Viscosity of Fe-Ni-C Liquids up to Core Pressures and Implications for Dynamics of Planetary Cores

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Key Points:

- Experimental and computational determinations of viscosity of Fe-Ni-C liquids up to core pressures
- Viscosities of Fe-Ni-C liquids increase by a factor of ~2 at ~3-5 GPa at which a liquid structural transition occurs
- Fe-Ni-C liquid has higher viscosity than Fe and Fe-S liquid, causing difference in core dynamo and percolative core formation

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Abstract

The viscosity of iron alloy liquids is the key for the core dynamo and core-mantle differentiation of terrestrial bodies. Here we measured the viscosity of Fe-Ni-C liquids up to 7 GPa using the floating sphere viscometry method and up to 330 GPa using first principle calculations. We found a viscosity increase at ~3-5 GPa, coincident with a structural transition in the liquids. After the transition, the viscosity reaches ~14-27 mPa·s, a factor of 2-4 higher than that of Fe and Fe-S liquids. Our computational results from 5 to 330 GPa also indicate a high viscosity of the Fe-Ni-C liquids. For a carbon-rich core in large terrestrial body, the level of turbulence in the outer core would be lessened as approaching the inner core boundary. It is also anticipated that Fe-Ni-C liquids would percolate in Earth's deep silicate mantle at a much slower speed than Fe and Fe-S liquids.

Plain Language Summary

Liquid cores of Earth and other terrestrial planets are composed of Fe-Ni metal with certain amounts of light elements (LE) such as H, C, O, Si, and S. Located at the center of the planets, these cores are under extremely high-pressure high-temperature conditions. The viscosity of Fe-Ni-LE liquids is the key to understanding the formation and life duration of core dynamo. In the present study, we used experimental and theoretical methods to study the viscosity of Fe-Ni-C liquid, a candidate liquid core component, at high pressures and temperatures. Our results show that the viscosity of Fe-Ni-C liquids first increase quickly at ~3-5 gigapascal (GPa), and then increases slowly up to Earth's core conditions at 330 GPa and 5530 K. The fast growth of viscosity concurs with a liquid structure transition at around 5 GPa, around the core pressure ranges of the Moon, Mercury and several Jupiter's satellites, and might result in difference in viscous forces in core dynamos between those cores with pressures above and below transition pressure. The viscosity of Fe-Ni-C liquids at Earth's core conditions is higher than Fe and Fe-S liquids from previous studies, which provides new clues for the operation of geodynamo in a carbon-rich outer core.

1 Introduction

The liquid cores of Earth and other terrestrial planets or moons are composed of Fe-Ni metal, alloyed with a range of potential light elements (LE). The dynamics of these liquid cores determines the formation and the duration of the planetary dynamo and magnetic field, which are crucial for planetary habitability. Knowledge on the viscosity of candidate Fe-Ni-LE alloys at high pressures and high temperatures (HP-HT) are required to understand and model the dynamics of these liquid cores. Additionally, viscosity is one of the key properties that determines the percolation velocity of iron alloy through silicate/oxide rocks, which is crucial for studying the time scale of core formation in Earth and other terrestrial planets. Carbon is among the top candidate light elements in Earth's and planetary cores due to its high cosmic abundance, siderophile nature and ubiquity in iron meteorites. The phase diagram and physical properties of Fe-(Ni)-C solids and liquids at HP-HT have been widely investigated [*Chen et al.*, 2014; *Lai et al.*, 2017; *Nakajima et al.*, 2009; *Nakajima et al.*, 2015; *Prescher et al.*, 2015; *Terasaki et al.*, 2010; *Wood*, 1993], but studies on the viscosity of Fe-(Ni)-C liquid are limited due to the challenges of performing *in-situ* radiography observation and viscometry measurements on liquids at HP-HT conditions.

The development of *in-situ* X-ray radiography techniques enables the real-time monitoring of falling/floating sphere experiments at HP-HT conditions, which can be employed to study the viscosity of Fe-Ni-LE alloys melts (Table S1). Previous X-ray radiography studies have successfully measured the viscosity of Fe and Fe-S liquids [Dobson et al., 2000; Kono et al., 2015; LeBlanc & Secco, 1996; Perrillat et al., 2010; Rutter et al., 2002a; Rutter et al., 2002b; Secco et al., 2002; Terasaki et al., 2006; Terasaki et al., 2002; Terasaki et al., 2001; Urakawa et al., 2001], but the viscosity of other LE-bearing iron alloys at high pressure, including those with C, have only been sparsely studied to date. At ambient pressure, the viscosity of Fe-C was measured to be around 4-10 mPa·s at ~1600 K [Lucas, 1964], and the HP-HT viscosities of Fe-C liquids were only reported up to pressures of 4.5 GPa [Terasaki et al., 2006]. The viscosity of Fe₈₄C₁₆ scattered around 2-5 mPa·s at 3-4.5 GPa and 1843 K, which is similar to those of Fe and Fe-S liquids in this pressure/temperature range [Terasaki et al., 2006]. However, only three data points were reported on this Fe-C alloy, and as such there are insufficient constraints for fitting or for extended extrapolations to higher pressures. In addition, recent liquid structure studies found a polyamorphic liquid structural transition in Fe-(Ni)-C alloy melts at ~5 GPa [Lai et al., 2017; Shibazaki et al., 2015]. With a similar liquid structure transition in liquid Fe around its δ - γ liquid triple point [Sanloup et al., 2000], Terasaki et al. [2002] reported a strong decrease in viscosity through the transition pressure around 5 GPa, while Kono et al. [2015] reported no discontinuious change. It remains unclear whether such liquid structural transition can have a significant influence on their viscosity. In this study, we measured the viscosity of two Fe-Ni-C alloys (Fe₉₀Ni₁₀-3wt.% C and Fe₉₀Ni₁₀-5wt.% C) up to 7 GPa. In addition, we carried out first principles molecular dynamics (FPMD) calculations for the viscosity of Fe-Ni-C alloys with 1wt.%, 3wt.%, and 5wt.% C up to the Earth's core conditions, to investigate the influence on the viscosity of iron alloy from the incorporation of carbon and from the liquid structural transition.

2 Materials and Methods

2.1 Paris-Edinburgh Press Experiments

The starting materials were mixtures of Fe (99.9+%, Aldrich), Ni (99.99%, Aldrich) and graphite (99.995%, Alfa Aesar) powders. The mixtures were ground in acetone using an agate mortar for more than 1 h to achieve compositional homogeneity and then dried in a vacuum oven at high temperature of ~383 K overnight before being sealed in glass vials.

The viscosity measurement was conducted in a Paris-Edinburgh Press at 16-BM-B, Advanced Photon Source (APS), Argonne National Lab (ANL). A packed cylindrical sample of 1 mm diameter and 2 mm height was placed in a BN capsule, and a ruby sphere was placed near the bottom of the sample (Figure S1). The sample was first compressed to the target pressure and heated at 1073 K for 30-120 min to sinter the sample and capsule. The sample was then rapidly heated to 1973 K by applying a corresponding heating power, which is at least ~300-400 K above the melting temperature of Fe₉₀Ni₁₀-3wt.% C and Fe₉₀Ni₁₀-5wt.% C alloys at a given pressure for all the investigated pressure points, to ensure the sample was fully molten and the ruby can float freely [*Nakajima et al.*, 2009]. The movement of the ruby sphere upon melting was recorded by a high-speed X-ray camera at a frame rate of 500 frame/s for most experiments, except for two points at lower frame rate (Figure S2 and Figure S3). The viscosity of the molten sample was calculated by the Stokes' equation with correction factors for the effect of the wall (W) and the end (E) of a cylindrical sample container:

 $\eta = \frac{g d_s^2 (\rho_s - \rho_l)}{18\nu} \frac{F}{E} \quad (1)$ $F = 1 - 2.104 \left(\frac{d_s}{d_l}\right) + 2.09 \left(\frac{d_s}{d_l}\right)^3 - 0.95 \left(\frac{d_s}{d_l}\right)^5 \quad (2)$ $E = 1 + \frac{9}{8} \frac{d_s}{2Z} + \left(\frac{9}{8} \frac{d_s}{2Z}\right)^2 \quad (3)$

where d and ρ are the diameter and density, and subscripts s and l stand for the solid and liquid. Z is the height of cylindrical sample and v is the terminal velocity of falling/floating sphere. The diameter of the ruby sphere and the dimension of the sample was determined by the X-ray image with 2 μ m resolution. The densities of ruby sphere [Fei et al., 1995] and Fe-Ni-C liquid [Zhu et al., 2021] are determined by their equation of states at corresponding pressure and temperature. The density error of the Fe-Ni-C liquid can be significant [Zhu et al., 2021], but the density contrast between Fe-Ni-C and ruby of ~3 g/cm³ is large enough to minimize the error introduced by the errors in density of each phase. The large density contrast results in a large terminal velocity, where the high-speed camera is essential to precisely determine the terminal velocity. The pressure was determined by measuring the unit-cell volume of MgO ring by energy dispersive X-ray diffraction after the floating sphere experiment before the temperature is quenched and then adopted a calibrated pressure correction between sample and MgO ring [Kono et al., 2014]. The temperature was estimated from the power curve in calibration experiments with an estimated standard deviation of 50 K. The details of the experimental setup and viscometry method can be found in Kono et al. [2014].

2.2 First-principle Molecular Dynamics (FPMD) simulations

Details of the FPMD simulations were reported in previous publications and summarized here for completeness [Lai et al., 2017; Wang et al., 2019; Zhu et al., 2021]. Density Functional Theory (DFT) with plane wave basis sets as implemented in Vienna Ab initio Simulation Package (VASP) [Kresse & Furthmüller, 1996]. The Projector-Augmented Wave (PAW) method and exchange-correlation with the Perdew-Burke-Ernzerhof (PBE) functional were utilized in the Generalized Gradient Approximation (GGA) [Blöchl, 1994; Perdew et al., 1992]. The kinetic energy for the plane-wave basis was cut-off at 520.00 eV for equilibration simulation and 400.00 eV for equilibrium simulation. The Brillouin zone was sampled at Γ point. The FPMD simulations were carried out with the NVT ensemble and Nosé-thermostat for up to 3 ps for the system to relax in the equilibration run. After equilibration run, each of the FPMD simulations was run ~17 ps (17,000 steps) for statistical analysis of the viscosity. The time step was set to be 1 fs. The details of viscosity calculation was summarized previously [Wang et al., 2019]. The calculated pressures were benchmarked based on the experimental data on iron at similar conditions [Anderson & Ahrens, 1994]. The FPMD simulations were performed using a cubic computation supercell of 192 atoms. The ratio between Fe and Ni remains approximately 9:1 with C at 1%, 3%, and 5%, which is corresponding to 165 Fe +18 Ni + 9 C, 151 Fe + 17 Ni + 24 C, and 139 Fe +15 Ni + 38 C atoms respectively. A system of pure iron with 192 atoms was also simulated using the same method. The computational supercells were gradually adjusted to different volumes.

3 Results and Discussion

3.1 Floating sphere experiments on the viscosity of Fe-Ni-C liquid

Figure 1a and Movie S1 show examples of the movement of the ruby sphere in Fe-Ni-C melts, while Figure 1b & 1c convert these results into velocity as a function of time. The ruby sphere sped up quickly after sample melting to reach the terminal velocity, and then kept the terminal velocity for a distance before it decelerated and approached the top of the sample capsule. Ideally, with only the drag force from liquid, the gravity, and the buoyancy, the sphere should either accelerate to or move uniformly at terminal velocity. The long deceleration distance in Figure 1b before reaching the sample end implies additional force on the sphere due to the wall and the end. Therefore equations (1-3) with correction factors for those effect from the wall and the end were used to calculate the viscosities. The terminal velocity was determined by fitting the slope of distance-time curve in the range where the single-step velocities inside reached a nearly constant highest value before deceleration of the probe sphere as demonstrated by *Kono et al.* [2015], e.g. between 718 and 750 ms in Figure 1b & 1c. This terminal velocity, together with the dimensions of sample and ruby sphere, was then used to calculate the viscosity by equations (1-3), which is shown in Figure 1c.

Figure 2a and Table S2 summarizes the viscosity of the two Fe-Ni-C compositions in the experimentally investigated pressure range. The viscosity of Fe₉₀Ni₁₀-3wt.% C was in the range of about 8-12 mPa·s up to 4.2 GPa, which is slightly higher but still consistent with the estimates of previous studies at ~4-6 mPa·s within the uncertainties [Terasaki et al., 2006]. However, our results show marked increase of viscosity of Fe₉₀Ni₁₀-3wt.% C by a factor of ~2, reaching ~14-18 mPa·s at >4.8 GPa. Similar viscosity increase is also observed in the Fe₉₀Ni₁₀-5wt.% C liquid from 3.0 to 5.0 GPa. This substantial increase in viscosity at ~3-5 GPa coincides with the polyamorphous liquid structural transition in Fe-Ni-C liquids around 5 GPa [Lai et al., 2017], suggesting that it is likely related to the transition. The increased fraction of the 3-atom shared polyhedral motifs was inferred as the main mechanism behind this transition. Unlike solid phase transition, this growth is a gradual change in a range of pressure from FPMD calculation [Wang et al., 2019], indicating that the viscosity increase may also be gradual. It is difficult to precisely determine the initial transition pressure from our data: For Fe₉₀Ni₁₀-3wt.% C, it is possible that the viscosity increases steeply from 4.2 to 4.8 GPa; But if excluding the viscosity at 4.2 GPa, it is more likely that the viscosity undergoes an gradual growth from 2.7 to 4.8 GPa. Considering the nature of liquid structural transition, here we conclude that the major viscosity change occurs between ~3 to 5 GPa.

A previous study has reported an abrupt drop in viscosity of liquid Fe around 5 GPa near the δ-γ-liquid triple point in Fe phase diagram [*Terasaki et al.*, 2002]. The viscosity drop was considered to be associated with the liquid phase transition. However, a later study using a higher speed camera reported no obvious discontinuious change in the visocity of liquid Fe [*Kono et al.*, 2015]. Thus, whether liquid structure transitions can induce substantial viscosity changes in iron alloys remains unclear. In this study, the viscosity increases by about 2 times through the polyamorphous transition around ~3-5 GPa [*Lai et al.*, 2017], and remains more or less constant after the transition, showing a different trend from the previous observations in liquid Fe [*Kono et al.*, 2015; *Terasaki et al.*, 2002]. FPMD calculation also shows the viscosity of Fe₉₀Ni₁₀-5wt.% C substantially increases through the transition around 5 GPa before it flattens out at >8 GPa [*Wang et al.*, 2019]. This trend agrees with the positive correlation we observed in

the experiments, although the absolute value of the computational viscosity was slightly lower than those observed in our experiments.

3.2 Viscosity of Fe-Ni-(C) liquids to Earth's core conditions by FPMD calculations

The experimental results show a plateau of viscosity for both Fe-Ni-C liquids after the transition at > ~5 GPa, but the pressure range is limited. It is still unknown whether this trend will persist to higher pressures only based on experimental results at low pressures. The previous FPMD calculations calculated the viscosity of Fe-Ni-C liquids to 67 GPa, and found the viscosities are very similar between 8 and 48 GPa along an isothermal compression path [Wang et al., 2019]. However, the viscosity showed another discontinuity at 67 GPa, which might result from the 1673 K temperature too low to keep Fe₉₀Ni₁₀-5wt.% C fully molten. Without considering both the core temperature and pressure, it is difficult to extrapolate these results at low temperatures and pressures to Earth's core conditions.

In this study, we directly calculated the viscosity of Fe, Fe₉₀Ni₁₀-1wt.% C, Fe₉₀Ni₁₀-3wt.% C, and Fe₉₀Ni₁₀-5wt.% C at the core mantle boundary (CMB) and inner core boundary (ICB) conditions, using the method as described in Wang et al. [2019] (Figure S4). The viscosity of Fe is consistent with the previous calculation of Fe and Fe₉₀Ni₁₀ [Cao & Wang, 2017; de Wijs et al., 1998]. Our results show that the viscosity of Fe₉₀Ni₁₀-3wt.% C and Fe₉₀Ni₁₀-5wt.% C are similar at both calculated conditions, and they are ~2 mPa·s (30%) and ~7 mPa·s (55%) higher than the viscosity of Fe at CMB and ICB conditions (Figure 3). But the increase is not linearly correlated with the amount of carbon incorporated. The viscosity of Fe₉₀Ni₁₀-1wt.% C is slightly lower than Fe at CMB condition, but it becomes higher at ICB conditions. The viscosity of Fe₅₂S₁₂ from a previous FPMD calculation is lowest at ICB conditions among all Fe alloys [Alfè & Gillan, 1998]. Combined with the results from Wang et al. [2019], the viscosity of Fe-Ni-C liquids increases with pressure slightly faster than that of Fe-(Ni) liquid, but the increase rate is much lower compared with the substantial increase by a factor of 2 from 4.1 to 8.4 GPa. The viscosities at 8-48 GPa and 1673 K are only about 4-5 mPa·s lower than those at CMB conditions (135 GPa, 4050 K), and about 11-12 mPa·s lower than those at ICB conditions (330 GPa, 5530 K). The results suggest that either the pressure does not have a significant impact on the viscosity at > 5 GPa after the polyamorphous transition, or its effect on elevating viscosity is largely mitigated by the elevated temperature towards the ICB, which lowers the viscosity.

Compared with the experimental results at < 10 GPa, our calculation seems to systematically underestimate the viscosity by approximately a factor of 1.5-4.5. Due to the challenges in *ab initio* calculations, it is not uncommon that there exist discrepancies between computational and experimental results, especially for the liquid properties such as viscosity. However, our FPMD results are consistent with the experiments in terms of the pronounced viscosity increase around 5 GPa and the viscosity plateau after the transition (Figure 3). If the factor of 3 difference is uniformly applied to core pressures, the viscosities of Fe-Ni-C at Earth's core conditions are roughly estimated at ~33 mPa·s at CMB and ~58 mPa·s at ICB. Nevertheless, the incorporation of 1-5 wt.% C in liquid Fe could significantly enhance the rate at which the viscosity increases with pressures and Fe-Ni-C liquids generally have higher viscosity than Fe, FeNi, and Fe-S liquids.

3.3 Implications for the dynamics of planetary cores

The pressure of 4-5 GPa at which a substantial increase in the viscosities of the Fe-Ni-C liquids lies at a pressure close to core pressures within a number of small planets and moons, such as Moon, Mercury, Ganymede, Io, and Europa [Garcia et al., 2012; Hauck et al., 2013; Kronrod & Kuskov, 2006] (Figure 2b). The liquid cores of the Moon, Mercury, Ganymede and Io are estimated to be under pressures close to or above this transition pressure, while Europa has a central pressure of ~4 GPa. If carbon is one of the major light elements in the cores of some of these terrestrial bodies, their core viscosity may be distinctly different due to the liquid structural transition. These differences will result in distinct viscous forces within their cores and thus dynamos. The variation of viscosity due to the effect of pressure and/or liquid structural transition in Fe-C liquids as revealed in this study or some other liquid alloys such as Fe and Fe-S liquids may render the extrapolations to higher pressures unreliable. Therefore, the viscosity of these planetary liquid cores may be significantly different from those inferred based on low pressure (0-5 GPa) measurements, depending on the core pressure and the light element species.

The viscosity of Fe-Ni-C after the transition at ~4.5 GPa becomes higher than Fe-Ni in both experiments and calculations, and this relationship persists over the pressure range of Earth's core, as shown from the FPMD calculations. This indicates that a C-rich liquid core in planets or moons with core pressures higher than 4.5 GPa would be more viscous in conjunction with generating core dynamos than Fe-Ni cores. However, the magnitude of viscosity is still in the range of tens of mPa·s, which remains small. In particular, Coriolis and inertial forces will continue to dominate within the core, as illustrated by the low Ekman number of the outer core (O~10⁻¹⁵: [Christensen & Aubert, 2006]). Nevertheless, the transition to a higher viscosity near 4.5 GPa, and the positive increase in viscosity across the depth range of the outer core for carbon-bearing alloys implies that Earth's outer core may lie modestly closer to the Ekman number range of current core geodynamo simulations. These simulations typically have, because of computational limitations, artificially high viscosities — often ~9 orders of magnitude larger than the core itself [e.g. King & Buffett, 2013; Schwaiger et al., 2019]. Our results suggest that the Earth's core might lie modestly closer to the domain in which current geodynamic simulations are conducted.

3.4 Implication for the percolative core formation and growth

The percolation of Fe-Ni-LE liquids through a solid silicate matrix is a possible mechanism for the formation of cores of terrestrial planets and moons in their early history. The segregation velocity, V_m , for a connected melt can be expressed as

$$V_m = \frac{k\Delta\rho g}{\varphi\mu} \quad (4)$$

in which k is permeability, $\Delta \rho$ is the density difference between solid and liquid, g is the acceleration of gravity, φ is the melt fraction, and μ is the viscosity of melt [Faul, 1997; McKenzie, 1989]. The viscosity has an inverse relation to the segregation velocity, which is vital in evaluating the time scale over which core-forming melts drained through early Earth's or protoplanetary silicate mantle. Considering the observed viscosity increase of Fe-Ni-C liquids at ~3-5 GPa, the segregation velocity could drop by several times at the corresponding depth. At higher pressures, the viscosity increases slowly and remains consistently higher than pure Fe-Ni, which means the segregation velocity of Fe-Ni-C will be consistently slower than Fe-Ni. A segregation velocity drop will delay the rate at which C-bearing Fe-Ni alloy enters the core

through percolation. Based on previous viscosity estimates (~3-6 mPa·s), assuming the melt fraction $\varphi = \sim 0.1$, $\Delta \rho = 3500 \text{ kg/m}^3$ and $k = \sim 10^{-15} \text{ m}^2$ and the viscosity remains constant in the entire mantle, the V_m is roughly estimated ~2 m/yr and melt would need 1.5 Myr to traverse through a 3000 km radius mantle, which is significantly faster than the 30 Myr core formation process estimated from mantle Hf-W isotopic ratios [*Kleine et al.*, 2002]. However, if the viscosity is increased to ~20-30 mPa·s, the time to drain the melt increases to ~10 Myr. On the other hand, whether the enhanced viscosity of C-bearing alloys might also affect their wetting behavior is unknown, but scaling relations have been proposed between viscosity and surface energies [e.g. Egry, 1993]. Hence, it is anticipated that the viscosity change may influence the percolation velocity of the C-rich liquids in a silicate matrix.

4 Conclusions

We have carried out viscosity measurements of Fe-Ni-C liquids up to 7 GPa and calculations up to 330 GPa. A pronounced viscosity increase was observed at ~3-5 GPa, concurring with a liquid structural transition in the Fe-Ni-C liquids. The transition pressure locates around the core pressures of many small planets and moons such as the Moon, Mercury and satellites of Jupiter. If their cores contain significant amount of carbon, the viscosity variation may result differences of the viscous forces within their cores at pressures above or below 4.5 GPa. Considering similar liquid structural transitions exist in other Fe alloys, whether the viscosity discontinuity through the structural transition is ubiquitous needs further investigation.

The viscosity after the transition is higher than the viscosity of Fe and Fe-S liquids from both experiments and calculations. The calculated viscosity at pressures greater than 5 GPa increases progressively and slowly up to inner core boundary conditions at 330 GPa and 5530 K. The high viscosity of Fe-Ni-C liquids at > 5GPa implies that a C-rich core in any large terrestrial body would have both a larger Ekman number and smaller Reynolds number than generally inferred by about half an order of magnitude. As such, the level of turbulence in the outer core would be lessened as approaching the inner core boundary, and the higher viscosity of a C-rich core of the Earth would imply that current artificially high-viscosity geodynamo simulations are modestly closer to the potential conditions within Earth's core. Fe-Ni-C liquids would also be expected to percolate in Earth's deep silicate mantle at a much slower speed compared with Fe and Fe-S liquids. This effect will expand the timeframe over which late-descending C-rich iron alloys will reach the core, and will allow for more time for isotopic fractionation during the coremantle differentiation.

Acknowledgments, Samples, and Data

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References

- Alfè, D., & Gillan, M. J. (1998), First-principles calculation of transport coefficients. *Physical Review Letters*, 81(23), 5161.
- Anderson, W. W., & Ahrens, T. J. (1994), An equation of state for liquid iron and implications for the Earth's core. *Journal of Geophysical Research: Solid Earth*, 99(B3), 4273-4284.
- Blöchl, P. E. (1994), Projector augmented-wave method. *Physical Review B*, 50(24), 17953.
- Cao, Q. L., & Wang, P. P. (2017), Stokes- Einstein relation in liquid iron- nickel alloy up to 300 GPa. *Journal of Geophysical Research: Solid Earth*, 122(5), 3351-3363.
- Chen, B., Li, Z., Zhang, D., Liu, J., Hu, M. Y., Zhao, J., Bi, W., Alp, E. E., Xiao, Y., & Chow, P. (2014), Hidden carbon in Earth's inner core revealed by shear softening in dense Fe7C3.

 Proceedings of the National Academy of Sciences, 111(50), 17755-17758.
- Christensen, U. R., & Aubert, J. (2006), Scaling properties of convection-driven dynamos in rotating spherical shells and application to planetary magnetic fields. *Geophysical Journal International*, 166(1), 97-114.
- de Wijs, G. A., Kresse, G., Vočadlo, L., Dobson, D., Alfe, D., Gillan, M. J., & Price, G. D. (1998), The viscosity of liquid iron at the physical conditions of the Earth's core. *Nature*, 392(6678), 805-807.
- Dobson, D. P., Crichton, W. A., VOCadlo, L., Jones, A. P., Wang, Y., Uchida, T., Rivers, M., Sutton, S., & Brodholt, J. P. (2000), In situ measurement of viscosity of liquids in the Fe-FeS system at high pressures and temperatures. *American Mineralogist*, 85(11-12), 1838-1842.
- Egry, I. (1993), On the relation between surface tension and viscosity for liquid metals. *Scripta Metallurgica et Materialia*; (United States), 28(10).
- Faul, U. H. (1997), Permeability of partially molten upper mantle rocks from experiments and percolation theory. *Journal of Geophysical Research: Solid Earth*, 102(B5), 10299-10311.
- Fei, Y. (1995), Thermal Expansion, In: *Mineral Physics and Crystallography, A Handbook of Physical Constants*, Vol. 2 (AGU Reference Shelf), pp. 29-44.
- Garcia, R. F., Gagnepain-Beyneix, J., Chevrot, S., & Lognonné, P. (2012), Erratum to "Very Preliminary Reference Moon Model", by RF Garcia, J. Gagnepain-Beyneix, S. Chevrot, P. Lognonné [Phys. Earth Planet. Inter. 188 (2011) 96–113]. *Physics of the earth and planetary interiors*, 202, 89-91.
- Hauck, S. A., Margot, J. L., Solomon, S. C., Phillips, R. J., Johnson, C. L., Lemoine, F. G.,
 Mazarico, E., McCoy, T. J., Padovan, S., & Peale, S. J. (2013), The curious case of
 Mercury's internal structure. *Journal of Geophysical Research: Planets*, 118(6), 1204-1220.

- King, E. M., & Buffett, B. A. (2013), Flow speeds and length scales in geodynamo models: the role of viscosity. *Earth and Planetary Science Letters*, *371*, 156-162.
- Kleine, T., Münker, C., Mezger, K., & Palme, H. (2002), Rapid accretion and early core formation on asteroids and the terrestrial planets from Hf–W chronometry. *Nature*, 418(6901), 952-955.
- Kono, Y., Park, C., Kenney-Benson, C., Shen, G., & Wang, Y. (2014), Toward comprehensive studies of liquids at high pressures and high temperatures: Combined structure, elastic wave velocity, and viscosity measurements in the Paris–Edinburgh cell. *Physics of the earth and planetary interiors*, 228, 269-280.
- Kono, Y., Kenney-Benson, C., Shibazaki, Y., Park, C., Shen, G., & Wang, Y. (2015), High-pressure viscosity of liquid Fe and FeS revisited by falling sphere viscometry using ultrafast X-ray imaging. *Physics of the earth and planetary interiors*, 241, 57-64.
- Kresse, G., & Furthmüller, J. (1996), Efficiency of ab-initio total energy calculations for metals and semiconductors using a plane-wave basis set. *Computational materials science*, 6(1), 15-50.
- Kronrod, V., & Kuskov, O. (2006), Chemical differentiation of the Galilean satellites of Jupiter:

 4. Isochemical models for the compositions of Io, Europa, and Ganymede. *Geochemistry International*, 44(6), 529-546.
- Lai, X., Chen, B., Wang, J., Kono, Y., & Zhu, F. (2017), Polyamorphic Transformations in Fe-Ni- C Liquids: Implications for Chemical Evolution of Terrestrial Planets. *Journal of Geophysical Research: Solid Earth*, 122(12), 9745-9754.
- LeBlanc, G., & Secco, R. (1996), Viscosity of an Fe- S liquid up to 1300° C and 5 GPa. *Geophysical Research Letters*, 23(3), 213-216.
- Lucas, L.-D. (1964), Viscosite du fer pur et du systeme Fe-C jusqu'a 4.8% C en poids, C. R. Acad. Sci., 259, 3760–3767.
- McKenzie, D. (1989), Some remarks on the movement of small melt fractions in the mantle. *Earth and Planetary Science Letters*, 95(1-2), 53-72.
- Nakajima, Y., Takahashi, E., Suzuki, T., & Funakoshi, K.-i. (2009), "Carbon in the core" revisited. *Physics of the earth and planetary interiors*, 174(1), 202-211.
- Nakajima, Y., Imada, S., Hirose, K., Komabayashi, T., Ozawa, H., Tateno, S., Tsutsui, S., Kuwayama, Y., & Baron, A. Q. (2015), Carbon-depleted outer core revealed by sound velocity measurements of liquid iron–carbon alloy. *Nature communications*, 6, 8942.
- Perdew, J. P., Chevary, J. A., Vosko, S. H., Jackson, K. A., Pederson, M. R., Singh, D. J., & Fiolhais, C. (1992), Atoms, molecules, solids, and surfaces: Applications of the generalized gradient approximation for exchange and correlation. *Physical Review B*, 46(11), 6671.
- Perrillat, J.-P., Mezouar, M., Garbarino, G., & Bauchau, S. (2010), In situ viscometry of highpressure melts in the Paris–Edinburgh cell: application to liquid FeS. *High Pressure Research*, 30(3), 415-423.
- Prescher, C., Dubrovinsky, L., Bykova, E., Kupenko, I., Glazyrin, K., Kantor, A., McCammon, C., Mookherjee, M., Nakajima, Y., & Miyajima, N. (2015), High Poisson's ratio of Earth's inner core explained by carbon alloying. *Nature Geoscience*, 8(3), 220.
- Rutter, M. D., Secco, R. A., Uchida, T., Liu, H., Wang, Y., Rivers, M. L., & Sutton, S. R. (2002a), Towards evaluating the viscosity of the Earth's outer core: An experimental high pressure study of liquid Fe- S (8.5 wt.% S). *Geophysical Research Letters*, 29(8), 58-51-58-54.

- Rutter, M. D., Secco, R. A., Liu, H., Uchida, T., Rivers, M. L., Sutton, S. R., & Wang, Y. (2002b), Viscosity of liquid Fe at high pressure. *Physical Review B*, 66(6), 060102.
- Sanloup, C., Guyot, F., Gillet, P., Fiquet, G., Hemley, R., Mezouar, M., & Martinez, I. (2000), Structural changes in liquid Fe at high pressures and high temperatures from synchrotron X-ray diffraction. *EPL* (*Europhysics Letters*), 52(2), 151.
- Schwaiger, T., Gastine, T., & Aubert, J. (2019), Force balance in numerical geodynamo simulations: a systematic study. *Geophysical Journal International*, 219(Supplement_1), S101-S114.
- Secco, R., Rutter, M., Balog, S., Liu, H., Rubie, D., Uchida, T., Frost, D., Wang, Y., Rivers, M., & Sutton, S. (2002), Viscosity and density of Fe–S liquids at high pressures. *Journal of Physics: Condensed Matter*, *14*(44), 11325.
- Shibazaki, Y., Kono, Y., & Fei, Y. (2015), Microscopic structural change in a liquid Fe- C alloy of~ 5 GPa. *Geophysical Research Letters*, 42(13), 5236-5242.
- Terasaki, H., Suzuki, A., Ohtani, E., Nishida, K., Sakamaki, T., & Funakoshi, K.-i. (2006), Effect of pressure on the viscosity of Fe-S and Fe-C liquids up to 16 GPa. *Geophysical Research Letters*, 33(22).
- Terasaki, H., Kato, T., Urakawa, S., Funakoshi, K., Sato, K., Suzuki, A., & Okada, T. (2002), Viscosity change and structural transition of molten Fe at 5 GPa. *Geophysical Research Letters*, 29(8), 68-61-68-63.
- Terasaki, H., Nishida, K., Shibazaki, Y., Sakamaki, T., Suzuki, A., Ohtani, E., & Kikegawa, T. (2010), Density measurement of Fe3C liquid using X- ray absorption image up to 10 GPa and effect of light elements on compressibility of liquid iron. *Journal of Geophysical Research: Solid Earth*, 115(B6).
- Terasaki, H., Kato, T., Urakawa, S., Funakoshi, K.-i., Suzuki, A., Okada, T., Maeda, M., Sato, J., Kubo, T., & Kasai, S. (2001), The effect of temperature, pressure, and sulfur content on viscosity of the Fe–FeS melt. *Earth and Planetary Science Letters*, 190(1-2), 93-101.
- Urakawa, S., Terasaki, H., Funakoshi, K., Kato, T., & Suzuki, A. (2001), Radiographic study on the viscosity of the Fe-FeS melts at the pressure of 5 to 7 GPa. *American Mineralogist*, 86(4), 578-582.
- Wang, J., Chen, B., Williams, Q., & Manghnani, M. (2019), Short-and intermediate-range structure and dynamics of Fe-Ni-C liquid under compression. *Frontiers in Earth Science*, 7, 258.
- Wood, B. J. (1993), Carbon in the core. Earth and Planetary Science Letters, 117(3-4), 593-607.
- Zhu, F., Lai, X., Wang, J., Amulele, G., Kono, Y., Shen, G., Jing, Z., Manghnani, M. H., Williams, Q., & Chen, B. (2021), Density of Fe- Ni- C liquids at high pressures and implications for liquid cores of Earth and the Moon. *Journal of Geophysical Research: Solid Earth*, *126*(3), e2020JB021089.

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Figure captions

Figure 1. A representative example of the movement of the ruby sphere upon sample melting. The sample was Fe₉₀Ni₁₀-3 wt.%C and the experimental conditions were 1.8 GPa and 1973 K. (a) The movement of the ruby sphere at a time interval of 75 ms between each frame, from 475 ms to 925 ms. The diameter of the Al₂O₃ sphere is ~250 μ m. (b) The distance-time relation was converted to a velocity-time relation to determine the terminal velocity. (c) The terminal velocity was used to calculate the liquid viscosity.

Figure 2. Viscosities of Fe and Fe alloys (a) Viscosities of Fe₉₀Ni₁₀-3 wt.%C and Fe₉₀Ni₁₀-5 wt.%C in comparison with previous results. The viscosities of both components increases from ~5-8 mPa·s at <3 GPa to ~15-27 mPa·s at ~5-7 GPa, which may result from the polyamorphic transition around 5 GPa. Open red and blue symbols represent an upper limit of viscosities at those pressures, due to a possibility of not reaching terminal velocity. (b) The polyamorphic transition pressure of Fe (light gray bar) and FeNi-C (light red bar) liquids in comparison with the core pressure ranges of small planet and moons (black lines, arrow means central pressure higher than 7 GPa).

Figure 3. Viscosity of Fe-Ni-C liquids to Earth's core pressure from FPMD calculations. The viscosity increases quickly from 3.3 to 6.0 mPa·s at 4.1-8.4 GPa, corresponding to a polyamorphic transition in the liquid (light blue arrow), and then flattens at 8.4-48 GPa (light red arrow) [*Wang et al.*, 2019]. The viscosities from CMB (4050 K) to ICB (5530 K) conditions increase slightly for all Fe-Ni-C species (light red band) and remain similar for Fe liquid. The gray band shows the calculated viscosity of Fe₉₀Ni₁₀ at 50-300 GPa from melting point (T_m , 2200 K at 50 GPa and 6500 K at 300 GPa) to 1.85 times of T_m [*Cao & Wang*, 2017]. Previous results of Fe [*Alfè & Gillan*, 1998; *de Wijs et al.*, 1998] and Fe-S [*Alfè and Gillan*, 1998] are also plotted for comparison. Some uncertainties for our calculated viscosities are within the symbols and can be found in Table S3.





