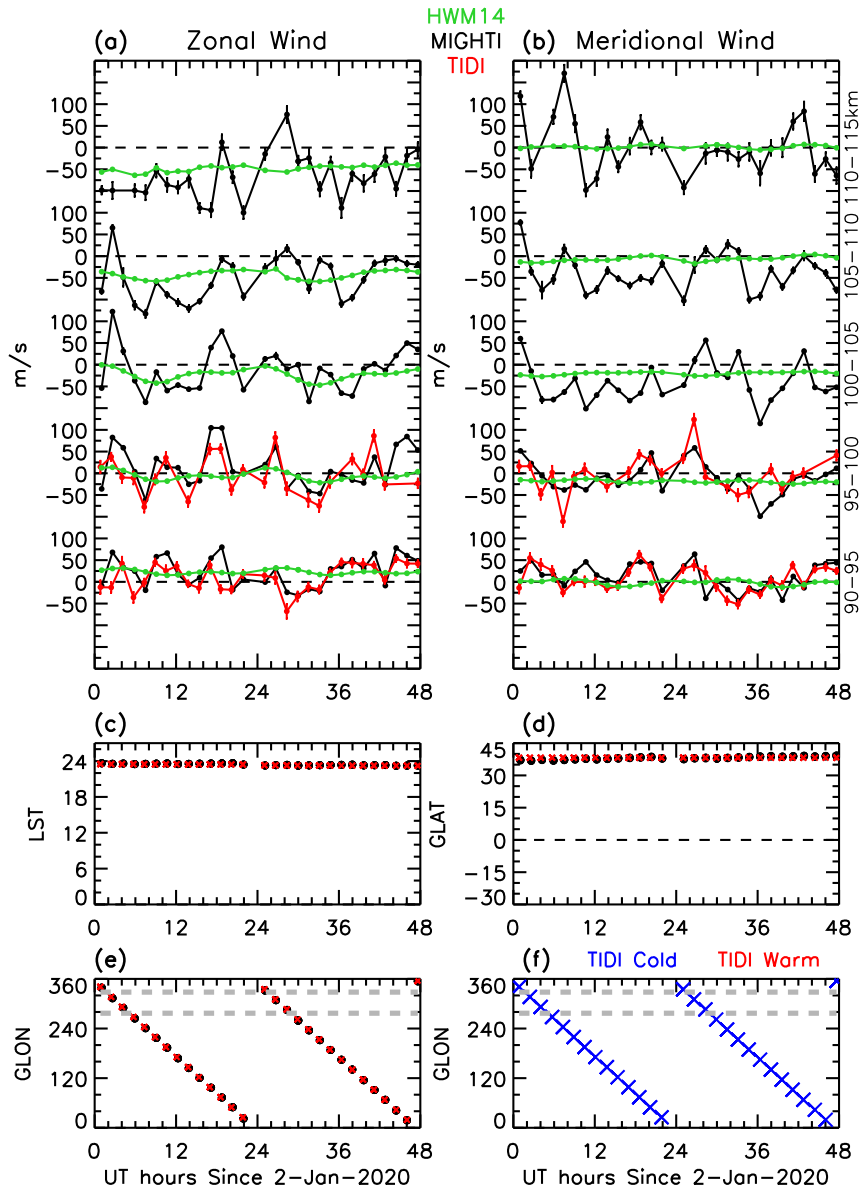


Figure 4. Same as Figure 3, but at nighttime and middle latitude.

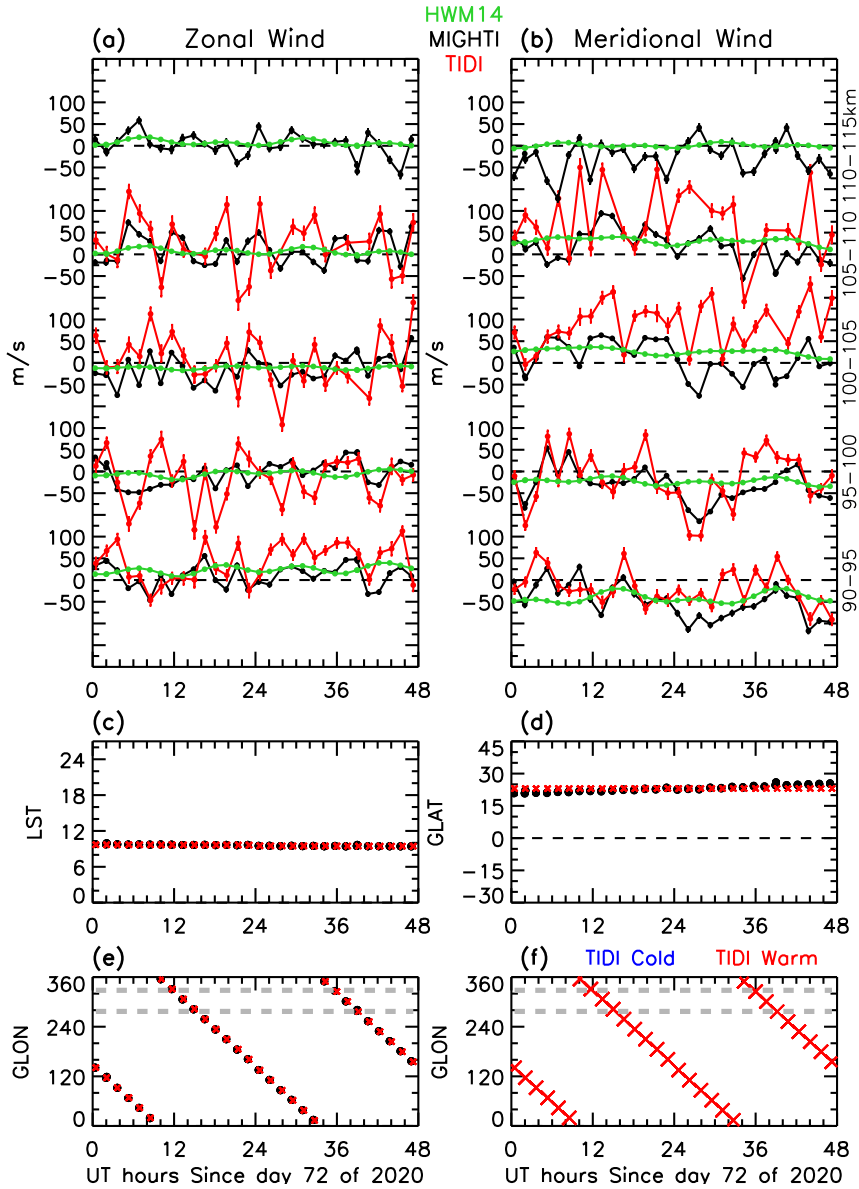


Figure 5. Same as Figure 3, but showing MIGHTI and TIDI wind comparisons at their conjunctions on days 72 and 73 days of 2020. TIDI measurements are from warmside. Days 72 and 73 are in TIDI yaw cycle 2 (days 56-119).

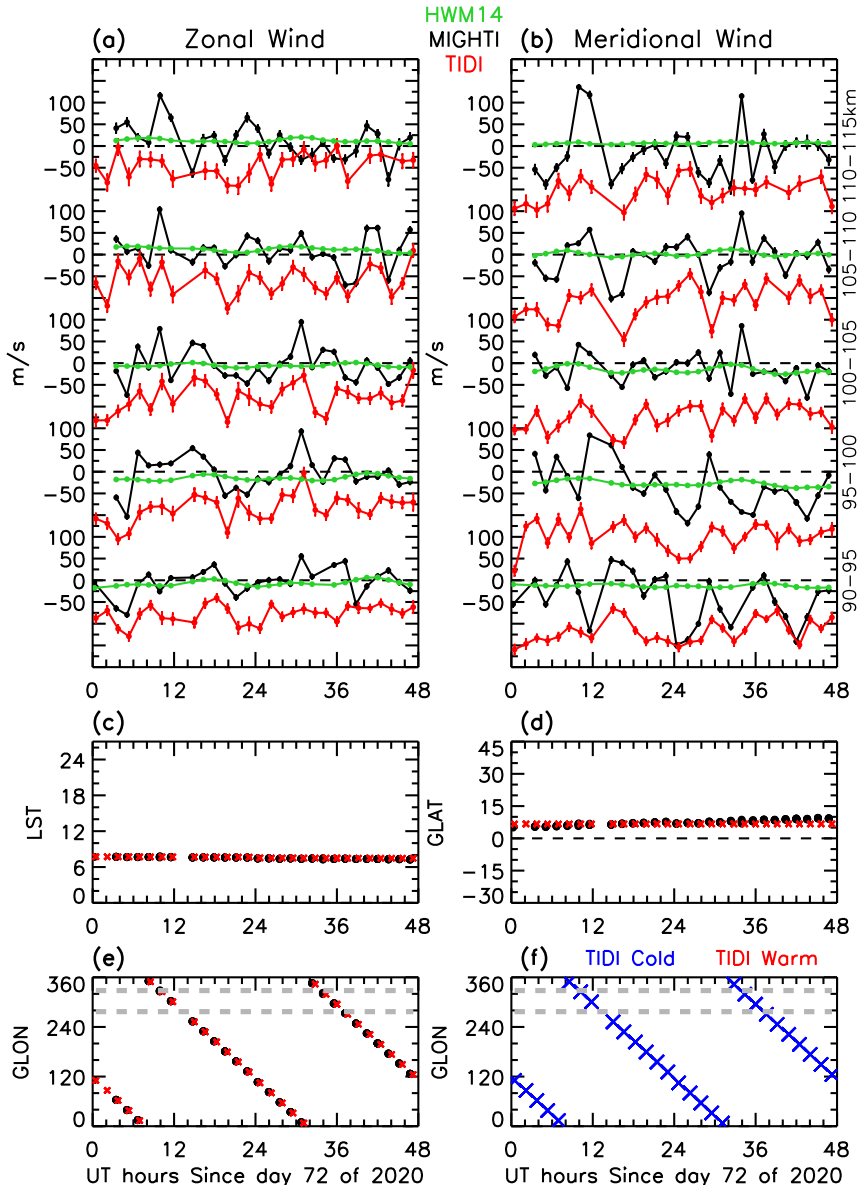


Figure 6. Same as Figure 5, but comparison are shown for lower latitude and different LST.

TIDI measurements are from coldside telescopes.

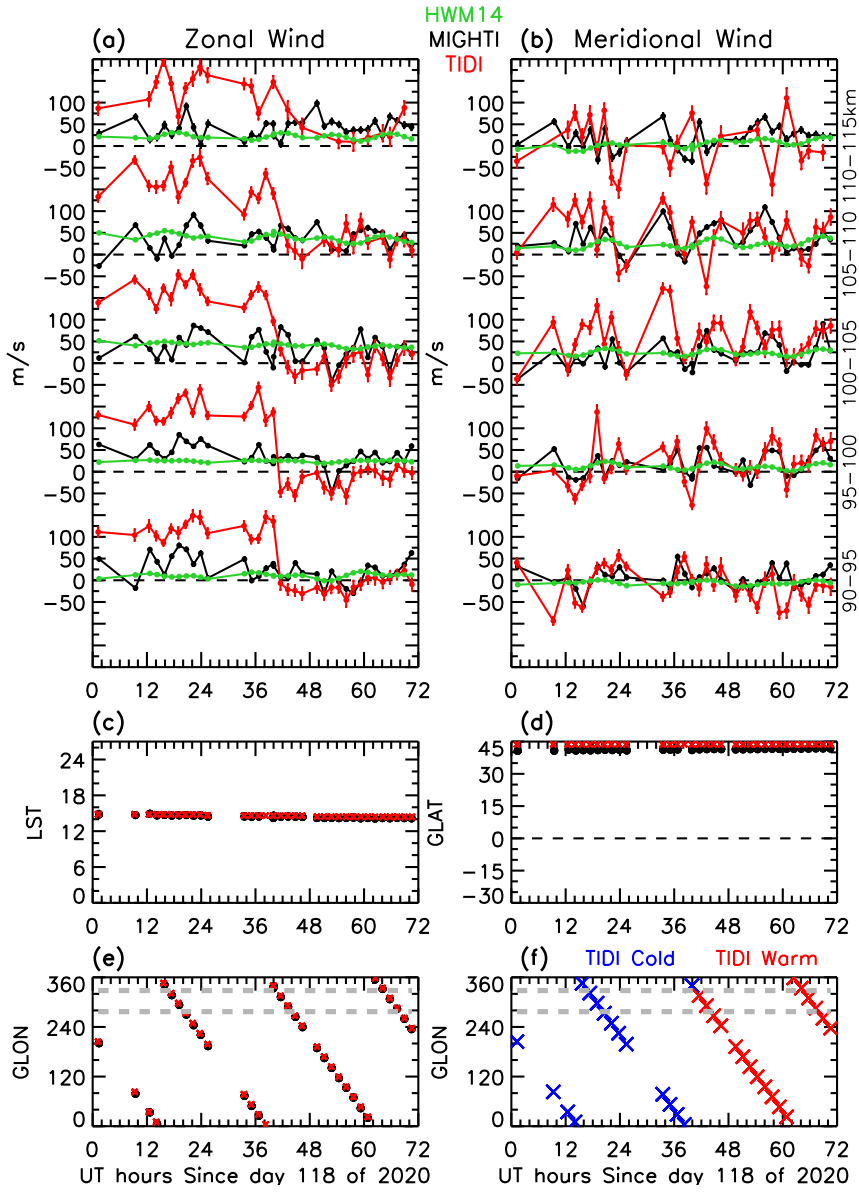


Figure 7. Same as Figure 5, but showing days from 118 to 120 that cover yaw transition from cycle 2 to 3 on day 119. The shown transition is from coldside to warmside.

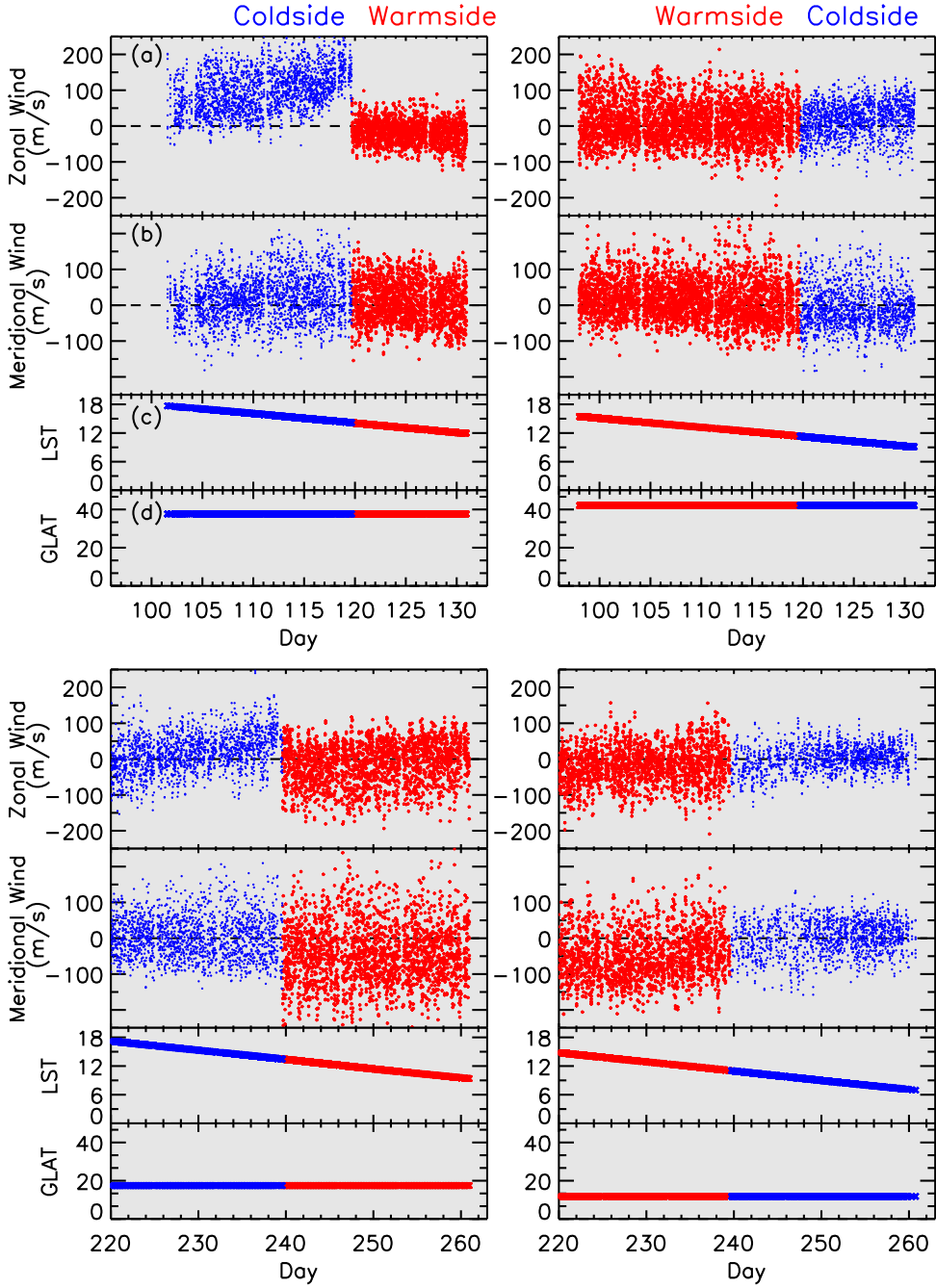


Figure 8. TIDI zonal (panel (a)) and meridional (panel (b)) winds (all longitudes, not just at MIGHTI and TIDI conjunctions) at the shown local solar times (panel (c)) and latitudes (panel (d)). The top two subplots show days between 96 and 133, and the bottom two subplots show days between 220 and 262. These days cover TIDI yaw cycle transitions occurring on days 119 and 239. TIDI covers two local times each day; the left column show TIDI one local time and the right column show other local time. The blue and red colors indicate measurements from coldside and warmside telescopes, respectively. The left and right columns panels show transitions from coldside to warmside and warmside to coldside telescopes, respectively.

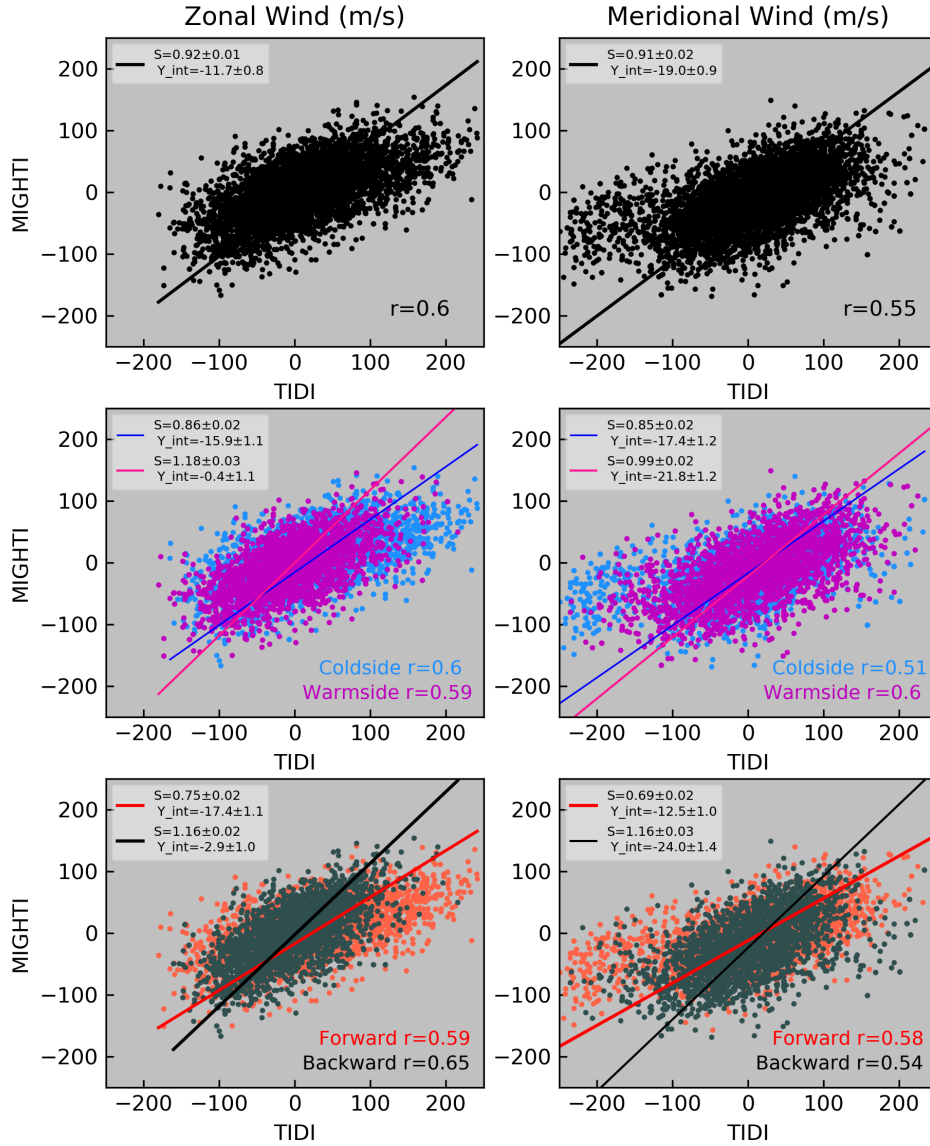


Figure 9. Comparison of MIGHTI and TIDI zonal winds (left column) and meridional winds (right column) at conjunctions between 95 km and 110 km. The top panels show winds at conjunctions collectively. The middle panels show winds at conjunctions from coldside and warmside telescopes separately. The bottom panels show winds at conjunctions from forward and backward traveling configurations of TIDI separately. Each panel show slopes (S) and intercepts (Y-int) obtained with orthogonal distance regression with involving wind errors in the fitting process. The correlation coefficient is also shown on the bottom of each panel. The color label descriptions are given in each panel.

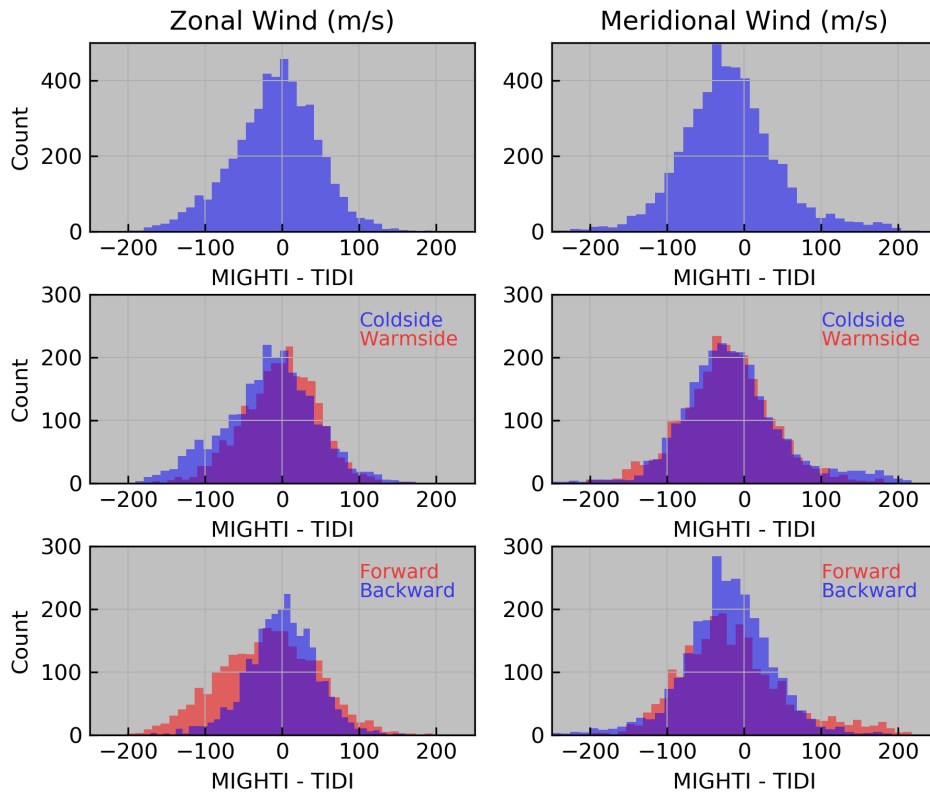


Figure 10. Histograms of MIGHTI and TIDI wind differences between 95 and 110 km. The top panels include all MIGHTI and TIDI conjunctions between 95 and 110 km. The middle panels include measurements from TIDI (and MIGHTI) coldside and warmside telescopes separately. The bottom panels show measurements from TIDI (and MIGHTI) forward and backward traveling configurations separately. The color label descriptions are given in each panel.

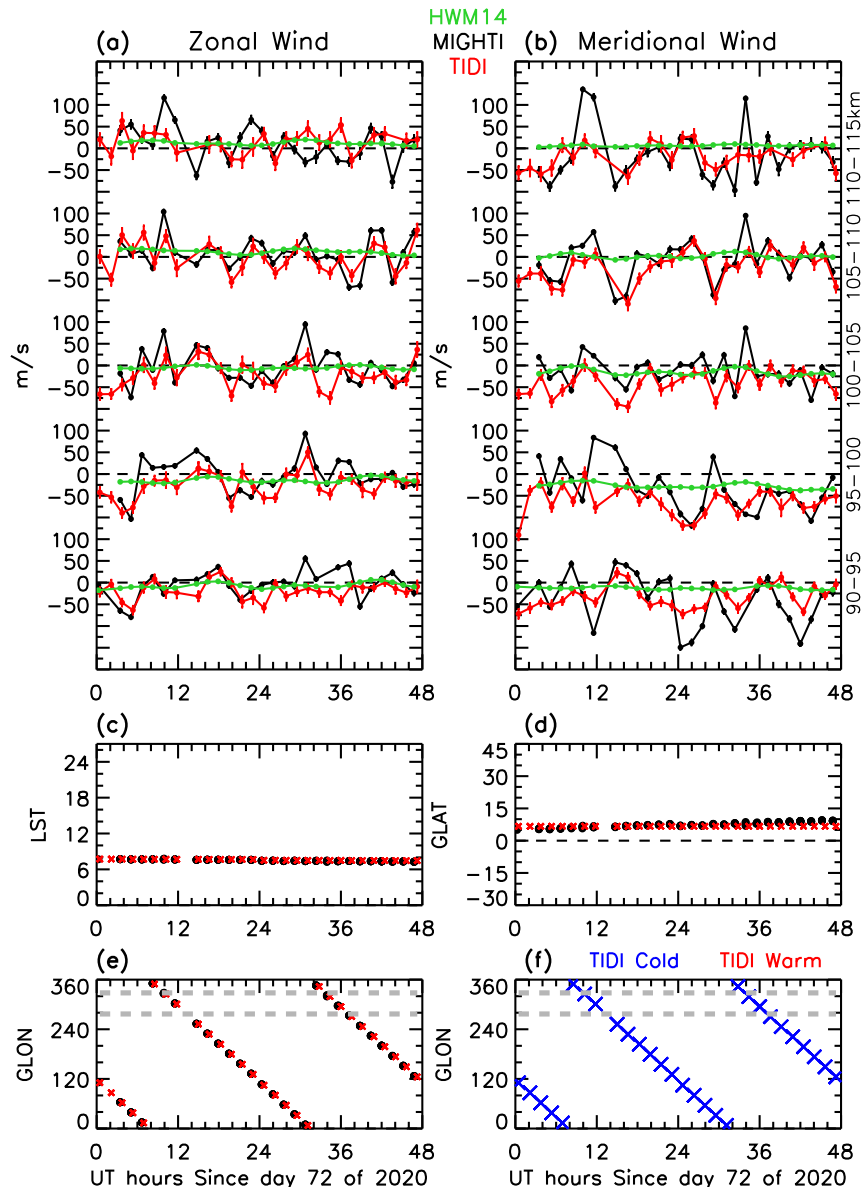


Figure 11. Same as Figure 6, but here using corrected TIDI winds by removing HWM14 baseline.