

## **The Chicxulub Impact Produced a Powerful Global Tsunami**

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### **Contents of this file**

Text S1

Figures S1-S4

Tables S1-S4

Captions for Movies S1-S4

## **Additional Supporting Information (Files uploaded separately)**

Tables S5 and S6

Movies S1-S4

### **Introduction**

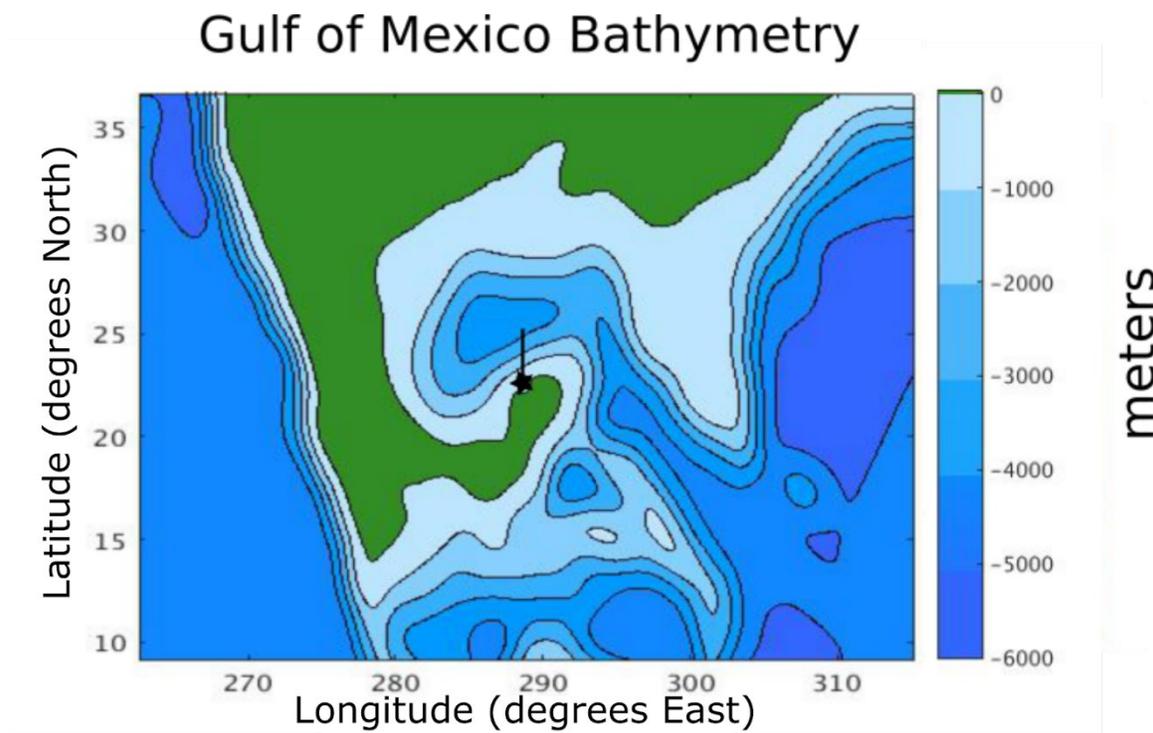
In this file and in related Supporting Information, we include results of the hydrocode simulation and the shallow water ocean model simulations run with parameters different from those of the 'fiducial' runs. For instance, we display results from MOM6 simulations with coarser grid spacing ( $1/5^\circ$ ) than the  $1/10^\circ$  grid spacing in the fiducial run, and discuss sensitivities to handoff times and other parameters of our coupled hydrocode/shallow-water code simulations. We also include descriptions of the construction of paleo-bathymetries, the paleobathymetry near the impact site, the importance of vaporization in hydrocode modeling, the procedure for blending hydrocode results with the paleo-bathymetry to produce initial conditions for the shallow-water model, the sensitivities of the shallow-water model to different aspects of the handoff from the hydrocode (e.g., the presence or absence of a rim wave, amongst other aspects), and the comparison of energies in the impact tsunami with energies in recent historical tsunamis. Finally, we include four movies: two of the iSALE hydrocode simulation over the first ten minutes of the impact, and one animation each of the MOM6 and MOST shallow-water simulation showing the tsunami propagation throughout the global ocean.

### **Text S1.**

#### ***Development of paleobathymetry and initial conditions***

The paleobathymetry was created by merging two datasets. Müller et al. (2008) uses a basin-age depth relation to create a deep-ocean bathymetric product. Basin age-depth relationship (Schroeder, 1984; Smith & Sandwell, 1997) refers to the relationship of the depth of the oceanic crust to the age of the oceanic crust. As the oceanic plate moves away from the divergent boundary, it cools and sinks. This relationship can then be used to construct bathymetries. The second dataset, from Scotese (PALEOMAP), is reconstructed through "backtracking" the formation of existing seafloor and lithospheric plate movement. When there is no modern sea floor that can be used to reconstruct the paleobathymetry, the PALEOMAP dataset sets a static water level. Because of this incomplete PALEOMAP bathymetry, in particular the lack of mid-ocean ridge positions in pre-modern static seafloor areas provided by PALEOMAP, we needed a supplemental dataset. We used mid-ocean ridge information from Müller et al. (2008) to supplement the PALEOMAP data, and used the shelf bathymetry from Scotese. A Blackman filter was used to smooth the combined bathymetries, thus preventing large bathymetric discontinuities. The paleobathymetry near the impact site is displayed in Figure S1.

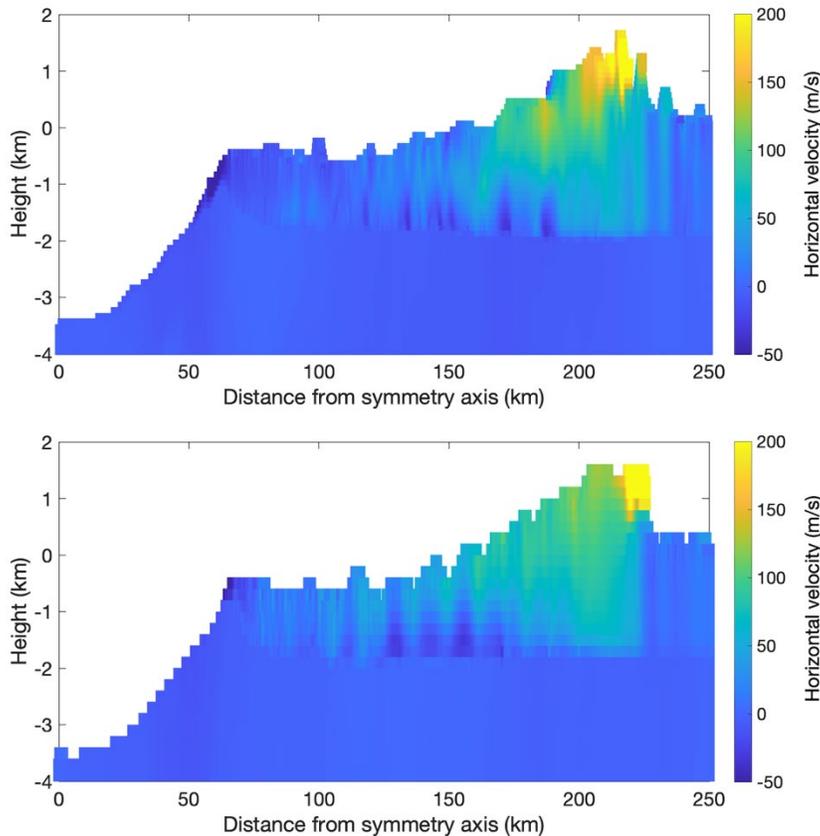
Our iSALE setup is described in Table S1. The model setup is the same as in Collins et al. (2008), but with higher resolution and larger model extent. The grid spacing of 100 m is the same grid spacing used by Bahlburg et al. (2010) to simulate the formation of the Chicxulub crater in a 1 km deep ocean. The simulations of Bahlburg et al. (2010) were focused on whether the crater rim would inhibit flow back into the recently formed Chicxulub crater. Thus, the high-resolution zone of Bahlburg et al. (2010) had a horizontal extent of less than 100 km. Because our simulations are focused on formation and evolution of the rim wave, the high-resolution zone in our baseline simulation has a horizontal extent of 250 km.



**Figure S1.** Pre-impact K/Pg bathymetry of Gulf of Mexico in meters. The impact occurred in shallow water of 100-200 m depth, at the location indicated by a star. At >50 km from the impact, the Gulf of Mexico seafloor depth is > 1 km, the depth used in the hydrocode simulation shown in Figure 1. The black line shown above has a length of 250 km, the length of the domain used in the hydrocode simulation.

We also ran an iSALE simulation, out to 1100 s, with the horizontal mesh extending to 300 km, to test sensitivity to handoff time. In particular, the longer run allows us to ensure that the wave is done with active plunging breaking, the process that cannot be modeled with the shallow water wave models, once the handoff occurs. However, the wave has clearly reached a virtually steady-state phase of propagation as a bore-

like wave. Bore-like waves are processes that can be modeled with the nonlinear shallow water wave approximation well, as the hydrocode simulation that runs out to 1100 s shows. In this longer hydrocode run we also limited the vertical extent of the domain to 50 km above the impact point. Ejecta leaving the mesh is therefore removed from the simulation. This choice removes the discontinuous clumps of ejecta seen in Figure 1b. The clumps landing ahead of the rim wave producing transient waves in the water make it unclear if the wave is experiencing strong plunging breaking or simply interacting with these lower amplitude waves. Movie S2 shows the results of this simulation. Out to about 200 s the simulation is nearly identical to the simulation shown in Figure 1. Without the added 'clumps' of ejecta landing ahead of the rim wave it is clear the wave is done with plunging breaking by ~600 s. At 600 s, the wave is more energetic in the case that includes the discontinuous ejecta, but this wave dissipates more quickly than the case that ignores the discontinuous ejecta (Figure S4). Figure S4 also shows that handoffs to MOM6 at 600 s vs. 850 s yield nearly identical results, as quantified by globally integrated energies. The simulation with vertical extent of the mesh limited to 50 km produces a wave that is 1.2 km high at 600 s, whereas the wave shown in Figure 1 has an amplitude of 1.4 km at 600 s. Therefore, the height of the rim wave is not overly sensitive to the presence of late-arriving ejecta.



**Figure S2.** iSALE resolution test. Material is colored according to horizontal velocity as indicated by the color scale. Both frames are plotted 600 s after an impact into a target with a 2 km thick ocean layer. The top frame has a resolution of 100 meters while the bottom frame has a resolution of 200 meters. Resting sea level is represented by a y-axis value of “0”.

Figure S2 shows that the waveform produced in a 200 m grid spacing iSALE simulation is very similar to that produced in a 100 m simulation. This similarity provides confidence that resolving the ocean depth by at least 10 cells is sufficient to capture the initial generation of the impact tsunami.

### ***The importance of vaporization in hydrocode modeling***

Gisler et al. (2011) pointed out the importance of accounting for vaporization of water when considering impact-generated tsunamis. We expect in simulations where the crater is much deeper than the ocean the effect of ocean vaporization will have a modest effect compared to impacts where the transient crater size is much smaller than the ocean depth. To accurately track the response of material to an impact, shock physics codes must include equations of state and accurately track the thermal state of materials including the effect of vaporization. To limit computational expense, vaporized material is removed from the mesh when densities drop below

10 kg/m<sup>3</sup>. This is why no vapor plume appears in Figure 1. In addition to producing an airblast, expansion of the vapor plume may also produce Lamb waves as observed in the recent Tonga event. The Lamb waves would likely enhance the impact-generated tsunami.

### ***Importing hydrocode outputs to the tsunami model***

The hydrocode outputs two quantities that will be used in the initial conditions for elevation in our shallow-water models; sea surface elevation perturbations, and perturbations to the seafloor depth. Because our impact simulations are axisymmetric, we write the output of the hydrocode in terms of a radial distance from the impact point to a target point of interest. Let the distance from our target point to the impact origin be  $r$ . Then  $\eta_{hydrocode}(r)$  is the perturbation (positive upwards) to resting sea level given by the hydrocode, in columns where water exists. Of course, in columns near the impact, there is no water at the end of the hydrocode simulation (600-1100 seconds post-impact). Let  $H_{hydrocode}(r)$  be the vertical distance between the seafloor prior to impact and the seafloor at the hand-off from the hydrocode simulation. Again, we take positive values of  $H_{hydrocode}(r)$  to denote a deepening of the seafloor. We note that the seafloor can be either sediment or crystalline basement.

For the fiducial “Half Crater” case, we impose the simulated crater on the pre-impact bathymetry only at cells that were initially water (in other words, the crater was not imposed on points that were land in the pre-impact bathymetry). More precisely, if  $H_{before}(x,y) + H_{hydrocode}(r) > 0$ , then we declare the point  $(x,y)$  as being water, with resting depth  $H_{before}(x,y) + H_{hydrocode}(r)$ , where  $H_{before}$  is the pre-impact bathymetry, discussed in more detail below. Negative values of this sum imply that the target point is land.

For the rim wave, the water was only added onto cells that had water, such that the rim wave is not emplaced on land.

For the post-impact points that are water, including those in the crater region, the new water column thickness is  $H_{before}(x,y) + H_{hydrocode}(r) + \eta_{hydrocode}(r)$ . In water points in the crater region, the initial water column thickness just after impact may be zero. This water is given the depth averaged velocity according to iSALE.

In all regions where there was water, the water column velocity was vertically averaged and also placed in each grid cell. The velocities (as well as sea surface heights and crater depths) are then interpolated to a 1/10th degree resolution and smoothed. For the ‘Crater Only’ simulation water was vacated from the crater, but no rim wave was initialized. Initial water velocities were zero. For the ‘Full Crater’ simulation the crater was imposed on the bathymetry including cells that were previously land.

**Table S1.** Model setup for fiducial iSALE simulation

<b>Description</b>	<b>Value</b>
Size of high-resolution cell	100 m
Number of high-resolution cells, horizontal direction	2500
Number of high-resolution cells, vertical direction	600
Physical dimension of entire mesh, horizontal direction	0 to 344.3 km
Physical dimension of entire mesh, vertical direction	-505.1 to 142.9 km
Adaptive time step	0.78 to 2.7 ms

**Table S2.** Setup for MOM6 1/10<sup>th</sup> degree simulations

<b>Description</b>	<b>Value</b>
Spatial resolution	1/10 degree
North wall boundary	82.05° N
South wall boundary	86.05° S
Temporal resolution	10 seconds

**Table S3.** MOST model setup

<b>Description</b>	<b>Value</b>
Spatial resolution	1/10 degree
North wall boundary	82.05° N
South wall boundary	86.05° S
Temporal resolution	6 seconds

### ***Initial Seafloor Topography***

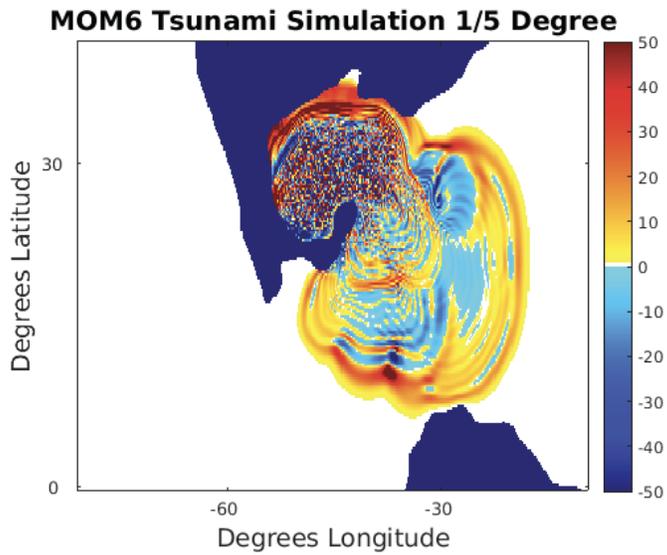
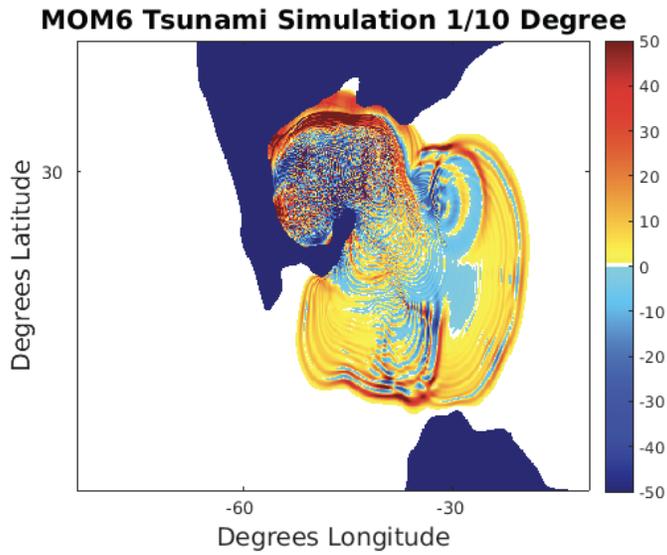
Because the fiducial hydrocode simulation covers points within 250 km of the impact origin, all such points will be part of the initial condition, set by the outputs of the hydrocode, for the shallow-water model. Let the coordinates of a target point, within 250 km of the impact origin, be  $(x,y)$ . Let  $H_{before}(x,y)$  be the topography (resting water depth) of the target point before impact (defined by the merging of the Scotese and Müller datasets, described earlier). The sign of  $H_{before}$  informs us about whether the target point was land (negative  $H_{before}$  value) or ocean (positive  $H_{before}$  value) prior to impact.

### ***Shallow-water model results and sensitivities***

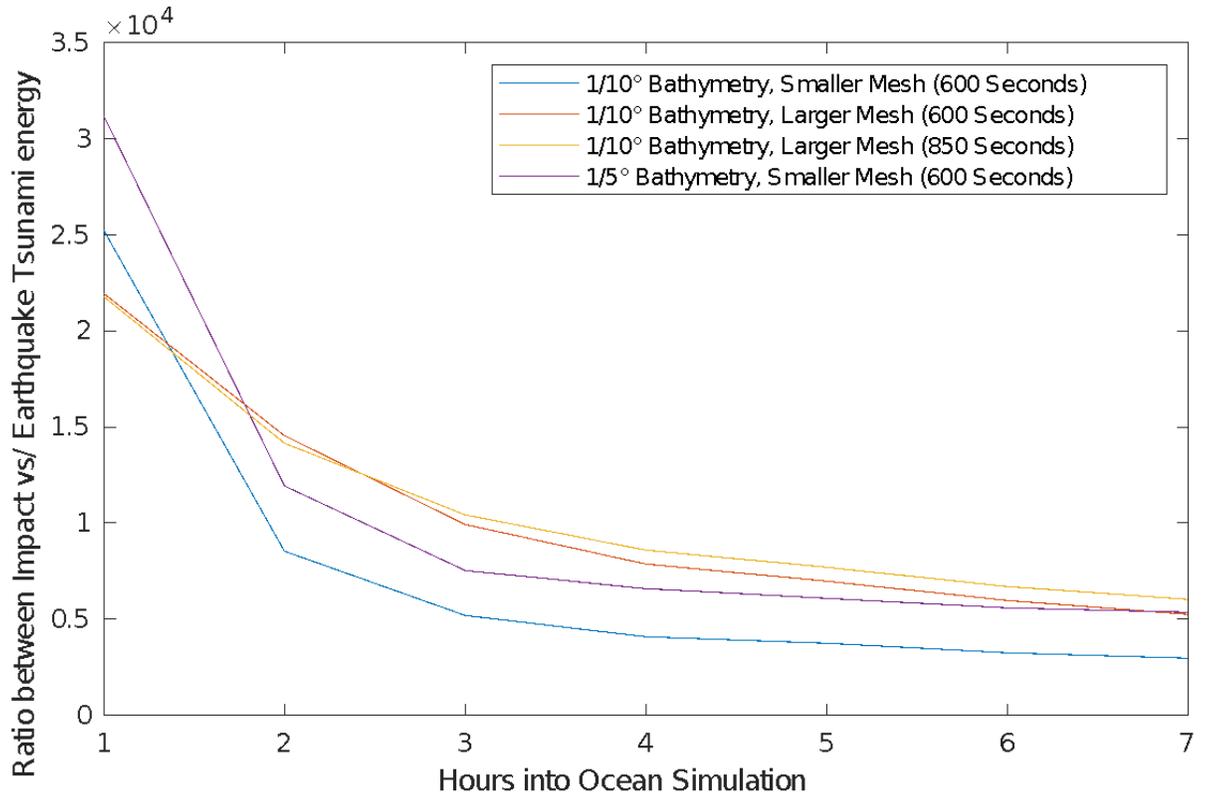
The sea surface height perturbation field four hours after handoff from the hydrocode in a MOM6 simulation with  $1/5^\circ$  grid spacing are quite similar to the sea surface height perturbation field four hours after the handoff in the fiducial  $1/10^\circ$  MOM6 run (Figure S3). This similarity indicates that the shallow water grid spacing of  $1/10^\circ$  is reasonably robust for our purposes.

In the ‘Half Crater’ simulation, the crater is only on grid cells that were originally water before impact. When the crater is placed over the pre-impact bathymetry, there are portions of the crater in water cells, and portions that are on the Yucatan Peninsula. When placing the crater into the model, two initial conditions were made. The ‘Full Crater’ replaces these land cells on the Yucatan Peninsula with water to a depth of the post-impact crater. The ‘Half Crater’ initial condition leaves the land cells as land. In the ‘Crater Only’ simulation, a full crater is made over the land and water cells, but the rim wave is removed, such that the tsunami is due only to water rushing inwards to fill the impact crater.

The ratios of the energies of different MOM6 simulations of the impact tsunami, relative to the 2004 Indian Ocean tsunami (as simulated in Smith et al., 2005) and plotted as a function of time into the shallow-water simulations, are given in Figure S4. The total energy of the Chicxulub tsunami and of three historical tsunamis is provided in Table S4, which also displays the energy of the tsunami source (e.g., Chicxulub asteroid impact in the case of the impact tsunami, and earthquake energies in the case of the 2004 Indian Ocean tsunami and 2011 Tohoku tsunami).



**Figure S3.** Sea surface height perturbations (m) four hours after handoff from the fiducial hydrocode simulation, in (top) fiducial 1/10° MOM6 simulation and (bottom) 1/5° MOM6 simulation.



**Figure S4.** Ratio of Chicxulub impact total (kinetic plus potential) energy to 2004 Indian Ocean earthquake-generated tsunami total energy, as a function of time into the respective shallow-water simulations. In the legend, ‘smaller mesh’ refers to our fiducial model while the ‘larger mesh’ refers to models that use our iSALE simulation that was run out to later times, had a larger horizontal extent of the high-resolution zone, and later arriving discontinuous ejecta. The initial energy of the fiducial model (blue curve) is  $5.1 \times 10^{19}$  J. For comparison the 2011 Tohoku Tsunami had an initial energy of  $3.0 \times 10^{15}$  J and  $1.5 \times 10^{15}$  J four hours later (Tang et al., 2012). The 2004 Indian Ocean Tsunami had an initial energy of  $1.7 \times 10^{16}$  J and  $9.1 \times 10^{15}$  J four hours later. Thus, the impact generated tsunami dissipates more quickly than earthquake generated tsunamis.

**Table S4.** Maximum energy values of the fiducial Chicxulub impact tsunami model runs compared to the modeled energy of several historical tsunamis

	<b>Tsunami Energy</b> $E_T$ (J)	<b>Source Energy</b> $E_S$ (J)	<b>Ratio</b> $E_T/E_S$ (%)
Full Crater, with Rim Wave (600 seconds)	$5.1 \times 10^{20}$ (max)	$2.7 \times 10^{23}$	0.19%
1883 Krakatau Tsunami (Maeno and Imamura, 2011)	$\sim 1.2 \times 10^{13}$ (max)	$\sim 1.0 \times 10^{17}$	0.01%
2004 Indian Ocean Tsunami (Tang et al., 2012)	$1.7 \times 10^{16}$ (max)	$6.2 \times 10^{18}$	0.27%
2011 Tohoku Tsunami (Tang et al., 2012)	$3.0 \times 10^{15}$ (max)	$2.8 \times 10^{18}$	0.11%

### ***Run times and computing hardware***

The hydrocode simulations employed here take of order one month on a high-end desktop. The 1/10<sup>o</sup> MOM6 simulations take of order two days to run on about 60 processors, on the University of Michigan Great Lakes supercomputer system (<https://arc.umich.edu/greatlakes/>).

### **Movie Captions**

The four movies that we made for this paper are posted with the Supporting Information and can also be found on our data repository <https://doi.org/10.7910/DVN/GWOFIO> where they are referred to as “SI\_Video1.mp4, SI\_Video2.mp4, SI\_Video3.mp4, SI\_Video4.mp4”.

**Movie S1.** Animation of the iSALE hydrocode fiducial simulation for the first 600 s post asteroid impact shown in Figure 1. Time series with material colored according to material type (crustal material is brown, sediments are yellow, and the ocean is blue). The origin marks the point of impact. Black curves mark material interfaces (e.g., sediment-crust interface).

**Movie S2.** Animation of the iSALE hydrocode model for first 1100 s post asteroid impact. This simulation is the same as the simulation shown in Figure 1 but has vertical extent of mesh limited to 50 km and the high-resolution zone extended to a radial distance of 300 km. Time series with material colored according to material type (crustal material is brown, sediments are yellow, and the ocean is blue). The origin marks the point of impact. Black curves mark material interfaces (e.g., sediment-crust interface).

**Movie S3.** Animation of the change in sea surface height tsunami propagation over 48 hours from the fiducial MOM6 model. Shown in m at ten-minute increments.

**Movie S4.** Animation of the change in sea surface height tsunami propagation over 48 hours from the fiducial MOST model. Shown in m at five-minute increments.

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